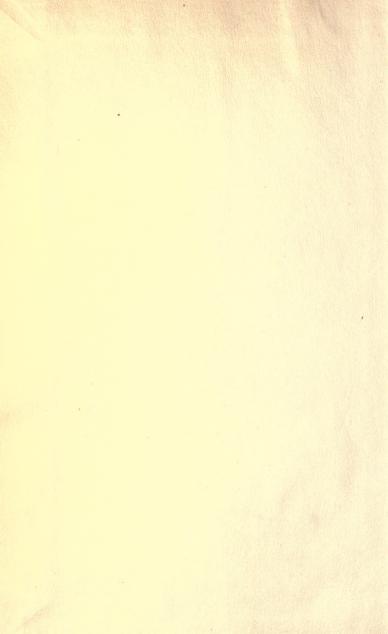


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THE ARTESIAN WATER SUPPLY OF EASTERN AND SOUTHERN FLORIDA.

BY E. H. SELLARDS AND HERMAN GUNTER.

From the Fifth Annual Report of the Florida State Geological Survey, pp. 103-290, 1913.



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1138	PAGE.
Introduction	113
The area treated	
Geology	114
Oligocene	114
Vicksburg group	114
Apalachicola group	117
Miocene	118
Pliocene	119
Pleistocene	119
Earth movements during the Pleistocene	120
Topography and Drainage	
Elevations	121
· Rivers	122
Climate	123
Temperature	
Precipitation	125
Soils	
General discussion of underground waters	
Source	
Annual rainfall	
Disposition of rainfall	
Amount available for the underground supply	
Underground circulation of water	133
Cause of movement	
Rate of movement	
Depth of underground water	
Hydrogen sulphide in underground water	
Sulphur water not evidence of beds of sulphur	
Sulphur deposits formed from hydrogen sulphide	
Absence of hydrogen sulphide from certain waters in Florid	
Amount of hydrogen sulphide influenced by preessure	
Artesian water	
Artesian water defined	
Conditions necessary to obtain artesian water	
Artesian basin	
Artesian slope	
Artesian water from solution passages	
Source of artesian water in Florida	144

	L'UGT
formations supplying artesian water	. 144
Depth of artesian water	. 145
Cost of wells	. 145
Increased flow with increased depth	
Increased head with increased depth	. 146
Increased temperature with increased depth	
Loss of head and reduction in flow	
Table showing loss of flow of artesian wells	
Cause of the loss of flow	
Waste of artesian water	
Method of measuring flow of artesian wells	
Tables for determining yield of artesian wells	
Area of artesian flow in Florida	
Discussion by counties.	
Nassau County	
Location and surface features.	
Water-bearing formations	
Area of artesian flow	
Local details	
Callahan	
Crandall	
Evergreen	
Fernandina	
Hilliard	
Italia	
King's Ferry	
Lessie	
Lofton	
Duval County	
Location and surface features	
Water-bearing formations	
Area of artesian flow	
Local details	
Baldwin	
Bayard	
Jacksonville	
Mandarin	
Manhattan Beach	
Maxville	
Mayport	
St. Johns County	
Location and surface features	
Water-bearing formations	
Area of artesian flow	. 185

CONTENTS.	107
	PAGE.
Local details	
Anastasia Island	
Armstrong	
Bunnell	
Dinner Island	
Elkton	
Espanola	
Federal Point	
Hastings	
Holy Branch	190
Hurds	191
Moultrie	191
Picolata	192
Riverdale	192
Roy	.193
St. Augustine	193
Switzerland	196
Yelvington	
Clay County	
Location and surface features	
Water-bearing formations	
Area of artesian flow	
Local details	
Doctors Inlet	
Green Cove Springs	
Hibernia	202
Leno	202
Magnolia Springs	203
Middleburg	203
Peoria	
Russell	205
Walkill	205
West Tocoi	205
Williams Crossing	205
Putnam County	206
Location and surface features	206
Water-bearing formations	206
Area of artesian flow	207
Local details	207
Bostwick	207
Crescent City	207
Orange Mills	
Palatka	209
Penial	211

Rice Creek	211
Rodman	211
San Mateo	212
Satsuma	212
Welaka	212
Woodburn	213
Orange County	
Location and surface features	214
Water-bearing formations	215
Area of artesian flow	215
Local details	
Chuluota	
Geneva	216
Orlando	217
Oviedo	217
Sanford	218
Volusia County	
Location and surface features	
Water-bearing formations	
Area of artesian flow	222
Local details	
Daytona	222
DeLand	
Enterprise	
Lake Helen	
New Smyrna	
Oak Hill	
Orange City	
Ormond	231
Pierson	
Seville	232
Brevard County	
Location and surface features	
Water-bearing formations	
Area of artesian flow	
Local details	
Chester Shoals	
City Point	
Cocoa	
Eau Gallie	
Frontenac	
Grant	
Malabar	
Melbourne	237

CONTENTS.	109
	PAGE.
Merritts Island	
Micco	
Rockledge	
Sharpes	
Tillman	
Titusville	
Valkaria	
St. Lucie County	
Location and surface features	
Water-bearing formations	
Area of artesian flow	
Local details	
Ft. Pierce	
Narrows	
Orchid	
Roseland	
Sebastian	
Pinellas County	
Location and surface features	
Water bearing formations	
Area of artesian flow	
Local details	250
Clearwater	250
Dunedin	251
Espiritu Santo Springs	251
Largo	
Ozona	
Pass-a-Grille	
Pinellas Park	
St. Petersburg	
Seminole Sutherland	
Tarpon Springs	
Wall Springs	
Hillsboro County	
Location and surface features.	
Water-bearing formations	258
Area of artesian flow	258
Local details	259
Plant City	259
Tampa	260
Polk County	
Location and surface features	202

	210141
Water-bearing formations	263
Local details	263
Bartow	
Carters	263
Lakeland	264
Mulberry	264
Osceola County	264
Location and surface features	264
Water-bearing formations	264
Area of artesian flow	266
Local details	266
Kissimmee	266
Narcoossee	266
Manatee County	267
Location and surface features	267
Water-bearing formations	268
Area of artesian flow	268
Local details	268
Bradentown	
Manatee	268
Palmetto	269
Sarasota	269
DeSoto County	269
Location and surface features	269
Water-bearing formations	270
Area of artesian flow	270
Local details	
Arcadia	
Ft. Odgen	
Nocatee	
Punta Gorda	
Palm Beach County	272
Location and surface features	272
Water-bearing formations	272
Area of artesian flow	
Local details	273
Gomez	273
Hobe Sound	273
Palm Beach	273
West Jupiter	277
Yamato	277
Lee County	278
Location and surface features	278
Water-bearing formations	278

CONTENTS.	111
	PAGE.
Area of artesian flow	. 279
Local details	. 279
Boca Grande	. 279
Ft. Myers	. 279
Labelle	. 280
Dade County	. 281
Location and surface features	. 281
Water-bearing formations	. 281
Area of artesian flow	. 281
Local details	. 281
Dania	. 281
Homestead	. 283
Miami	. 285
Monroe County	. 286
Location and surface features	. 286
Water-bearing formations	. 286
Area of artesian flow	. 287
Local details	. 287
Key Vaca	. 287
Key West	900

ILLUSTRATIONS.

FIGURES.

- Fig. 1. Artesian basin.
- Fig. 2. Artesian slope.
- Fig. 3. Artesian water from unconfined horizontal beds.
- Fig. 4. Artesian water from solution passages in limestone.
- Fig. 5. Method of measuring flow of artesian well.
- Fig. 6. Map showing area of artesian flow in Nassau and Duval Counties.
- Fig. 7. Map showing the area of artesian flow in St. Johns County.
- Fig. 8. Map showing the areas of artesian flow in Clay and Putnam Counties.
- Fig. 9. Map showing the area of artesian flow in Orange and Seminole Counties.
- Fig. 10. Flowing artesian well.
- Fig. 11. Map showing the area of artesian flow in Volusia County,
- Fig. 13. Map showing the area of artesian flow in Pinellas and Hillsboro Counties.
- Fig. 14. Map showing the area of artesian flow in Polk County.
- Fig. 15. Map showing the area of artesian flow in Osceola County.
- Fig. 16. Map showing the area of artesian flow in Manatee County.
- Fig. 17. Map showing the area of artesian flow in DeSoto County.

PLATES.

No.

- 10. Fig. 1. Palmetto flatwoods, Amelia Island.
 - Fig. 2. Palmetto flatwoods, Ft. Myers.
- 11. Fig. 1. Scrub, east side of Lake Kingsley, Clay County.
 - Fig. 2. Sandy pineland, DeLeon Springs.
 - Fig. 3. Open flatwoods, three miles east of DeLeon Springs.
- 12. Fig. 1. Everglades west of Ft. Lauderdale.
 - Fig. 2. Small prairie, four miles west of Sebastian.
 - Fig. 3. Turnbull Hammock, one mile west of Daytona.
- 13. Fig. 1. Sand dune near Mayport.
 - Fig. 2. Ancient sand dune, two miles west of Daytona.
- Fig. 3. Exposure at Saw Pit landing, St. Marys River.
- Fig. 1. Exposure of hardpan at Black Bluff on Clark's Creek, eight miles from Fernandina.
 - Fig. 2. Artesian well used for power, Melbourne, in Brevard County.

THE ARTESIAN WATER SUPPLY OF EASTERN AND SOUTHERN FLORIDA

E. H. SELLARDS AND HERMAN GUNTER.

INTRODUCTION.

A study of the water supply of Florida was begun in 1907 as co-operative work between the Florida State Geological Survey and the National Geological Survey. The first paper was issued in 1908 as Bulletin No. 1 of the Florida State Geological Survey, and relates to the underground water of Central Florida. The second paper of the series was published by the State Survey in 1910 and related to the water supply of the counties of Eastern Florida. A third paper included in the Fourth Annual Report, 1912, extended the study of the water supply to the counties of West Central and West Florida. The present paper includes a reprint of the paper on the water supply of Eastern Florida, published in 1910, revised to include a report on the water supply of Southern Florida.

The writers are indebted to the many well drillers and well owners who have contributed data regarding wells. Among the many who have given assistance the following should be especially mentioned: Messrs. Bellough & Melton, J. M. Chambers, C. I. Cragin, Dr. E. S. Crill, Capt. R. N. Ellis, Hughes Specialty Well Drilling Co., W. E. Holmes, John McAllister, Dr. J. N. Mac-Gonigle, McGuire & McDonald, W. J. Nesbitt, Hugh Partridge, H. Walker, Dr. DeWitt Webb, J. W. Wiggins, H. Van Dorn, W. D. Holcomb, G. A. Miller, and Mr. Holmes of the water supply department of the Florida East Coast Railway, J. C. Danielson, T. J. Zimmerman, F. S. Gilbert, W. F. Hamilton, Dibble and Earnest, The Artesian Well Co., D. W. Brown, F. J. White & Co., Sydner Pump and Well Co., E. J. Pettigrew, J. O. Edson, F. B. Bradley, and C. E. Reid.

Extensive well records made in 1907-1908 in cooperation with the U. S. Geological Survey by Messrs. Geo. C. Matson and F. G. Clapp have been utilized in the preparation of this report. Data regarding climate and rainfall have been supplied by Hon. A. J. Mitchell, Director of the Florida section of the U. S. Weather Bureau.

Many of the analyses included have been made in the office of the State Chemist especially for this report. Others have been made at various times by other chemists. Credit is given with each analysis.

THE AREA TREATED.

The area considered in detail in this report includes the following counties: Brevard, Clay, Dade, DeSoto, Duval, Hillsboro, Lee, Manatee, Monroe, Nassau, Orange, Osceola, Palm Beach, Pinellas, Polk, Putnam, St. Johns, St. Lucie, Seminole, Broward, and Volusia. This section borders the Atlantic and Gulf coasts and comprises the principal artesian areas of Peninsular Florida.

GEOLOGY.

A knowledge of the geologic structure is essential to a clear understanding of the underground water conditions. The prevailingly level country of Florida renders geologic observations difficult. Some favorable exposures occur, however, and these together with data obtained from well samples and well records permit a reasonably full understanding of the structure of the State.

The formations found in Florida belong to the: Oligocene, Miocene, Pliocene, and Pleistocene. Of these divisions the Oligocene is the oldest; the Pleistocene the most recent.

OLIGOCENE.

VICKSBURG GROUP.

The oldest or deepest formations reached in well drilling in Peninsular Florida are the Vicksburg limestones. The Vicksburg

is an extensive deposit underlying all of Florida and extending into adjacent States. In Central Peninsular Florida, from Columbia to Sumter Counties, these limestones are frequently exposed at the surface. Passing to the east and south from Central Florida they dip beneath the surface, and while nowhere exposed at the surface they are reached by all the deeper wells. It is in fact from these limestones that the principal water supply of Eastern and Southern Florida is obtained. The Vicksburg is very characteristic in appearance and structure, and when once seen is not likely to be mistaken for any other formation in this part of the State. The limestone as seen in well samples has a granular appearance and may contain many small shells. This phase of the limestone is frequently spoken of by the drillers as the "coral" formation. As a matter of fact, however, the formation contains relatively few corals. After passing one or two hundred feet into this formation a more compact limestone is encountered. This part of the formation often has a slightly pinkish cast, the rock being very hard, and the drilling difficult. While these are the general characteristics of the Vicksburg, yet its texture is not uniform. Hard layers usually alternate with soft layers, the water supply as a rule increasing as each hard layer is penetrated. Not infrequently masses of flint are found imbedded in the limestone which in some instances have given much difficulty in drilling.

While, as already stated, the Vicksburg limestones dip on passing to the east and south, yet the dip is not uniform and the depth at which it is encountered varies from place to place.

In the wells at Jacksonville the Vicksburg is reached at a depth of from 500 to 525 feet. At Callahan and at Fernandina, in Nassau County, although no samples have been obtained, the Vicksburg is believed, from well records, to be reached at about the same depth as at Jacksonville.

Along the St. Johns River the Vicksburg maintains a similar depth for some distance. At Ortega, seven miles south of Jacksonville, the limestone was reached at a depth of about 500 feet. At Magnolia Springs, and Green Cove Springs, thirty miles south of Jacksonville, and on Black Creek, while no well samples were

obtained, the Vicksburg is believed from well records to occur at a depth of from 325 to 400 feet.

Passing to the south the Vicksburg lies nearer the surface. Samples of drillings from wells at St. Augustine and at Hastings in St. Johns County and at Orange Mills in Putnam County show that the Vicksburg in this section lies at a depth of 130 to 225 feet, the greater depth being at St. Augustine and the minimum depth at Orange Mills. Passing to the south the Vicksburg lies, so far as well records indicate, at a fairly uniform level for a distance of 150 miles. At Sanford, 75 miles south of Orange Mills, the Vicksburg is reached at a depth of from 113 to 125 feet. At Daytona, although samples are lacking, the depth of this formation is believed, judging from well records, not to exceed 150 feet. At Cocoa the Vicksburg is reached at a depth not exceeding 190 feet, while at Melbourne Beach, 150 miles south of St. Augustine, its depth in one well was found to be 221 feet.

Passing to the south from this point the Vicksburg dips rapidly. At Palm Beach, 100 miles farther south, this limestone was reached at a depth of approximately 1,000 feet, *a dip of about 750 feet in 100 miles or 7½ feet per mile. The Vicksburg was not reached in a well 700 feet deep drilled by the Florida-East Coast Railway Company at Marathon Key, 175 miles south of Palm Beach.† At Key West, however, the formation is believed to have been reached at a depth of 700 feet.‡

It is thus seen that the Vicksburg forms a broad arch extending from central Florida to the Atlantic Ocean. St. Augustine lies near the north slope of this arch, while Melbourne, as nearly as can be determined, lies near the south slope. On either side of the arch the limestone dips at a moderate rate. On the north side of the arch the maximum depth recorded in Florida is 500 feet. Passing to the south the maximum of approximately 1,000 feet is recorded at Palm Beach.

In view of the importance of the Vicksburg as an artesian water reservoir, the depth at which it is to be expected is a matter

^{*}Darton, N. H.; Amer. Journ. Sci. (3) XLI, pp. 105-6, 1891.

[†]Florida Geol. Survey. Second Annual Report, p. 206, 1909. ‡Hovey, E. O.; Mus. Comp. Zool. Bull. XXVIII, pp. 65-91, 1896.

of very great importance and it is to be hoped that well drillers will find it possible to keep accurately labeled well samples in order to determine more definitely the distribution of this formation.

APALACHICOLA GROUP.

The Apalachicola group of formations is of a much less uniform character than the Vicksburg and is also of less importance in connection with the water supply. A full description of this group of formations will be found in the Second Annual Report of this Survey, pp. 67-106.

The formations which make up the Apalachicola group include the Chattahoochee and Alum Bluff formations, well exposed along the Apalachicola River; the Hawthorne formation in Central Florida; and the Tampa formation in Southern Florida. The relative position of three of these, the Chattahoochee, the Hawthorne and the Tampa formations, has not been definitely determined, and they may be largely contemporaneous. The Alum Bluff formation lies above the Chattahoochee formation. The limestone of this group consists largely of impure clayey material which upon decay weathers to a sticky blue green clay. The Chattahoochee Limestone is difficult to recognize in well samples. Fossils in this formation are comparatively rare and such as occur are largely destroyed in drilling. In surface exposures it may be recognized by its lithologic character and by the characteristic cubical blocks into which some of the strata break upon exposure.

The Apalachicola group has not been recognized from well drillings in East Florida. Clays taken by Mr. S. L. Hughes from the new city well at Jacksonville at the depth of 320 feet have a very close resemblance to the fuller's earth clays which occur in the Apalachicola group, above the Chattahoochee Limestone. On the other hand, Matson obtained from Jacksonville a Miocene shark's tooth from a well sample supposed to come from the depth of 496 feet. In order to determine more fully the area and extent of the Apalachicola group of formations in Eastern Florida it will be necessary to obtain large and carefully collected well samples. In Southern Florida the Apalachicola group is recog-

nized at Tampa and thence south along the Gulf coast as far as Sarasota Bay.

MIOCENE.

The Miocene deposits are well developed in Eastern Florida. At the city water works at Jacksonville this formation was encountered in excavating for the basin for the city water supply,* and was also reached in the city wells at a depth of from 35 to 36 feet. At Jacksonville this formation has a considerable, although undetermined thickness. It consists of a buff limestone grading to a lighter color, more or less phosphatic, grading below to phosphatic sands and sandy marls. The formation is in places fossiliferous, although the shells are usually preserved as casts

In Clay County the Jacksonville formation is extensively exposed along Black Creek. The exposure of this formation appears along both the South and North Fork of Black Creek, some miles above Middleburg, and may be observed for five or six miles below Middleburg. The following section was observed at High Bluff, on the South Fork of Black Creek, about five miles above Middleburg:

Covered and sloping	5	feet
Sloping, some sticky clay exposed	5	feet
Yellow sand	8	feet
Buff colored sandy limestone, containing a small propor-		
tion of black phosphatic pebbles	12	feet
Same, with greater amount of phosphate		
Same, with some phosphate		

This is the thickest exposure of the Jacksonville formation observed at any one place along Black Creek.

The following section was observed in the pit of the Jacksonville Brick Company, two miles southwest of Jacksonville:

Incoherent sand and soil	2.4 feet	
Sandy clays, the top 5 or 6 feet oxidizes yellow	16 feet	
Bluish fossiliferous marl	4 feet	

The marl obtained from test holes in the bottom of the pit is

^{*}Dall, W. H., U. S. Geol. Surv. Bull. 84, 124-125, 1892.

similar in character to the Choctawhatchee marl of West Florida, and the clay used for brick making in Duval, Nassau and Putnam Counties is probably of Miocene age. Beneath this marl, as shown by numerous well drillings, the sandy limestones of the Jacksonville formation occur.

Miocene deposits in Florida were first recognized by Dr. E. A. Smith,* at Rock Springs, in the northwestern part of Orange County. The limestone exposed here is a light, sandy, fossiliferous limestone and is probably of the Jacksonville formation.

PLIOCENE.

Pliocene is known to occur in Eastern Florida, although the extent and distribution of the deposits have been but imperfectly determined. The shell deposits of this period occurring in the St. Johns valley and along the East Coast have been described by Messrs. Matson and Clapp.† Localities mentioned by them are Nashua, on the St. Johns River, in Putnam County, and at DeLand and near Daytona, in Volusia County. Other localities at which these deposits were observed to be exposed are one-half mile above the Atlantic Coast Line bridge over the St. Johns River, in Putnam County; on the east side of the St. Johns River, about five miles north of the Atlantic Coast Line bridge, in Volusia County. Pliocene beds were also recognized from a well near Kissimmee. From the exposures thus recognized it is evident that Pliocene beds underlie a considerable area of Eastern Florida. In Southern Florida the Pliocene is well developed in the valley of the Caloosahatchee River. The land pebble phosphate deposits are also believed to be Pliocene.

PLEISTOCENE.

The marine Pleistocene deposits have been recognized at seval localities in Eastern and Southern Florida. Messrs. Matson

^{*}Smith, E. A., On the Geology of Florida. Amer. Journ. Sic. 3d Ser., Vol. XXI, pp. 302-303.

Pla. Geol. Surv. Sec. Ann. Rpt., pp. 128-133, 1909.

and Clapp obtained collections from Eau Gallie, Titusville and Mims in Brevard County. It is probable that marine Pleistocene shell deposits are somewhat widely distributed along the coast and perhaps in the St. Johns River valley. Here, again, satisfactory determination can be made only from large and carefully kept samples obtained in well drilling. The coquina rock which occurs extensively at St. Augustine and extends along the coast to the south for 250 miles is also to be placed with the Pleistocene. Some of the older sand dunes of the coast also probably belong to the Pleistocene. In southern Florida Pleistocene limestones are extensively developed in Palm Beach, Dade and Monroe Counties, bordering and underlying the Everglades and on the keys.

The following is an analysis of a sample of the Miami Limestone from near Miami, Florida. Analysis given by John B. Reilly. Name of analyst not recorded.

Silica	6.42
Alumina and iron oxides	0.94
Carbonate of lime	91.23
Carbonate of magnesium	1.08
	99.67

EARTH MOVEMENT DURING PLEISTOCENE.

Changes in the relation of land and water have occurred recently along the East Coast, probably during Pleistocene time. The best evidence of these changes is that offered by the sand dunes and the coquina rock bordering the East Coast. The line of sand dunes along the coast is well developed and largely continuous. From Daytona south these dunes occur, not on the present beach, but back from the beach a variable distance, depending upon the configuration of the country. At Daytona the sand dune lies back from the Halifax River about two miles. From Daytona to Titusville the dunes are to be seen lying mostly to the west of the East Coast Railway at a distance of one or two miles from the coast. At Titusville the dunes lie back from the Indian River two to two and one-half miles. At Rockledge the

dunes approach closer to the coast. They recede again, however, to the south and at no place directly face the ocean. The dunes are now quiescent and are covered with a thick growth of trees, indicating that they have been undisturbed for a long time. In the same way the coquina rock, found facing the ocean at Anastasia Island, in St. Johns County, falls back from the coast to the south, extending at places a few miles inland. The presence of this ledge of coquina rock bordering the coast together with the sand dunes lying back clearly indicates that the land level formerly stood lower than at present, the coquina rock and sand dunes having accumulated along what was then the beach.

Conrad as early as 1846 noted the occurrence of marine shells of post-Pliocene age along the bank of the St. Johns River at an elevation of from ten to fifteen feet above the present high tide.

Matson has described* what he believes to be a Pleistocene terrace bordering the St. Mary's River, in Nassau County. A similar abrupt rise in passing onto the upland may be observed in many places bordering the coast and the valley of the St. Johns River. It may be observed that a subsidence of 25 feet would submerge the entire St. Johns valley and would allow the sand dunes once more to face the ocean.

TOPOGRAPHY AND DRAINAGE.

The section of the State to which this report relates borders the Atlantic Ocean and the Gulf of Mexico. From sea level the rise in elevation is as a rule gradual and the country in general level or rolling. It is probable that with the exception of sand dunes all of Monroe, Lee, Dade, Palm Beach, St. Lucie and Brevard Counties as well as the eastern one-half or more of Nassau, Duval, Clay, Putnam, Volusia and Orange Counties and the entire St. Johns River Valley lie below the 50-foot contour line. Elevations exceeding 50 feet occur in the western part of Nassau, Duval, Clay, Putnam and Orange Counties and as a ridge extending from northwest to southeast through Volusia County.

^{*}Florida Geol, Survey, Second Annual Report, p. 39, 1909.

The maximum elevation for Eastern Florida is found in the northwestern part of Clay County, approaching "Trail Ridge." On this ridge are found, according to levels made in 1911 by the United States Geological Survey, a maximum elevation of 246 feet. In Polk County elevations approximating 250 feet are also reported. (See map.)

RIVERS.

The St. Johns River rises from the lakes of southern Brevard County, within a few miles of the Atlantic coast. From this point it flows north or slightly west of north about 200 miles, entering the Atlantic Ocean within 25 miles of the north line of the State. The elevations along this river at no point exceed 25 feet above sea, the entire valley lying within the artesian flow area of the State. The principal tributaries of the St. Johns are Black Creek and Ocklawaha River. The former heads in the uplands of Clay County, while the latter is fed from numerous lakes of Lake County and receives tributaries from Silver Springs in Marion County and from the lakes of southeastern Alachua County.

The St. Mary's River, forming a part of the northern boundary of the State, rises in or near Okefenokee Swamp, in Georgia. From its origin it flows south until on a parallel with the mouth of the St. Johns river. From this point it bends abruptly and flows north for thirty miles, then, turning again, flows a little south of east to the Atlantic Ocean. Nassau is one of the smaller rivers and with its tributary, Thomas Creek, forms part of the boundary between Nassau and Duval Counties. The Withlacochee, Hillsboro, Peace and Caloosahatchee rivers flow into the Gulf.

Bordering the streams, both the main rivers and their tributaries, are found in many places, open, flat, imperfectly drained pine lands. These lands are classed in the section treating of soils as open flatwoods. A somewhat different and more extensive type of country is that designated as palmetto flatwoods. An essential difference in these two types of country is the presence or absence of the saw palmetto, the pine forest being common to both. In Nassau and Duval Counties and along the tributaries

of the St. Johns River extensive areas of open flatwoods occur.

Along the border of the uplands, back from the river and from the coast, a different type of topography has developed, consisting largely of the sandy or rolling pine type of soil although scrub hammock lands occur. These several types of country are due to a considerable extent to the drainage conditions. On the summit of the plateau, in the interior of Florida, palmetto flatwoods and to some extent open flatwoods are again encountered.

CLIMATE.

The counties of Florida, covered by this report, lie bordering or near the Atlantic Ocean and the Gulf, and are favorably located for a mild and equable climate. The heat of summer, as elsewhere in Florida, is tempered by the proximity to the ocean. By varying the crops, the growing season can be made to extend practically throughout the year.

TEMPERATURE.

As the total length of the section covered by this report extends north and south fully 425 miles, the temperature varies appreciably between northern and southern points. At Jacksonville, in Duval County, within about 25 miles of the north line of the State, the mean annual temperature is 69 degrees Fahrenheit. The means for the four seasons of the year are as follows: Winter, 56; Spring, 69; Summer, 81; Fall, 70. The absolute maximum for summer heat recorded at Jacksonville is 104, although temperatures above 100 are rare. The lowest temperature recorded is 10 above zero. The mean temperatures for the several months of the year at Jacksonville are as follows: January, 55; February, 58; March, 63; April, 68; May, 75; June, 80; July, 82; August, 82; September, 78; October, 71; November, 62; December, 56.*

At New Smyrna, in Volusia County, a station about 100 miles

^{*}United States Weather Bureau Bull. Q, Climatology of the Eastern United States, by Alfred Judson Henry, p. 352, 1906.

south of Jacksonville, as shown by the same report, the annual mean temperature is 70 degrees F. The means for the four seasons are: Winter, 58; Spring, 68; Summer, 79; Fall, 72. The absolute maximum for summer heat recorded at New Smyrna is 100 degrees F. The lowest temperature recorded is 16 above zero. The mean temperatures for the several months of the year (Fahrenheit) are as follows: January, 57; February, 59; March, 65; April, 67; May, 73; June, 78; July, 80; August, 80; September, 78; October, 73; November, 66; December, 58.

At Tampa, in Hillsboro County, the annual mean temperature is 72 degrees F. The means for the four seasons are: Winter, 61; Spring, 71; Summer, 81; Fall, 73. The absolute maximum for summer recorded at Tampa is 96 degrees F. The lowest temperature recorded is 19 above zero. The mean temperatures for the several months of the year (Fahrenheit) are as follows: January, 59; February, 62; March, 67; April, 70; May, 76; June, 80; July, 81; August, 82; September, 80; October, 74; November, 67; December, 61.

At Miami, in Dade County, the annual mean temperature is 75 degrees F. The means for the four seasons are: Winter, 67; Spring, 73; Summer, 82; Fall, 78. The absolute maximum for summer recorded at Miami is 96 degrees F. The lowest temperature recorded is 29 above zero. The mean temperatures for the several months of the year (Fahrenheit) are as follows: January, 65; February, 67; March, 71; April, 74; May, 76; June, 81; July, 82; August, 82; September, 81; October, 78; November, 74; December, 69.

At Key West, in Monroe County, the annual mean temperature is 77 degrees F. The means for the four seasons are: Winter, 70; Spring, 76; Summer, 83; Fall, 79. The absolute maximum for summer recorded at Key West is 100 degrees F. The lowest temperature recorded is 41 above zero. The mean temperatures for the several months of the year (Fahrenheit) are as follows: January, 70; February, 71; March, 73; April, 76; May, 79; June, 82; July, 84; August, 84; September, 85; October, 79; November, 74; December, 70.

At Jacksonville, in the northern part of the State, there is

little or no danger of frost before the latter part of October. Light frosts, however, may occur as early as the latter part of October. The earliest killing frost recorded, at this station, is November 2, while the average date of the first killing frost for the past fifty-three years is December 4. The latest date of a killing frost in the spring, at Jacksonville, is April 6, and the average date of the last killing frost is February 14. Light frosts, however, have been known to occur as late as April 28.

At New Smyrna the earliest date of a killing frost in the fall is November 28, while the average date of the first killing frost for the past sixteen years is December 23. The latest date of a killing frost at this place in the spring is March 22. The average date of the last killing frost is February 16.*

At Tampa the earliest date of killing frost recorded is November 28, while the average date of the first killing frost is January 9. The latest date of killing frost in the spring recorded at Tampa is March 19. The average date of the last killing frost is February 8.

At Miami the earliest recorded date of the killing frost in autumn is December 26, and the latest date in the spring is February 19. The killing frost at this locality is so infrequent that no attempt is made to determine the average date.

At Key West, at the extreme southern end of Florida, frosts do not occur.†

PRECIPITATION.

The season of heavy rainfall in Eastern Florida includes the summer and early fall months. As a rule approximately one-half of the rainfall of the year comes during the four months, June, July, August and September.

^{*}U. S. Dept. Agri. Summary of the Climatological Data for the United States by sections: Section 83.—Northern Florida, A. J. Mitchell, Section Director. Also Climatology of Jacksonville, Fla., and Vicinity, Monthly Weather Review for December, 1907, by T. Frederick Davis.

[†]United States Weather Bureau, Summary of the Climatological Data for the United States by Sections: Section 84.—Southern Florida, A. J. Mitchell

The average rainfall at Jacksonville for the 32 years ending with 1903 was 53.4 inches annually. The mean for the four seasons of the year is as follows: Winter, 9.4 inches; Spring, 10.4 inches; Summer, 17.9 inches; Fall, 15.7 inches. The mean for the several months of the year at Jacksonville is as follows: January, 3 inches; February, 3.4 inches; March, 3.5 inches; April, 2.9 inches; May, 4 inches; June, 5.5 inches; July, 6.2 inches; August, 6.2 inches; September, 8.1 inches; October, 5.1 inches; November, 2.5 inches; December, 3 inches.

At New Smyrna the annual rainfall as shown by the same report is 51.1 inches. The mean for the four seasons is as follows: Winter, 8.4 inches; Spring, 6.8 inches; Summer, 17.4 inches; Fall, 18.5 inches. The mean precipitation for the several months of the year at this station is as follows: January, 2.8 inches; February, 3.6 inches; March, 2.6 inches; April, 1.6 inches; May, 2.6 inches; June, 6.2 inches; July, 5.6 inches; August, 5.6 inches; September, 9.2 inches; October, 6.7 inches; November, 2.6 inches; December, 2 inches.*

At Tampa the annual rainfall is 53.1 inches. The mean for the four seasons is as follows: Winter, 8.1 inches; Spring, 7.4 inches; Summer, 24.9 inches; Fall, 12.7 inches. The mean precipitation for the several months of the year at Tampa is as follows: January, 2.8 inches; February, 3.5 inches; March, 2.9 inches; April, 2.1 inches; May, 2.4 inches; June, 8.5 inches; July, 8.0 inches; August, 8.4 inches; September, 8.2 inches; October, 2.8 inches; November, 1.7 inches; December, 1.8 inches.

At Miami the annual rainfall is 58.3 inches. The mean for the four seasons is as follows: Winter, 8.1 inches; Spring, 11.1 inches; Summer, 20.6 inches; Fall, 18.5 inches. The mean precipitation for the several months of the year at Tampa is as follows: January, 4.0 inches; February, 2.5 inches; March, 3.1 inches; April, 3.5 inches; May, 4.5 inches; June, 8.2 inches; July, 7.0 inches; August, 5.4 inches; September, 9.1 inches; October, 7.1 inches; November, 2.3 inches; December, 1.6 inches.

At Key West the annual rainfall is 37.9 inches. The mean

^{*}United States Weather Bureau, Bull. Q.

for the four seasons is as follows: Winter, 5.3 inches; Spring, 5.5 inches; Summer, 12.6 inches; Fall, 14.5 inches. The mean precipitation for the several months of the year at Key West is as follows: January, 2.0 inches; February, 1.6 inches; March, 1.2 inches; April, 1.2 inches; May, 3.1 inches; June, 4.2 inches; July, 3.7 inches; August, 4.7 inches; September, 7.0 inches; October, 5.4 inches; November, 2.1 inches; December, 1.7 inches.

SOILS

The geologic, topographic, climatic and drainage conditions have much to do with the character of soils. Since the inorganic constituents of soils are derived primarily from the decay of pre-existent formations, the character of the soil is determined to a considerable extent by the formation from which it is derived. The thickness and manner of accumulation of the residual material as well as accumulation of the organic constituents is affected by the topographic, climatic and drainage conditions. The following are the more prominent soil types in the part of Florida covered by this report:

Rolling pine lands: This type includes light, sandy, well-drained soils. The native vegetation is pine and wire grass. Oaks and other hard wood trees occasionally occur. The saw palmetto is for the most part absent. This type of soil predominates in the lake region of Florida.

Palmetto flatwoods: The palmetto flatwoods occur over an extensive area in Florida. This type of country is flatter than the sandy pine land and not so well drained. The native vegetation of these lands consists chiefly of pine, saw palmetto and wire grass. The sand is dark at the surface, becoming lighter below. As a rule the so-called "hardpan" underlies the palmetto flatwoods. This "hardpan" consists of sand stained with organic matter and has the appearance of being partly cemented with iron. When dry it is fairly well indurated, but as a rule it may be penetrated with the soil auger. The transition in the bore hole from the light colored sand to "hardpan" is abrupt. The "hardpan" itself is very dark colored at the top and grades into chocolate colored sands below.

The "hardpan" is very objectionable in farming lands as it prevents free movement of water by capillary attraction. The lands underlaid by "hardpan" are not resistant to droughts. However, where an abundance of water can be obtained cheaply, as in the section of flowing artesian water, such lands may be used to advantage by keeping them saturated with water.

Open flatwoods: The open flatwoods are much less extensive than the palmetto flatwoods. The native vegetation of the land of this type is chiefly pine and wire grass with little or no underbrush. The saw palmetto is absent or nearly so and there is little or no "hardpan." The soil to a depth of from one to three feet is dark ashy gray owing to the presence of organic matter mixed with the sand. A clay sub-soil is usually found at the depth of from one to four feet. This type of land when drained and irrigated has been used with great success in growing Irish potatoes, sweet potatoes and other trucking crops and in general farming.

Prairie lands: The word "prairie" is applied to open lands devoid of trees. The native growth is largely grasses.

Muck lands: The term "muck soils" is applied in ordinary usage to lands on which organic matter from decay of vegetation has accumulated to some depth. Vegetable matter accumulates in this way only on such lands as are overflowed during a considerable part or all of the year. The largest tract of muck lands in the State is the Everglades. Many smaller tracts occur, however, throughout the State.

Clay lands: The clay soils are usually of limited extent, occurring at places where the superficial sands have been removed by surface wash. The clay soils are lacking in organic matter and before being farmed must be broken up and organic matter incorporated.

Hammock lands: The term "hammock land" is most frequently applied to lands underlaid by marl or limestone and supporting a thick growth of vegetation, including hardwood trees and cabbage palmetto. These lands when cleared make excellent farming lands. Other hammock lands occur, however, which have no evident relation to marl deposits. These likewise support

a heavy growth of hardwood trees. The soil consists of a rich humus due to the accumulation of leaves. Beneath the humus is usually found several feet of orange yellow sand.

Sandy hammock lands: The sandy hammock lands as developed in the sections bordering the coasts are of wind-blown sands or low dunes on which vegetation has gained a foothold. Various hardwood trees grow on this type of land. It has been found in many instances desirable for orange culture. It is used also to some extent in vegetable growing. The open nature of the soil, however, results in a heavy loss of fertilizer from leaching.

Scrub: Scrub is a term applied to very sandy lands which support a dense growth of shrubby plants. The sandy pine lands often pass very abruptly and with no apparent reason into scrub. Few attempts have been made to utilize the scrub lands for farming purposes.

Sand dunes: The sand dunes both of recent and of earlier formation occur frequently in Florida particularly along the coast. The sand dune soil has been found especially adapted to the growing of pineapples, the extensive pineapple farms of St. Lucie County being largely located on quiescent dunes.

River swamp: The river swamp lands support a dense growth of hardwood trees. On the smaller streams where the elevation is sufficient to permit of successful drainage these lands if cleared would furnish desirable trucking and farming land.

Salt marsh: Extensive salt marshes occur along the Atlantic coast and bordering the streams entering the ocean.

UNDERGROUND WATER: GENERAL DISCUSSION.

SOURCE.

Rainfall: The chief source of underground water is the rainfall. Water vaporized through the energy of the sun passes into the atmosphere and is precipitated over the land as rain or condensed as dew or fog. The vapor is supplied to the atmosphere by evaporation, principally from the ocean, which, occupying three-fourths of the earth's surface, is continuously exposed to the sun's rays. To the vapor from the ocean is added that arising from inland waters, from the dry land surface to the earth, and from the leaves of plants.

Small additions to the underground water supply may come through any one of a number of other possible sources, but the total amount thus added is relatively small and may be omitted in a general discussion.*

ANNUAL RAINFALL.

The annual rainfall is the measure of the column of water that would accumulate at any spot in the course of a year, if all that falls should be preserved. The measurement is commonly stated in inches. The average rainfall for the State as a whole for the fifteen years, from 1892 to 1906, inclusive, as deduced from the U. S. Weather Reports, was 53.17 inches, annually. The year 1907 was a year of less than average rainfall, 49.15 inches, and if this year is included the average for the sixteen years, 1892 to 1907, falls below 53 inches, being 52.92 inches. If longer periods be considered the variation from this average is not sufficient to materially change the result. The area covered by this report lies in that part of the State supplied with about the average rainfall, and 53 inches may be safely assumed as a close approximation to the annual rainfall for this section.

DISPOSITION OF RAINFALL.

Of the total rainfall of any area, (1) a part is returned as vapor to the atmosphere without having entered the earth; (2) a part is carried off by streams and rivers to the ocean without penetrating the earth; (3) a part is absorbed into the earth.

(1) WATER EVAPORATED WITHOUT ENTERING THE EARTH,
Immediately following a rain the atmosphere is nearly or quite

^{*}A recent discussion of possible sources of underground water other than rainfall will be found in Bulletin 319, U. S. Geol. Surv., by M. L. Fuller.

saturated. The evaporation at this time is slow, and the part returned to the atmosphere directly from the land is an almost negligible amount. This is especially true of a soil into which the water enters quickly. Some of the water clinging to the leaves of plants is re-evaporated, as well as a part of that which falls into lakes, ponds and temporary pools. While an estimate of the amount evaporated must be regarded as only in the roughest way approximate, yet it is probably safe to assume that not more than two or three per cent. of the total rainfall is returned to the atmosphere by direct evaporation without having entered the earth.

(2) SURFACE RUN-OFF.

The relative proportion between the surface run-off and the surface in-take of water is dependent upon the character of the surface and the deeper formations and upon the topography. The former affects rapidity of in-take of water into the earth; the latter the rapidity of surface run-off.

With regard to topography Peninsular Florida is either flat or rolling. Rarely can a locality within this section be described as hilly. The elevation increases gradually from sea level at the coast to a maximum of scarcely more than 200 feet inland, while large sections are so flat as to present no perceptible slope. Topographically the conditions are, therefore, very unfavorable to surface run-off. On the other hand, the conditions are exceptionally favorable to large surface in-take. The soils are sandy and receive and store the rainfall with great readiness.

(3) RAINFALL ENTERING THE EARTH.

Of the water which enters the earth, a part is ultimately returned to the atmosphere by evaporation. The water retained in soils is slowly given up through evaporation during dry weather. As the evaporation takes place near the surface, the capillary attraction draws a new supply from beneath, thus maintaining to some extent the moisture content of the soil. The amount of water thus brought to the surface and evaporated, while varying with climate and with soils, is, in the course of a year, considerable.

To the evaporation from the surface of the soil must be added that from the leaves of plants. This in turn varies greatly with the different plants and with different climatic conditions. King, in 1892, in one experiment, found that a crop of peas evaporated 477 pounds of water for each pound of dry matter formed, while corn under the same conditions evaporated in one instance 238 pounds of water per pound of dry matter.* Assuming that a citrus tree evaporates approximately as much as the European oak (Ouercus cerris), the water evaporated from the leaves of a fifteen-year-old orange tree is estimated, by Hilgard, at 20,000 pounds a year, or about 1,000 tons of water per acre of 100 trees.† This is equivalent to about nine inches annual rainfall over the same area. Water is the chief vehicle for conveying plant food absorbed from the soil by the roots. This enormous evaporation from the leaves is in part for the purpose of disposing of the water thus taken up by the plant. It serves chiefly, however, the purpose of preventing, through the conversion of water into vapor, an injurious rise of temperature during the hot sunshine and dry weather

It is impossible to estimate within even approximate limits the loss of water by evaporation from the surface of the ground, and from the leaves of plants in the area under consideration. The atmosphere in Florida is relatively humid. On the other hand, the temperature throughout most of the year is high. Much of the country is uncultivated, and practically all of the soil is of medium coarse texture.

It is probable that almost one-half of the rainfall entering the earth is re-evaporated from the surface of the ground and from the leaves of plants, and that not more than one-half of the total rainfall in Florida passes through the soil and surface material to join the underground water supply.

^{*20}th Ann. Report Wis, Agriculture Experiment Station, p. 320,1904. †Based on weighings made by R. H. Loughridge of the leaves of a citrus tree at Riverside, Calif. Soils, by E. W. Hilgard, p. 263, 1906.

AMOUNT OF WATER AVAILABLE FOR THE UNDERGROUND SUPPLY.

An annual rainfall of 53 inches is found by computation to amount to 921,073,379 gallons per square mile. Of this amount it is estimated that in Central Florida about one-half is added each year to the underground water supply.

UNDERGROUND CIRCULATION OF WATER.

Underground water is found usually to be in motion, threading its way through pores, breaks, crevices, joints and other openings in the rocks. Its movement is ordinarily slow and varies with different rocks and under different conditions.

CAUSE OF MOVEMENT.

The chief cause of movement of underground, as of surface water, is gravity. Capillarity is an additional force which, under special conditions, may become the controlling factor. The water returned to and evaporated from the surface of the ground, as well as that carried to and evaporated from the leaves of plants, is moved by capillarity in opposition to gravity. Gravity, however, is the controlling force in the movement of water through the deep zones of the earth. Pressure, which is an important secondary cause of the movement in the earth, is the expression of gravity. Except in the case of capillarity, the movement of water, apparently in opposition to gravity, is, upon closer observation, found to be in reality movement in response to gravity. The water, which rises in a boring or flows from an artesian well or spring, is forced up by pressure, due principally to the weight of water lying at a higher level. The familiar observation that water seeks its own level has the same explanation.

RATE OF MOVEMENT.

The chief factors affecting the rate of movement of water through a porous medium, as given by Slichter, are as follows:*

^{*}Water Supply Paper, U. S. Geol. Surv., No. 67, p. 17, 18, 1902.

- (1.) Porosity of the material.
- (2.) Size of the pores in the water-bearing medium.
- (3.) Pressure.
- (4.) Temperature of the water.
- (1.) Rocks contain pores which, in the absence of a liquid, are ordinarily filled with air. The relative proportion of these spaces in the rock to the whole volume is the measure of the porosity. Thus, if a cubic foot of sandstone will hold in its pores one-fourth cubic foot of water, its porosity is 25 per cent. The greater the porosity, the more water absorbed by the rocks.
- (2.) The size of the pores in the rock affects the rate of flow. Rocks having large pores receive and conduct water many times more rapidly than those having small pores.
- (3.) The greater the pressure, other conditions remaining the same, the more rapid the flow. A pressure of one pound per square inch is required to support each 2.31 feet of a column of distilled water at the temperature of 60 degrees F. The weight of water from the deep zones is increased by solids in solution and in suspension, and is affected by changes in temperature. Something more than a hundred pounds pressure to the square inch is required to cause a flow from the bottom of a well 231 feet deep. Something more than 500 pounds pressure to the square inch is required to cause the rise of water in a boring, a distance of 1,150 feet. Pressure of this magnitude must materially assist in forcing water through the rock.
- (4.) The temperature of the water is found to influence the rate of flow. Slichter finds that a change from 50 to 60 degrees F. increases the capacity to transmit water, under identical conditions, by about 16 per cent.†

DEPTH OF UNDERGROUND WATER.

The limit of the downward extent of water has not been reached by borings or tunnels, some of which exceed a mile in

[†]Water Supply and Irrigation Paper, U. S. Geol. Surv. No. 140, p. 13, 1905.



Fig. 1.—Palmetto flatwoods. View taken on Amelia Island in Nassau County.



Fig. 2.—Palmetto flatwoods. View taken five miles east of Ft. Myers, Lee County.

EXPLANATION OF PLATE 11.

Fig. 1.—Scrub. This type of soil consists of white sand and is not adapted for farming. Photograph by R. M. Harper, View taken on east side of Lake Kingsley, Clay County.

Fig. 2.—Well drained pine lands. This type of soil is well drained, and consists of a sandy loam. The prevailing vegetation is pine, wire grass and oaks. The soil is light, and is suitable for early vegetables, and for orange growing. As a farming soil it requires building up and fertilizing. View taken near DeLeon Springs, in Volusia County.

Fig. 3.—Open flatwoods. The soil consists of a dark sandy loam underlaid at the depth of one to five feet by clay subsoil. The prevailing vegetation is pine and wire grass. These flatwoods are naturally poorly drained. When properly drained, however, the soil is good and suitable for trucking and general farming. View taken three miles east of DeLeon Springs.





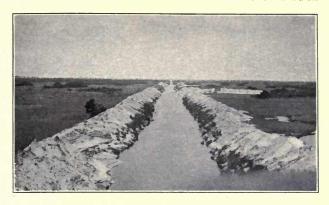


EXPLANATION OF PLATE 12.

Fig. 1.—Muck soil. The Everglades of Florida along the drainage canal, west of Fort Lauderdale. The soil here consists of muck to a depth of three to five feet, underlaid by sands which, in turn, rest upon oolitic limestone.

Fig. 2.—Prairie soil. One of the typical small prairies. View taken 10 miles west of Sebastian. The soil consists of light colored sands to a depth of several feet, underlaid by clay or hardpan. The small prairie shown in the foreground is surrounded by palmetto flatwoods.

Fig. 3.—Calcareous hammock soil. A view in Turnbull Hammock, one mile west of Daytona. Shell marl here lies at or very near the surface. The native vegetation includes cabbage palmetto and various deciduous hardwood trees. The calcareous soils are desirable, particularly for vegetable growing.







EXPLANATION OF PLATE 13.

Fig. 1.—Sand dune. This view illustrates one of the recent sand dunes near Mayport, at the mouth of the St. Johns River.

Fig. 2.—Ancient sand dune. This view is taken at the crossing of the public road across the dunes, about two miles west of Daytona. The dune here consists of light colored sand to a depth of four or five feet, underlaid by ochre yellow sands.

Fig. 3.—Clay soil. Exposure at Saw Pit Landing on the St. Marys River, in Nassau County. The soil here is a sticky clay soil residual from the decay of the clayey limestone.







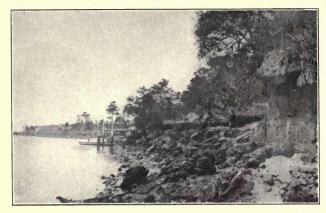


Fig. 1.—Exposure of hardpan along Black Bluff on Clarks Creek, eight miles from Fernandina.



Fig. 2.—Artesian well used for power belonging to H. T. Bowden, Melbourne, Brevard County. The water from the artesian well affords power by which water is pumped from a nearby shallow well.

depth. Water, while thus known to penetrate to a depth greater than a mile, probably does not reach beyond five or six miles at the most. The movement, as has been stated, is through natural openings in the rock. Pressure increases in the earth with depth, and it is estimated that, at a depth of approximately six miles, the pressure is so great that the pores and cavities of even the strongest rocks are completely closed,‡ making it impossible for water to penetrate beyond this depth. Most of the water, however, returns to the surface after a comparatively short underground course, only a small part of it reaching to this great depth.

HYDROGEN SULPHIDE IN UNDERGROUND WATER.

The underground water of Florida is very generally impregnated with hydrogen sulphide (H2S), also known as sulphuretted hydrogen, and hydro-sulphuric acid. Water containing hydrogen sulphide is commonly known as "sulphur water." Sulphur water is especially characteristic of the areas of artesian flow. In those sections in which open, porous limestone is the surface formation, hydrogen sulphide is usually absent from the first water encountered, although, even here, it is found to exist in the water from the deep wells and in some springs.

Source:—Hydrogen sulphide may originate, in nature, in any one of several ways. The following have been suggested: (1) The decay of organic matter containing sulphur; (2) the reaction of organic matter upon sulphides or sulphates; (3) the reaction of acids upon sulphides; (4) partial oxidization of sulphides; (5) steam passing over sulphur.

The decay of organic matter is an obvious source of hydrogen sulphide in the underground waters of Florida. Chemical analysis shows that sulphur is very generally present in Florida soils,* and apparently invariably present in muck soils. Analyses of samples of peat, which is, like muck, a vegetable accumulation, will be found in the paper on peat deposits published in 1910. The amount of sulphur in the Florida peat, in the dried samples, varies from less than 1 per cent. to over 4 per cent.

[‡]L. M. Hoskins, 16th Ann. Rept. U. S. Geol. Surv., Part I, p. 859, 1896.

Hydrogen sulphide is formed in connection with the decay of eggs. In this case the albumen of the egg, according to Ostwald, contains the sulphur.† H2S is also found escaping from sewer drains and cesspools, and is formed, during the decomposition, both of animal and vegetable substances. The H2S occurring in shallow springs from marsh lands is, doubtless, supplied largely from organic material.

The sulphur in soils is, probably, often present as sulphates. Thorpe states that the decay of organic matter in contact with sulphates results in the formation of H₂S.‡ The reaction in this case, probably, results from reducing properties of decaying organic matter, the sulphates being first reduced to sulphides, according to the following reaction: Na₂ S O₄+C₂ (carbon of organic matter) =₂CO₂+Na₂S. The sulphide is then acted upon by the carbonic acid to form H₂S as follows: Na₂S+H₂CO₃= H₂S+Na₂CO₃. The reaction of organic matter upon the sulphides is regarded, by Van Hise, as another important source of H₂S in underground water.*

The formation of hydrogen sulphide, as a result of the action of acids upon metallic sulphides, is one of the most familiar of laboratory experiments. This suggests the possibility of the formation of this gas, as the result of the action of acids upon metallic sulphides, contained in the rocks. Sulphides, especially those of iron, are widely scattered in the earth's crust, and occur in sufficient quantity to account for the formation of H2S gas in water. Hydrogen sulphide is a weak acid, and its salts are decomposed by a stronger acid. Sulphuric and other mineral acids should certainly react upon sulphides liberating H2S. Carbonic acid, when abundant, reacts upon alkali sulphides to produce hydrogen sulphide. It is true that the alkali sulphides are normally not abundant in the crust of the earth. Stokes has shown, however, that the reaction of sodium carbonate within the

^{*}Bulletin 43, Florida State Experiment Station, pp. 653, 657, 659, 1897. †Ostwald, Principles of Inorganic Chemistry, page 274, 1904.

Dictionary of Chemistry, Vol. III, p. 697, 1900.

^{*}A Treatise on Metamorphism, Mon. XLVII U. S. Geol. Surv., p. 1112, 1904.

earth, upon pyrite or marcasite, produces sodium sulphide. The reaction given by him is as follows: (L. C. page 1107.)

8FeS2+15Na2CO3=4Fe2O3+14Na2S+Na2S2O3+15CO2.

It is a well-known fact that the carbon dioxide, which unites with water to form carbonic acid, is abundant in the deep waters, especially in the limestone formations; the pressure existing at considerable depth enabling the water to hold great quantities of carbonic acid. The series of reactions given by Stokes accounts for the presence of alkali sulphides in solution in the deep waters. It may be added that all sulphides are soluble, to some extent, in water and, in that condition, may be acted upon by carbonic acid.†

The partial oxidation of sulphides is, according to Van Hise, a possible additional method of formation of hydrogen sulphide, the reaction being as follows: (L. C. p. 1113.)

 $3\text{FeS}_2 + 4\text{H}_2\text{O} + 4\text{O} = \text{Fe}_3\text{O}_4 + 4\text{H}_2\text{S} + 2\text{SO}_2.$

The oxidizing processes are the most rapid near the surface, especially above the underground water level, and H₂S derived from this source, probably, supplies relatively shallow rather than deep waters.

The formation of H₂S by steam passing over sulphur, which occurs in connection with volcanoes, may be dismissed in considering the sulphur waters of Florida, since Florida has no volcanoes and no indications of volcanic activity.

SULPHUR WATER NOT EVIDENCE OF BEDS OF SULPHUR.

There is a widespread belief that the presence of sulphur water must necessarily indicate the existence of beds of the mineral sulphur. This conclusion does not follow. The probable sources of the sulphur in sulphur waters, as indicated above, is organic matter, together with metallic sulphates and sulphides scattered through sedimentary rocks.

[†]Inorganic Chemistry. International Library of Technology. Sec. 12, p. 11.

SULPHUR DEPOSITS FORMED FROM HYDROGEN SULPHIDE.

As stated in the last paragraph, sulphur waters are not to be regarded as resulting from beds of pure sulphur. On the contrary, it is probably true that these waters may, in some instances, result in the formation of such deposits. Hydrogen sulphide, when acted upon in the water by oxygen, breaks up, forming water and sulphur; the reaction being H2S+O=H2O+S. It is thus possible that H2S in the underground water, or escaping from the underground water, may become disassociated, forming deposits of pure sulphur. Such deposits of economic value have not been reported in Florida. It is a noteworthy fact, however, that one large mass of sulphur has been found underneath phosphate beds in Citrus County.* The formation of this mass of sulphur is probably due to hydrogen sulphide. A flocculent white coating of sulphur, or a sulphur compound invariably forms around sulphur springs and flowing sulphur wells.

ABSENCE OF HYDROGEN SULPHIDE FROM CERTAIN WATERS IN FLORIDA.

The absence of hydrogen sulphide from the first water obtained from areas in which the open porous limestone is the surface formation, has already been stated. It is a well-known fact that if sulphur water is allowed to stand in the open air the gas will escape. This method of freeing water from an excess of H₂S gas is a common practice wherever sulphur water is used for domestic purposes. Wherever porous limestone lies at or near the surface the sulphur gas, which the water may have contained, will find a ready means of escape. In other parts of the State, where compact and impervious formations rest upon the limestone, the gas is prevented from escaping and sulphur water is obtained.

^{*}Florida Geological Survey, First Annual Report, p. 44, 1908.

AMOUNT OF HYDROGEN SULPHIDE INFLUENCED BY PRESSURE.

The quantity of H2S gas, which the water is able to hold in solution under these conditions, is determined by the pressure. The law of the solubility of gases in liquids is as follows: The quantity of the gas which the liquid is able to dissolve is directly proportional to the pressure on the gas. In the open, porous limestone with no confining stratum above, the water at the top of the underground water level is merely under atmospheric pressure. After passing the underground water level, however, the pressure increases rapidly. The increase of pressure is not simply that due to the atmosphere, but that due to the weight of the overlying column of water plus the atmosphere. According to Van Hise: "The pressure, which really is determinative as to the amount of gas which may be held in solution, is that of a column of water extending to the free surface, plus the atmospheric pressure." From this law it follows that water, at a great depth and under great pressure, is capable of holding a large quantity of hydrogen sulphide in solution. When brought to the surface the pressure is relieved and the gas rapidly escapes. The artesian waters, in the flowing areas of the State, are under considerable pressure, thus enabling them to hold a large quantity of hydrogen sulphide as well as a high proportion of mineral solids in solution.

In order that the deep waters may hold large quantities of H₂S in solution, it is necessary that the gas be available. This implies that the gas in the artesian and other deep waters originates at some considerable depth rather than at or near the surface.

ARTESIAN WATER.

The term "artesian" has been variously used by different writers. Flowing wells first became well known in the province of Artios, France, and hence were called "artesian wells," and

^{*}L. c., page 70.

their water "artesian water." The first meaning of "artesian well" was, therefore, a flowing well; and of "artesian water," water under sufficient pressure to cause it to flow. With the extension into other areas of the use of deep wells as a source of water supply, many instances were found in which the water, although under pressure and rising almost to the surface, would not flow. In some cases the water will flow in areas of low surface elevation, and yet fail to flow in a slightly elevated area nearby. Artesian water thus came to mean water under pressure, causing it to rise in a boring when tapped, regardless of whether or not the pressure was sufficient to cause the water to rise above the surface level, and hence to flow. In the same way, and for similar reasons, the term "artesian well" came to include not only flowing wells, but also wells in which the water rises when the water-bearing stratum is tapped, regardless of whether or not the rise is sufficient to cause a flow. Occasionally, in popular usage, the term "artesian well" has been applied to any deep boring, and "artesian water" to water from such a well. In this report the term artesian is applied to water under pressure, and hence rising in a boring when tapped. The water may, or may not, rise to or above the surface. An "artesian well" is any well reaching to and tapping a stratum bearing such water; a "flowing well" is an "artesian well" that gives a surface flow. Artesian pressure is the pressure causing the water to rise in the boring when tapped. This is essentially the usage of these terms as adopted by the Division of Hydrology of the U.S. Geological Survey.*

CONDITIONS NECESSARY TO OBTAIN ARTESIAN WATER.

As essentials for artesian water, it is necessary to have (1) an adequate source of water, and (2) the proper structural conditions to retain the water under hydrostatic or artesian pressure. It will be convenient to discuss first the structural conditions.

^{*}Water Supply and Irrigation Paper, U. S. Geological Survey No. 160.

ARTESIAN BASIN.

A variety of conditions in the arrangement and structure of the underlying deposits may bring about artesian pressure. The simplest, although probably not the most common, is that of a basin-like arrangement of successive relatively pervious and impervious strata. This typical structure, known as an artesian basin, is shown in the accompanying diagram. It consists of a pervious layer (a), out-cropping at the surface on either side and sagging at the middle, above which is an impervious or water-



Fig. 1.—Illustrating Structure of an Artesian Basin.

tight confining layer (c), and below which is also an impervious layer (b). Water enters the pervious layer at its surface exposures at the sides. The water collecting in the central part of the basin is under pressure from the weight of the additional water entering from the sides. Therefore, a well put down to the water stratum in any part of the basin will obtain artesian water, or water which will rise in the boring. The rise in the boring is determined by the elevation of the in-take area, and can in no case rise above the elevation of the exposed edges of the stratum. As a matter of observation, it is found in all cases to rise not quite so high as the exposed edge of the stratum, the loss being due to the friction of movement through the rock. This loss of head due to friction necessarily varies with the texture of the stratum through which it passes, the passage being more free through the coarse material, and hence meeting with less friction than through fine. Whether or not wells put down in the basin will obtain flowing or non-flowing artesian water, depends upon the surface elevation of the mouth of the well. The diagram illustrates a basin in which flowing artesian wells may be obtained.

ARTESIAN SLOPE.

The basin arrangement of strata is not the only possible structure resulting in artesian pressure. The same result may, among other ways, be brought about quite effectively by an inclined porous stratum wedging out between two impervious strata. This condition is illustrated by the accompanying simple sketch, in which the pervious stratum (a) is represented as pinching out and disappearing between impervious strata. A pervious stratum grading into an impervious, or less pervious condition resulting

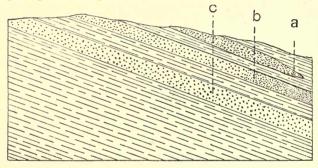


Fig. 2.—Illustrating structures that may prevail in an artesian slope; a. a pervious water-bearing stratum which pinches out between impervious strata; b. a pervious water-bearing stratum which grades into a less pervious stratum; c. a pervious water-bearing stratum in which the artesian pressure is due merely to the friction of water moving through the pores of the rock.

in artesian pressure, is represented by (b) of the same drawing. These conditions are often met with in the strata of the coastal plain. Not infrequently, a sandstone formation grades off shore into a finer sandstone, and ultimately into a shale. This condition comes about naturally through the sorting power of water acting along what was the coastal line at the time of formation of the strata under consideration. The coarser sand particles are dropped near the shore and form the sandstone; the finer sandgrains, together with more or less clay, are carried farther out.

and form a finer grained sandstone grading ultimately into a clay. Similarly, a sandstone, or other pervious formation, may pinch out as a result of the thickening of a shale or clay bed. The term "artesian slope" has been applied to such an area to distinguish it from an artesian basin.

The friction of water threading its way long distances through the pores of an inclined pervious formation may result in an appreciable artesian pressure. That this is true, may be demonstrated by the following very simple experiment: Fill a tube of any length with sand, and incline at a convenient angle. The sand of the tube represents the pervious water-bearing stratum; the tube itself, the impervious confining strata. Let smaller tubes placed vertically be welded into the larger tube. These vertical tubes represent bored wells. The water will be found to rise in the vertical tubes, exhibiting an appreciable artesian pressure due to the friction of flow through the sand.

ARTESIAN WATER FROM UNCONFINED HORIZONTAL BEDS.

It is, doubtless, possible to obtain artesian water in some instances from unconfined horizontal beds. This condition is illus-

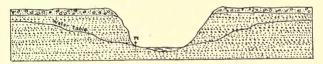


Fig. 3.—Illustrating artesian water from unconfined horizontal beds. The pressure in this case is due to the friction of water moving through the pores of the rock.

trated by the following sketch taken from the report of M. L. Fuller.* It is possible that some of the small local flows obtained in the lake region of interior Florida are due to similar conditions.

ARTESIAN WATER FROM SOLUTION PASSAGES.

Solution passages through limestones undoubtedly facilitate the free movement of water. If limestones should be otherwise relatively water tight, flows might still be obtained, in some instances, from water conducted through the cavities in the limestone. Such possible conditions are illustrated by the accompanying sketch, also taken from Mr. Fuller's paper.* Several other possible structural conditions that may give rise to artesian flows are described and illustrated in the paper referred to. Those illustrated above, however, include the structural conditions which seem likely to prevail in Florida.

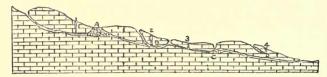


Fig. 4.—Sketch illustrating artesian flow obtained from solution passages in the limestone. After Fuller.

SOURCE OF ARTESIAN WATER OF FLORIDA.

The idea is rather prevalent that the artesian waters of Florida are in no sense local but are derived from the Appalachian Mountains, or some other remote inland point. This is an error, which, if not corrected, may prove detrimental. That the supply is local is evidenced by the fact that the artesian wells of the State are affected by local rains. Many of the well owners have recognized the effect of local rains on their wells; others who have observed less closely recognize no such variation. That the rainfall is sufficient to supply the large quantities obtained has already been demonstrated.

FORMATIONS SUPPLYING THE ARTESIAN WATER OF EASTERN AND SOUTHERN FLORIDA.

As explained in the chapter on Geology, the principal artesian reservoir of the eastern and southern part of Florida is the Vicks-

^{*}U. S. Geological Survey, Bull. 319, p. 39, 1908. Summary of the Controlling Factors of Artesian Flows.

burg group of limestones. In some localities, however, formations lying above the Vicksburg group supply a flow, although the flow from these more shallow formations is rarely ever so strong as from the deeper or Vicksburg limestones.

DEPTH OF THE ARTESIAN WATER.

The depth at which the artesian water is obtained is variable in different parts of the area. To find the depth for any particular locality, it will be necessary to refer to the subsequent chapters in which the several counties are treated individually.

COST OF WELLS.

It has been only within the past few years that artesian wells have begun to supplant shallow, open dug wells in the rural districts. One cause of the rapid increase of artesian wells in these districts is the necessity of irrigation in order to safeguard trucking and general crops against droughts. Again, from a health standpoint, the water from these deeper wells is less liable to contamination than is the water from the shallower or surface wells.

The cost of an artesian well depends upon the depth to which it is necessary to drill, the size of the well desired, the amount of casing used and the character of the material that will probably be penetrated in drilling. With a knowledge of the nature of the underlying formations in a given area well drillers know approximately the time and labor it will take to complete a certain size well. In such an instance it is frequently the case that a well is completed for a stipulated amount, regardless of the depth. It is more customary, however, to let a contract for a certain size well at a given price per foot. These prices vary in different sections of the State, but on the average two-inch wells are sunk for from \$1.00 to \$1.25 per foot; three- and four-inch wells from \$1.50 to \$2.00 per foot. The larger wells range in proportion, a ten-inch well costing about \$3.50 per foot. The driller, at these prices, furnishes the casing.

INCREASED FLOW OF ARTESIAN WELLS WITH INCREASED DEPTH,

As a rule, the amount of flow or yield of wells in Eastern Florida increases with depth. To this rule there are, doubtless, exceptions, since the amount of flow, in all cases, depends upon the variable structure of the rock through which the drill passes. As illustrations of increased flow with increased depth, the following may be cited:

In the new city well at Jacksonville, well No. 10 of the city water supply, the first flow obtained was a light flow of 5 gallons per minute at a depth of 270 feet. At a depth of 498 feet the flow increased to 112 gallons per minute. Upon reaching the Vicksburg Limestone, at a depth of 510 feet, the flow increased to 200 gallons per minute. The flow at the depth of 635 feet was found to be 500 gallons per minute. At 900 feet the flow was about 900 gallons per minute. At 980 feet, the full depth of the well, the flow was from 1,500 to 2,000 gallons per minute. For the detailed measurements of flow on this well the Survey is indebted to the drillers, the Hughes Specialty Well Drilling Company of Charleston, South Carolina.

A like increase of flow is shown by the Ponce de Leon well in St. Johns County, the measurements of which were kept, and have been kindly supplied by Messrs. McGuire and McDonald, contractors. The first flow in this well of 50 gallons per minute was obtained at a depth of 170 feet. At 177 feet the flow increased to 350 gallons per minute. At 410 feet the flow was 2,083 gallons. At 520 feet the flow had increased to 4,860 gallons. At 1,110 feet the flow was 6,075 gallons. The well was continued to a total depth of 1,440 feet. The record of the well, however, contains no mention of increased flow below 1,110 feet. While exact measurements, like those given above, are seldom made; the drillers, with few exceptions, report increased flow with increased depth.

INCREASED HEAD WITH INCREASED DEPTH.

Not only does the amount of flow of the water in this section of the State increase with increased depth, but the head or pressure, or height above the ground to which the water will rise likewise increases. The head is, in reality, only a measure of the pressure. The amount of flow is within limits dependent upon the amount of pressure. Other conditions remaining the same, an increased pressure will result in an increased flow. For the records regarding pressure, it is necessary to rely chiefly upon the Jacksonville and St. Augustine wells already referred to.

At 680 feet the pressure of the artesian water in the Jackson-ville well was 12 pounds per square inch, or sufficient pressure to cause the water to rise vertically in a pipe 27.72 feet. At 900 feet the pressure, as shown by the gauge, was 15 pounds, or sufficient to cause the water to rise 34.65 feet.

The Ponce de Leon Hotel well, at St. Augustine, affords valuable information as to the possibility of obtaining increased head, in this section of the State, by drilling to greater depths. This well was drilled to a total depth of 1,440 feet. A measure of the head was made at frequent intervals while drilling. The first considerable flow obtained at St. Augustine is under a pressure, causing it to rise about 32 feet above sea. At the depth of 350 feet the head was found to have increased to 38 feet above sea. At the depth of 520 feet the head had increased to 42 feet, a total gain of 10 feet. The head at greater depths than 520 feet is not specifically recorded.

INCREASED TEMPERATURE WITH INCREASED DEPTH.

The temperature of the water at St. Augustine was found to increase with the depth. The temperature of the water in the Ponce de Leon well, at the depth of 35 feet, is reported as 62 degrees F. At approximately 100 feet the temperature was 72 degrees F. At 170 feet the temperature was 74 degrees F. The increased flow obtained at 177 feet showed a temperature of 76 degrees F. At 520 feet the temperature of the water in the pipe was found to be 79 degrees F. At 1,110 feet the temperature was 80 degrees F. Between 1,170 and 1,225 feet the water taken from the sand pump showed a temperature of 85 degrees F. Water taken from the sand pump, between 1,340 and 1,390 feet, showed a temperature of 86 degrees F.

This record of the Ponce de Leon well, at St. Augustine, is supplemented by the record from the new city well at Jackson-ville. In the Jacksonville well the following temperatures were recorded: At a depth of 498 feet, the temperature of the water flowing from the pipe was 71 degrees F. At 635 feet the temperature was 74 degrees F. At 900 feet the temperature still registered 74 degrees F. These measurements made, as the water escapes from the pipe, are necessarily approximate measurements. Not only does the water lose in temperature in moving to the mouth of the pipe, but it mingles with the higher and colder waters entering the pipe, which necessarily equalizes the temperature of the whole. They show, however, increase of temperature with increase of depth.

TABLE SHOWING PROGRESSIVE LOSS RECORD OF JACKSONVILLE

Number of Well	Size of Well, inches	Date When Completed	Flow of Well When Completed. Gal- lons in 24 Hours.	1886	1888	1889	1891 May 30	1892 Nov. 1	1893 Jan. 1	1894 Jan. 1	1895
1	6	Nov. 1885 Dec.	864,000				309,096		,	,	200,232
4	12	1896 April					•••••				
6	10	1901	2,095,639								
7	10	Aug. 1904									
2	6	Mch. 1886 Feb.		1,296,000	1,167,360	808,485	458,784	412,128	381,024	332,424	309,096
3	10	1889 April	3,360,052			3,360,652	1,995,840	1,829,952	1,752,192	1,440,152	1,347,840
5	8	1899	590,676								
*****					••••••						
Total Flow											
Loss in Flow Gain by New Well											

LOSS OF HEAD AND REDUCTION IN FLOW.

Exact measurements of loss of head and reduction in flow in artesian wells are usually difficult to obtain. In the case of the Jacksonville city water supply, fortunately, measurements of flow have been taken at intervals from the time the first wells were put down in 1885 to the present time. These measurements kept through a period of 24 years afford records of especial interest and value. The following table of flow was supplied by Capt. R. N. Ellis, Superintendent of the Jacksonville city water supply. Two basins are used to receive the flow known as the north basin and the south basin. The wells are grouped in the table according to the basin into which they flow. The wells are

OF FLOW OF ARTESIAN WELLS. CITY WELLS, 1885-1904.

1896	1897	1898	1899 April	1900	1901	1902 Jan.	1902 Nov. 29	1903	1904 April 1	1904 Oct. 26
188,568			208,640		1					
1,354,320)		881,280	'			2,287,440				
)							
285,693			(••••
1,093,456	1,368,576	1,368,576	1,322,220	1,829,947	1,441,147	1,418,907	1,368,576	1,368,576	1,347,840	1,099,080
••••••										
			••••••		••••••			•••••		***************************************
2,922,037	2,476,656	2,249,876	2,784,176	2,639,947	3,935,947	3,706,347	3,079,296	3,079,296	3,032,640	3,385,380
289,451	445,381	226,800				229,600				352,740
1,064,869	•••••		534,320	••••••	1,296,000	***************************************			•••••	248,760

numbered chronologically in the order of the date when completed.

This table shows conclusively that, although the rate of flow is variable for different wells and for the same wells at different periods, yet in this group of wells there is a continuous and progressive loss of flow. That the same is true of other wells throughout this area, there can be no reasonable doubt. Those who give no special attention to their wells suppose, as a rule, that the flow remains unaffected indefinitely. Many other well owners, however, have observed this loss in flow with succeeding years. The reduced flow is best observed near the margin of the flowing area in wells located on somewhat elevated ground. Many of the wells from which the water will flow only a few feet above the surface when first drilled may, in time, cease to flow. In these cases the pressure which originally caused the flow having been partly relieved, the water no longer rises above the surface of the ground.

Exhaustion and ultimate failure of an artesian reservoir is not unknown. It is, probably, true that, in nearly all artesian sections, the original pressure gradient in the water-bearing rock is appreciably lowered by the drafts made upon the subterranean supply, with a consequent actual decrease in the capacity of the wells. In this connection, Professor C. S. Slichter states: "It must be kept well in mind that there is a limit to the amount of water that can be withdrawn from an artesian basin. There is no such thing as an inexhaustible supply in this connection. The amount of water available is limited on the one hand by the amount of rainfall upon the catchment area, and the facility with which the rainfall can obtain entrance to the porous stratum and. on the other hand, by the capacity of the water-bearing rock to transmit the water over long distances and diminution through leakage and seepage. These two limiting conditions are usually of sufficient magnitude to render the overdrawing of the supply a practical and present danger, which should be constantly kept in mind."

With regard to the artesian basin at Denver, Colorado, the

^{*}U. S. Geol. Surv., Water Supply Paper, No. 67, p. 94, 95, 1902.

failure of which was unusually rapid, Slichter says: "This basin was discovered in 1884, and in a few years about 400 wells had been drilled within an area extending a distance of 40 miles, along South Platte River, in a strip about 5 miles wide on both sides of the stream. Most of the wells were within the limits of the city itself. Many of the wells had a good pressure and strong flow when first constructed. In 1886 it was not thought that any general decrease in the flow of the wells could be detected. Between 1888 and 1890, however, a continuous decrease in the flow of the city wells took place, and by the end of the latter year all but six of the city wells had to be pumped, while numerous wells in the basin were permanently abandoned."

CAUSE OF THE LOSS OF FLOW.

The loss of flow may be due to several causes. It is frequently the case that the life of an artesian well is limited. The escape of water through the well relieves the pressure, which results in a reduced flow. In some instances, pressure has so far been relieved that wells have ceased to flow entirely. This may be regarded as a natural and unavoidable loss of flow.

The second cause of reduced flow, which may have affected the Jacksonville and other wells, is interference of wells. Numerous instances are on record where one artesian well has affected surrounding wells.

A third possible cause is clogging of the wells through accumulation of sand or other material in the pipes, or in the formations through which the water comes. In addition to the accumulation of sand, it is not impossible that the porosity of the formation immediately around the well may have been more or less affected by chemical deposition since the well was drilled. It seems probable, however, that the clogging of the pores of the rock is more likely to be caused by material mechanically transported than by chemical deposition.

Improper casing is likewise a frequent cause of failure. It is frequently the case that an insufficient length of casing is used in the well. In such cases the sand gains entrance, or the well caves below the casing, clogging or partly clogging the opening, thereby reducing or entirely stopping the flow.

NECESSITY OF GUARDING AGAINST WASTE OF ARTESIAN WATER.

The records that have been given above indicate clearly that useless waste of water should not be permitted. An artesian well draws not on an inexhaustible supply of water from some remote source, but draws upon a relatively local supply which is appreciably affected by continued use. A well permitted to flow uninterruptedly draws not only on the supply of the land on which it is located, but affects also the supply of the adjacent land. A State, a community, or an individual that permits the useless and reckless waste of artesian water will ultimately find a most valuable asset impaired by extravagance, and possibly no longer adequate.

It is urged by some well owners that to cut off a well, or to stop the flow when not in use is unsafe as sand or other material may get into and clog the well. The flow of the well can be reduced to one-third or one-fourth its normal volume and the danger from the accumulation of sand, when there is such danger, guarded against. Moreover, where wells are cased, as they should be to the Vicksburg Limestone, it is doubtful if there is danger of clogging and reducing or stopping the flow. A law restricting the waste of artesian water is urgently recommended.

SIMPLE METHOD OF DETERMINING FLOW OF ARTESIAN WELLS.

A simple method for measuring approximately the flow from an artesian well has been devised by Professor J. E. Todd, formerly State Geologist of South Dakota. The following is Professor Todd's method in full:

"It is often desirable to know the amount of water delivered by an artesian well. Frequently a contract calls for a certain amount. It is also well to know whether the flow is diminishing and how much.

"When a well is small, its flow may be measured easily with

a watch and a gallon measure, or a keg or a barrel of known capacity, but for wells flowing over twenty or thirty gallons a minute, it is not so easy to determine with accuracy.

"If the well is large it may be measured with a weir, but that is constructed only with considerable trouble. If the water runs in a sluice or ditch of uniform width, its cross section may be estimated and its velocity taken. This method, however, is not very accurate. The following are methods which give fairly accurate results with little trouble and in short time. All that is necessary for the purpose is that the water be discharged through a pipe of uniform diameter, a foot rule, still air, and care in taking measurements.

"Two methods are proposed, one for pipes discharging vertically, which is particularly applicable before the well is permanently finished, and one for horizontal discharge, which is the most frequent way of finishing a well. For the measuring a vertical flow we have extended a method which was first used by Mr. P. E. Manchester, C. E., of Chamberlain, who published a table adapted to large wells, in the Chamberlain Register, December, 1895.

"The table below is adapted to wells of moderate size as well as to larger. In case a well is found of other diameter than that given in the table, its discharge may be obtained without much difficulty from the table by remembering that other things being equal the discharge varies as the square of the diameter of the pipe. If, for example, the pipe is one-half inch in diameter its discharge will be one-fourth of that of a pipe one inch in diameter, whose stream reaches the same height, so also a pipe eight inches in diameter may be obtained by multiplying that of the four-inch pipe by four.

"In the first case the inside diameter of the pipe may be measured, then the distance from the end of the pipe to the highest point of the dome of water above, in a strictly vertical direction—a to b in the diagram. Then these distances may be found in the table and the corresponding figure will give the number of gallons discharged per minute. The blowing of the wind

need not interfere in this case as long as the measurements are taken vertically.

"The method for determining the discharge of horizontai pipes requires a little more care. First, measure the diameter of the pipe as before, then the vertical distance from the middle of the opening of the pipe, or some convenient point corresponding to it on the side of the pipe, vertically downward six inches—a to

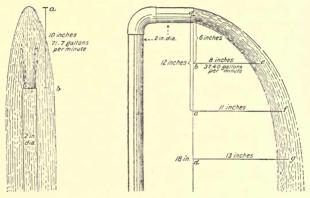


Fig. 5.—Illustrating method of measuring the flow of an artesian well from horizontal and vertical pipes. After Todd.

b, then from this point strictly horizontally to the center of the stream—b to e. With these data, the flow in gallons per minute may be obtained from the table. It will be readily seen that a slight error may make much difference in the discharge. Care must be taken to measure horizontally and also to the middle of the stream.

"Because of this difficulty, it is desirable to check the first determination by a second. For this purpose, columns are given in the tables for corresponding measurements twelve inches below the center of the pipe. Of course, the discharge from the same pipe must be the same in measurements of the same stream. In this case, the occurrence of wind, blowing either with or against

the water, may vitiate results to an indefinite amount, therefore measurements should be taken while the air is still.

"The flow of pipes of diameters not given in the Table II, may be easily obtained for corresponding measurements, as follows: For ½ inch, multiply discharge of 1-inch pipe by .25; for ¾-inch, by .56; for 1¼-inch, by 1.56; for 1½-inch, by 2.25. For 3-inch, multiply 2-inch pipe by 2.25; for 4-inch, by 4; for 5-inch, by 6.25; for 6-inch, by 9; for 8-inch, by 16.

TABLES FOR DETERMINING YIELD OF ARTESIAN WELLS.

I. Flow from Vertical Pipes						II.	Flow from Horizontal Pipes				
Helght of Jet.	Discharge in Gallons per Minute from Respective Pipes of Diameter given in Inches.					Jet.	Flow in Gallous per Minute for Pipes.				
Heigh	1	11/4	1½	2	3	Horizontal Length of J	1 inch in	1 inch in Diam.		2 inch in Diam.	
In. ½	3.96 5.60 7.99	6.2 8.7 12.5	8.91 12.6 18.0	15.8 22.4 32.0	30.6 51.4 71.9	Hon	6 in. level.	12 in. level.	6 in. level.	12 in. level.	
½ 1 2 3 4 5 6 7 8	9.81 11.33 12.68	15.3 17.7 19.8	22.1 25.5 28.5	39.2 45.3 50.7	88.3 102.0 113.8	In. 6 7 8	7.01 8.18 9.35	4.95 5.77	27.71 32.33 36.94	19.63 22.90 26.18	
9	17.01	21.7 23.6 25.1 26.6	31.2 33.7 36.0 38.3	55.5 59.8 64.0 68.0	124.9 134.9 144.1 153.1	9 10 11	10.51 11.68 12.85	6.60 7.42 8.25 9.08	41.56 46.18 50.80	29.45 32.72 35.99	
10 11 12 13	18.80 19.65	28.1 29.5 30.7 31.8	40.3 42.3 44.2 45.9	71.6 75.2 78.6 81.8	161.3 169.3 176.9 184.1	12 13 14 15	14.02 15.19 16.36 17.53	9.91 10.73 11.56 12.38	55.42 60.03 64.65 69.27	39.20 42.51 45.81 49.08	
14 15 16	21.22 21.95 22.67	33.0 34.2 35.2	47.6 49.3 50.9	84.9 87.8 90.7 93.5	190.9 197.5 203.9 210.3	16 17 18 19	18.70 19.87 21.04 22.21	13.21 14.04 14.86 15.69	69.27 73.89 78.51 83.12 87.74	52.35 55.62 58.90 62.17	
17 18 19 20	24.06 24.72 25.37	36.3 37.5 38.6 39.6	52.5 54.1 55.6 57.0	96.2 98.9 101.6	216.5 222.5 228.5	20 21 22	23.37 24.54 25.71	16.51 17.34 18.17	92.36 96.98 101.60	65.44 68.71 71.98	
21 22 23 24	26.02 26.66 27.28 27.90	40.6 41.6 42.6 43.5	58.4 59.9 61.4 62.8	104.2 106.7 109.2 111.6	234.3 240.0 245.6 251.1	23 24 25 26	26.88 28.04 29.11 30,38	18.99 19.82 20.64 21.47	106.21 110.83 115.45 120.07	75.26 78.53 81.80 85.07	
24 25 26 27	29.59	44.4 45.3 46.1 46.9	64.1 65.3 66.4 67.5	114.0 116.2 118.2 120.3 121.9	256.4 261.4 266.1 270.4	25 26 27 28 29 30	29.11 30.38 31.55 32.72 33.89 35.06	21.47 22.29 23.12 23.95 24.77	124.69 129.30 133.92 138.54	88.34 91.62 94.89 98.16	
28 29 30 36	30.94 34.1	47.5 48.2 53.2	68.5 69.4 76.7	123.4 136.3	274.1 277.6 306.6	31 32 33	36.23 37.40 38.57	25.59 ,26.42 27.25	143.16 147.78 152.39	101.43 104.70 107.98	
48 60 72 84	43.8 48.2 51.9	61.0 68.4 75.2 81.0	88.0 98.6 108.0 116.8	156.5 175.2 192.9 207.6	352.1 394.3 434.0 467.0	34 35 36	39.64 40.45 41.60	28.08 28.64 59.46	157.01 161.63 166.25	111.25 114.52 117.19	
96 108 120 132	55.6 58.9 62.2	86.7 92.0 98.0 102.6	125.0 132.6 139.9 146.5	222.2 235.9 248.7 260.4	500.0 530.8 559.5 585.9		Continue 1.15	by addir	for eac	3.27	
144		106.4	153.1	272.2	612.5	11					

"In both these tables it has not been thought necessary to make any allowance for the resistance of the atmosphere. Doubtless, when the velocity of the stream is great, the resistance is considerable; but as the pressure checks the flow, and our object is simply to measure the amount of flow, it need not be taken into consideration. In case pipes are found of diameters not corresponding to the table, the same rule may be applied as in the first case.

"Whenever fractions occur in the height or horizontal distance of the stream, the number of gallons may be obtained by dividing the difference between the readings in the table for the nearest whole numbers, according to the size of the fraction. For example, if the distance from the top of the pipe to the top of the stream, in the first case, is nine and one-third inches, one-third of the difference between the readings in the table for nine and ten inches must be added to the former to give the right result. In case one measures the flow of his well according to both methods, he may think that they should correspond, but such is not the case. In the vertical discharge, as there is less friction, the flow will be larger, so also difference will be found according to the length of horizontal pipe used in the second case. The longer the pipe, the more friction and less the flow.

"As pipes are occasionally at an angle, it is well to know that the second method may be applied to them, if the first measurement is taken strictly *vertically* from the center of the opening, and the second from that point *tarallel with the axis* of the pipe to the center of the stream as before. The measurements may then be read from the table as before.

"This method is also applicable to measuring the discharges of different pipes when water is distributed about a farm or in

a city.

"Pipes which have been cut in the usual way are frequently diminished in diameter by the incurving of the edge of the pipe. This will diminish the flow, but how much can only be roughly estimated. It will be greater than that of a straight pipe having the exact diameter of the opening as reduced."

THE AREAS OF ARTESIAN FLOW IN FLORIDA.

The accompanying map indicates, in a general way, the flowing and non-flowing areas of the State. In using such general maps it should be borne in mind that artesian water depends primarily upon the structure of the underlying formations, and these are subject to variations of which there may be no surface indications. Moreover, local elevations which affect flow can not be indicated on a small scale map. Thus while the map indicates approximately the limits of flow, the exact limits can be determined in most cases only by drilling.

The shading on the map indicates those parts of the State in which flowing wells have been obtained, or may be expected. There are, as will be seen, three principal areas of flow as follows: the Atlantic Coast area, the Southern Gulf Coast area and the Western Gulf Coast area.

THE ATLANTIC COAST AREA.

The Atlantic Coast area includes much of Nassau and Duval Counties, and, with the exception of local elevated areas, all of St. Johns County; it follows the valley of the St. Johns River almost if not quite to the head waters, while a narrow strip reaches south along the Atlantic Coast for 250 to 300 miles. The artesian water-bearing formation dips in passing to the south, being reached at Palm Beach at the depth of about 1,000 feet. In addition to its increased depth the water at Palm Beach was found to be too salty to be used for household purposes. Between Palm Beach and Key West no wells have been drilled deep enough to reach this formation. The deep well drilled on Key Vaca by the Florida East Coast Railway terminated at 700 feet in quartz sands, with sandstones and clay in streaks, not having reached the Vicksburg Limestone.*

At Key West two wells have been drilled to the Vicksburg, which is reached at that locality at a depth of about 700 feet.

The first of these wells, drilled in 1895, is reported to have

^{*}Florida Geol. Survey, Second Annual Report, p. 205, 1909.

reached a depth of 2,000 feet. The well was non-flowing and the water salty. No adequate record of this well was kept, and it is not known to what depth the well was cased, nor whether or not there was any attempt made to drill beyond and case off the salty water. The second well was drilled, in 1909-10, by S. O. Johnson and reached a total depth of 1010 feet. This well is cased about 150 feet. It is non-flowing and salty. Two samples of water from this well have been received from Mr. Johnson. One is said to have been taken from the water near the top of the well; the other from near the bottom of the well. The first of these samples contains chlorine 2,340 parts per million parts water. The sample said to have come from the bottom of the well contains 1358 parts chlorine per million parts water.

THE SOUTHERN GULF COAST AREA.

Flowing wells have been obtained in areas of low elevation at Tampa. St. Petersburg and elsewhere, along the Gulf Coast for some distance north of St. Petersburg. It is only near the sea level in this northward extent of the area that a flow is to be expected. In Manatee County, along the Manatee River, strong flowing wells have been obtained; some of them having a pressure of eight or more pounds. The wells in this county are used extensively for irrigation. In DeSoto County flowing wells occur at Punta Gorda, and along Peace Creek into Polk County. Some of the wells at Punta Gorda have a head of about fifty feet. In Lee County flowing wells have been obtained at Ft. Myers, along the Caloosahatchee River to Labelle, and in the interior southeast of Ft. Myers. In the well of A. P. Miller, of Ft. Myers, having a depth of 535 feet the water was found to be under a pressure of 17 pounds, giving it a head of 39 feet above the surface. The southward extent of this flowing area has not been determined. Approaching the southern limit the amount of salt in the water increases, certain of the wells toward the southern part of Lee County becoming too salty for use. The Vicksburg Limestone is probably the water bearing formation in Southern as in Eastern Florida.

Whether or not flowing wells can be obtained in the Ever-

glades, east and south of Lake Okeechobee, has not been determined as no wells have been drilled in this part of the State. While definite information is lacking, it is considered probable that flowing wells will be obtained within the Everglades; particularly toward the western side. Subsequent records may show that the Atlantic Coast and Gulf Coast flowing areas are connected by way of the Everglades and around Lake Okeechobee.

While the northern limit of the Southern Gulf Coast area has been given as the Pinellas Peninsula, from recent well records it seems probable that a flow may be obtained north of this limit, and possibly entirely around the Gulf Coast. Two wells have reached this deeper flow, one at Crystal River, in Citrus County, and one at Perry, in Taylor County. The well in Taylor County reached a depth of 1,199 feet. The total dissolved solids in this water, as shown by analysis made by the State Chemist, is 5,650 parts per million parts water. The chlorine alone amounts to 590 parts per million parts water. The water is reported to have medicinal qualities. The well in Citrus County reached a depth of 1,900 feet. The following is an analysis of the water from this well made for the State Survey by the State Chemist in 1907:

Ingredients. Parts per million	
Calcium oxide (CaO)	
Magnesium oxide (MgO) 480.6	
Sulphate (SO ₄) 2,684.0	
Chlorine (C1) 903.9	
Silica (SiO ₂) 30.0	
Total solids 6,474.0	

WESTERN GULF COAST AREA.

The Western Gulf Coast area begins at Carrabelle, in Franklin County, and extends to the western line of the State. The flow along this westward extension of the State is evidently due to the rapid southward dip of the formations exposed along the northern line of the State, and in southern Georgia and Alabama. Both the Oligocene and the Miocene formations exposed

along the Ocklocknee, Apalachicola and other rivers crossing Western Florida, from north to south, dip and pass from view in approaching the coast. It is doubtless from these or from later formations that the flowing water of this section is obtained. At Apalachicola the artesian water has a head bringing it only a few feet above the surface. The wells at this locality vary in depth from 350 to 620 feet. A number of deep wells have been drilled along St. Andrews Bay, in Washington County. The artesian water in this section will rise several feet above sea level. One of the city wells at Panama City is reported to flow 13.02 feet above the surface, or about 15 feet above sea level. A second city well, located on higher ground, is non-flowing although drilled to a depth of 630 feet.

Several wells, ranging in depth from 181 to 210 feet, have been drilled along Choctawhatchee Bay, in Walton County. A strong flow is obtained in this section. A well 210 feet deep, 3 miles south of Freeport, owned by the Baker-Wingfield Company, had a pressure when measured September 22, 1910, of 15 pounds, equivalent to a head of 34.65 feet above surface. Another we'l near by, 189 feet deep, belonging to the Choctawhatchee Lumber Company, had a pressure on the same date of 12½ pounds, equivalent to a head of 28.87 feet above the surface. Both of these wells are located on low ground, near sea level. A well, 181 feet deep, belonging to Messrs. J. C. Blackburn and J. N. McLain, located on higher ground, in the town of Freeport, had a pressure of 6½ pounds, equivalent to a head above the surface of 15 feet.

At Pensacola, and generally along the coast in Escambia County, good flowing wells are obtained. A well at Northrop, 1,030 feet deep, belonging to Stephen Lee, is reported to have a head of 60 feet above the surface. At Muscogee a well, 175 feet deep, belonging to the Southern States Lumber Company, is reported to have a head of 38 feet above the surface. A well on Bayou Grande, near Pensacola, belonging to Messrs. Stephen and W. F. Lee, is reported to be 1,000 feet deep and to have a pressure of 24 pounds, equivalent to a head of 55.44 feet above the surface. The temperature of the water is given as 92 degrees F. and the flow as 225,000 gallons per day.

Among the isolated flowing wells in the State two at Graceville, in Jackson County, are of especial interest. The first well at this locality was drilled some years ago by Mr. F. J. White. When first drilled, Mr. White says, the well flowed slightly above the surface, but soon afterwards ceased to flow. On the day following the great San Francisco earthquake of 1906, however, the well was observed to be flowing, and it has continued flowing from that date. The second well at Graceville was drilled in 1910 for the city by Mr. C. D. Williams. This well is 287 feet deep. The water has a head sufficient to rise about 2 feet above the surface. The well is eight inches in diameter for 161 feet. and six inches to the bottom. The flow is estimated at 20 gallons per minutes. Although no well samples have been obtained it seems probable from the driller's notes that the wells at this locality pass through the Vicksburg Limestone and enter an underlving formation.

A well drilled as a test well for oil about six miles south of Chipley, in Washington County, is said to have flowed at a depth of about 1.250 feet.

During 1912 flowing wells were obtained at and near Ponce de Leon, in Holmes County. These wells vary in depth from 200 to 213 feet. The water rises 5 to 6 feet above the surface. After passing through about 100 to 130 feet of sands, sandstone, and blue marl, limestone is reached from which the artesian water is obtained. The following is a log of one of these wells drilled for the town of Ponce de Leon. This well flows 65 gallons per minute and has a head of six feet above the surface. The record is by the drillers, M. J. Gray & Company.

	Feet.
Coarse yellow sand	0- 10
White sandy clay	10- 39
Yellow sand	39- 43
Sandstone	43- 60
Blue marl	60-130
White limestone	130-203

DISCUSSION BY COUNTIES

NASSAU COUNTY.

LOCATION AND SURFACE FEATURES.

Nassau County lies bordering the Atlantic Ocean in extreme northeastern Florida. The St. Mary's River, taking its source in Okefenokee and other swamps along the Florida-Georgia boundary line, after flowing south and southeast until approximately on a parallel with the mouth of the St. Johns River, turns abruptly and flows directly north for a distance of 30 miles. From this point the river flows slightly south of east to the Atlantic. Nassau County occupies the northern and western part of the peninsula-like extension of Florida formed by the northward bend of this river, the northern and western boundaries of the county being formed by the river.

The surface is in general level or rolling. The highest elevation found within the county is near the western side, where a flat-topped ridge extends north and south, lying only a few miles distant from the St. Marys river. Towns lying on this ridge are as follows: Boulogne, elevation 70 feet; Hilliard, elevation 66 feet; Crawford, elevation 85 feet; Kent, elevation 70 feet Some places on this ridge may exceed 100 feet in elevation. Aside from this ridge no points are recorded in Nassau County having an elevation reaching 50 feet.

That part of the county east of this ridge, including fully two thirds of the county, is lower in elevation and is prevailingly of the open flatwoods type of soil.

WATER-BEARING FORMATIONS

Up to the present time the identification of the age and character of the different strata encountered in drilling in Nassau County has been difficult owing to the fact that no complete set of well samples from any well in this county has been obtained.

From an incomplete set of samples from the J. R. Wilson well at Callahan, kindly saved by the driller, Mr. H. C. Russell, it is seen that limestone was encountered at a depth of from 212 to 255 feet. The limestone was very hard and massive and no fossils were observed in the sample. Just above this stratum of rock is reported a twelve foot layer of sand and black pebbles, and in fact these black pebbles were seen imbedded in the underlying limestone. Water is reported to flow from this depth. Below this stratum of rock 100 feet of blue marl with inclusions of several thin strata of shells is reported. In a sample from this stratum the sand was gray in color and the grains were round in outline. The black pebbles, smaller than those in the above stratum, occur also at this depth but may have dropped down from above. At a depth of from 355 to 364 feet a very hard rock is reported, but no further notes were made of this and no samples kept. From 364 to 418 feet indurated gray sand and blue marl are reported and immediately below this is encountered a rock, apparently limestone, in which the water is reported to increase in head and in volume of flow as each hard layer is penetrated. From all information that could be gathered it seems probable that this limestone is the Vicksburg.

Exposures of clayey, impure limestones are found along the St. Marys River, at High Bluff, about six miles and at Saw Pit Bluff, about two miles above the Atlantic Coast Line Railroad bridge; also at Chalk Bluff and at Orange Bluff, near King's Ferry.

The section at Saw Pit Bluff is as follows

	Feet.
Sticky blue clay with some soil	5
Impure limestone	э

At Chalk Bluff, about two miles above King's Ferry, the following section was observed:

	 eet.
Sticky blue clay with some soil at top	 2
Calcareous clay resembling fuller's earth	 2
White chalky material	 1
Clay resembling fuller's earth	 2

Going down the river from Kings Ferry no rock or shell exposures are seen until Reeds Bluff, near Crandall, is reached. This bluff, which lies on the Florida side of the St. Mary's River, is semi-circular in shape and is about three-fourths of a mile long. The following section was made near the middle of this bluff:

	Feet.
Incoherent pale yellow sands	20-40
Oyster shell reef imbedded in fine, sandy clay	10-15
Blue sands and sandy clays oxidizing yellow	10-20

The oyster reef in this section rests irregularly upon the underlying sands, the base of the reef being 10 to 20 feet above low tide. The oyster reef extends about two hundred feet along the face of the bluft.

The unusual thickness of the loose yellow sands at the top of the bluff is due to the fact that the upward moving currents of air carry sand as it is loosened along the face of the bluff to the top, where it accumulates as a sand dune.

Roses Bluff, also on the Florida side of the river, about two miles below Crandall, is semi-circular in shape and is fully two miles long. The following section was made near the middle of this bluff:

	Feet.
Dark colored sand and soil	. 4
Dark iron-stained sand (hardpan)	. 7
Ochre yellow sand	. 8
Sand with some clay	. 5
Sandy shell bearing marl, blue, oxidizing yellow	. 4
Sloping to water's edge at low tide	. 5
	_
	20

AREA OF ARTESIAN FLOW IN NASSAU COUNTY.

That part of Nassau County in which flowing wells can be obtained is indicated on the accompanying map by shading. Flowing wells may be obtained as shown by the map, Fig. 6, in approximately the eastern two-thirds of the county. A relatively small area, including the ridge already mentioned, lying near the western part of the county and extending north and south, parallel with the St. Marys River, stands too high to obtain

flowing wells. In this section, however, non-flowing artesian water may be obtained which will stand within a few feet of the surface.

LOCAL DETAILS.

CALLAHAN.

There are several flowing wells at and in the vicinity of Callahan, varying from 410 to 489.7 feet in depth. Three different water-bearing strata are reported in all the deeper wells at Callahan, the first occurring at about 50, the second at from 160 to 200, and the third at 400 to 425 feet. The water from the first stratum does not flow, but rises to within 6 to 10 feet of the surface, and is found in a shell formation. The water from the other two strata rises from 28 to 48 feet above the surface.

The first deep or artesian well at Callahan was drilled in 1904. This well was put down at the instance of several of the residents, by D. C. Stafford. It is a three-inch well and reported to be about 400 feet deep. The main source of domestic water supply at Callahan until the completion of this well had been shallow wells. These wells, which vary in depth from 25 to 60 feet, obtain their water supply chiefly from the underlying sands and clays. The water from these sands and clays, while soft and very desirable for domestic purposes, seemed to be contaminated by surface impurities as was indicated by the many cases of typhoid fever. Several of the citizens suspected that this sickness was due to the drinking of this surface water and their combined efforts resulted in the completion of this first artesian well. Since the completion of this and other deep wells the healthfulness of the locality has greatly improved.

A three-inch well drilled for J. R. Wilson in 1908 by H. C. Russell reached a total depth of 412 feet. It is reported cased 188 feet and has a pressure of 21 pounds, as shown by the pressure gauge February 3, 1910, or a head of 48.51 feet above the surface. The elevation of the depot at Callahan, as given by the Atlantic Coast Line Railroad, is 20 feet above sea. The location of the above well is approximately 2 feet lower than the depot,

or about 18 feet above sea, thus making a total head of 66.51 feet above sea.

Another three-inch well was drilled by H. C. Russell for T. R. Wells & Brother. This well reached a total depth of 420 feet and is cased 192 feet. The pressure of this well, as shown by the pressure gauge, February 3, 1910, was 19 pounds or a head of 43.89 feet above the surface. The elevation of the well is approximately 3 feet higher than the depot or 5 feet higher than the Wilson well. The head would thus be 66.89 feet above sea or about the same as that of the Wilson well.

In February, 1910, H. C. Russell completed a second well for J. R. Wilson. This well is located about three-fourths of a mile east of Callahan. It is a three-inch well and reaches a total depth of 489.7 feet. 212 feet of 3-inch casing was used. The first flow in this well was encountered at 200 feet, the second at 275 feet and the third at 425 feet. Although the drilling in this well was continued to a depth of 489.7 feet it is reported that no increase of water was obtained below 460 feet. The following is a log of this well as constructed from the notes kept by the driller and from samples of the drillings saved by him:

G 4	
Sand 0-	2
Red clay 2- 1	0
Blue clay and sand 10- 4	5
Shell deposit, including a thin layer of hard rock at 52 ft.	
Water above and below this rock comes to within ten	
feet of surface 45- 6	60
Blue marl with occasional beds of shells 3 or 4 feet thick	
and containing black to dark gray water-worn pebbles. 60-20	00
Medium coarse sand with numerous very small black grains	
or pebbles. A flow was obtained at this depth200-21	2
Limestone (sample)	
Blue marl and fine sands with inclusions of several thin	
strata of shell. (Sample)255-38	55
Very hard rock355-36	
Indurated gray sand and blue marl	
Rock, hard and soft strata with increase of flow upon pene-	
trating each hard stratum. No increase reported below	
460 feet. Driller reports the rock to be closer grained	
from 460 to 489.7 feet, and not containing much water418-48	39.7

CRANDALL.

Two wells are reported at Crandall, both of which are owned by Messrs. L. A. Davis & Brother. These wells are three inches in diameter and both are reported cased to a depth of 80 feet. One was drilled to a depth of 480 feet; the other to a depth of 450 feet. The water is reported to rise 35 feet above the surface. The water from one of the wells is used for the boiler supply at the sawmill and is said to form a hard scale. The other well is used for general drinking purposes.

EVERGREEN

Flowing wells are obtained at Evergreen postoffice, a village about four miles distant from Evergreen station on the Seaboard Air Line Railway. A well owned by Mr. L. L. Owens and drilled by Mr. D. C. Stafford in 1909 is about 500 feet deep. It is two inches in diameter and is reported cased 270 feet. The water is reported to rise 25 feet above the surface.

FERNANDINA

Fernandina, the county seat of Nassau County, is located in the northeastern part of the county, on Amelia Island. This island is thirteen miles long and is from one to three miles wide. The greater portion is low and flat, while other parts are gently undulating. The highest elevation on the island is to be found along the line of sand dunes bordering the ocean. The dune on which the lighthouse is placed reaches an elevation of about 55 feet above the sea.

The first flow of water in and near Fernandina is reported to be encountered at a depth of from 400 to 500 feet after drilling through a considerable thickness of sand and blue to greenish clay or marl. The water at this depth, as indicated by notes obtained from well drillers, comes from a sand stratum confined there by the overlying, very compact, blue to greenish clays.

The second water bearing stratum or chief source of supply is obtained at or about the depth of 600 feet. In the log of the

new well at the city water works limestone or what was termed by the driller, Mr. H. Walker, "water rock" was encountered at a depth of 556 feet. This was reported to consist of alternating hard and soft strata and the flow of water to increase with depth as each hard stratum was penetrated.

The first well drilled on Amelia Island was put down for the City of Fernandina by Messrs. Wade and Hampton in 1888. This well is located 5 blocks east of the city postoffice and is eight inches in diameter and was drilled to a total depth of 640 feet. It is reported cased 618 feet. At this depth an abundance of flowing water was obtained but as the well subsequently became filled with sand the flow decreased to such an extent that in order to get a sufficient amount of water to supply the city pumping had to be resorted to. Later the well was drilled deeper to a depth of 731 feet. The flow, however, is reported not to be as great as it was originally, although the deepening of the well increased the amount of flow to such an extent that the pumping of the water became unnecessary. This well is reported to have had a pressure of 14 pounds when first drilled in 1888. The following record of measurements of the flow of this well were kindly supplied by Mr. R. V. Nolan, superintendent of the City Waterworks

	Flow of well.
Date.	Gallons per day.
1890	. 1,152,000
1902,	641,832
1904	495,408
1905	. 440,564
1907	425,952
1909	408,000

In 1906 a second well was drilled for the city by Mr. H Walker. This well contains 120 feet of 10-inch casing; 356 feet of 8-inch casing; and 455 feet of six-inch casing and is drilled to a total depth of 733 feet. The head of the water in this we'l as shown by the pressure gauge January 28, 1910, was 14 pounds to the square inch or 32.3 feet above the surface elevation of the well, which is about 29 feet above sea, thus making a total head

of 61.3 feet above sea. The flow of this well in 1909 was 672. 000 gallons per day.

The following is a log of the new well at the City Waterworks as given by Mr. H. Walker, the driller:

Feet.
Sand 0-110
Medium hard rock110-126
Sand and clay
'Clay
Sand400-450
Green clay
Rock512-517
Blue clay
Limestone, termed "bed rock," with alternating hard and
soft strata556-733

A well three and one-fourth miles south of Fernandina owned by the Nassau Truck & Farm Company was drilled by J. W. Wiggins in 1909. This is a six-inch well, 650 feet deep and cased 442 feet. The first hard rock is reported at a depth of 500 feet. The pressure of this well was taken January 14, 1910. and was found to be 201/2 pounds or a pressure sufficient to cause the water to rise 47.3 feet above the surface.

The following is a log of this well as constructed from the notes kept and kindly made available by Mr. Walter Schucht, Superintendent of the company:

	Feet.
Muck	0- 3
Hardpan. A small flow just below this	3- 9
Sand	9-100
Blue clay. A good flow of water reported	00-200
Sand	00-400
Coarse sand and black pebbles4	00-500
Hard rock5	00-630
Limestone, hard and soft strata. Increase of flow upon	
breaking through each hard stratum	30-650

The following is an analysis of the water drawn from this well January 14, 1910. Analysis made for the State Survey in the office of the State Chemist, A. M. Henry, analyst:

Constituents.	Parts per million
Silica, (SiO ₂)	24.0
Chlorine, (Cl)	30.0
Sulphates, (SO ₄)	133.0
Phosphates, (PO ₄)	0.0
Carbonates, (CO ₃)	0.0
Bicarbonates, (HCO ₃)	195.0
Sodium and Potassium (Na & K)	30.0
Magnesium (Mg)	13.0
Calcium (Ca)	55.0
Iron and Alumina, (Fe & Al)	Trace
Loss on Ignition	130.0
Total dissolved solids	500.0

A well just across Amelia River and about two miles southwest of Fernandina was driven by James Jones for L. G. Hirth. The well is 94 feet deep, two inches in diameter and the water stands 7 feet below the surface.

The following is an analysis of the water from this well made by Dr. E. R. Flint, Chemist, University of Florida, Gainesville, Fla.:

Constituents.	Parts per million.
Free Ammonia	None
Albuminoid Ammonia	Slight Trace
Nitrites	Slight Trace
Nitrates	None
Chlorine	20.40
Total Solids	192.01
Organic and Volatile Solids	30.00
Hardness (CaCO ₃)	54.86
Permanent Hardness	None

HILLIARD.

Hilliard is located in northwestern Nassau County, on the Atlantic Coast Line Railroad, and about eight miles distant from the St. Marys River. No flowing wells have been reported in this part of the county, the elevation being too great. The elevation of the depot at Hilliard as recorded by the Atlantic Coast Line Railroad is 66 feet. Mr. D. W. Griffing has kindly fur-

nished several points of elevation covering the property of the Cornwall Farm Land Company

The only deep well reported at Hilliard is owned by The Cornwall Farm Land Company and was drilled by J. W. Wiggins in 1909. It is an eight-inch well, $648\frac{1}{2}$ feet in depth and cased about 400 feet. The elevation at the well is somewhat above the depot and the water is reported to rise to within 12 feet of the surface. Hard rock was encountered at 300 feet and the principal supply of water is reported as being obtained from the depth of 400 feet. The following is an analysis of the water from this well. Analysis by the Chemical and Engineering Company, 35 Kinzie Street, Chicago, Ill.:

Constituents.	Parts per million.
Organic Matter	37.0
Silica	36.0
Calcium Carbonate (Lime 91. parts per mil.)	151.0
Calcium Sulphate	16.0
Magnesium Sulphate	105.0
Magnesium Chloride	40.8
Sodium Chloride (common salt)	20.3

ITALIA.

One deep well is reported at Italia. This well is now owned by McLeod Bros, & Airth and was drilled in 1905. It is a 2-inch well and reached a total depth of 430+ feet. It is reported cased 40 feet and to have a head of 30 feet above the surface.

KING'S FERRY.

Kings Ferry is located on the St. Marys River, about 30 miles up the river from Fernandina. One deep well owned by W. J. Carlton is reported from Kings Ferry. This well is two inches in diameter and about 400 feet deep and was drilled in 1909 by D. C. Stafford. The pressure of this well could not be ascertained but it furnishes a strong flow and was reported to rise more than 31 feet above the surface in a one-inch pipe.

LESSIE.

A deep well at Lessie, owned by J. R. Wilson & Company and drilled by D. C. Stafford, is reported to have a depth of 450 feet. It is a two-inch well and furnishes an abundant supply of water.

LOFTON.

The well of J. W. Rodgers at Lofton was bored in 1906 and is reported to have a depth of 510 feet. It is two inches in diameter and gives a good flow, but the height to which the water would rise above the surface was not learned. The water from the well is used for general domestic purposes and to supply the turpentine still.

DUVAL COUNTY.

LOCATION AND SURFACE FEATURES.

Duval County joins Nassau County on the south, and is separated from it by the Nassau River and its tributary, Thomas Creek. The St. Johns River flows through Duval County. The surface drainage from this county is carried off largely through these rivers and their tributaries.

The surface is in general flat or but slightly rolling. The surface elevation rises gradually from sea level. The highest elevation reached is found in the southwestern part of the county, where the "Trail Ridge" forms part of the boundary. A narrow strip along this part of the county exceeds 100 feet in elevation. With this exception practically all parts of this county lie below the 100-foot contour line, while much of the area lies below the 25-foot contour line.

The elevations in Nassau and Duval Counties have been obtained from various sources. An important line of levels extending from Trout Creek across Nassau and Duval Counties in a southwesterly direction, made during the summer of 1909, in connection with a preliminary survey for a ship canal across Florida, were kindly made available for this purpose in the office of the

United States Engineer at Jacksonville. Similar surveys made by the same office in 1879 supplied elevations from Fernandina

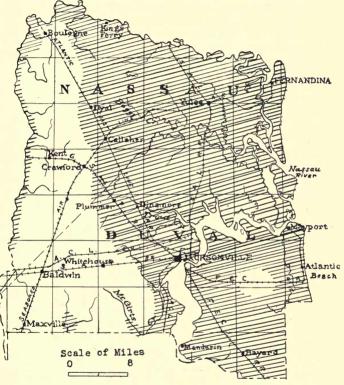


Fig. 6.—Map of flowing area of Nassau and Duval Counties. The area in which flowing wells can be obtained is indicated by shading.

to Maxville and at various points along the St. Marys River.*
In addition much information as to elevations has been obtained

^{*}Annual Report of the Chief of Engineers for 1880, pp. 973-1010.

from the profiles of the several railroads crossing this section, particularly the Seaboard Air Line from Jacksonville to Maxville, the Florida East Coast from Jacksonville to Mayport and the Atlantic Coast Line from Jacksonville to the St. Marys River.

From Jacksonville westward the rise in elevation, as shown by the profile of the Seaboard Air Line Railway, is very gradual to a point three miles west of Jacksonville where an elevation of 27 feet is reached. From this summit the elevation drops off slightly, the elevation of Cedar Creek being 17 feet. Beyond Cedar Creek the elevation rises more rapidly. Marietta station is approximately 60 feet above sea. The summit of this rise is reached two miles west of Marietta where the elevation is 94 feet. White House station is 82 feet above sea. Beyond McGirts Creek one and one-half miles an elevation of 91 feet is reached. From this point there is a very gradual slope to Baldwin, this latter place being 86 feet above sea. South from Baldwin the contour rises in general, reaching an elevation of 93 feet at Maxville and 100 feet one-half mile beyond the county line.

The line of levels run by United States Engineers extends from Trout Creek, passing just to the south of Brandy Branch station, or Bryceville postoffice. The summit elevation in Nassau and Duval Counties along this line occurs about four miles northeast of Brandy Branch, where an elevation of 90 feet is recorded.

WATER-BEARING FORMATIONS.

The deeper wells in Duval County reach and terminate in the Vicksburg Limestone. This is known to be the case at Jacksonville, at which place the Vicksburg is reached at approximately five hundred feet from the surface. The wells at Jacksonville, the deepest of which reach a total depth of something over a thousand feet, do not, so far as the records show, pass entirely through the Vicksburg.

The formations lying above the Vicksburg are less characteristic lithologically and are not easily differentiated. The surface deposits include both recent and Pleistocene material. During a part of Pleistocene time this section of the State stood at a

lower level than at present, permitting the ocean to extend inland some distance beyond the present coast line. Conrad* has re-

*Conrad, T. A., Am. Journ. Sci. (2) 11, 38, 1846. corded the occurrence of marine shell deposits of post-Pliocene age along the banks of the St. Johns River at an elevation of from ten to fifteen feet above the present high tide. Conrad also reports a similar post-Pliocene deposit about one-half mile from the bank of the river near the ancient village of Hasard. Marl deposits are said to occur near the mouth of the St. Johns River, on the banks of Ft. George Inlet. That the depression of the coast during Pleistocene time was general is indicated by the records from several other localities.

Beneath the Pleistocene, Pliocene deposits probably occur over some parts of the county. The total thickness of the Pleistocene and Pliocene, if both are represented, is, however, not great, as the fossiliferous Miocene limestone was reached at Jacksonville, in the boring at the city well, at a depth of 33 feet.

AREA OF ARTESIAN FLOW IN DUVAL COUNTY.

The area of artesian flow in Duval County is indicated on the accompanying map by shading. As will be observed the flowing area borders the Atlantic coast, Nassau and St. Johns Rivers and extends some distance inland, following each smaller stream and tributary. The wells in western Duval County are nonflowing. A topographic map of this section would assist in determining flowing and non-flowing sections, since the flow is to a large extent correlated with elevation. It is to be borne in mind, however, that artesian water depends primarily upon the structure of the underlying formations and these formations are liable to variations of which there is no surface indication. For this reason, while the map indicates the area of probable flow the exact limits of the area are best determined by drilling.

LOCAL DETAILS.

BALDWIN.

Baldwin is located on the Seaboard Air Line Railway, nineteen miles west of Jacksonville. The elevation is approximately 86 feet above sea. Three wells have been drilled at or near Baldwin. The deepest of these, located at the Atlantic Coast Line Railroad crossing, one-half mile north of Baldwin, is reported to reach a total depth of 580 feet and is cased 511 feet. A second well nearby reaches a depth of 100 feet. A third well located at Baldwin reaches a depth of 92 feet. All of these wells are non-flowing, although the water rises within a few feet of the surface. The distance at which the water stands from the surface in the deep well is not reported beyond the statement that the well is nonflowing.

BAYARD,

Bayard is located on the Florida East Coast Railway, fifteen miles south of Jacksonville. The elevation of this place is approximately 22 feet above sea. Flowing water is obtained at Bayard, one well having been put down for the Carter-Lucas Co. This is a three-inch well, reported to have been drilled to a depth of 280 feet. The water here will rise at least fifteen feet above the surface.

JACKSONVILLE.

The large number of wells occurring at Jacksonville precludes the possibility of listing or describing all. Probably not less than five hundred flowing wells occur in or near this city.

The first flow obtained at Jacksonville, according to the records of the city well, was a light flow from a depth of 487 feet. A large flow, however, is not obtained until the drill enters the Vicksburg limestones, at a depth of about 524 feet. After reaching the Vicksburg the flow increases upon breaking each compact layer. At a depth of 632 feet the flow in the new city well was found to be one million gallons per day. At a depth of 980 feet

the same well supplied a flow of two million gallons per day.

The material penetrated in the drilling at Jacksonville, for a depth of about 500 feet, consists largely of clays, sandy clays, and sands with some fossiliferous limestone and some shell deposits. From about 500 to 524 feet the record shows considerable dense hard rock. After penetrating this stratum the limestones of the Vicksburg group are reached.

The water supply for the city of Jacksonville is obtained from artesian wells. At present ten artesian wells are in use. Details as to the depth and construction of these wells will be found in the table of well records Nos. 1 to 10. The log of well No. 6 was given in the Second Annual Report, p. 109. The samples from which this log was made were obtained by Superintendent Ellis by first drilling an eight-inch well, and afterwards reaming it out to a ten-inch well.

The following is the record of the new city well at Jacksonville. Sample of drillings from this well, together with notes on the materials penetrated, were kindly kept by Mr. S. L. Hughes of the Hughes Specialty Well Drilling Company, of Charleston, South Carolina:

Filled ground and sand	. 0	- 15
Sand with some clay	15	- 33
Sandy limestone, yellowish or light buff in color	. 33	- 37
Light colored clay marl	37	- 70
Blue sticky clay with black phosphatic pebbles	70	-100
Marls, usually green or olive green in color containing variable amount of sand, and clay. Black phosphatic		
pebbles together with some shell fragments occur		
throughout the marl. Occasional thin layers of		
light colored limestone are reported within this		
interval. First flow of water at 270 feet 5 gallons		
per minute	100	-320
Buff clay resembling fuller's earth mixed as seen in the		
sample, with green sandy marl		-340
Greenish and sandy clayey marl		-390
Indurated sands or sandstones	390	-396
Greenish sandy marls	396	-415
Light colored limestone	415	-420
Greenish calcareous sandy clay	420	-434
Dark colored hard sand rock	434	-435

435 -455
Olive green calcareous sandy Clay
Tight condy mar! 455 -45572
Green sandy marl
Dark sandy clay
Very hard dark or gray sand rock
Silicified and very hard shell rock with siliceous phos-
Silicified and very hard shell lock with sinceous phos
phatic pebbles. After passing through this rock the
flow is increased to 112 gallons per minute, tem-
perature 71 degrees F
Light colored mail
Hard rock 500 -506
Light gray sandy calcareous rock with black phosphatic
pebbles 506 -510
Feet.
Light colored fossiliferous limestone (Vicksburg). Upon
reaching this formation the flow is increased to 200
gallons per minute. At 625 to 635 feet the harder
gallons per limite. At 025 to 055 feet the natural
stratum was drilled through, which flowed 500 gallons
per minute, temperature 74 degrees F. At 680 feet the
water pressure measured, as shown by the gauge, 12
pounds510-680
Limestone, prevailing brownish in color, and as a rule hard-
er than above. Occasional thin layers of marl and
shell. Slight increase of flow at 780, water pressure at
900 feet 15 pounds; flow about 900 gallons per minute;
temperature 74 degrees F680-900
Limestone similar in character to above, but as a rule not
so hard. Flow at 980 feet, 1,500 to 2,000 gallons per
minute900-980

The Vicksburg Limestone was reached in this well at a depth of about 510 feet. The first 170 feet of the Vicksburg is prevailingly light colored or white and fossiliferous. Below 680 feet the limestone is as a rule brownish in color, compact and harder in texture and not so fossiliferous. The amount of flow, the pressure and the temperature increased as the deeper layers of the Vicksburg Limestone were penetrated.

The formations lying above the Vicksburg Limestone can scarcely be differentiated. The Jacksonville formation, Miocene, is reached at the depth of 33 feet. At about 320 feet some clays resembling fuller's earth were obtained. At from 415 to 420 feet light colored clayey limestones were encountered. With these

exceptions the interval from 37 feet to 510 feet consists largely of an olive green sandy mark.

An analysis of the water of the public supply at Jacksonville was made in 1898. Analyst, Albert Leeds, Stevens Institute of Technology. The analysis is as follows:

87	•	Grains per	Parts per
		Grains per	rarts per
Constituents.		U. S. gallon.	million.
Silica and insoluble :	matter	0.729	12.497
Alumina		0.047	8.057
Carbonate of lime .		3.866	66.274
Sulphate of lime		4.053	69.480
Sulphate of magnesia	a	2.927	50.177
Sulphate of soda		5.843	100.166
Chlorides of soda		4.811	82,474
Free ammonia			0.143
Albuminoid ammonia	a		0.044

The following is an analysis of the water from the well of the Florida East Coast Railway, at South Jacksonville. The well is 651 feet deep. The analysis is by the American Water Softener Company, Philadelphia, Pa.

	Grains per	Parts per
Constituents.	U. S. gallon.	million.
Calcium carbonate	32	5.48
Calcium sulphate	15.00	257.14
Calcium chloride	1.23	21.08
Magnesium carbonate	5.94	101.82
Sodium chloride	0.69	11.82
Free carbon dioxide	0.41	7.02
Iron, aluminum and silica	0.09	1.54
Incrusting solids	22.59	387.26
Non-incrusting solids	0.69	11.82
Total solids	25.90	444.00

The following is a log of this well obtained through Mr. G. A. Miller, as reported by the driller, Mr. H. Walker.

	Feet.
Dark sand	0- 6
Clay	6- 7
White sand	7- 9
Gravel	9- 13

White clay 13- 17
White clay and sand 17- 31
Hard rock, clay and rock
Blue clay 35- 50
Rock 50- 56
White clay and sand
Sand 89- 90
White clay and sand
Soft rock
Blue clay and sand
Loose sand
Tough clay and sand
Sand
Loose sand
Clay and sand
Clay
Clay and gravel
Rock
White clay
Rock and clay
Hard rock
Clay with thin strata of soft rock414-451
Clay and sand
Blue clay
Sand
Soft sandy rock
Sand
Loose sand
Hard rock 501-510
501t rock 510-526
Limestone

MANDARIN.

Mandarin lies within the flowing area which borders the St. Johns River. Several wells have been put down in this section. A well near Mandarin, drilled by H. Walker for J. D. Mead, reached a total depth of 600 feet. This well is cased 377 feet and the water is reported as rising 60 feet above the surface.

MANHATTAN BEACH.

The following is a log of a well drilled at Manhattan Beach by H. VanDorn for the Florida East Coast Railway. This well flows 15,000 gallons per hour through a two-inch pipe. The pressure at the surface is 20.5 pounds. The record has been obtained through Mr. G. A. Miller.

Feet.	
Sand 0- 35	,
Clay 35- 47	,
Clay resembling soapstone 47- 90)
Clay 90-140)
Seft rock	,
Clay)
Soft rock)
Sand and clay	,
Sand)
Clay	,
Rock)
Clay)
Rock	2
Sand and clay)
Rock	
Clay)
Sand and clay)
Clay340-350)
Sand350-357	1
Clay357-361	
Rock	}
Clay)
Rock)
Clay370-385	5
Rock385-387	,
Sand387-390)
Rock390-391	L
Clay391-395	í
Rock395-396	5
Clay396-398	3
Rock398-404	
Water-bearing rock404-450	
Soft rock450-490	
Hard rock490-520	
Water-bearing rock520-540)

Hard and soft rock in thin layers	5
Soft rock	555-576
Hard and soft rock in thin layers	576-600

MAXVILLE.

Maxville is located on the Seaboard Air Line Railway, near the southwestern corner of Duval County. The elevation at this point is, according to the profiles of the railroad, about 93 feet above sea. A well drilled at this place in 1902 for Mr. R. V. Douglass is reported to have reached the depth of 650 feet. This well is non-flowing.

MAYPORT.

The following is an analysis of the water of the well of the Florida East Coast Railway at Mayport. The well is 600 feet deep and has a pressure of 22 pounds. Analysis by the American Water Softener Company, Philadelphia, Pa.:

Constituents,	Grains per U. S. gallon.	Parts per million.
Calcium carbonate	3.57	60.20
Calcium sulphate	5.33	91.37
Magnesium carbonate	4 46	76.45
Sodium carbonate	70	12.00
Sodium chloride	2 45	42.00
Free carbon dioxide	32	5.48
Iron, aluminum and silica	33	5.65
Incrusting solids	13 69	234.68
Non-incrusting solids	3.13	53.65
Total solids		310.11

The following is a log of a well drilled at Mayport by B, S. Partridge for the Florida East Coast Railway. The record has been made available by Mr. G. A. Miller:

Sand and muck	Feet.
Sand and muck	0- 57
Rock Sand	57- 61
Sand Rock	61- 85
	Q# 079

Clay 87-160
Rock
Clay
Sand
Clay
Rock
Sand280-350
Rock
Clay
Rock
Clay
Rock
Sand
Clay
Soft rock
Soft water-bearing rock
Hard rock

ST. JOHNS COUNTY.

LOCATION AND SURFACE FEATURES.

St. Johns County lies in northeastern Florida, bordering the Atlantic Ocean. On the north it joins Duval County and on the south Volusia County. The western boundary is formed by the St. Johns River. The county has a total length of sixty miles. In width it varies from eighteen to twenty-four miles. The total area is approximately 1,000 square miles.

Owing to the location of St. Johns County between the St. Johns River, on the west, and the Atlantic Ocean, on the east, no great variation in elevation is to be expected. It is probable, however, that small areas in the interior of the county lie above the fifty-foot contour. In passing from St. Augustine to Jacksonville, levels made by the Florida East Coast Railway show near the county line an elevation over a small area of 57 feet. The greatest elevation recorded between St. Augustine and Hastings is in the vicinity of Hurds. A line of levels run from the coast at St. Augustine, at the instance of Mr. B. A. Carter, gave for Hurds an elevation of thirty-eight feet. Levels obtained from the U. S. Engineers' Office, Jacksonville, Florida, give, for a point

a short distance east of Hurds, a level of thirty-six feet. From East Palatka south information regarding elevation is unfortunately very deficient. From the fact that such wells as have been put down at Dinner Island, Espanola, Bunnell and Dupont, are non-flowing, it is probable that this part of the county is above the twenty-five-foot contour line, and parts of this area may, in fact, approach or exceed the fifty-foot contour. Along the west side of the county bordering the St. Johns River areas varying in width from 3 to 10 or more miles lie below the twenty-five-foot contour line.

WATER-BEARING FORMATIONS.

The Vicksburg Limestone is the chief source of the artesian water supply of St. Johns County, although a small flow is probably obtained before reaching this formation. The Vicksburg Limestone consists of alternating hard and soft fossiliferous strata and is usually easily recognized. At St. Augustine, according to determinations made by Dr. W. H. Dall,* fossils characteristic of this formation were obtained from a depth of 224 feet. At Hastings, 17 miles southwest of St. Augustine, well records indicate that a limestone similar in character to the Vicksburg is reached at a depth of from 175 to 200 feet. At Orange Mills, in Putnam County, 3 miles southwest of Hastings, Orbitoides, apparently representing some member of the Vicksburg group, were obtained at a depth reported at 110 feet. At the time the sample was received the well was drilled to a total depth of only 130 feet. Toward the northern part of St. Johns County the Vicksburg Limestone probably dips deeper, since, at Jacksonville, this formation is first reached at a depth of about 524 feet.

The superficial material in this county is largely Pleistocene and recent sands together with Pleistocene and recent shell deposits. Oscillations of level have affected the surface elevation, and consequently the relative extent of land and water area in this county within comparatively recent time. That this part of the State stood at a lower level during a part of Pleistocene time is evident from the occurrence of marine shell deposits of Pleisto-

cene age at some distance inland and at an elevation of several feet above the present sea level. Oyster banks, probably of Pleistocene age, are exposed along a small drainage ditch on the farm of A. W. Corbett, four miles southwest of St. Augustine, at an elevation of at least 15 to 20 feet above the present sea level. That this depression during Pleistocene time was general for this part of the State is indicated by the evidence already given.

The identification of the formations lying above the Vicksburg limestones and beneath the superficial sands, from well records alone is a matter of difficulty. This interval in St. Johns County is occupied largely by clays, although some sand, shell and rock strata occur.

AREA OF ARTESIAN FLOW IN ST. JOHNS COUNTY.

The areas of flowing and non-flowing wells in St. Johns County are indicated on the accompanying map.

The shaded lines on the map indicate the area in which flowing artesian wells can be obtained in this county. As will be seen from the map the flowing area borders the Atlantic coast and the St. Johns River, and has a width along the coast and also along the St. Johns of from two or three to eight or ten miles. The flowing area extends inland following the streams. So far as present records show, a narrow strip extending north and south through the central part of the county is non-flowing. A fresh water spring is reported to occur in the ocean opposite Matanzas. Springs of this character represent the natural escape of the underground waters into the ocean.

LOCAL DETAILS.

ANASTASIA ISLANĐ.

A six-inch well, drilled in 1895, at South Beach, on Antastasia Island, reached a total depth of 260 feet. A strong flow of sulphur water was obtained from this well.

^{*}U. S. Geol. Surv. Bull. 84, p. 125, 1892,

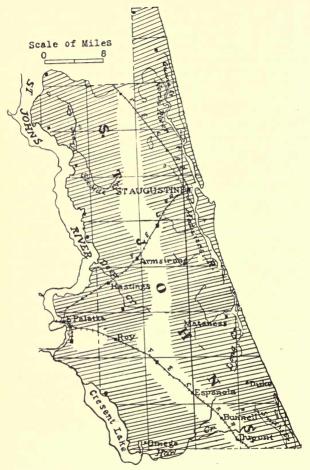


Fig. 7.—Map showing the area of artesian flow in St. Johns County. The area in which flowing wells can be obtained is indicated by shading.

ARMSTRONG.

Flowing wells have been obtained in the vicinity of Armstrong. A four-inch well, drilled in 1908, for J. W. Williams by N. H. Monck, reached a total depth of 200 feet. This well is cased 70 feet and the water is reported to rise 12 feet above the surface.

BUNNELL.

An effort was made in 1909 to obtain a flowing well at Bunnell. A five-inch well was drilled at this place by Mr. N. H. Monck for Messrs. Lambert & Moody. This well was cased to a depth of 130 feet and is reported to have been drilled to a total depth of 300 feet. A flow is not obtained in this well, although the water rises to within about two feet of the surface.

A second well owned by Messrs. Lambert & Moody, drilled by Bellough & Melton in 1910, is 128 feet deep. The following log of this well was supplied by the drillers:

		Pe	eet.
Surface material and sand		. 0	- 45
Blue clay		. 45	- 90
Black material looking like gunpowder or pepper	r	. 90	-109
Blue clay		.109	-119
Shell and sand		.119	-124
Blue hard rock		.124	-1241/2
Cavity 6-inch, sand and shell. Water rises to v	vithin		
1.4 feet of surface		.1241	2-125
Blue hard rock, more water, with same head; dr	rilling		
stopped in second cavity		.125	-128

DINNER ISLAND

A record of one well has been obtained at Dinner Island. This is a three-inch well drilled by Mr. H. Mervin for Padgett & Company. It has a total depth of 200 feet and does not flow, although the water is reported to rise to within two feet of the surface.

ELKTON.

Flowing wells are obtained at Elkton. A five-inch well drilled by N. H. Monck, in 1908, on the Middleton farm, reached a total

depth of 260 feet. The well is cased 100 feet and the principal supply of water comes from a depth of 200 feet. The water is reported to rise five feet above the surface.

ESPANOLA.

A few wells occur in or near Espanola. The wells immediately in the town do not flow. Flowing wells are obtained, however, from one to five miles south, along Haw Creek.

FEDERAL POINT

Federal Point lies within the flowing area bordering the St. Johns River. A considerable number of wells have been drilled in the vicinity of this place. The material encountered here, to the depth of about 125 feet, consists largely of clays. Water is obtained at a depth of from 200 to 250 feet, the wells terminating in limestone.

The following is a partial log of the well of Messrs. Hubbard and Hart, one-fourth mile northwest of Federal Point. This is a six-inch well drilled by Lloyd Crary in 1889. The well has a total depth of 225 feet and is cased 60 feet. The water is reported to rise twenty feet above the surface or about thirty feet above sea level. The principal supply is obtained at a depth of two hundred feet.

A mixed sample contained material similar to above
with addition of gray sandy clay160-168
Buff colored sandy clay
White granular fossiliferous limestone

This well probably reaches the Vicksburg group of limestones, as indicated by sample, from the depth of 180 to 225 feet. The material obtained between the depth of 168 and 180 feet may represent the Upper Oligocene, as it has certain lithological resemblances to parts of the Alum Bluff formation. The conglomerate material from 145 to 160 feet together with a part of the overlying clays probably represents the Jacksonville formation of the Miocene.

HASTINGS.

Hastings is in the western part of St. Johns County, and is located on Deep Creek, a tributary to the St. Johns River. The town site is inland about three miles from the river. The elevation at Hastings, at the residence of T. H. Hastings, is, according to the U. S. Coast and Geodetic Survey, 8 feet above sea.

A considerable number of artesian wells have been put down at and in the vicinity of Hastings. Record has been obtained of fifty-one wells within a radius of three miles of the town.

Wells at Hastings are largely used for irrigating purposes. The average depth of the wells now in use is 148 to 272 feet, although some reach a greater depth. Most of the wells are 4 to 6 inches in diameter. The length of casing used in the wells is variable, ranging from 65 to 170 feet.

Aside from the superficial soil and sand the material penetrated at Hastings to a depth of about 170 feet consists largely of clays although some water-bearing sands are reported and a shell stratum at a depth of 60 to 62 feet is specially mentioned.

At a depth of 170 to 180 feet a dark colored, very hard stratum occurs. This rock appears from the well records to be similar in character to the rock found at St. Augustine at a depth of 170 to 180 feet. After passing through this stratum the wells penetrate limestone consisting of alternating hard and soft strata, the

flow increasing as each hard stratum is penetrated. This limestone, probably representing the Vicksburg group, has been penetrated at Hastings about 200 feet or to a total depth of 365 feet.

Of the many wells at Hastings it is possible to give an individual record of only a few. The following is a log of the well of F. R. Allen, kindly supplied by the driller, Mr. H. Walker. This is a 6-inch well, located three miles southeast of town. It was drilled in May, 1908, and is used for irrigating purposes.

	F	eet.
Yellow clay	0	- 6
Blue clay	6	- 60
Shell stratum	60	- 64
Clay	64	-160
Soft rock	160	-165
Clay		-171
Rock supplying small flow	171	$-171\frac{1}{2}$
Limestone		/ ₂ -183
Shell and limestone	183	-245
Material not reported	245	-300

The following is a partial log of the well of Henry Bugbee taken from the notes kept by I. C. Peck. This is a four-inch well drilled in 1902 and located two and one-half miles south of Hastings. The well has a total depth of 257 feet and is cased 178 feet. It is used for irrigating purposes.

	Feet.
Surface material, soil and sand	0- 6
Mostly clay, some sand at 32 feet. Material from 38	
to 70 feet not reported	6-186
Seven feet of very hard rock through which it was	
possible to drill only a few inches a day1	86-193
Porous limestone from which flowing water is obtained1	93-208
Soft limestone, flow increasing with depth2	08-257

HOLY BRANCH.

Flowing wells are obtained at Holy Branch. A four-inch well drilled in 1908 for Charles Slater by N. H. Monck reached a total depth of 240 feet. This well is cased 200 feet and the water is reported to rise 12 feet above the surface.

The following is a log of the well of Mr. G. A. Beach, supplied by the driller, Mr. Frank Bartlett. This is a 4-inch well, 257 feet deep, and is cased 184 feet:

	Fe	et.
Surface sand and soil	0	- 6
Red clay	6	- 20
Hardpan, black	20	- 24
White sand	24	- 30
Blue clay and marl	30	- 33
Sand and shell	33	- 53
Blue clay and marl	53	- 59
Shell and sand, water rises to within nine feet of surface	59	- 80
Blue clay and marl	80	-130
Black quicksand, water plentiful	30	-146
Very hard blue marl and clay	46	-180
Black quicksand, water-bearing	.80	-186
Blue marl	86	-196
Very hard black flint, water flows	96	-1971/2
Hard rock, flint and more water	971/	2-2011/2
Softer limestone, more water with increase of depth2		

HURDS.

Hurds is located on the Florida East Coast Railway, seven miles southwest of St. Augustine. The elevation at Hurds, according to levels made for Mr. B. A. Carter, is 38 feet above sea. The deepest well recorded at this point is 385 feet. This is a 4-inch well and was drilled in 1906. It was cased to a depth of 160 feet. This well does not flow, although the water rises to within five feet of the surface. The well was drilled for B. A. Carter by I. C. Peck.

MOULTRIE.

Flowing wells are obtained at Moultrie. A six-inch well put down here for the St. Augustine Industrial School reached a total depth of 300 feet. The water at this locality is reported to rise 32 feet above sea level. The surface elevation in the vicinity of Moultrie varies from 0 to about 30 feet above sea.

PICOLATA.

Picolata is in the extreme western portion of St. Johns County, almost due west of St. Augustine, on the St. Johns River. A four-inch well, drilled about the year 1890, is now owned by R. H. Bohn. The depth was reported to be about 300 feet. The pressure of this well was taken January 10, 1910, and was found to be 15 pounds. The elevation of the well is approximately 8 feet above the river. This, together with a pressure of 15 pounds, would give the well a head of 42.65 feet above the level of the water in the St. Johns River.

RIVERDALE.

Riverdale is a settlement along the St. Johns River, in southwestern St. Johns County. At this place several artesian wells have recently been drilled. A well 302 feet deep was sunk in 1909 by Mr. R. C. Walker for the Riverdale Land Company. This is a six-inch well and is cased 107 feet. The well is reported to have a head of 33½ feet above the surface and the surface elevation above the St. Johns River is estimated to be 8 feet, which gives the well a total head of 41½ feet. The first rock encountered was at a depth of 175 feet, and at this depth the water was found to be under sufficient pressure to rise to the surface. An increase in the flow of water was reported at a depth of 190 feet.

Mr. R. C. Walker completed on February 1, 1910, a well for Mr. J. D. Clark. This well is six inches in diameter, 318 feet deep, and is cased 136 feet. At the depth of 174 feet a one-foot stratum of bluish, clayey limestone was encountered. An increase in water is recorded at the depth of 200 feet, from which depth the first flowing water is reported. The well samples indicate that this flow comes from a very hard, bluish colored rock and water-worn small pebbles. Immediately on passing through this stratum, which was 19 feet in thickness, the Vicksburg Limestone was reached, as is shown by the presence of Nummulites. This determination was made from a very complete set of samples of the drillings from this well, kindly saved by the driller, Mr.

R. C. Walker. This limestone was penetrated for nearly 100 feet, the total depth of the well being 318 feet. The following is a log of this well, constructed from the notes and the samples sent in by Mr. Walker:

	Feet.
Surface sand, yellow in color. Soft water	0- 18
Light gray sands	18- 30
Dark gray sands, partly indurated; some clay	30- 44
Shell, sand and gravel	44- 55
Very dark (almost black) marl, similar in appearance	
to Miocene marls, including shell fragments	55- 63
Light greenish sandy marl	63-80
Dark green marl, small shark's tooth observed	80-100
Gray sand and shell fragments; water	100-112
Gray sand and shell, water, shark's tooth, also minute	
black phosphatic pebbles	112-133
Blue clayey marl	133-135
No sample	135-153
Blue marl with inclusions of black phosphatic pebbles1	153-174
Blue clayey limestone; water-bearing	74-175
Dark green marl with some black phosphatic pebbles1	75-200
Very hard bluish colored rock, and water-worn small	
pebbles; water commenced to flow upon pene-	
trating this stratum2	200-219
Limestone, Vicksburg as indicated by the presence of	
Nummulites2	219-318

ROY.

Roy is located on the Florida East Coast Railway, about six miles inland from the St. Johns River. One deep well is reported from this place. This is a four-inch well drilled by Mr. S. I. Killingsworth for Mr. L. J. Campbell. The well has a total depth of 298 feet and is cased 150 feet. The flow is reported to rise four feet above the surface.

ST. AUGUSTINE.

St. Augustine, the county seat of St. Johns County, is located on Matanzas Bay. An abundance of flowing water is obtained at this place. Probably not less than 100 wells occur in and near St. Augustine. Of this large number it is possible to mention only a few.

The first considerable flow in and near St. Augustine is obtained at a depth of from 170 to 180 feet after drilling through a five- or ten-foot stratum of dense hard rock. The material penetrated before reaching this hard rock stratum consists largely of sand near the surface, followed by blue clays with some shell and occasional thin layers of rock. A shell stratum often described as "coquina" occurs at a depth of about 60 feet.

The material below the depth of about 180 feet consists of alternating hard and soft strata, largely limestones, with probably occasional flints. The flow of water increases as the limestone is penetrated. The chief large increase of flow occurs at a depth of about 520 feet and most of the wells at St. Augustine terminate at this depth.

Water for the city of St. Augustine is obtained from two artesian wells located about one mile north of the city. Well No. 1 was drilled in 1897 by Mr. Hugh Partridge and had originally a depth of 371 feet. About 1903 this well was deepened to a total depth of 550 feet. The well is 12 inches in diameter for 354 feet; 9 inches for 17 feet, and four inches for 179 feet. It is reported cased to a depth of 100 feet. The head of the water is given as 33 feet above the surface or about 38 feet above sea level. The flow of the well when first drilled in 1897 was 2,396,000 gallons per day (1,664 gallons per minute).

Well No. 2 is a 10-inch well and has a total depth of 500 feet. It is cased about 140 feet. The head of the water is the same as well No. 1 or about 38 feet above sea. The total flow of this well is not recorded. This well was drilled in 1903 by Mr. Horace Walker.

The water system at St. Augustine is now owned by the city. Formerly the city was supplied by five artesian wells, the system then being under private ownership. These wells were located in various parts of the city. They vary in depth from 250 to 500 feet and range from 6 to 8 inches in diameter. The first of these wells was drilled in 1884. They are now in use as private wells.

Several wells have been drilled at St. Augustine to supply water to the Ponce de Leon and other hotels of the Florida East Coast Hotel Company. One of these, commonly known as the Ponce de Leon well, reached a total depth of 1,440 feet, and is the deepest well in St. Johns County. The following log of this well has been made up from records kindly supplied by Messrs. McGuire & McDonald, under whose direction the well was drilled, supplemented by a partial set of samples from the boring. The original intention was to go to a depth of about 3,000 feet in the expectation of obtaining warm water. The well was begun November 27, 1886, and drilling continued until February 24 of the following year. Owing to delay caused by the loss of the drill, boring was finally discontinued at the depth of about 1,440 feet.

	Fe	et.
Sand. Temperature of the water at 35 feet, 60 degrees F.	0-	35
Sand, with some shell	35-	50
Blue clay	50-	57
Shell	57-	65
Sand	65-	76
Indurated clay and sand	76-	95
Blue clay and black sand, pieces of hard stone. Tem-		
perature of the water 72 degrees at 110 feet, 74		
degrees at 170 feet. Head 32 feet above sea. Sul-		
phur water, 50 gallons per minute at 170 feet	95-	170
Hard rock. Temperature of water 76 degrees at 177		
feet. Flow 350 gallons per minute at 177 feet	170-	177
Limestone. Flow 1,800 gallons per minute at 350 feet	177-	350
Limestone. Temperature of water 76 degrees at 410		
feet. Flow of 2,083 gallons per minute at 410 feet	350-	410
Limestone	410-	495
Dense light brown limestone. Temperature of water 79		
degrees at 520 feet. Head 42 feet above sea at 520		
feet. Flow of 4,860 gallons per minute at 520 feet	495-	520
White "chalk," green clay, dark porous limestone	520-	557
Limestone	557-	675
Hard rock	675-	685
Limestone	685-	770
Limestone, grav to light vellow	770-	960
Thin stratum of hard limestone, followed by limestone		
similar to above. Temperature of water 80 degrees		

at 1,110 feet. Flow of 6,075 gallons per minute at	
1,110 feet	960-1110
Hard rock, said to be sandstone, with some flint	1110-1140
Material not recorded	1140-1170
"Sandstone," followed by limestone. Temperature of	
water 85 degrees at 1,225 feet	1170-1225
Limestone, as above	
"Sandstone." Sample not seen	1278-1293
Fossiliferous limestone	
Fossiliferous limestone, easily penetrated. Temperature	
of water 86 degrees at 1,340 feet	1340-1390
Denser limestone	

The following is a log of the well of Mr. W. J. Sherman. This well was drilled by the owner in 1886 and is 210 feet deep. It is two inches in diameter and is cased 110 feet. The head is reported to be 32 feet above sea and the flow about 80 gallons per minute:

]	Feet.	
Sand	. 0	- 5	
Clay	5	- 6	
White quicksand		- 0	
Clay	ь	- 11	
Clay	11	- 11	1/2
Coarse pebbles and some shells.	115	6- 43	
Coarse gray to greenish sands, water-bearing; slight flow	43	- 45	
white plastic clay and fine sand	45	- 90	
Greenish clay, very compact	90	-142	
Hard rock	140	-143	
Greenish clay with a mixture of black sand.	142		
Hand and a diskure of black sand	143	-172	
Hard rock; water rises 32 to 37 feet above sea	172	-180	
White chalk rock (probably Vicksburg)	180	-210	

SWITZERLAND

Switzerland is located in the area of artesian flow on the St. Johns River, in the northwestern part of St. Johns County. Wells at this locality reach a depth of from 350 to 500 feet, and the water is reported to rise 29 to 30 feet above the surface.

YELVINGTON.

Records of two wells have been obtained from and near Yelvington. Well No. 1 is located near Yelvington depot and is

owned by E. L. Campbell. This well was drilled by Frank Bartlett in 1909 and reached a total depth of 352 feet. It is reported as having 95 feet of four-inch casing. The head of this well was measured December 11, 1909. The water was found to stand at this time 7½ feet below the surface.

Well No. 2 is located one mile west of Yelvington depot. It is a four-inch well and is owned by Campbell & Killingsworth. This well was drilled in 1907 by S. I. Killingsworth and is reported to be 300 feet deep and cased 180 feet. The water is said to stand two feet below the surface.

CLAY COUNTY.

LOCATION AND SURFACE FEATURES.

Clay County has a varied topography. The eastern portion, bordering the St. Johns River, is low and flat and consists largely of open pine woods. Extending westward from the river the elevation rises and the country becomes more rolling. The county is intersected by a number of streams, the largest of which is Black Creek, a tributary to the St. Johns River. This stream is navigable for small boats to or above Middleburg, at which point it divides, forming the north and south forks. The north fork rises in Lake Kingsley, and with its tributaries drains the northwestern part of the county. The south fork rises in Blue Pond and other lakes and drains the central part of the county. In the southwestern part of the county many small lakes occur.

The elevations in this county have been obtained from the levels made by the railroads crossing the county, including the Seaboard Air Line, the Atlantic Coast Line and the Georgia Southern and Florida Railway. In addition levels made during 1909 by the U. S. engineers in connection with a preliminary survey for a ship canal have been available. These levels show that the water level in Lake Kingsley stood at the time the levels were made 170 feet above sea. The measurements of depth show that this lake averages 58 to 60 feet, although one place was found at which the depth exceeded 78 feet, the full length of the sound-

ing line. The country surrounding this lake stands at or about 175 feet above sea. According to the levels made by the Seaboard Air Line Railway the town of Highland, in the northwestern part of the county, stands 210 feet above sea. Newburg and Brooklyn, in the lake region of the southwestern part of the county, have elevations, as recorded by the Georgia Southern and Florida Railway, of 155 and 157 feet, respectively.

WATER-BEARING FORMATIONS.

Most of the flowing wells of Clay County terminate in the Vicksburg Limestone. The first flow at Green Cove Springs, in the eastern part of the county, is obtained at a depth of from 325 to 400 feet.

The Miocene formations underlie much if not all of Clay County. In the pit of Union Brick Company, at Middleburg, the following section was observed:

F	eet.
Loose sand and soil	1
Sandy clays oxidized red	7
Blue sticky clay, comparatively free from sand	10
Light-colored sands	3

The clay exposed in this pit is probably the same as the clays in the clay pit near Jacksonville. Beneath these clays, as indicated by well borings, calcareous and phosphatic Miocene rocks are encountered. This part of the Miocene, the Jacksonville formation, is exposed at many localities along Black Creek and its tributaries. The section exposed at High Bluff, on the south fork of Black Creek, about five miles above Middleburg, has already been given.

Other exposures of this formation were noted at the following localities along the river. At Fowler's Landing, on the south fork of Black Creek, three miles above Middleburg, fifteen feet of the Jacksonville formation is exposed. At Buddington's Landing, one and one-half miles above Middleburg, seventeen feet of the Jacksonville formation is exposed. Hogan's Landing, just below Middleburg, shows twenty-eight feet of the Jacksonville

formation. A bluff at the mouth of the south fork shows twenty-five feet of the Jacksonville formation. A bluff on the north bank

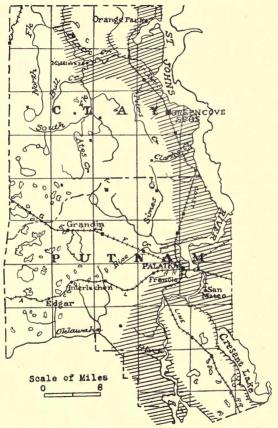


Fig. 8.—Map showing the areas of artesian flow in Clay and Putnam Counties. The area in which flowing wells can be obtained is indicated by shading.

of the north forks, one and one-half miles from Middleburg, shows three feet of the Jacksonville formation.

AREA OF ARTESIAN FLOW IN CLAY COUNTY.

The area of artesian flow in Clay County is confined to that portion bordering the St. Johns River and its tributaries. As has already been stated, upon leaving these streams the elevation soon becomes too great for a flow to be obtained. The location of successful flowing wells, together with the consideration of the elevation, will aid in the determination of the flowing and non-flowing sections in the county. The flowing area in this county is outlined on the accompanying map:

LOCAL DETAILS.

DOCTORS INLET.

A well owned by D. D. Denham and drilled in 1908 by D. C. Stafford is located near Doctors Inlet. This is a four-inch well, 372 feet deep, in which the water is said to rise twelve to fifteen feet above the surface.

A second well, two and a half miles east of Doctors Inlet, was drilled by H. Mervin for Messrs. DeLoach & Edwards in 1907. This is a three-inch well and is 400 feet deep. It is reported cased 120 feet and the water is said to rise twelve feet above the surface. Blue marl or clay from the depth of 198 to 398 feet is reported as encountered in this well. Immediately below this blue marl or clay the first hard rock was struck.

GREEN COVE SPRINGS.

Green Cove Springs, the county seat of Clay County, is supplied with water from two artesian wells. These wells are under private ownership. One is owned by N. B. Ivey, the other by O. A. Buddington. The well owned by Mr. Ivey is 815 feet deep, four inches in diameter, and cased 556 feet. The well is reported to have a head of 23 feet above the surface. The eleva-

tion of the well above the St. Johns River is given as 24 feet, thus giving the well a total head of 47 feet above the level of the water in the St. Johns River. The first flow in this well was encountered at a depth of 400 feet.

The following is an analysis of the water from this well drawn January 6, 1910. Analysis made for the State Survey in the office of the State Chemist, A. M. Henry, analyst:

Constituents.	Parts per million.
Silica (SiO ₂)	7
Chlorine (C1)	9
Sulphates (SO ₄)	7
Phosphates (PO ₄)	0
Carbonates (CO ₃)	0
Bicarbonates (HCO ₃)	107
Magnesium (Mg)	4
Calcium (Ca)	16
Iron and Alumina (Fe and Al)	Trace
Loss on Ignition	67
Total dissolved solids	155

Aside from the above well, the following two records of wells have been obtained: A well on the property of Mrs. George Halliday (known as the Borden estate), is 825 feet deep and six inches in diameter. The head is reported as 25 feet above the surface. A little southeast of this well is one owned by L. A. Hamilton. This has a reported depth of 785 feet, is six inches in diameter and is cased 100 feet. The head is given as 25 feet above the surface. A well four and one-half miles southwest of Green Cove Springs, drilled by H. Mervin in 1907 for the LaVilla Turpentine Company, is non-flowing. This well contains 128 feet of three-inch casing and 320 feet of two-inch casing. It is 406 feet deep and the water stands 17 feet below the surface. The first rock noted in this well was at a depth of 170 feet.

A well directly east of Green Cove Springs and across the St. Johns River is owned by W. A. Hallows. This well was drilled by N. B. Ivey and is used for irrigation and general domestic purposes. It is 500 feet deep, six inches in diameter and is cased about 200 feet. The water is reported to rise 35 feet above the surface.

Another well owned by N. B. Ivey is located about two miles southwest of Green Cove Springs. This well is used for irrigation and was sunk by the owner in 1907. It is a four-inch well and is reported to be 500 feet deep. At this depth the water is reported to rise five feet above the surface.

HIBERNIA.

One well is reported from Hibernia. This well was commenced July 20, 1885, and was finished in October of the same year. It was drilled by O. H. Wade for F. A. Fleming. The well is 468 feet deep, four inches in diameter and is cased 377 feet. This well when first drilled, in 1885, had a pressure of 23 pounds. Unfortunately, when visited in January, 1910, the pressure could not be obtained. The elevation of the well is about 25 feet above the St. Johns River. A pressure of 23 pounds will cause the water to rise 53.1 feet above the surface, or about 68.1 feet above the St. Johns River. The first water-bearing stratum in this well was reported at a depth of 400 feet, and the first rock noted was at a depth of 120 feet.

The following is an analysis of the water from this well drawn December 17, 1909. Analysis made for the State Survey in the office of the State Chemist, A. M. Henry, analyst:

Constituents.		Pa	rts	ре	r mill	ion.
Silica (SiO ₂)					. 9	
Chlorine (Cl)					7	
Sulphates (SO ₄)						
Phosphates (PO ₄)					. 0	
Carbonates (CO ₃)					0	
Bicarbonates (HCO ₃)					98	
Sodium and potassium (Na and K)					23	
Magnesium (Mg)					5	
Calcium (Ca)			٠.		14	
Iron and alumina (Fe and Al)	٠.			.Ί	`race	
Loss of ignition					15	
Total dissolved solids					122	

LENO.

There are two deep wells at Leno, owned by the Leno Turpentine Company, and drilled in 1903 by H. Mervin. One well, 404 feet deep, is four inches in diameter and the water stood when measured January 6, 1910, 12.5 feet from the surface. The second well is two inches in diameter and 220 feet deep. The water is reported to stand at about the same level.

MAGNOLIA SPRINGS.

Magnolia Springs, a station on the Atlantic Coast Line Railroad, one mile north of Green Cove Springs, takes its name from a small spring located along the western bank of the St. Johns River. A four-inch well owned by O. D. Seavey, proprietor the Magnolia Springs Hotel, was sunk by W. J. Sherman in 1882. This well is said to be 325 feet deep and flows several feet above the surface, although the exact head could not be obtained. This water is bottled and sold as a medicinal and table water. The following analysis shows the mineral constituents. Analysis by C. F. Chandler, Ph. D., School of Mines, Columbia College, New York, N. Y.:

Constituents.	Parts per million.
Sulphate of potash	Trace
Sulphate of lime	21.3
Chloride of sodium	14.4
Carbonate of soda	26.1
Carbonate of lime	40.4
Oxide of iron and alumina	Traces
Silica	31.0
Organic and volatile matter	16.4
	190.4

Two other wells occur on this same property, but a record of these was not obtained. They are both reported to furnish an abundant supply of water and are used for general household purposes.

MIDDLEBURG.

Middleburg lies in the north-central portion of Clay County, just at the point where Black Creek divides, forming the north and south forks. There are several flowing wells in the vicinity of Middleburg. The wells vary in depth from 355 to 498 feet.

The 498-foot well is owned by George A. Chalker and was drilled in 1907 by D. C. Stafford. The well is six inches in diameter at the top and one and a quarter inches at the bottom. The pressure of this well as indicated by the pressure gauge, January 10, 1910, was 18.5 pounds, or a pressure sufficient to cause the water to rise 42.7 feet above the surface. The elevation of the well is approximately 24 feet above the level of the water in Black Creek; thus, with the head of 42.7 feet above the surface, would give the well a total head of 67.7 feet above the water in Black Creek. The temperature of the water at the point of overflow was reported as 72° F. The first rock of which note was made was at a depth of 68 feet.

The well of C. C. Howard, two miles northeast of Middleburg, has a depth of 490 feet. The well was bored by D. C. Stafford in 1907, is cased 80 feet, and is four inches in diameter. The pressure of this well could not be taken, but it is reported to have a head of 21 feet above the surface.

Another well, two and a half miles northwest of Middleburg, was sunk by D. C. Stafford for Messrs. Long & Buddington, in 1907. The exact depth of this well could not be obtained, but it was reported to have a depth of about 370 feet. The well flows and gives an abundant supply of water, but measurement of the head could not be made.

In addition to the above wells is one eight and one-half miles northwest of Middleburg, or six miles southeast of Maxville, on the west bank of Yellow Water Creek, a tributary of the north fork of Black Creek. This well is located in the northwest part of the northwest quarter of the southwest quarter of Section 17, Township 4, Range 24 east. It is owned by Messrs. Long & Buddington, and is said to be 370 feet deep. It is a three-inch well and was drilled in 1907 by D. C. Stafford. The head of this well is reported to be 30 feet above the surface and the first flow encountered was at a depth of 44 feet in a stratum of black pebbles.

PEORIA.

A deep well was put down by Mr. Joseph Doyle at Peoria. This well was drilled to a total depth of 498 feet. The water rises to the surface, giving a slight flow. The well is located about one-half mile west of Peoria station and on the ridge probably 40 or 50 feet above the St. Johns River.

RUSSELL.

One flowing well is reported from Russell. This well is now owned by the Florida Farmers' Land Company and was drilled by L. J. Campbell. The well flows several feet above the surface, but a measurement could not be made and information in regard to the depth and size was not procured. It is used for general drinking purposes.

WALKILL.

A deep well at Walkill, drilled by H. Mervin in 1903 for E. B. Willcoxon & Company, reached a total depth of 352 feet. This well contains 128 feet of three-inch casing and 330 feet of two-inch casing. The water is reported to rise 25 feet above the surface.

WEST TOCOL

A record of one well has been obtained from West Tocoi. This is a three-inch well, reported to have a depth of 313 feet, and is owned by the R. W. Mattox Company. The head of this well is given as 21 feet above the surface.

WILLIAMS CROSSING.

Messrs. De Loach and Edwards have one deep well at Williams Crossing. This well is 395 feet deep and is three inches in diameter and was sunk by H. Mervin in June, 1907. The pressure of this well, as shown by the pressure gauge, January

6, 1910, was eight and one-half pounds or a pressure sufficient to cause the water to rise 19.6 feet above the surface.

PUTNAM COUNTY.

LOCATION AND SURFACE FEATURES.

Putnam County lies bordering the St. Johns River. On the north it joins Clay County, and on the south Marion and Volusia Counties. The total area of the county is 772 square miles. The elevation increases inland from the St. Johns River. At Florahome, in the northern part of the county, along the line of the Georgia Southern and Florida Railway, an elevation is reached of 150 feet. On the Rochelle branch of the Atlantic Coast Line Railroad an elevation of 105 feet occurs at Interlachen, in the central part of the county. That part of the county bordering the St. Johns River includes palmetto flatwoods and some open flatwoods. Much of the southern and western part of the county is occupied by the lake region, many small, beautiful lakes occurring in this section.

WATER-BEARING FORMATIONS.

The data regarding the formations reached by the wells in Putnam County is very meager, owing to the fact that few well samples have been preserved.

After passing through the superficial sands in this county, calcareous clay and sands are reached, in which are imbedded black phosphatic pebbles and water-worn gravels. From such imperfect information as has been obtained it seems probable that some of the wells terminate in this formation and do not reach the Vicksburg Limestone. The log of a well at Orange Mills, which terminated in loose, clear-grained sand at a depth of 160 feet, is given on a subsequent page. A second well within a half-mile of this well apparently reached the Vicksburg Limestone at or about the depth of 160 feet. Samples from the well of B. F. Dotney, at San Mateo, drilled in 1909, by H. Mervin, show the presence of black phosphatic pebbles as deep at least

as 175 or 180 feet. At a depth of 315 feet light-colored calcareous sands were penetrated. It is probable, as these wells seem to indicate, that the Vicksburg Limestone here, as at some other localities, has a very irregular top surface.

AREA OF ARTESIAN FLOW IN PUTNAM COUNTY.

The flowing area of Putnam County includes a relatively narrow strip bordering the St. Johns River and its tributaries. Upon leaving the river the elevation rises and flowing wells are not obtained. The flowing area in this county is indicated by shading on the map.

LOCAL DETAILS.

BOSTWICK.

Flowing wells are obtained at Bostwick. A three-inch well, drilled in 1904 for J. W. Glisson by H. Mervin, reached a total depth of 248 feet. This well is reported cased 60 feet and the water is reported to rise 18 feet above the surface.

Another well three and one-half miles northeast of Bostwick was drilled in 1906. This well is now owned by the R. W. Mattox Company and is used for the general supply around the turpentine camp. It is a three-inch well and reached a total depth of 215 feet.

CRESCENT CITY

Crescent City lies in southeastern Putnam County, on the western shore of Crescent Lake. Immediately along this western border flowing wells are obtained.

The first flow of water at this locality is obtained from a shell stratum lying from 30 to 60 feet below the surface. Most of the wells at Crescent City terminate at this depth. In some instances this shell stratum is reported absent and in such cases the water is reported as coming from a very fine sand. The water from this depth is usually more or less hard and is impreg-

nated with hydrogen sulphide gas. These wells are reported to have a head of about 15 or 16 feet above the surface.

The second flow in and near Crescent City is obtained at a depth of about 300 to 316 feet. From the immediate vicinity of Crescent Lake westward to the St. Johns River flowing wells are not obtained. The intervening country includes rolling, sandy hills. Surface wells, terminating in the sands and sandy clays furnish an abundant supply of soft water.

Aside from the use of private wells, Crescent City is supplied with water from four artesian wells. The water supply system is under private ownership. Two of the wells are two inches in diameter, while one is six inches in diameter. They are all reported as reaching a depth of approximately 316 feet, and cased about 100 feet. The wells are located on Crescent Lake and have approximately the same elevation. The head is reported 26 feet above the surface or about 27 feet above the level of the water in Crescent Lake. In addition to supplying the town the flow from one two-inch well is used for condensing purposes and for the manufacture of ice. Part of the flow from the other three wells is used for power to run an overshot wheel, which in turn runs a pump, pumping the surplus flow of water to a reservoir or tank where the water is distributed to different parts of the city by gravity.

ORANGE MILLS.

Orange Mills is located on the Florida East Coast Railway, midway between Hastings and East Palatka. The wells in this vicinity are used for the purpose of irrigation. The depth of the wells range from 143 to 200 feet. All of the wells of which record has been obtained are four inches in diameter. The length of casing used in the wells averages 60 feet.

Four wells drilled for J. H. Bahrenberg & Brother by N. H. Monck in December, 1909, gave the following pressure: Well No. 1 is 143 feet deep and is cased 65 feet. The pressure of this well as shown by the pressure gauge December 4, 1909, was 534 pounds. Well No. 2 is 160 feet deep and is cased 74 feet. The

pressure December 4, 1909, was 5½ pounds. Well No. 3 is 219 feet deep and is cased 54 feet. The pressure of this well on the same day was 5¼ pounds. Well No. 4 is 160 feet deep and is cased 58 feet. This well was not finished at the time the pressure of wells Nos. 1, 2 and 3 was taken. As will be seen from the above records the pressure in the case of these three wells diminished with depth. In this respect the wells are exceptional. The amount of flow of these three wells was not obtained. The following is the record of well No. 4, made from the samples kindly kept by the driller:

	Fe	et.
Sand	0-	5
Olive green calcarcous clay, with black phosphatic pebbles and fragments of shell, and flattened water-worn		
gravels	5-	40
No sample	40-	45
Similar or somewhat more calcareous green clay or clayey marl. This sample contains occasional fragments		
of chert	45-	80
This sample contains the black phosphatic water-worn		
pebbles in greater number than the above sample.		
Clear quartz grains are numerous. Flattened, water- worn siliceous pebbles up to size 1x½ inches occur	80-	90
In this sample clear quartz grains predominate. These are		
mixed with gray sand grains. Calcareous gray sand nodules occur, water-worn chert gravels are present,		
also numerous large, water-worn chert fragments	90-1	13
No sample		15
Loose, clear-grained sand in mass appearing light gray and		
contains a small amount of calcareous matter in the		
form of fragments of shell	115-1	.60

PALATKA.

Palatka, the county seat of Putnam County, is located on the St. Johns River, 55 miles south of Jacksonville. The elevation of the Atlantic Coast Line depot, as recorded by the U. S. Coast and Geodetic Survey, is thirteen feet above sea. Records from 35 wells have been obtained from and in the vicinity of Palatka.

The first flowing water encountered at Palatka is obtained from a shell stratum at a depth varying from 30 to 60 feet. A

great many wells in the city terminate at this depth. The water from this formation is more or less hard, but is not so strongly impregnated with hydrogen sulphide gas as is the water from the deeper water-bearing formations.

These more shallow wells at one time ceased to flow and pumps had to be resorted to. When the deeper wells were put in, the shallow wells in this vicinity commenced flowing again. As an instance of this, the well now owned by Messrs. L. H. and W. A. Merryday and located in the yard of the Putnam House, may be cited. This is a two-inch well and is 50 feet deep. It is reported as being cased the total depth. The well flowed when first put in, but in subsequent years had ceased to flow. During the year 1908 Mr. H. Mervin drilled a four-inch well for Dr. G. E. Welch about two blocks to the north. This well reached a total depth of 220 feet and is reported cased 120 feet. Immediately on the completion of this well the Merryday well commenced to flow. This seems to indicate that these wells are supplied with water through leakage from the wells reaching the deeper water-bearing strata.

The principal flow in and near Palatka is obtained from a depth of 175 to 250 feet. At this depth an abundance of water is obtained having a head varying from 18 to 26 feet above sea. A measurement was made of the pressure in the well of A. D. Curry, about three-fourths of a mile southwest of Palatka, in December, 1909. The well at this time was found to have a pressure of eleven and one-half pounds. The pressure was taken at the top of the pipe which stands about two feet above the surface of the ground.

A number of wells have been put down across the river and in the vicinity of East Palatka. The elevation of the depot at East Palatka, as given by the Coast and Geodetic Survey, is seventeen feet above sea level. A four-inch well drilled for H. Hanna at this place by N. H. Monck in 1909 reached a depth of 225 feet. It is reported cased 135 feet and the water is reported as rising fifteen feet above the surface. A second well drilled for the Florida East Coast Railway by N. H. Monck in 1909 was

drilled to a depth of 256 feet. This is a four-inch well and is reported cased 135 feet. The water is said to rise fifteen feet above the surface.

The following is an analysis of the water from the city well at Palatka. The water was sent in by Dr. E. S. Crill. Analysis made in the office of the Chemist, B. H. Bridges, analyst:

Constituents.	Parts per million.
Silica (SiO ₂)	18.0
Chlorine (Cl)	156.0
Sulphates (SO ₄)	76.9
Carbonates (CO ₃)	7.3
Bicarbonates (HCO ₃)	156.1
Magnesium oxide (MgO)	43.3
Calcium oxide (CaO)	
Total solids	

PENIAL.

A three-inch well was drilled at Penial by H. Mervin in 1904. This well is now owned by E. L. Parker and is used for general supply around the turpentine camp. This well reached a total depth of 235 feet and is reported cased 110 feet. The water is reported to rise 16 feet above the surface.

RICE CREEK

A two-inch well drilled at Rice Creek in 1904 reached a total depth of 175 feet. This well is reported cased 60 feet. It has a small flow of sulphur water, perhaps 12 to 15 gallons a minute. The head as shown by the pressure gauge December 8, 1909, is 25.1 feet above the surface.

RODMAN.

An attempt was made in 1909 to obtain a flowing well at Rodman. Two four-inch wells were drilled by H. Mervin for the Rodman Lumber Company. Well No. 1 reached a total depth of 139 feet and is reported cased 110 feet. Well No. 2 has 110

feet of four-inch casing, 200 feet of three-inch casing and 420 feet of two-inch casing, and was drilled to a total depth of 507 feet. The head did not increase with depth in this well, as is shown by the level of the water in either well, the head being three and one-half feet below the surface.

Approximately one mile east of Rodman a flow is obtained. A well drilled by H. Mervin for J. P. Buie in 1909 at this point has a head of four feet above the surface. It is a three-inch well and has a depth of 270 feet. The flow as measured December 9, 1909, is twelve gallons per minute.

SAN MATEO.

Flowing wells are not obtained at San Mateo, the surface elevation of the town, according to barometic readings, being approximately sixty feet above the St. Johns River. A four-inch well drilled for B. F. Dotney in 1900 by H. Mervin reached a total depth of 365 feet. The water in this well rises to within 48 feet of the surface. A number of flowing wells have been obtained, however, along the river, near San Mateo.

SATSUMA

No artesian wells are in use at Satsuma. The water used at this place comes from surface sands or clays at a depth varying from 25 to 46 feet. Flowing wells have been obtained along the river west of town.

WELAKA.

Welaka is located on the St. Johns River, about twelve miles south of Palatka. Records of two wells have been obtained at this place. One of these is the well now owned by the Welaka Mineral Water Company, a three-inch well, drilled in 1906. The first water under pressure was encountered at a depth of 160 feet. Below 160 feet the size of the boring was reduced to two inches, and was continued to a total depth of 329 feet, at which depth a highly mineralized water is obtained. The well has 98 feet of three-inch casing and 213 feet of two-inch casing. The

elevation of the well above the St. Johns River is reported to be 22 feet. The water in the well comes to within 16 feet of the surface or stands 6 feet above the level of the water in the St. Johns River.

The following is an analysis of the water from this well. Analysis by Robert Spurr Weston, 14 Beacon street, Boston, Mass.:

Constituents.	Parts per million.
Silica	12.00
Alumina	8.57
Iron carbonate	12.00
Calcium chloride	
Calcium sulphate	697.75
Calcium nitrate	
Magnesium bromide	
Magnesium chloride	507.45
Magnesium carbonate	
Sodium chloride	8808.52
Potassium chloride	

A second well at Welaka is owned by Mrs. Franklin Swift and was drilled by H. Mervin in 1909. This is a four-inch well and has a total depth of 151 feet. It is reported to be cased 104 feet and the water is said to stand eight feet below the surface.

WOODBURN.

A well was drilled one and one-half miles northeast of Woodburn in 1905 by H. Mervin for J. E. Edmonson. This is a four-inch well and has a depth of 185 feet. It is reported cased 120 feet and to have a head of five feet above the surface.

ORANGE COUNTY.*

LOCATION AND SURFACE FEATURES.

Orange County lies in South Central Florida, bordering the St. Johns River. This county has an area of 1,250 square miles and presents considerable diversity in soil and topography. The northwestern one-half of the county is included within the lake region of Florida and is dotted with innumerable small and large lakes. This part of the county has a rolling surface topography, the uplands rising considerably above the lakes. The eastern and southeastern part of the county bordering the St. Johns River is of lower elevation and consists largely of pine lands of

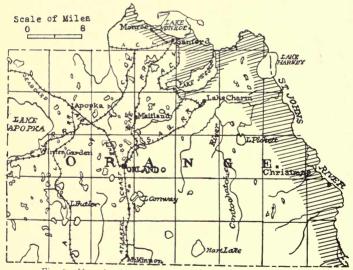


Fig. 9.—Map showing the area of artesian flow in Orange County. The area in which flowing wells can be obtained is indicated by shading.

^{*}Including Seminole County, which was created from Orange County after this paper was set in type.

the palmetto flatwoods type. The surface elevation in this county varies from about 20 feet above the sea in the northern part of the county to elevations of from 100 to 150 feet at points in the interior.

WATER-BEARING FORMATIONS.

The deep wells in Orange County terminate in the Vicksburg Limestone. At Sanford, in the northern part of the county, this formation lies comparatively near the surface, being reached at a depth of from 113 to 125 feet. Owing to the lack of a complete set of well samples the depth at which the formation is to be expected in other parts of the county has not been accurately determined. The formations lying above the Vicksburg have not been fully differentiated. It is probable that the Miocene occurs over the county, as the surface exposure of this formation has been recognized at Rock Springs, in the northwestern part of the county.*

AREA OF ARTESIAN FLOW OF ORANGE COUNTY.

The flowing area of Orange County is confined to a narrow strip bordering the St. Johns River. At Sanford this strip has a width of from three to five miles. Passing inland these low lands quickly give place to the more elevated, rolling lands of the lake region. With the exception of a few wells immediately bordering some of the lakes, flowing wells in this upland section have not been obtained. The flowing area in this county is outlined on the accompanying map,

LOCAL DETAILS.

CHULUOTA.

A two-inch flowing well three miles east of Chuluota is owned by Mr. G. M. Jacobs. The well is 114 feet deep, is cased 75 feet,

^{*}Smith, E. A., On the Geology of Florida. Amer. Journ. Sci. (3) XXI, 292-309, 1881.

and has a head of eight feet above the surface. The water is used for stock.

GENEVA.

There are several non-flowing wells in Geneva, the elevation being too great for a flowing well to be obtained. Mr. H. H. Pattishall has a two-inch well 133 feet deep and cased 85 feet. This well was drilled by the Geo. H. Fernald Company in 1909. The water is said to rise to within 29 feet of the surface.

Mr. J. T. McLain owns a well one and one-half miles north of Geneva. This is a two-inch well and is 135 feet deep. The water is reported to rise to within 31 feet of the surface. The water from this well is hard and is charged with hydrogen sulphide. In addition to the above well Mr. McLain has two wells on Mullet Lake, on the St. Johns River, about four miles slightly west of north from Geneva. Both of the wells furnish salt water impregnated with hydrogen sulphide and are not used. One is seventy-five feet deep and is said to flow two feet above the surface; the other is 135 feet deep and the water is reported to rise within one foot of the surface. The apparent difference in head is due to the difference in the elevation of the two wells.

Mr. W. B. Raulerson owns a two-inch well five miles northwest of Geneva and near the St. Johns River. This well is 76 feet deep and is cased 72 feet and furnishes a small flow of salt water which rises a few inches above the surface. The first flow in the well was encountered at a depth of 70 feet. An increased flow was obtained at 72½ feet. The first water was reported to be more salty than the second, as was indicated when the first flow was cased off. Owing to inability to drill deeper with the light drilling outfit used, the boring was discontinued. Mr. Raulerson states that the water is more salty in seasons of drought than in seasons of normal or heavy rainfall.

A two-inch well owned by Chase & Company, two miles southeast of Geneva, on Lake Harney, is 35 feet deep. This well was sunk by F. B. Bradley and is cased 34 feet. It has a head of four feet above the surface. The water is fresh and is only slightly charged with hydrogen sulphide.

ORLANDO.

Orlando, the county seat of Orange County, lies in the lake region of Florida. The elevation at the depot, as given by the Atlantic Coast Line Railroad, is 111 feet. Several wells have been drilled at Orlando. These are non-flowing wells, the elevation being too great to obtain a flow. The deep wells at this locality are used principally for drainage purposes and for irrigation, the city water supply being obtained from one of the small lakes. A few private wells in and around Orlando are used as a source of water supply.

A well near the north edge of the city owned by Mr. F. A. Lewter, has a total depth of 216 feet and is cased 86 feet. The water is used for general purposes.

A second well at the ice plant is used in cooling pipes in the manufacture of ice. This well is 470 feet deep.

The use of wells to carry off surface waters at this locality was described in the Third Annual Report. One of these drainage wells has developed at intervals the unusual phenomenon of spouting. An account of this well, together with an explanation of its unusual behavior is given in the report referred to, page 72.

OVIEDO.

Oviedo lies on the eastern edge of the lake region of Orange County. The region is sandy and the topography is flat to gently undulating. The country east of Oviedo is of the prevailing flatwoods type bordering the St. Johns River and Lake Jessup, and flowing wells are here obtained at comparatively shallow depths. Both flowing and non-flowing wells occur at Oviedo, depending on the local elevation.

Mr. N. J. Tanner's well, about one-eighth of a mile east of the postoffice at Oviedo, located in a depression, is about 114 feet deep, two inches in diameter, and is cased 75 feet. The water from this well flows just above the surface. It is a hard, sulphur water and is used for irrigating purposes.

The well of Mr. A. J. McCulley is 75 feet deep, two inches in diameter and is cased 70 feet. This well was sunk by the owner in 1907. The water is reported to rise to within 14 feet of the surface. Mr. McCulley owns another two-inch well which is 73 feet deep, and is cased 68 feet. The water in this well is said to rise to within three feet of the surface. This apparent difference in head is due largely to a difference in elevation of the wells.

A two-inch well, 117 feet deep, one and one-quarter miles west of Oviedo, was completed for Mr. D. W. Curry in 1910 by Mr. A. J. McCulley. This well gives a good flow of sulphur water and had, when measured in April, 1910, a pressure of four and one-quarter pounds, the measurement being made about five feet above the ground. The first flow in this well was encountered at a depth of 79 feet.

A well fourteen miles east of Oviedo, on the Econlockhatchee Creek, furnishes a flow of salt water. This well was drilled in 1907 by A. J. McCulley and is 114 feet deep, two inches in diameter, and is cased 75 feet. The first flowing water, which was salty, was found at a depth of 70 feet.

SANFORD.

Probably not less than 1,000 wells occur in and around Sanford. These wells are used for irrigating purposes and obtain flowing artesian water at a comparatively shallow depth, the average being from 125 to 200 feet. Bordering Lakes Monroe and Jessup and the St. Johns River, the wells are more shallow and terminate at a depth of from 66 to 85 feet.

The first flow in the wells at Sanford is encountered at a depth of from 100 to 125 feet, after drilling through a rock more or less hard and penetrating the characteristic "water rock" or the Vicksburg Limestone. In some instances a light flow is obtained above this harder rock immediately overlying the Vicksburg. When such is the case it seems the water comes from a quicksand or sometimes from a stratum of sand and shell. In order to get a sufficient and permanent flow, however, the boring is continued until the Vicksburg Limestone is reached.

In a well owned by Mr. L. E. Morrow, four miles south of Sanford, on the Sanford-Orlando public road, and drilled by Mr. W. E. Holmes in April, 1910, the Vicksburg Limestone was reached at a depth of 113 feet. The first flow was obtained at a depth of 110 feet, coming from a light yellow sand. The following is an approximate log of this well constructed from notes given by the driller and from a partial set of samples kindly saved by him:

	Feet.	
Surface soil	- 0- 5	
Yellow sand	5- 40	
Shell and sand, water, no flow	40- 60	
Sand	60- 91	
Shell and sand with shark's teeth	91- 95	
Dark blue rock with black phosphatic pebbles	95-100	
Very dark rock	100-101	n'
Light yellow sand	101-113	
Vicksburg Limestone	113-	

The principal supply of water for the city of Sanford is drawn from Lake Ada, about four miles southeast of the city. The soft water from the lake is preferred to the hard, sulphuretted artesian water. However, the city has four artesian wells, which serve as a source of supply when the lake is low. These wells are all four inches in diameter and are reported to have an average depth of 130 feet. Measurements in regard to the volume of flow of these wells could not be obtained.

Several flowing wells occur at Cameron City, on Lake Jessup, about six miles southeast of Sanford. The wells here are of about the same depth as those in and near Sanford and good flows are obtained. The principal use of the water is for irrigating purposes.

At Monroe, a station four miles northwest of Sanford, on the Atlantic Coast Line Railroad, a number of wells have been sunk. According to reports from drillers the artesian conditions here are essentially the same as at Sanford. A well about one-fourth of a mile southwest of the depot was drilled for the Title, Bond and Trust Company by W. E. Holmes & Son. This is a two-inch well, 180 feet deep and cased 120 feet. The pressure of this

well April 19, 1910, was eight and one-half pounds, the measurement being made one and one-half feet above the surface. About one-fourth mile beyond the above is a second well. This well indicated a pressure of eight pounds, the measurement in this instance being made three feet above the surface. Unfortunately the total depth of this well could not be learned. A third well about one and one-fourth miles beyond this second well indicated



Fig. 10.—Artesian well of E. Hy. Palmer on the west side of Lake Jessup.

a pressure of one pound. This well has a total depth of 201 feet, is two inches in diameter and is cased 154 feet. As will be seen these wells decrease in pressure on leaving the river. This decrease in pressure is due to the increase in elevation. All of the above mentioned wells are along the grade of the now abandoned railroad from Paola to Monroe.

Another well four miles southwest of Sanford and owned by Mr. J. V. Weeden, terminated in the Vicksburg Limestone, as is shown by a mixed sample of the drillings gathered after the well was completed. Unfortunately neither the total depth of the well nor the depth at which the Vicksburg Limestone was reached could be learned. This well is two inches in diameter and furnishes a good flow of water.

The well of Mr. E. Hy. Palmer, seven miles south of Sanford,

near the western shore of Lake Jessup, is 75 feet deep and was drilled in 1907. This is a four-inch well and is cased 40 feet. The pressure of this well as indicated by the pressure gauge, April 26, 1910, was nine and one-half pounds, or a pressure sufficient to cause the water to rise 21.9 feet above the point of connection of the gauge, which was three feet above the surface. The well is estimated to be about 12 feet above Lake Jessup, which estimation will give the well a total head of 36 feet and 9 inches above the surface of the lake.

The deepest well at Sanford is the well owned by Mr. J. E. Pace. This well is located just outside of the known flowing area and was sunk in the hopes of obtaining a flow. The well is six inches in diameter to a depth of five hundred feet, below which depth the size of the drill hole was reduced to four inches. It has a total depth of 670 feet and the water rises to within one and one-half feet of the surface. The well is reported cased only 94 feet. A detailed record of the well could not be obtained, but it was stated by Mr. Pace that no apparent increase in head resulted from the increased depth, although no exact measurements regarding this were made.

VOLUSIA COUNTY.

LOCATION AND SURFACE FEATURES.

Volusia County lies between the St. Johns River and the Atlantic Ocean. It joins St. Johns County on the north and Brevard County on the south. The area of the land surface of this county is approximately 1,281 square miles. Much of the eastern part of the county is level and consists largely of palmetto flatwoods. Bordering the Atlantic Ocean, however, is an extensive strip of hammock known as Turnbull Hammock. Back of the hammock is found the line of sand dunes. Bordering the St. Johns River is found some open flatwoods. Running in a general north and south direction through the western part of the county is a ridge including much sandy pine land. Numerous lakes occur in this upland section which forms a part of the lake region of Florida. Elevations above sea level recorded by the Atlantic

Coast Line Railroad which traverses this ridge are as follows: Seville, 52 feet; Pierson, 78 feet.

WATER-BEARING FORMATIONS.

No complete set of well samples having been obtained from any one well in Volusia County the information regarding the underlying formations is very meager. In the city well at De-Land the first water was obtained at a depth of 113 feet after passing through eight feet of clay and entering a twelve-foot shell stratum. The stratum of shell overlies a bed of rock reported to be 24 feet thick. The next rock encountered is at a depth of 237 to 247 feet. At Daytona the Vicksburg Limestone, as shown by the comparatively shallow depths of the wells, lies close to the surface and is presumably reached at from 125 to 150 feet.

AREA OF ARTESIAN FLOW IN VOLUSIA COUNTY.

The area of artesian flow in Volusia County is confined to a strip bordering the Atlantic Ocean on the east and a strip on the west bordering the St. Johns River. This area is indicated on the accompanying map. There are no doubt areas not mapped where flowing wells can be obtained. The area mapped, however, is based on definite information and on well records. In the northern portion of the county flowing wells are obtained as far west as Crescent Lake. This part of the county is flat and of low altitude.

LOCAL DETAILS.

DAYTONA

Daytona lies in the flowing artesian section in eastern Volusia County, along the western bank of Halifax River. The city is supplied with water from four artesian wells, all of which are six inches in diameter. These wells were drilled in 1909, but in order to obtain an increased flow were deepened in 1910 and now range in depth from 165 to 260 feet. The 260-foot well on April 7, 1910.

had a head of 9.3 feet above the surface or approximately 13.3 feet above sea. The wells now furnish an abundant supply of hard sulphuretted water.



Fig. 11.—Map showing the areas of artesian flow in Volusia County. The areas in which flowing wells can be obtained are indicated by shading.

In addition to the city wells above mentioned numerous private wells occur in and near Daytona. Of these it is possible to list only a few. Mr. Paul Petion owns a two-inch well about two and one-half miles south of the city. The well was drilled by Mr. H. VanDorn in 1910. It is 145 feet deep and is cased 85 feet. The first flowing water is reported to have been encountered at a depth of 85 feet after drilling through about one foot of hard rock.

Messrs. Bellough and Melton completed a two-inch well for Mr. Chas. Lee about two miles southwest of Daytona in April, 1910. This well is 130 feet deep and has a head of five feet above the surface. The first flow is reported from a depth of 88 feet just below a hard rock upon which the casing was landed. The following is a log of this well as given by Mr. Melton:

	Feet.
Dark sandy soil	0- 6
White marl	6- 15
Sand and shell	15 20
Blue clay	10- 00
Sand and shell	30- 63
Limestone medium 1 at Times	65- 87
Limestone, medium hard. First flow at 88 feet, increase of	
water with depth	87-130

The following is a log of Mr. H. VanDorn's well. The well is one-half mile west of the postoffice and was completed by Mr. VanDorn in April, 1910. It is a four-inch well, 205 feet deep, and is cased 83 feet:

Dark sandy soil	Feet.
Dark sandy soil	0- 3
Hardpan	3- 5
White sand	5- 40
Coquina and shell	40- 45
Blue clay Hard rock Light 9	45- 65
Hard rock. Light flow just above this rock.	65-83
Light-colored limestone, with harder and softer layers. In-	83- 84
crease of water with increase of depth	
merease of depth	84-205

The wells listed are representative of the wells surrounding Daytona. Flowing water is obtained at a comparatively shallow

depth. From the above two logs it will be seen that hard rock was encountered at the depth of 87 and 84 feet, respectively. Immediately under this hard rock a softer limestone is reported and in this limestone the first flowing water is obtained. The description of this formation given by the drillers characterizes it as the Vicksburg which is apparently reached in this section at a depth of not more than 125 to 150 feet.

DE LAND.

The city of DeLand, the county seat of Volusia County, lies in the southwestern portion of the county. There are a number of non-flowing artesian wells in and near DeLand. The city is at present supplied by two deep wells located at the pumping station. The six-inch well is 406 feet deep and was sunk in 1895. This well was reduced in diameter in the process of drilling and is cased as follows: Six-inch casing to 100 feet, four-inch casing to 290 feet, two-inch casing to 390 feet. The second well, which was drilled in 1906 by W. F. Hamilton, is ten inches in diameter and is 269 feet deep. At the depth of 191 feet hard rock was encountered upon which the casing was landed. The head of the wells, regardless of the difference in depth, was reported to be 27 feet below the surface in both cases. The following log and analysis of the water from this well were kindly made available by Mr. E. D. McLeod:

	Feet.
Sand	0- 25
Clay	25- 45
Shell	45- 50
Rock	50- 55
Sand	55-105
Clay	105-113
Shell, water-bearing	113-125
Rock	125-149
Sand	149-157
Rock	157-197
Sand and shell	197-237
Rock	237-247

Clay 247	-257
Sand	-265
Rock	-277
Clay	-292
Shell and clay	-372
Rock 372	-392
Cavity with water 392	-406

The following is an analysis of the water from the six-inch city well at DeLand. Analysis by H. Herzog, Jr., Gainesville, Fla.

* *		
Constituents,	Parts 1	er million.
Total solids		136.29
Residue after ignition (mineral matter)		76.11
Gas and ignition (organic matter)		60.17
Sodium chloride		11.31
Free ammonia		
Albuminoid ammonia		
Oxygen (consuming power)		
Nitrates		.34
Nitrites		
Sulphates		
Phosphates		Trace

ENTERPRISE

Flowing wells are obtained at Enterprise, along the shore of Lake Monroe, and in areas where the elevation does not exceed more than ten or twelve feet above the level of the water in the lake. The depth of the wells in this vicinity ranges from 20 to 200 or more feet, the average depth being between 90 and 110 feet. The water is hard and is charged with hydrogen sulphide, in some instances containing a large amount of salt. A well owned by Mr. William S. Thayer was drilled to a depth of 98 feet. It is two inches in diameter and is cased 45 feet. The estimated elevation of this well is 15 feet above the level of the water in Lake Monroe. The water is reported to rise to within three feet of the surface of the ground. An analysis of the water from this well made in the office of the State Chemist showed it to contain 140 parts total solids to 1,000,000 parts water. The total solids are reported to be composed of calcium carbonate

(lime), sodium chloride (common salt), and magnesium sulphate (Epsom salts).

The following is an analysis of the water of the Benson Mineral Spring, located about one-fourth mile west of town, and owned by the Misses Emma and Tina Tucker. Analysis made at Vanderbilt University, Nashville, Tenn., by W. H. Hollenshead:

Constituents.	Parts per million.
Potassium	27.104
Sodium	1805.046
Magnesium	213.047
Calcium	321.619
Iron	
Chlorine	3389.640
Bromine	103.206
Carbon dioxide	559.234
Sulphuric acid (radical)	541.132
Silica	16.989
Phosphoric acid (radical)	702
Boric acid	Heavy trace
Organic matter	Small amount
Hydrogen sulphide	Slight trace

The above are probably combined in the water as follows:

Constituents.	Parts	per million.
Potassium sulphate		60.346
Calcium sulphate		720.043
Sodium bromide		133.722
Magnesium chloride		819.787
Sodium phosphate		.994
Iron chloride		1.594
Sodium chloride	4	504.371
Calcium chloride		76.701
Calcium bicarbonate		330.928
Silica		16.989
Carbonic acid		379.624
Sodium biborate	. Heav	y trace
Hydrogen sulphide	Sligh	it trace
Organic matter	Small a	amount

LAKE HELEN.

Lake Helen lies in the lake region of southern Volusia County. The land here is high, rolling pine woods. The elevation of the depot at Lake Helen, as recorded by the Florida East Coast Railway, is 70 feet. The wells recorded from this place range in depth from 130 to 238 feet. The Bond Sand-Lime Brick Company own several three-inch wells ranging in depth from 130 to 140 feet. The water is reported to rise within 28 feet of the surface. A well for Mr. G. W. Webster was drilled in 1897 by Mr. H. C. Haven. This well is 238 feet, four inches in diameter and cased 158 feet. The first rock is reported at a depth of 78 feet. The principal water supply is obtained from a depth of 210 feet. The water is hard and is only slightly charged with hydrogen sulphide.

NEW SMYRNA.

The artesian conditions at New Smyrna are essentially the same as those given for Daytona. The wells in this vicinity range in depth from 108 to 144 feet. The water is hard and is charged with hydrogen sulphide and is used to a large extent for irrigating purposes.

The following is an analysis of the water from the well of Mr. W. L. Widmeyer, made in the office of the State Chemist, B. H. Bridges, analyst:

	er million.
Silica (SiO ₂)	27.0
Chlorine (Cl)	836.6
Sulphates (SO ₄)	7.8
Carbonates (CO ₃)	12.0
Bicarbonates (HCO ₃)	209.8
Magnesium oxide (MgO)	108.6
Calcium oxide (CaO)	197.7
Total solids	1980.0

The following is a log of a four-inch well drilled by R. C. Walker for the Florida East Coast Railway. The record is obtained through the courtesy of Mr. G. A. Miller:

	Feet.
Coal cinders (filled land)	0- 5
Coquina rock	5- 14
Sand	14- 16
Sand and shell	16- 42
Blue clay	42- 45
Fine shell	45- 64
Fine shell and sand	64-80
Coarse shell	80- 91
Rock	91- 92
Clay and shell	92- 96
Hard rock	96-100
Soft white limestone	100-156

The following is a log of a three-inch well drilled by H. Van-Dorn, two miles west of New Smyrna, for the Florida East Coast Railway. The record is obtained through the courtesy of Mr. G. A. Miller:

	reet
Sand	0 - 16
Rock	16 - 201/2
Shell	201/2- 24
Clay	. 24- 40
Rock	40- 42
Clay	42-44
Rock	44- 46
Clay	46- 79
Rock	. 79- 81
Shell	81-85
Rock	85-87
Rock, bearing salty water	87-103
Rock, bearing fresh water	

OAK HILL.

Oak Hill is eleven miles south of New Smyrna, on the Florida East Coast Railway and about four miles north of the head of Indian River. Several flowing wells occur in the vicinity of this place. These wells are reported to be about 130 feet deep. The water is hard and sulphuretted. Approaching the head of Indian River, some four or five miles south of Oak Hill, flowing wells of brackish water are obtained. Mr. T. J. Murray owns four

wells, all near the head of Indian River, which are used for stock. One of these wells was never satisfactorily completed. Two of the wells give a brackish flow while the water from the other well, which is located about one mile south and west of the head of the river, is reported to be fresh. This well, however, is not as deep as the other two wells, being only 82 feet deep and terminating before passing through the "bed" or hard rock which was encountered at that depth. The two brackish wells are reported to have a depth of 110 feet and to have a head of about seven feet above the surface. According to well records this seems to be the northern extent of the shallow brackish flowing wells, fresh water wells being obtained just a few miles to the north. Eastward this salt area presumably extends to the Atlantic Ocean. In 1907 Mr. J. W. Griffis had a well sunk one mile northwest of Shiloh, to a depth of 149 feet. The well at this depth flowed just above the surface and furnished a very strong salt water. The well is now capped and is not used. The character of the artesian water westward in this part of the county is not known, records of wells not having been obtained.

ORANGE CITY.

The Orange City wells vary in depth from 117 to 890 feet. The 890-foot well is owned by Mr. Albert Dickinson and is not used. Salt water was encountered at the depth of 890 feet and the well was plugged up below 660 feet. The depth of the well as now used is 660 feet. The principal use of the artesian wells in this vicinity, aside from general domestic purposes, is that of irrigation, the Orange City Mineral Spring Company, however, have a well 117 feet deep, the water from which is bottled for sale. This is a ten-inch well and is reported cased to a depth of fifteen feet. The water is said to rise to within twenty feet of

the surface. The following is an analysis of the water from this well.* Analyst unknown:

Constituents.	Parts per million.
Free ammonia	0.00
Albuminoid ammonia	0.05
Oxygen consumed	1.05
Nitrites	0.00
Nitrates	1.00

ORMOND.

Several deep wells have been sunk at Ormond. These deep wells all furnish a salt water which cannot be used except in some instances where it is used for bathing purposes. A four-inch well was drilled by Mr. H. Walker in 1900 at the Hotel Ormond. This well reached a depth of 752 feet and is cased 360 feet. At a depth of 320 feet salt water was encountered. The water from the well is used for bathing purposes. Another well at the Hotel Ormond reached the same depth. This well is eight inches in diameter and is cased 400 feet, at which depth salt water is reported, continuing to 550 feet. From the depth of 550 feet to the total depth of the well, 752 feet, no water was encountered.

The average depth of the wells surrounding Ormond and vicinity is from 160 to 225 feet. At this depth a hard sulphuretted water is obtained. However, in some instances salt water at this shallow depth is reported. Mrs. A. M. Watson owns a three-inch well which is 180 feet deep and cased 90 feet. The water from this well is not used because it contains salt. This well is the only one of this depth on record that contains salt, other wells of medium depth furnishing an abundant supply of fresh water, which is used for domestic and irrigating purposes. The head of the wells range from eight to nine feet above the surface or about fourteen to fifteen feet above sea.

^{*}U. S. Geological Survey, Bull. 102, p. 263, 1904.

PIERSON.

Pierson is located on the sandy ridge running through the west central portion of Volusia County. The elevation of the depot at this place, as recorded by the Atlantic Coast Line Railroad, is 78 feet. Records of two deep wells occurring here have been obtained. The N. L. Pierson well is three inches in diameter and 150 feet deep. The water is reported to rise to within forty feet of the surface. Its use is general domestic and irrigation purposes. The second well was drilled at the public school house and is used for general drinking purposes.

SEVILLE.

The Atlantic Coast Line Railroad owns four artesian wells at Seville, used for the railroad boiler supplies. One well is four inches in diameter and is reported to be 126 feet deep. The other three wells are two inches in diameter. The exact depth is not known. The water is said to rise to within 18 feet of the surface.

About two miles south of Seville and west of the Atlantic Coast Line Railroad is a flowing artesian well. This well is owned by J. W. Whitner, and was drilled in 1909. This is a two-inch well, 140 feet deep and is cased 90 feet. The elevation at the well, as determined by measurement, is sixteen feet above Lake George. The well on April 25, 1910, as indicated by the pressure gauge, had a pressure of four and one-quarter pounds, equivalent to a head of 9.8 feet above the surface or 25.8 feet above the level of the water in Lake George. The first flowing water was reported at the depth of 80 feet, at which depth hard rock was encountered.

BREVARD COUNTY.

LOCATION AND SURFACE FEATURES.

Brevard County lies between the St. Johns River and the Atlantic Ocean. It has a total length of 66 miles and, including Merritts Island, is about 25 miles wide. It joins Volusia County

on the north and St. Lucie County on the south. Aside from the line of sand dunes running parallel with the coast this county is prevailingly of the palmetto flatwoods type of country, although extensive prairie and muck lands occur in the interior of the county. Lake Washington, in the central part of this county, has an elevation of 15.74 feet while Lake Wilmington, the head waters of the St. Johns River, in St. Lucie County, has an elevation of 23.37 feet above mean sea level at Indian River Inlet.*

WATER-BEARING FORMATIONS.

The deep wells in Brevard County enter the Vicksburg Limestone. At Melbourne this limestone, as indicated by well samples kept from the well of Mr. Oliver Gibbs, was reached at the depth of 221 feet. At Cocoa, in the well of Mr. H. Bradford, the Vicksburg Limestone was recognized at a depth not exceeding 190 feet.

AREA OF ARTESIAN FLOW IN BREVARD COUNTY.

Although the interior of this county is but thinly settled and but few wells have been put down, it is probable that the greater part of this county lies within the area of artesian flow. On the high sand dune ridge, which lies out three or four miles from the coast, a flow is not to be expected. This is probably also true of points within the interior of the county, particularly in the southwestern part.

LOCAL DETAILS.

CHESTER SHOALS.

Some fifteen miles from Titusville, through Banana Creek, is the Chester Shoals Life Saving Station and Canaveral Club House. At this club house an artesian well was drilled about 1890. It is a three-inch well and the original depth was 222 feet. The amount of casing used could not be learned. The well in

^{*}Survey made in 1903, under the direction of Captain F. R. Shunk, U. S. Army.

subsequent years decreased in flow, and in order to get a sufficient amount of water for general use it became necessary to deepen the well. In 1905 Captain Alex. Near continued the drilling to 297 feet. The well now gives an abundance of water strongly impregnated with hydrogen sulphide and tasting slightly brackish, although not so much so as to condemn it for general purposes.

CITY POINT.

Flowing wells are obtained along the shore of Indian River at City Point. Between the village on the river and the City Point depot, on the Florida East Coast Railway, there is quite an elevation, evidently an old sand dune. The elevation of this ridge, according to barometric readings, is about fifty feet above the level of the water in the river. A well sunk here some years ago failed to flow, although the water rose to within a few feet of the surface. A well owned by S. Hendry is reported to have a depth of about 160 feet. The elevation of the well is approximately twenty feet above the water in Indian River. The pressure of this well, as indicated by the pressure gauge March 5, 1910. was five pounds, or sufficient pressure to cause the water to rise 11.5 feet above the surface, or approximately 31.5 feet above the river level.

The following is an analysis of the water from this well. Analysis made for the State Survey in the office of the State Chemist, A. M. Henry, analyst:

Constituents.	Parts per million.
Silica (SiO ₂)	raits per minion.
Chloring (Ct)	17.00
Chlorine (Cl)	2248.00
Sulphates (SO ₄)	207.00
Phosphates (PO ₄)	201.00
Carbonates (CO)	8.00
Carbonates (CO ₃)	0.00
Dicarbonates (HCO ₃)	160 00
Sodium and potassium (No and E)	100.00
Sodium and potassium (Na and K)	
magnesium (Mg)	116.00
Calcium (Ca)	440.00
Iron and alumina (Fe and Al)	1.00
Loss on ignition	1.00
Loss on ignition	960.00
Total dissolved solids	5053 00

COCOA.

The number of artesian wells in and around Cocoa renders it impossible to specifically mention more than a few representative ones. The artesian wells in this section terminate at a medium depth and are sunk without encountering great difficulty in drilling, thus making the cost comparatively slight. The wells terminate in the Vicksburg Limestone, as indicated by the mixed samples of drillings from the well of H. Bradford, one mile southwest of Cocoa. The water is reported in some instances to contain a trace of salt, but only in a very few cases was it found to be injurious to vegetation.

The well of O. K. Key was sunk by the owner in 1908. It is a three-inch well and has a depth of 202 feet. The well is cased 140 feet. The pressure of the well, as indicated by the gauge, March 10, 1910, was ten pounds, or a head of 23.1 feet above the surface. The elevation of the well above the level of the water in the Indian River, as shown by barometric readings, is 15 feet, thus giving the well a total head of 38.1 feet above the water level in the river. The water has a slight trace of salt and is impregnated with hydrogen sulphide gas.

About one-fourth mile southwest of the city postoffice is the well of the Cocoa Ice Company. This well is reported to have been drilled in 1888. It is a four-inch well, 325 feet deep, and cased about 125 feet. The pressure of this well in 1908 was reported to be twelve and one-quarter pounds. This pressure would give the well a head of 28.2 feet above the surface. The estimated surface elevation is about 10 feet above the river, making a total head of 38.2 feet above the level of the water in Indian River.

An artesian well one mile southeast of Cocoa was completed in February, 1910. This well was drilled by J. A. Coward and is owned by H. Bradford. It is three inches in diameter, 190 feet deep and is cased to a depth of 80 feet. A mixed sample of the drillings taken after the completion of the well indicates that the Vicksburg Limestone was encountered. The exact depth at which this limestone was reached could not be learned. The volume of

flow, as measured March 10, 1910, was 60 gallons per minute and the pressure as indicated by the pressure gauge on the same date was five pounds or a pressure sufficient to cause the water to rise 11.5 feet above the surface. The elevation of the well above the level of the water in Indian River, as shown by barometric readings, is 20 feet. This elevation, together with a head of 11.5 feet above the surface, gives the well a total head of 31.5 feet above the river level. The water is the characteristic sulphur water common to most of the artesian wells of the State,

The following is an analysis of the water from this well. Analysis made for the State Survey in the office of the State Chemist, A. M. Henry, analyst:

Constituents. Parts 1	er million.
Silica (SiO ₂)	
Chlorine (Cl)	. 1082
Sulphates (SO ₄)	. 201
Phosphates (PO ₄)	. 0
Carbonates (CO ₃)	. 0
Bicarbonates (HCO ₃)	159
Sodium and potassium (Na and K)	536
Magnesium (Mg)	77
Calcium (Ca)	167
Iron and alumina (Fe and Al)	1
Loss on ignition	470
Total dissolved solids	2546

EAU GALLIE.

The first artesian well in Eau Gallie was drilled, in 1887, by John McAllister. This well is now owned by George F. Paddison, and is 337 feet deep. It is one and one-fourth inches in diameter and cased 136 feet. The depth to the water rock was reported by the driller, Mr. McAllister, to be 237 feet. The head of this well is given as 42 feet above the surface, or approximately 52 feet above the level of the water in Indian River. Since the completion of the above test well, many wells have been sunk in and around Eau Gallie, varying in depth from 315 to 500 feet. The principal water supply is obtained at a depth of from 230 to 315 feet.

The East Coast Lumber and Supply Company use two artesian wells as a source for power in running a planing mill. They are both six-inch wells and are about 500 feet deep. The pressure of the wells could not be obtained, but they are reported to have a head of 50 feet above the river. The principal use of the surrounding artesian wells is for general domestic purposes and irrigation.

FRONTENAC.

Mr. Josiah Thompson owns an artesian well at Frontenac. This well was reported to be 190 feet deep, and is four inches in diameter. The water is strongly impregnated with salt, and is used for power to pump water from a shallow fresh water well. The pressure of the well could not be obtained, but the head and flow were reported to be very good.

GRANT.

A four-inch well, now owned by Mr. Charles Christiancy, at Grant is the only flowing well in the vicinity. The well is 350 feet deep and is cased 90 feet. It was drilled, in 1896, by Messrs. Near & Taylor. The principal supply of water is said to come from a depth of 256 feet.

MALABAR.

Several deep wells have been sunk at Malabar. They vary from 300 to 390, or more, feet in depth. The principal use of the water is for irrigation purposes.

MELBOURNE.

At Melbourne, a record of several deep wells was obtained. Mr. W. T. Wells owns an artesian well, which was sunk by Capt. Alexander Near in 1898. This well is 389 feet deep and four inches in diameter. The pressure, as shown by the pressure gauge on March 15, 1910, was eleven and one-quarter pounds. The surface elevation was given as about 26 feet above the level of the water in Indian River, and this elevation, together with a pressure of

eleven and one-quarter pounds, would give the well a head of 51.9 feet above the river.

The six-inch well of Capt. J. S. Sammis is 400 feet deep and is cased about 73 feet. The pressure of this well was taken on March 15, 1910, but since all connections to the well could not be shut off, the full pressure could not be obtained. The reading, however, was 11 pounds, which was a sufficient pressure to cause the water to rise 25.4 feet above the surface, or about 47.4 feet above the river; the well being about 22 feet above the river.

A three-inch well, owned by Mr. Wm. R. Campbell, near Melbourne is used for power purposes and for irrigation. The water from the well turns an overshot wheel, which runs a pump, pumping water from a surface well. The surface water is soft and is preferred to the hard sulphur water of the deeper well. The well is 385 feet deep and was sunk by Messrs. Near & Taylor in 1895.

A well, one mile west of Melbourne, owned by Mr. H. P. Bowden, is six inches in diameter and is 400 feet deep. The well was sunk by Capt. Alexander Near in 1907. The pressure, as indicated by the pressure gauge March 14, 1910, was 12 pounds, or a head of 27.7 feet above the surface. The surface elevation of the well, shown by barometer, was 22 feet above the water level in Indian River. This would give the well a total head of 49.7 feet above the river. The water from this well, besides being used for general domestic purposes, is used for bathing and for power. Two large concrete bathing pools have been built and the water flows continually into them. The temperature of the water is said to be 77 degrees F. A water wheel, connected near the well, is used to pump water from a shallow, soft water well.

The following is an analysis of the water from this well. Analysis made for the State Survey in the office of the State Chemist, A. M. Henry, analyst:

Constituents,	Parts per million.
Silica (SiO ₂)	18
Chlorine (Cl)	573
Sulphates (SO ₄)	150
rnosphates (PO ₄)	
Carbonates (CO ₃)	0

Bicarbonates (HCO ₃)	156
Sodium and potassium (Na and K)	269
Magnesium (Mg)	68
Calcium (Ca)	123
Iron and alumina (Fe and Al)	8
Loss of ignition	375
Total dissolved solids	1555

Mr. M. B. Rhodes' well, near the postoffice, at Melbourne, is 45 feet deep and furnishes a flow, which rises about three feet above the surface. The elevation of the well is about three feet above the water level in Indian River. The well is of interest in that the water flows at such a shallow depth. The materials penetrated in the sinking of this well were approximately as follows:

	Fee	t.
Sand	0 -	10
"Hardpan"	10 -	11
Sand, water	11 -	-20
"Hardpan," water	20 -	201/2
Sand	201/2-	-35
Sandy, clay, water, flowing 3 feet above the surface of the		
well	35 -	-45

The water is soft and very desirable for all domestic purposes. Another such well as the above is owned by Dr. L. A. Peek. This well is 52 feet deep, one and one-fourth inches in diameter and furnishes a good supply of soft water.

The well owned by Mr. Oliver Gibbs is located at Melbourne Beach, across the Indian River from Melbourne. This is a four-inch well drilled, in 1907, by Capt. Alexander Near. It reached a total depth of 318 feet and is cased 100 feet. The pressure of the well, as indicated by the pressure gauge March 15, 1910, was 17½ pounds. This gives the well a head of 40.4 feet above the surface, or estimating the surface elevation of the well to be 12 feet above the river level, a total head of 52.4 feet above the level of the water in Indian River. From an examination of a mixed sample from the drillings of this well, it is seen that the Vicksburg Limestone is reached. From Mr. Gibb's record made at the time the well was drilled, it would appear that this formation was

encountered at a depth of 221 feet. The log, as made out by Mr. Gibbs, is as follows:

	F	eet.
Surface sands and soil	0	- 3
Yellow sand		- 11
Coquina rock	11	- 21
Fine gray sand	21	- 51
Shell and sand	51	- 56
Hard shell rock	56	-119
Greenish clay	119	-173
Dark colored rock; sharks' teeth	173	-1731/2
Greenish clay	1731/	2-1741/4
Dark colored rock; sharks' teeth	1741/	4-175
Greenish clay		
Vicksburg Limestone. Increase of flow with depth. A		
pressure of 171/2 pounds at this depth was shown by		
the gauge March 15, 1910. Mild sulphur water	221	-318

MERRITTS ISLAND.

From the well records obtained in this locality, it is probable that flowing artesian wells can be obtained at any point on Merritt Island. Record of wells are on file from every postoffice on the island, bordering the Indian River. Also, records have been obtained from Artesia, Cape Canaveral Light House, and Canaveral Club House, on the peninsula, east of the island; good flows being reported from all of these localities. The pressure of two of the wells on the southern end of the island was obtained, one at Lotus and one at Tropic. The well of L. D. Hancock, one mile south of Lotus, has a depth of about 300 feet. The pressure of this well March 12, 1910, was 16 pounds. The elevation of the well, according to barometric readings, is 10 feet. This, together with a pressure of 16 pounds, gives the well a total head of 46.9 feet above the level of the water in Indian River. The following is an analysis of the water from this well. Analysis made for the State Survey in the office of the State Chemist, A. M. Henry, analyst:

Constituents		Parts per million.
Silica (SiO2) .	• • • • • • • • • • • • • • • • • • • •	12
Chlorine (Cl) .		642

Sulphates (SO ₄)	178
Phosphates (PO ₄)	0
Carbonates (CO ₃)	0
Bicarbonates (HCO ₃)	149
Sodium and potassium (Na and K)	309
Magnesium (Mg)	63
Calcium (Ca)	132
Iron and alumina (Fe and Al)	3
Loss on ignition	370
Total dissolved solids	1710

At Tropic Mrs. John W. Merrill has two artesian wells, two and three inches respectively. These wells were drilled about 1885. The depth was not learned. The gauge on the two-inch well, March 12, 1910, indicated a pressure of 16½ pounds, or a head of 38.1 feet above the surface, or about 48 feet above the water level in Indian River. The wells are used for general purposes and give an excellent flow of sulphur water.

From the records obtained it appears that the pressure of the wells on the island increases in passing from north to south. At Lotus the pressure was 16 pounds; at Tropic 16½ pounds, and at Melbourne Beach 17½ pounds. No measurements of the pressure of the wells north of Lotus were obtained.

MICCO.

The wells at Micco have, for the most part, been drilled a number of years and, for this reason, no satisfactory records could be obtained. The principal use of the water is for irrigating purposes. One well, drilled in 1908 for Peter Bertleson by J. L. Mobley, was never completed. The well is 3 inches in diameter and is cased 180 feet. At a depth of 300 feet the drill was broken off and was never recovered. A flow coming just over the top of the casing was obtained at this depth.

ROCKLEDGE.

The Rockledge wells vary in depth from 150 to 480 feet. These wells are the principal source of domestic water supply, as well as being used for irrigating purposes. In a few instances the artesian

wells are used for power purposes, such as for generating electricity by means of a water turbine. A ten-inch well, drilled in 1893 and now owned by Mr. G. M. Houston, about one and one-half miles south of Rockledge, is used for this purpose. The well has a reported depth of 480 feet. A gauge on the well indicated a pressure of 12½ pounds, March 10, 1910, or a head of 28.8 feet above this point. The gauge was estimated to be ten feet above the level of the water in the river, thus giving the well a total head of 38.8 feet above the river level. The water contains a trace of salt, as is common to the wells in this vicinity.

The well of Mr. H. S. Williams is of particular interest, in that it is the only well in this vicinity, of which a log has been obtained. It was drilled about 1890 and is 304 feet deep. It is three inches in diameter and is cased 130 feet. The following is a log of this well, as reported by Mr. Williams:

	Feet.	
Sand and soil	0 - 10	
Coquina rock	10 - 30	
Sand	30 -100	
Sand rock	100 -140	
Blue clay	140 -170	
Hard flint rock. At this depth water rose to the surface,		
small stream	170 -173	
Rock in layer from 3 to 18 inches thick	173 -269	
Hard rock		
Soft rock		
Hard rock, good flow of water		

The first flow in the well, as will be seen by consulting the log, was obtained from a depth of 170-173 feet. At this depth 3 feet of hard flint rock was encountered and on penetrating this stratum the first water-bearing formation was reached.

SHARPES.

Several flowing wells occur in and near Sharpes. The water here contains salt to such an extent that it can not be used for irrigation. The well of J. W. Spafford furnished the following record. The well is four inches in diameter and 200 feet deep. It is reported cased only about 40 feet, and to have a head of 10

feet above the surface. The first flow was encountered at 70 feet and it is reported by the driller, Capt. W. H. Sharpes, that neither the head nor the volume increased with the depth. As indicated from the well records and from all obtainable information, only a small amount of casing was used in the wells in this vicinity, and a knowledge as to whether or not fresh water was encountered below the stratum of salt water is, therefore, lacking.

The following is an analysis of the water from the well of J. J. Ollif, Sharpes, Fla. This well is near the Spafford well and approximately one mile north of the Hendry well, at City Point, analysis of which is given on page 234. Analysis made for the State Survey in the office of the State Chemist, A. M. Henry, analyst:

Constituents.	Parts per million
Silica (SiO ₂)	16
Chlorine (Cl)	3120
Sulphates (SO ₄)	302
Phosphates (PO ₄)	0
Carbonates (CO ₃)	0
Bicarbonates (HCO ₃)	165
Sodium and potassium (Na and K)	1634
Magnesium (Mg)	286
Calcium (Ca)	262
Iron and alumina (Fe and Al)	4
Loss on ignition	974
Total dissolved solids	6520

TILLMAN.

The only deep well at Tillman, of which record has been obtained, was drilled by John McAllister, in 1890, and is owned by R. A. Conkling. It is 350 feet deep and furnishes an excellent flow of water, which is used for general domestic purposes.

TITUSVILLE.

Titusville, the county seat of Brevard County, is located on the Indian River. Several artesian wells have been sunk at this locality, but up to the present time principally salt water has been obtained. A test well, put down about 1890, was drilled to a total depth of 864 feet. A salt water stratum was reached at a depth of about 100 feet. The well was cased to a depth of about 110 feet, but no attempt was made to case off the salt water. Both the flow and the head is reported to have increased with increase of depth. Several other wells have been subsequently drilled in and near the city. One of these located at the Dixie Hotel, is said to have been drilled to a depth of about 400 feet. Another, located at the Grand View Hotel, drilled about 1895, is believed to have reached the depth of about 200 feet. Two other wells, one located at the old plant of the Florida Extract Company, the other at the plant of the Titusville Ice Company, were drilled to a depth of 150 and 145 feet, respectively. Salt water was obtained from all of these wells, and in none of them was an attempt made, so far as the records indicate, to go through or to case off the salt water stratum. Fresh water is obtained from shallow driven wells, none of which exceed 100 feet in depth. The water obtained from these wells, as a rule, does not flow. In at least one instance, however, a small flowing fresh-water well has been obtained at a depth of less than 100 feet. The wells, which exceed 100 feet in depth, as stated above, have yielded only salt water.

The following is an analysis of the water of the well of the Titusville Ice Company. Analysis made for the State Survey in the office of the State Chemist, A. M. Henry, analyst:

Constituents.	Parts per million.
Silica (SiO ₂)	8
Chlorine (Cl)	
Sulphates (SO ₄)	547
Phosphates (PO ₄)	0
Bicarbonates (HCO ₃)	177
Sodium and potassium (Na and K)	6542
Magnesium (Mg)	669
Calcium (Ca)	637
Iron and alumina (Fe and Al)	
Loss on ignition	1380
Total dissolved solids	

VALKARIA.

A record of one deep well, at Valkaria, has been obtained. This well was drilled by Mr. W. J. Nesbitt, in 1892, for Mr. E. A. Svedelius. It is a 3-inch well, 350 feet deep, and is cased to a depth of 90 feet. The water is reported to rise 15 feet or more above the surface. At a depth of 320 feet hard rock was encountered and, immediately below this rock, the first water, under sufficient pressure to cause it to rise to the surface, was obtained.

ST. LUCIE COUNTY.

LOCATION AND SURFACE FEATURES.

St. Lucie County lies south of Brevard County. It is 42 miles long and from 24 to 42 miles in width. Ft. Drum ridge in this county has an elevation of 66.74 feet above the mean sea level.* The eastern part of the county, aside from the line of sand dunes near the coast, consists largely of palmetto flatwoods. Towards the west border the land is more rolling and numerous small lakes occur. Some muck lands are found near the headwaters of the St. Johns River.

WATER-BEARING FORMATIONS.

The wells of this county, as elsewhere along the East Coast, reach the Vicksburg Limestones. These limestones, however, dip in passing to the south and lie at a greater depth in St. Lucie County than in the adjoining counties to the north. The wells of the St. Lucie Ice Company, at Ft. Pierce, are 812 feet deep and, probably, reach the Vicksburg Limestone. The first flow from the wells, at Ft. Pierce, is reported to have been obtained from the depth of 725 to 750 feet.

AREA OF ARTESIAN FLOW OF ST. LUCIE COUNTY.

Owing to the few wells that have been drilled, the area of artesian flow in St. Lucie County is imperfectly determined.

^{*}Survey made in 1903, under the direction of Captain F. R. Shunk, U. S. Army.

Along the East Coast wells are in use as far as the southern line of the county. It is probable that flowing wells can be obtained for some miles inland from the coast.

LOCAL DETAILS.

EDEN.

A four-inch well at Eden, owned by Mr. Chas. Edison, was sunk by Messrs. Fee & Nesbitt and is 870 feet deep. The water is used for general and irrigation purposes. It rises 25 feet above the surface. It is a hard water and is impregnated with hydrogen sulphide.

FT. PIERCE.

Two artesian wells occur at Ft. Pierce, the county seat of St. Lucie County. These are owned by the St. Lucie Ice Company. The wells are reported to have a depth of 812 feet. One is six inches in diameter, the other 2 inches, and both are reported cased 200 feet. The first flow is said to have been obtained from limestone, at a depth of from 725 to 750 feet. The last 100 feet of the well is said to have been through this limestone. The following is an analysis of the water from one of these wells. Analysis by the Geo. W. Lord Company, 2238-2250 North Ninth Street, Philadelphia, Pa., Chester Alsmere, chemist, reported January 18, 1907:

Constituents.	Parts per million.
Organic and volatile matter	
Calcium oxide	
Magnesium oxide	
Sodium oxide	
Sulphur trioxide	
Chlorine	
Carbonic acid (combined)	
Silica	33.979

As will be seen in the above analysis this water contains a high percentage of sodium and chlorine. The water tastes brackish, and is used for cooling purposes in the manufacture of ice. The principal water supply for domestic purposes, in and around Ft. Pierce, is obtained from shallow wells, ranging in depth from 12 to 50 feet.

The following is a record of a well drilled at Ft. Pierce by H. Walker for the Florida East Coast Railway in 1912. The well is 814 feet deep and is cased with eight-inch casing 184 feet and 9 inches, and with six-inch casing 570 feet and 10 inches. The head above the surface is 28 feet and 6 inches. The head above Indian River is 46 feet. The well flows 800 gallons per minute at the surface. The record has been kindly supplied by Mr. G. A. Miller, of the Florida East Coast Railway.

	Feet.
Yellow sand	0- 55
Shell and sand	5575
Shell and gravel	75-85
Shell, sand and clay	85-120
Blue clay and sand	120-135
Soft blue clay and very fine sand	135-145
Blue clay and sand	145-165
Blue clay	165-190
Tough, dry blue clay	190-200
Soft sandy, blue clay	200-250
Hard sandy, blue clay	250-300
Smooth blue clay, no sand	300-400
Blue clay, very tough and sticky	400-460
Yellow clay, with black streaks in it	460-500
Yellow clay, with a few pebbles	500-520
Blue clay, tough and sticky	520-545
Very hard yellow clay	545-555
Blue clay, very sticky	555-585
Yellow clay in hard and soft layers	585-600
Yellow clay, very dry	600-647
Shell and soft rock	647-656
Tough white clay	656-662
Hard white rock	662-676
Soft rock, small flow	676-685
Soft limestone rock, flow increasing very slowly with	
depth	685-807
Hard rock	807-814

The following is an analysis of the water from this well made by the American Water Softener Company, Philadelphia, Pa.:

		D .
G	rains per	Parts per
Constituents. U.	S. gallon.	million.
Calcium carbonate	1.71	29.31
Calcium sulphate	8.34	142.97
Magnesium carbonate	9.26	156.73
Sodium sulphate	18.90	324.00
Sodium chloride	43.50	745.71
Free carbonic acid	1.00	17.14
Iron, aluminum and silica	0.28	4.70
Incrusting solids	19.59	335.83
Non-incrusting solids	62.40	1069.72
Total solids	83.00	1422.86

NARROWS.

Two deep wells are reported from Narrows. One is owned by Mr. F. Foster, the other by Mr. E. L. Gray. These wells were drilled by Mr. W. J. Nesbitt about the year 1892. Both are three inches in diameter and 420 feet deep. The height to which the water would rise above the surface was not obtained, but the wells are reported to have a head of several feet, and to furnish a strong flow of water.

ORCHID

Mr. S. K. Michael owns an artesian well at Orchid. This well was sunk by Capt. Alexander Near in 1896. It is 480 feet deep, four inches in diameter and is cased 85 feet. The well is reported to have a head of 40 feet above sea, and to furnish an abundant supply of hard, sulphur water.

ROSELAND

The artesian wells at Roseland have been drilled for a number of years and, for this reason, no very definite information could be obtained. Mr. L. C. Moore owns three wells, located about one and one-half miles north of Roseland, on the point between the Sebastian and Indian Rivers. These wells range in depth

from 350 to 453 feet. The water is hard and impregnated with hydrogen sulphide and is used for irrigating and general purposes.

SEBASTIAN.

There are several flowing artesian wells in and near Sebastian. They vary in depth from 365 to 500 feet. At this depth an abundance of hard sulphuretted water is obtained, rising from 16 to 25, or more, feet above the surface. A well owned by Mr. I. A. Groves, drilled by Mr. J. McAllister, was completed in 1896. This well is 460 feet deep, four inches in diameter, and is cased 100 feet. The water is reported to have a head of 16 feet above the surface. the surface elevation being estimated at 25 feet above the level of the water in Indian River. The total head of the well above the river is thus 41 feet. The water is used for general and domestic purposes and for irrigation. A four-inch well, drilled by Capt. Alexander Near, in 1901, owned by the Indian River Cooperage Company, is 365 feet deep. The water is reported to rise 25 or more feet above the surface. The elevation of the depot at Sebastian, according to levels run by the Florida East Coast Railway, is 19 feet. This well has approximately the same elevation as the depot, and this, in addition to head of 25 feet above the surface, gives the well a total head of 44 feet above sea. The well is now abandoned, but, when first sunk, was used for the manufacture of ice.

The Fellsmere Farms Company have recently completed a well, about ten miles west of Sebastian (Sec. 22, T. 31 S., R. 37 E.). The well is four inches in diameter, 370 feet deep, and is cased 146 feet. The head, tested September 23, 1910, by Mr. E. H. Every, manager, was found to be 25 feet above the surface, and the flow 185 gallons per minute.*

The following is the analysis of the water from this well made by the State Chemist:

Constituents.	Parts per million.
Chlorine (Cl)	257
Carbonates (CO ₃)	0

^{*}Letter to Capt. R. E. Rose, State Chemist, Sept. 23, 1910.

Bicarbonates (HCO ₃)	177
Loss on ignition	245
Total dissolved solids	905

PINELLAS COUNTY.

LOCATION AND SURFACE FEATURES.

Pinellas County lies on the Gulf Coast and includes the peninsula between Tampa Bay and the Gulf of Mexico. The area of the land surface of the county is approximately 260 square miles. The surface is prevailingly level, with a gradual rise in passing inland from the coast. The county is crossed by the Atlantic Coast Line Railroad, and by the Tampa and Gulf Coast Railroad. The elevations recorded by the Atlantic Coast Line Railroad are as follows: Belleair, 49 feet; Clearwater, 29 feet; Dunedin, 13 feet; Largo, 50 feet; St. Petersburg, 20 feet; Tarpon Springs, 14 feet.

WATER-BEARING FORMATIONS.

The deep wells in Pinellas County doubtless terminate in the Vicksburg Limestone.

AREA OF ARTESIAN FLOW.

The flowing area, in this county, includes a narrow strip bordering the coast and extending somewhat north of Dunedin. Flowing wells can probably be obtained along the shore entirely around Tampa Bay. The accompanying map shows the area in this county, in which it is believed that flowing wells can be obtained.

LOCAL DETAILS.

CLEARWATER.

Clearwater is near the center of the county, from north to south. The city water supply, at Clearwater, is taken from a well 250 feet deep. A second well, 270 feet deep, is held in reserve. Both wells are eight inches in diameter and are cased about 30 feet. When not in use the wells flow, but when either well is being

pumped the head is reduced, stopping the flow in the other well. The 270-foot well has brackish water

The Clearwater Ice Factory has three wells, 46, 52 and 80 feet deep. They are all cased 30 feet and the water is reported to stand 26 feet from the surface. These wells are located on high ground, the difference in elevation being probably sufficient to account for the difference in head between these and the city wells.

DUNEDIN

Flowing wells are obtained along the coast at Dunedin. The wells range in depth from 55 to 120 feet. C. B. Bowden has a six-inch well, 98 feet deep and cased 70 feet, in which the water stands 20 feet from the surface. This well is used as a public supply. T. J. Zimmerman has a well, at this locality, 68 feet deep, in which the water stands 12 feet from the surface. W. C. McLain has a flowing artesian well, about two miles north of Dunedin. This well is 202 feet deep and is estimated to flow 10 gallons a minute. This is the northernmost flowing well in this county.

ESPIRITU SANTO SPRINGS.

The Espiritu Santo Springs, located near the north end of Tampa Bay, include five springs. The following is an analysis of the water from the one known as the drinking spring: Analysis by the N. B. Pratt Laboratory, Atlanta, Georgia:

	Grains per	Parts per
Constituents.	U.S. gallon.	million.
Peroxide of iron and alumina	.1692	2.9007
Sodium chloride	137.8520	2363.2208
Magnesium chloride	25.8768	443.5292
Potassium sulphate	3.4815	59.6854
Calcium sulphate	19.7172	338.0297
Calcium carbonate	12.6145	216.2607
Silica	.9972	17.0958
Total solids by evaporation	254.9165	4370.2629

LARGO.

Several wells have been drilled at Largo. Lewis Johnson has a four-inch well, 200 feet deep, which is used as a public well. Joel McMullen has a well, about eight miles southwest of Largo, 227 feet deep, in which the water stands 15 feet from the surface.

OZONA.

The wells at Ozona are mostly 50 to 60 feet deep. A two-inch well, owned by Wm. Woods, is 80 feet deep and the water stands eight feet below the surface. C. R. C. Smith has a well 106 feet deep, but the water at this depth is said to be salty.

PASS-A-GRILLE.

The following is a log of a four-inch well, 256 feet deep, drilled by J. C. Danielson, and is used as a public well. The well is cased 204 feet and the water rises 14 feet above sea:

	Feet.
White beach sand	0- 3
Shells	3- 7
Fine sand	7- 35
Coquina	35- 41
Quicksand and blue clay	41- 80
Hard blue clay	80-200
Limestone, principal flow from 230 feet	05-256

The following is an analysis of the water from this well. Analysis made in the office of the State Chemist of Florida, A. M. Henry, analyst:

	Milligrams per liter. (Parts per million.)
Si0 ₂	46.2
Fe and Al	6.2
Ca	393.8
Mg	187.0
Na	611.9
K	10.9
Cl	
CO ₃	0.0

HCO_3	
SO ₄	
PO ₄	0.0
Total	3775.3
These may be combined as follows:	
, , , , , , , , , , , , , , , , , , , ,	
KCl.	20.8
NaCl	1555.3
MgCl ₂	777.9
CaCl ₂	25.0
CaSO ₄	1068.6
Ca (HCO ₃) ₂	271.6
CaSiO ₃	10.7
SiO ₂	
Fe and Al	

PINELLAS PARK.

A six-inch well, drilled at this locality by J. C. Danielson, reached a total depth of 325 feet. The water in the well stands five feet from the surface.

ST. PETERSBURG.

Flowing wells are obtained at St. Petersburg, along the coast, the water rising 10 to 14 feet above sea level. A well near the dock is four inches in diameter and 100 feet deep. The water from this well rises 10 feet above the surface. Another well near the bay, about one mile southeast of St. Petersburg, was drilled 480 feet deep. The water from this well rises about 10 feet above the surface and is salty.

The city supply at St. Petersburg is obtained from one seveninch and two ten-inch wells, variously reported at 135 and 235 feet deep. These wells are located on the upland, about 31½ feet above the level of the bay. When first drilled, the water is said to have stood 22 feet from the surface, but, after being used for some time, the water level was reduced to 36 feet from the surface. The following is an analysis of the water made by Dearborn Drug and Chemical Works, Chicago, Ill., December 11, 1911:

G	rains per	Parts per
Constituents. U.	S. gallon.	million.
Silica	.934	16.012
Oxides of iron and alumina	.117	2.005
Carbonate of lime	9.900	169.724
Chloride of lime	2.761	47.334
Sulphate of lime	.405	6.943
Carbonate of magnesia	Trace	
Chloride of magnesia	1.993	34.167
Sodium and potassium sulphates	Trace	
Sodium and potassium chlorides	4.964	85.102
Loss, etc.	.183	3.137
Total soluble mineral solids	21.257	364.427
Suspended matter	1.168	20.024
Organic matter	Trace	
Total soluble incrusting solids	16.110	276.188
Total soluble non-incrusting solids	5.147	98.239

The following is analysis of water from one of these wells. Analysis made in the office of the State Chemist, A. M. Henry, analyst:

Well water of 155-foot city well of St. Petersburg, Pinellas County, Florida:

	Milligrams per liter.
	(Parts per million.)
Silica (SiO ₂)	28.5
Iron and alumina (Fe and Al)	0.9
Calcium (Ca)	
Magnesium (Mg)	9.6
Sodium (Na)	
Potassium (K)	
Chlorine (CI)	120.6
Carbonates (CO ₃)	0.0
Bicarbonates (HCO ₃)	216.6
Sulphates (SO ₄)	
Phosphates (PO ₄)	1.5
Total	580.00

The following is a log of a six-inch well, 99 feet deep, drilled by J. C. Danielson, and owned by the St. Petersburg Investment Company. The well was drilled in 1912 and is cased 64½ feet.

	Feet.
Casing driven and no record	0 -641/2
Hard lime rock	641/2-691/2
Soft lime rock	691/2-71
Close grained lime rock	71 -86
Water-bearing rock	86 -88
Hard lime rock	88 -99

The following is an analysis of the water from this well made by the Bird-Archer Company, 90 West Street, New York City:

	Grains per	Parts per
Constituents.	U. S. gallon.	million.
Organic and volatile matter	4.717	80.867
Sodium chloride	3.244	55.614
Calcium carbonate	9.529	163.364
Magnesium chloride	2.332	39.979
Total solids	19.822	339.826
Free carbonic acid	9.415	161.409

The following is a log of a six-inch well, 155 feet deep, owned and drilled by J. C. Danielson. The well is located on the bay shore, two miles north of St. Petersburg. It is cased 76 feet, and the water rises about three feet above the surface. The flow is estimated at 200 gallons per minute.

	Fe	eet.
Soil	0	- 11/2
Dark colored sand	11/2	- 9
Hard pan	9	- 16
White water-bearing sand	16	- 50
White clay	50	- 60
Water-bearing rock	60	- 66
Light brown, sticky clay	66	- 76
Rock, alternating hard and soft strata	76	-156

The following is a log of a four-inch well, 230 feet deep, drilled by J. C. Danielson for R. S. Hanna, at Maximo Point, five miles southwest of St. Petersburg. The elevation at the well is about seven feet above sea and the water rises six feet above the surface, or a total head of about thirteen feet above sea. The well is cased 86 feet.

Feet. Soil	
$\begin{array}{c} \text{Milligrams per lite} \\ \text{(Parts per million} \\ \text{PO}_4 \\ \text{SiO}_2 \\ \end{array}$	
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	
Fe and Al. 2.6 Ca 328.6 Mg. 122.8 K 8.8	
Na 462.2 O (calculated) 7.8 Total 2830.6	
These may be combined as follows:	
$\begin{array}{ccc} KCl & 24.8 \\ NaCl & 1175.2 \\ MgCl_2 & 480.9 \\ CaCl_2 & 46.7 \end{array}$	
$\begin{array}{ccc} \text{CaSO}_4 & & 791.7 \\ \text{Ca} & (\text{HCO}_3)_2 & & 239.1 \\ \text{CaSiO}_3 & & 56.9 \\ \text{SiO}_2 & & 12.7 \\ \end{array}$	
Fe and A1 2.6	

SEMINOLE.

A number of wells have been drilled in the vicinity of Seminole. A four-inch well, owned by Frank Grable, near the locality, drilled by T. J Zimmerman, reached a depth of 270 feet. The water is reported to stand 16 feet from the surface.

SUTHERLAND.

Several wells have been drilled at Sutherland. Those exceeding about 100 feet in depth are reported to reach salt water. Fresh water is obtained from 50 to 100 feet.

TARPON SPRINGS.

The city supply at Tarpon Springs is obtained from three sixinch wells, 80, 108 and 126 feet deep respectively. The water stands 20 feet from the surface. The Polar Ice Company also have three wells, 82, 90 and 120 feet deep respectively. In the deepest of these salt water was reached at 120 feet, and the well was plugged at 100 feet.

Tarpon Springs, at this locality, comes up in a bayou from Anciote River. Although covered at high tide, the strong boil from the spring can be seen at medium and low tides.

WALL SPRINGS.

A well, drilled at Wall Springs by T. J. Zimmerman for W. W. Clark, reached a total depth of 313 feet. Fresh water was found in this well to a depth of about 90 feet. Below 90 feet the water is brackish. Three lines of casing were used in this well as follows: eight-inch, six-inch and four-inch. The four-inch casing is said to reach 312 feet. The water in the well stands 13 feet from the surface. There are a number of wells that have been drilled at this locality from 50 to 90 feet deep and yield a fresh water.

Wall Spring, at this locality, has an estimated flow of 3,000 gallons per minute. The water from this spring is used for medicinal purposes.

HILLSBORO COUNTY.

LOCATION AND SURFACE FEATURES.

Hillsboro County includes an area of 1,049 square miles. The county is crossed by the Atlantic Coast Line Railroad and by the Seaboard Air Line Railway, and their branches, and by the Tampa Northern Railroad. The elevation rises in passing inland from Tampa Bay and the Gulf. Plant City, near the east line of the county, is reported, by the Atlantic Coast Line Railroad, to be 137 feet above sea level. The level given by the Seaboard Air Line Railway for this locality is 125 feet above sea level. The elevation of other points in this county, along the Seaboard Air Line Railway, is as follows: Brandon, 74 feet; Knights, 117 feet; Turkey Creek, 87 feet. The elevation of points in this county, along the Atlantic Coast Line Railroad, is as follows: Hillsboro, 35 feet; Seffner, 74 feet, and Thonotosassa, 49 feet above sea. The Hillsboro and Alafia Rivers flow across this county and enter Hillsboro Bay.

WATER-BEARING FORMATIONS.

While no complete set of well drillings has been obtained, there is little doubt but that the deep wells of this county terminate in the Vicksburg Limestone. The surface exposures along Tampa Bay and along the Hillsboro River, for some miles above Tampa, are of the Tampa Limestone, Upper Oligocene, which overlies the Vicksburg formation or Lower Oligocene. A full description of the exposures of the Tampa formation in this county, by George C. Matson and F. G. Clapp, will be found in the Second Annual Report of this Survey, pages 84 to 91, 1909.

AREA OF ARTESIAN FLOW.

Flowing artesian wells are, probably, to be obtained entirely around Hillsboro and Tampa Bays. The head is sufficient to bring the water about ten to fifteen feet above sea level, and the wells will usually flow where the rise above sea does not exceed this elevation. The accompanying map shows the area in this county in which flowing wells can be obtained.

LOCAL DETAILS.

PLANT CITY.

The public water supply at Plant City is taken from a well 340 feet deep. This well is cased 260 feet, and the water stands 33 feet from the surface.

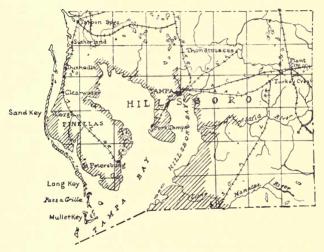


Fig. 13.—Map showing the flowing area in Hillsboro and DeSoto Counties. The area in which flowing wells have been obtained is indicated by shading.

The Plant City Ice and Power Company have a well about 600 feet deep. The water in this well stands 20 feet from the surface.

The following is a log of the well of the Warnell Lumber and Veneer Company, as kept by the drillers, the Hughes Specialty Well Drilling Company. The well is $266\frac{1}{2}$ feet deep and is cased with eight-inch casing to a depth of 105 feet. The water in this well stands 45 feet from the surface.

	Feet.
Sand and clay	1- 20
Sand and dark colored marl	20- 40
Marl and medium hard rock	40-100
Light colored hard rock	100-105
Medium hard rock	
Light colored hard rock	175-190
Shell-bearing medium hard rock	190-215
Soft shell-bearing rock	215-2661/2

TAMPA.

The water supply for the city of Tampa is obtained from drilled wells, of which there are twenty-eight at present. The wells range in depth from about 200 to 325 feet. Wells at a greater depth, as a rule, reach salty water. The wells are mostly 10 inches in diameter. The casing extends from 52 to 103 feet. The elevation above sea varies from 8 to $15\frac{7}{2}$ feet. The water in these wells will rise 15 to 17 feet above sea level, hence most of the wells flow at the surface. The following is a log of one of the wells taken from the Second Annual Report of this Survey, page 89:

Feet

	I CCL.
White Pleistocene sand	0- 2
Tough yellow clay with no sand, residual clay	2- 12
Soft limestone, which disintegrates readily, "Tampa lime-	
stone"	12- 26
Chert, "Tampa silex bed"	26- 30
Soft limestone, closely resembling that at 12 to 26 feet	30- 36
Tough, plastic, greenish sandy clay	36- 77
Base of the Tampa formation:	
Chert	77- 79
White marl	79- 85
Soft limestone	85- 90
Very light colored hard rock	90-105
Very hard dark yellow limestone	105-111
Gray, porous limestone with some water	111-126
Cherty beds	126-140
Darker limestone.	
Gray plastic clay.	
Hard yellow rock with chert.	
Gray, porous rock, water-bearing.	
Like preceding.	

The following is a log of the well, located at the southwest corner, Lot 2, Block 10, Bouquardez Sub-Division, 233 feet deep:

	Feet.
Sand	0- 8
Clay	8- 12
Hard gray rock	12- 36
Soft gray rock	36- 48
Very hard dark flint	48- 53
Blue clay	53-103
Gray rock with hard streaks	103-124
Dark, hard flint	
White lime rock	129-160
White clay	160-164
Brownish gray rock	164-180
Lime rock	180-233

The following is a log of a well at Ybor City. Northwest quarter, Section 17, Township 29, Range 19, 355 feet deep. Wellnow abandoned.

	Feet
Sand	0 - 19.4
Gray rock	19.4- 37.6
Tough gray clay	37.6- 41.6
Tough gray clay with streaks of rock	41.6- 48.2
Gray, hard, flinty rock	48.2- 95
Yellow sand rock with a little water	95 - 97
Gray rock, coarse with dark green streaks	97 -109.4
Yellow shell rock	109.4-110.4
Gray hard rock	110.4-223
White sticky clay	223 -228
Gray hard rock	228 -240.5
Soft gray rock, not porous	
Soft gray rock	315.5-355.5

Four wells at West Tampa, drilled by W. F. Hamilton, formerly used for the public supply, are now used for the manufacture of ice. These wells vary in depth from 360 to 760 feet. One of these is a six-inch well, 360 feet deep, cased 135 feet. Another is a six-inch well, 390 feet deep, cased 93 feet. A third well is 406 feet deep and is cased 90 feet. The fourth well was drilled at a depth of 760 feet and was cased 412 feet. The water

from this depth was salty and the well was subsequently filled to about 400 feet. The casing was broken and fresh water admitted at this depth. The water in the wells at this locality is reported to stand at about seven feet from the surface.

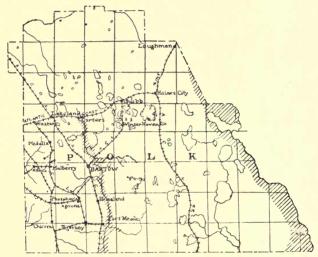


Fig 14.—Map showing of artesian flow in Polk County. The area in which flowing wells have been obtained is indicated by shading.

POLK COUNTY.

LOCATION AND SURFACE FEATURES.

Polk County includes a land area of 1,967 square miles. The Lake Region crosses the central part of this county, and the principal pebble phosphate deposits of the State are found in the western part of the county. The following elevations are recorded along the Atlantic Coast Line Railroad, which crosses the county from east to west and from north to south: Auburn-

dale, 167 feet; Bartow, 115 feet; Bartow Junction, 165 feet; Ft. Meade, 130 feet; Haines City, 157 feet; Homeland, 139 feet; Lakeland, 206 feet.

WATER-BEARING FORMATIONS.

The deep wells of this county reach the Vicksburg Limestone. Some of the more shallow wells, especially in the Lake Region, receive their water supply from formations lying above the Vicksburg.

ARTESIAN WELLS.

Artesian wells are obtained throughout this county. As a rule, however, the surface elevation is such that the wells do not flow at the surface. Some of the deep wells, in the vicinity of Mulberry, flowed when first drilled, but subsequently ceased to flow, owing to the heavy pumping from surrounding wells. Flowing wells may be obtained in the valley of Peace River and in the eastern part of the county, in the valley of the Kissimmee River. In no part of the world, perhaps, is water from deep wells more extensively used than in the pebble phosphate mining section of Polk County. The wells in this section range in depth from 500 to 800 feet. The water rises in the boring to within 20 to 40 feet of the surface, depending upon the elevation. Pumping is chiefly by air lift.

LOCAL DETAILS.

BARTOW.

The city water supply, at Bartow, is taken from a six-inch well, 720 feet deep. The well is reported cased to the bottom. The water stands 24 feet from the surface. Pumping from this well is direct, the pump being lowered in pit to about six feet of the water level.

CARTERS.

Flowing wells have been obtained at Carters. These flowing wells average in depth about 50 feet, and will flow a few feet above the surface.

LAKELAND.

The public water supply of Lakeland is taken from a sixinch well, 489 feet deep, drilled by C. E. Reed in 1904. The well is cased about 350 feet, and the water stands 79 feet from the surface. The well of the Lakeland Refrigerator and Ice Company is reported to be 336 feet deep. The water in this well stands about 100 feet from the surface.

MULBERRY.

The city water supply at Mulberry is taken from an eight-inch well, 385 feet deep. The water stands in this well 21 feet from the surface. The many wells in this locality, used as a source of water supply in phosphate mining, range in depth, as previously stated, from 500 to 800 feet. In size they vary from eight to fourteen inches. The water stands twenty to forty feet from the surface.

OSCEOLA COUNTY.

LOCATION AND SURFACE FEATURES.

Osceola County includes an area of 1827 square miles. Kissimmee River and the chain of lakes from which it takes its origin forms most of the western boundary of this county. The surface elevation of Kissimmee, at the head of Lake Tohopekaliga, according to levels made by the Atlantic Coast Line Railroad, is 63 feet above sea. Campbell, a few miles west of this lake, is 75 feet above sea. St. Cloud, on East Lake Tohopekaliga, is 63 feet above sea. Narcoossee, also on East Lake Tohopekaliga, is 72 feet above sea.

WATER-BEARING FORMATIONS.

Pleistocene shell deposits are found in the valley of the Kissimmee River, this formation having been recognized at a depth of 100 feet in the well of Captain H. Clay Johnson, at Kissimmee. The formations beneath the Pleistocene have not been determined from well samples, but it is probable that the deep

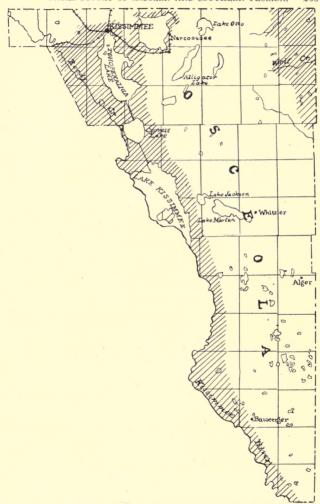


Fig. 15.—Map showing the area of artesian flow in Osceola County. The area in which flowing artesian wells have been obtained is indicated by shading.

wells in this county reach and obtain their chief supply from the Vicksburg Limestone.

AREA OF ARTESIAN FLOW.

Flowing artesian wells are obtained in this county in the valley of the Kissimmee River. It is probable, also, that flowing artesian wells can be obtained in the extreme northeastern part

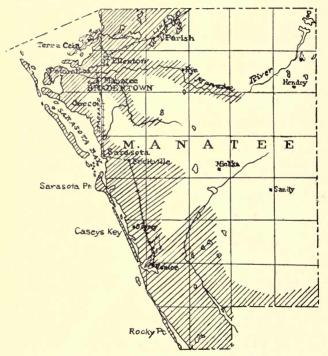


Fig. 16.—Map showing the area of artesian flow in Manatee County. The area in which flowing artesian wells have been obtained is indicated by shading.

of the county, near the St. Johns River. The artesian pressure in the wells in the Kissimmee River valley is sufficient to bring the water from 3 to 7 feet above the surface. These flowing wells vary in depth from less than 100 to 500 and 600 feet.

LOCAL DETAILS.

KISSIMMEE.

Numerous artesian wells have been drilled in and near Kissimmee. These vary in depth from less than one hundred to several hundred feet. The height to which the water will rise above the surface varies from one to three or four feet. The well of the Kissimmee Ice Factory is reported to be 309 feet deep. The water in this well will rise four feet above the surface. The well of H. W. Thurman, at the Granada Hotel, is 341 feet deep and flows at the surface, supplying water for a bathing pool and other domestic purposes. The well of F. Vans Agnew, two miles southeast of Kissimmee, used for domestic and irrigation purposes, is 300 feet deep and yields a strong flow of water at the surface. Many other wells have been drilled for stock, irrigation and domestic purposes in the Kissimmee River valley, and the number is being rapidly increased.

NARCOOSSEE.

Several wells have been drilled at Narcoossee. These vary in depth from 200 to 415 feet. These wells are non-flowing, the elevation here being somewhat greater than at Kissimmee.

MANATEE COUNTY.

LOCATION AND SURFACE FEATURES.

Manatee County lies, bordering the Gulf Coast, between Tampa Bay and Charlotte Harbor. The county includes an area of 1,275 square miles. The principal streams of the county are the Manatee River, which flows across the county from east to west, and enters Tampa Bay and the Myakka River, which flows to the south and enters Charlotte Harbor. It is probable that the northeastern part of the county, near the headwaters of these streams, reaches an elevation of 100 feet above sea. From this part of the county the elevation falls off gradually toward the coast.

WATER-BEARING FORMATIONS.

The deep wells of Manatee County are believed to enter the Vicksburg Limestone. The more shallow wells terminate in the sands and clays before reaching this formation.

AREA OF ARTESIAN FLOW.

Flowing artesian wells are obtained in Manatee County, along the coast and, for some distance inland, along the Manatee and the Myakka Rivers and other streams. The flowing artesian wells along the Manatee River, where a great many have been drilled, vary in depth from 200 to 600 feet. At Sarasota, on Sarasota. Bay, flowing water is obtained at 360 feet.

LOCAL DETAILS.

BRADENTOWN.

The city water supply at Bradentown is obtained from artesian wells, which vary in depth from 410 to 528 feet. The water from these wells will rise about thirteen feet above the surface, equivalent to a head of about twenty-nine feet above sea level. Numerous other wells have been drilled in and near Bradentown, which vary in depth from 200 to 600 feet.

MANATEE.

Numerous artesian wells have been drilled in and around Manatee for household use, irrigation and other purposes. The well of the Excelsior Ice Company, at this locality, is 540 feet deep, although a first flow was obtained at a depth of 360 feet. Mr. C. H. Davis has a four-and-one-half-inch well, 510 feet deep,

cased 150 feet. This well, when measured May 21, 1910, showed a pressure of eight pounds at the surface, which is equivalent to a head of eighteen and one-half feet above the surface. Another well at this locality, having a depth of 529 feet, cased 260 feet, owned by Mr. Tallant, was found, on the same date, to have a pressure of seven and one-half pounds, or a head above the surface of seventeen and three-tenths feet. The relative elevation of these two wells was not determined, but the surface elevation at the Tallant well is estimated at about six feet above sea.

PALMETTO.

The city water supply at Palmetto is taken from artesian wells. In addition to the city supply, several artesian wells have been drilled at this locality. These vary in depth from 370 to 600 feet. The water is reported to rise 20 to 25 feet above the surface.

SARASOTA.

The well from which the public water supply is taken at Sarasota is reported to have a depth of 450 feet. The water rises about twenty feet above the surface. Other wells drilled at this locality vary in depth from 360 to 400 feet. A flowing well, drilled on Sarasota Key, is reported to be 252 feet deep. The water from this well rises 15 feet above the surface.

DESOTO COUNTY.

LOCATION AND SURFACE FEATURES.

DeSoto County has an area of 3,755 square miles, and extends from the Gulf of Mexico to Lake Okeechobee and the Kissimmee River. The Lake Region extends into the north central part of this county. It is probable that local areas are found in the Lake Region of this county which exceed 150 feet in elevation. From these high lands the slope is gradual to the Gulf and to Lake Okeechobee and to the Kissimmee and the Caloosahatchee Rivers. The following elevations are recorded along the Atlantic

Coast Line Railroad, which crosses the county from north to south: Arcadia, 56 feet; Bowling Green, 116 feet; Ft. Ogden, 37 feet; Nocatee, 38 feet; Wauchula, 107 feet; Zolfo Springs, 61 feet.

WATER-BEARING FORMATIONS.

As elsewhere in Southern Florida, the deep wells obtain their water supply from limestones of the Vicksburg formation.

AREA OF ARTESIAN FLOW.

DeSoto County includes a considerable area, in which flowing artesian wells can be obtained. This flowing area surrounds Charlotte Harbor and in the valley of the Peace River extends entirely across the county. Flowing wells are also obtained along the Caloosahatchee River to Lake Okeechobee. It is also believed that flowing artesian wells may be expected along the west border of Lake Okeechobee and in the valley of the Kissimmee River,

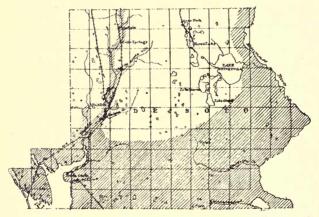


Fig. 17.—Map showing the area of artesian flow in DeSoto County. The area in which flowing artesian wells have been obtained is indicated by shading.

along the east border of this county. The deep wells at Punta Gorda show a pressure of 20 pounds or more, indicating a head of 45 to 50 feet above sea. In the interior of the county, where the elevation is greater, the rise of the artesian water above the surface is correspondingly less. The accompanying map shows approximately the area of artesian flow in the county. It is probable that flowing wells can be obtained over a somewhat larger area than is here indicated. Owing to the fact that no topographic map has been made of this county, and comparatively few wells have been drilled, it is impossible to closely outline the flowing area.

LOCAL DETAILS.

ARCADIA.

The city water supply at Arcadia is taken from an eight-inch well, 375 feet deep. The elevation at Arcadia is given by the Atlantic Coast Line Railroad as 56 feet above sea, and the water in the city well is reported to rise to within one foot of the surface. A number of other wells are reported from the vicinity of Arcadia ranging in depth from 215 to 380 feet. The water from these wells rises to within one to seven feet of the surface. In the valley of the Peace River, near Arcadia, flowing wells are obtained, the water rising from seven to ten feet above the surface.

FT. OGDEN.

The surface elevation at the depot at Ft. Ogden is given as 37 feet above sea. A well, 280 feet deep, located one-half mile west of Ft. Ogden, and belonging to Carr & Williams, flows six or more feet above the surface. The second well, 289 feet deep, belonging to Russell & Windsor, is said to flow 14 feet above sea.

NOCATEE.

Flowing wells are obtained at Nocatee. A well of the DeSoto Fruit Company, one-half mile east of Nocatee, 355 feet deep, flows eight feet above the surface. The well of the Nocatee Fruit

Company, a few miles east of Nocatee, 300 feet deep, also flows eight feet above the ground.

PUNTA GORDA.

The city water supply at Punta Gorda is taken from a sixinch well, 484 feet deep. The well is cased 240 feet. The water from this well is reported to rise about 40 feet above the surface. An eight-inch well, owned by the DeSoto Manufacturing Company, is 430 feet deep. Water from this well is reported to rise about 50 feet above the surface. Numerous other artesian wells have been drilled in and near Punta Gorda, varying in depth from 265 to 600 feet.

PALM BEACH COUNTY.

LOCATION AND SURFACE FEATURES.

Palm Beach County extends from the Atlantic Ocean to Lake Okeechobee, and includes an area of 2,809 square miles. The western part of the county extends into the Everglades of Florida.

WATER-BEARING FORMATIONS.

Samples obtained by N. H. Darton, many years ago, from the well of C. I. Craigin, at Palm Beach, afford practically the only information available regarding the deeper formations of this county. The Vicksburg Limestone is believed to have been reached in this well between 915 and 1,000 feet. The material above this level was scarcely determinable, although apparently the Miocene and, presumably, other formations are represented. The limestone, lying near the surface in the eastern part of this county, is of Pleistocene age, and is known as the Palm Beach Limestone.*

^{*}Second Annual Report, Florida Geol. Surv., p. 209, 1909.

AREA OF ARTESIAN FLOW.

Flowing artesian wells have been obtained in Palm Beach County, along the coast as far south as Palm Beach. The depth to the Vicksburg Limestone, which is the chief water-bearing formation, increases in passing south to east, owing to the dip of the formation in that direction. The Vicksburg at Palm Beach is reached, as previously stated, between 915 and 1,000 feet. In the northern and western parts of the county this formation may be expected at a lesser depth, and it is probable that flowing artesian wells may, ultimately, be obtained throughout all of the northern and much of the western parts of Palm Beach County.

LOCAL DETAILS.

GOMEZ.

A well drilled at Gomez in 1900, by John McAllister, is reported to have reached a depth of 1,200 feet. This is a four-inch well and is cased 300 feet. The water, which is slightly brackish, is reported to flow 20 feet above the surface.

HOBE SOUND.

A well near Hobe Sound, belonging to T. A. Snider, and drilled in 1895 by Near & Taylor, reached a depth of 1,100 feet. This is a four-inch well and the water, which is slightly salty, is reported to rise 12 feet above the surface.

PALM BEACH.

The following is a log of the artesian well of C. I. Cragin, two and one-fourth miles north of Palm Beach. The well is 1,212 feet deep, four inches in diameter and is cased 846 feet. The original four-inch casing having rusted out, is now replaced by a line of 2½-inch casing. At the depth of 1,140 feet the four-inch bore hole was reduced to three inches, making the well three inches in diameter from the depth of 1,140 to the bottom of the well, 1,212 feet. The well was commenced in 1889 and finished in 1890 by J. A. Durst, driller:

Feet from surface. Character of material.

0 - 5 .. Surface soil.

5 - 7 .. Rock,

7 - 8 .. First sand.

8 - 36 .. Mostly fine coquina rock.

36 - 57.2 .. Ouicksand and sharp pieces of stone.

57.2 - 58 .. First really hard rock.

58 - 76.10. Coquina, alternating with sandy strata.

76.10- 78 .. Hard rock.

78 - 78.6 .. Very hard flint.

78.6 - 84 .. Sand, white and solid, but not hard.

84 - 96 .. Quicksand bed, mixed with bits of coarser material.

96 - 96.6 . Flint rock, thin.

96.6 - 97 . Fine sand.

97 - 148 ... Ouicksand bed.

148 - 151 ... Solid limestone.

151 - 169.6 . . Soft gray limestone.

169.6 - 170 ... Hard rock.

170 - 171 ... Shell stratum.

171 - 171.3 .. Very hard sandstone.

171.3 - 175 . . Sandstone.

175 - 185 . . Alternately hard and soft limestone,

185 - 190 . . Straw colored sandstone.

190 - 238 ..Drab colored solid sandstone, gradually deepcning in its color to a final blue at 230 feet, with small delicate shells throughout.

238 - 238.8 ..Bed of small dainty shells. Water level is 3 feet
4 inches below wooden curb.

238.6 - 248 ..Very hard drilling, required to move casing in these alternations. Water level above 20 inches (near 242 feet). Very active quick-sand.

248 - 250 .. Took out loads of quicksand.

250 - 262 ... Sand. Water in this sand ran slowly out of pipe at 3 feet 6 inches above ground.

262 - 263 .. Coquina.

263 - 300 ..Broken shell and sand, more shell (white and pulverized), the last few feet. Water level just above ground level.

300 - 301.6 ..Rock, water stands 2 feet 4 inches above curb in

301.6 - 303.6 . . Brown clay, first seen in this well,

303.6 - 310 . . Sand.

310 - 312.4 .. Blue sandstone.

312.4 - 312.10.. Blue sand, shells and pieces of rotten sticks.

312.10- 315.2 .. Blue sandstone,

315.2 - 320 ...Sand, water in this sand stands 3 feet above wooden curb.

320 - 321 .. Blue sandstone.

321 - 340 ...Fine shell and sand, coarser broken shell toward bottom.

340 - 340.3 .. Rock.

340.3 - 350 ...Coarse broken shell, blue pebbles and pieces of coquina, water 2 feet above curb, runs freely at 1 foot above.

350 - 357 ... Yellow sandstone, water 2 feet above curb.

357 - 359 .. Broken shell, pebbles, pieces of coquina.

359 - 373 .. Pulverized shell.

373 - 374 ...Gray limestone, with some broken shell lying immediately beneath, water stands 2 feet 4 inches above on penetrating this rock.

374 - 392 ...Pulverized shell, water stands at level of wooden curb.

392 - 400 .. Alternations of rock and blue marl.

400 - 409 .. Blue marl.

409 - 432 ... Alternations of blue marl and sand which afforded
the greatest flow to date and the first fresh
water below 49 feet.

432 - 507 .. Blue marl.

507 - 510 .. Coquina.

510 - 542 .. Proportion of sand in the marl increases very much.

542 - 571 ...Quicksand, below casing, can not drill at all. Proportion of sand in the marl increases.

571 - 614 ... Marly sand. Head of water from 9 to 11 feet above ground. Water rises so as to dribble from a height of 11½ feet.

614 - 618 ..Quicksand bed.

618 - 618.6 . . Rock.

618.6 - 640 .. Sand or sandstone.

640 - 707 ...Lighter colored and runs to greenish marly sand all through here. At depth 678-688 more sand, water from 690-700, very many tiny spiral shells.

707 - 710 .. Brown, coarse material.

710 - 794.6 . . Sand with enough marl with it to give a green color to the slush as ejected.

794.6 - 809 ...Loose sand full of black specks and tiny bivalve and spiral shells.

809 - 826 .. Blue marl full of black specks.

826 - 828 .. Sand.

828 - 834 ... Sandstone.

834 - 839 .. Very fine, tough clay, thoroughly impervious.

839 - 860 ..Fine grained coquina, get dribble of water at depth of about 844 feet 4 inches, casing driven to depth of 846 feet, tight in rock.

860 - 867 .. Solid hard limestone.

867 - 874 ...Fine clay, devoid of grit.

874 - 876 .. Hard rock.

876 - 878 ...Lots of black specks here.

878 - 902 .. Clays, sandy and lots of black specks, no water.

902 - 905 ...Dark sand bed; here the water supply is 115,000 gallons per diem.

905 - 917.6 .. Thin block of stone 909 feet, about. This is the lowest sand bed with thin block of limestone at intervals. Water comes from between these thin flakes of limestone.

917.6 - 917.9 ..Limestone.

917.9 - 923 .. Coralline.

923 - 961 ..Hard limestone rock at 923, solid rock nearly 39 feet.

961 - 973 ..Gritty marl.

973 - 990.6 . . Solid rock.

990.6 -1009 ... Sandy marl, full of tiny spirals.

1009 -1012 ..Limestone.

1012 -1023 ..Yellow sandstone.

1023 -1025 .. Hard rock.

1025 -1088 ..Rock, first of the regular water strata. Alternating hard and soft strata. Increase of water with depth. At depth of 1042 feet 270,000 gallons, 1057 feet 300,000 gallons, 1075 feet 350,000 gallons; water strata found at frequent intervals.

1088 -1110 ...Gray limestone.

1110 -1116 ..Gray limestone interspersed with water strata,
but the flow increases but slightly. At 1160
feet flow total 400,000 gallons.

1116 -1174 ... Solid gray limestone.

1174 -1175 ..Blue limestone.

1175 -1193 .. All solid.

1193 -1195.6 .. Blue limestone (?). 1195.6 -1196 .. Six inches water stratum.

1196 -1212 ...Mostly gray limestone, with some hard and some water strata, flow increases but little.

The following is an analysis of the water from this well. Analysis made in the office of the State Chemist, A. M. Henry, analyst.

Colorless, odorless, slightly salty taste, no sediment.

Milligrams per	liter.
SiO ₂	
C1	1337
SO ₄	431
PO ₄	3
CO ₃	0
HCO ₃	195
Na and K	835
Mg	
Ca	102
Fe and Al	2
Loss on ignition	357
_	
Total dissolved solids	3000

WEST JUPITER.

The following is an analysis of the water from Weybrecht's well, at West Jupiter, 57 feet deep. Analysis by the American Water Softener Company, Philadelphia, Pa., July 23, 1908.

	Grains per	Parts per
	U. S. gallon.	million.
Total solids	. 62.50	1071.49
Calcium carbonate	. 15.75	270.01
Calcium sulphate	. 3.13	53.56
Calcium chloride	. 2.47	42.34
Magnesium carbonate	. 5.86	100.46
Sodium chloride	. 30.40	521.17
Free carbonic acid	. 1.22	29.48
Iron, alumina and silica	. 1.68	28.80
Incrusting solids	. 28.89	495.28
Non-incrusting solids	. 30.40	521.17

YAMATO.

The following is a log of a well at Yamato, drilled by the Florida East Coast Railway. The well is cased 65 feet and the water stands nine feet below the surface.

	Feet.
Sand	0 -24
Yellow clay	24 -34
Sand and shell	34 -40
Rock	40 -41
Gravel	41 -45
Rock	45 -461/2
Gravel and rock	
Quicksand	
Rock	
Sand	67 -74
Rock	75 -88

The following is an analysis of the water from this well made by the American Water Softener Company, Philadelphia, Pa., November 3, 1909:

, , , , , , , , , , , , , , , , , , ,		No.
	Grains per	Parts per
	U. S. gallon.	million.
Calcium carbonate	. 7.22	123.77
Calcium sulphate	. 0.54	9.25
Calcium chloride	. 0.78	13.38
Magnesium carbonate	. 0.73	12.51
Sodium chloride	. 0.81	13.78
Free carbonic acid	. 0.56	9.60
Iron, alumina and silica	. 0.23	3.94
Incrusting solids	. 9.50	162.87
Non-incrusting solids	. 0.81	13.78

LEE COUNTY.

LOCATION AND SURFACE FEATURES.

Lee County lies bordering the Gulf of Mexico and extends inland to Lake Okeechobee. The area of the county is 4,641 square miles. The surface elevation in the northeastern part of the county approximates 25 feet above sea level. No topographic map has been made of the county, but the surface is prevailingly level with, in general, a slope toward the coast.

WATER-BEARING FORMATIONS.

The artesian wells in this county are believed to obtain their chief supply from the Vicksburg formation.

AREA OF ARTESIAN FLOW.

Flowing wells have been obtained over an extensive area throughout the interior of Lee County, as well as along the Caloosahatchee River, along the northern border of the county. It is believed that almost the whole of this county may be included in the artesian flow area.

LOCAL DETAILS.

BOCA GRANDE.

Three deep wells have been drilled at Boca Grande, on Gasparilla Island. The first of these, drilled in 1910, is located 200 feet north of Boca Grande station, and was drilled by G. H. Southard. This well is 1.030 feet deep and is reported cased 800 feet. The well yields a heavy flow of salty water. The second deep well at this locality, drilled in 1911 by F. S. Gilbert, is located 600 feet south of Boca Grande station. This well is 1,220 feet deep and yields a flow of 450 gallons per minute of salty water. The temperature of the water at 1,220 feet was 89 degrees Fahrenheit. The driller, F. S. Gilbert, reports that he cased twenty-two times in drilling this well, the casing being driven and pulled at each show of water in order to test for fresh water. The well, as completed, was cased with six-inch casing to a depth of 1,200 feet. The third well, also drilled by F. S. Gilbert, is located 2.700 feet north of the station. This well is 1,812 feet deep and is cased 1,500 feet. The water is salty. The temperature was 90 degrees Fahrenheit at 1,800 feet. The flow from these wells rises about fifteen feet above sea level. These wells enter the Vicksburg Limestone, and the deepest of the wells apparently does not pass through the Vicksburg Limestone.

FT. MYERS.

The public water supply at Ft. Myers is taken from drilled wells, of which three are in use at present. Three additional wells are available as a reserve supply. These latter vary in depth from

487 to 587 feet. The water from these wells will rise about 45 feet above the surface. A well near Ft. Myers, belonging to Thomas A. Edison, reaches a depth of 648 feet. The water from this well will rise about 45 feet above the surface. Many additional wells have been drilled in and around Ft. Myers; these vary in depth from 400 to 960 feet. The water from these deep wells rises 40 to 50 feet above the surface.

LABELLE.

Flowing wells are obtained at Labelle and elsewhere, along the Caloosahatchee River. D. G. McCormick & Company have a flowing well, about a mile north of the east end of Lake Flirt (T. 42, R. 30, S. 19). This is a three-inch well, 490 feet deep. The well is cased 450 feet and the water is reported to rise 40 feet above the surface. The strong flow reported for this well indicates that flowing wells may be expected over a considerable area, north of the Caloosahatchee River and west of Lake Okeechobee.

THE KEYS.

A number of wells have been drilled on the keys in Lee County. Those at Boca Grande, on Gasparilla Island, have already been described. Two wells are reported to have been drilled on Sanibel Island. One of these belonging to F. P. Bailey, reached a depth of 500 or 600 feet. The second well, belonging to Harry Bailey, is 500 feet deep. The water from both of these wells is said to be brackish. On Useppa Island a fresh water well was obtained by W. H. Towles, at a depth of 250 to 300 feet. A second well on this island, reaching a depth of 400 feet, was said to have been somewhat brackish.

Two wells are reported from St. James Island. One of these is 184 feet deep, the other is 344 feet deep. Both yield fresh water. A well on Bucks Key reaches a depth of 600 feet. The water in this well is reported to rise 20 feet above the surface.

DADE COUNTY.

LOCATION AND SURFACE FEATURES.

Dade County lies in Southern Florida, bordering the Atlantic Coast. The county includes an area of 2,305 square miles. The western part of the county reaches into the Everglades of Florida. East of the Everglades the surface formation is chiefly the Miami Oolitic Limestone.

WATER-BEARING FORMATIONS.

The limestones exposed at the surface, in Dade County, are of Pleistocene age and it is probable that most of the wells terminate without passing through these Pleistocene formations. The deepest well recorded in Dade County is a well drilled recently at Homestead by the Florida East Coast Railway. This well reached a depth of 300 feet, but the age of the formation in which it terminated was not determined.

ARTESIAN WELLS.

The water in the wells at the city waterworks, at Miami, rises to within fourteen inches of the surface level and flows into the collecting basin excavated for that purpose. The possibility of getting flowing artesian water from the Vicksburg Limestone, which lies at a depth of several hundred feet, has not been tested by deep borings.

LOCAL DETAILS.

DANIA.

A well has been drilled at Dania by the Florida East Coast Railway, to a depth of 54 feet. The following is an analysis of the water from this well made by the American Water Softener Company, Philadelphia, Pa., November 3, 1909:

		Grains per U. S. gallon.	Parts per million.
Total solids		. 17.50	300.01
Calcium carbonate		. 13.70	234.87
Magnesium carbonate		77	13.20
Sodium chloride		. 2.56	43.88
Sodium carbonate		07	1.20
Free carbonic acid		. 1.56	26.74
Iron, alumina and silica		09	1.54
Incrusting solids		. 14.56	249.61
Non-incrusting solids	٠.	. 2.63	45.08

An analysis of the second sample of water from this well, made by the Dearborn Drug and Chemical Works, Chicago, Ill., July 2, 1910, is as follows:

· , · · · · · · · · · · · · · · · · · ·		
	Grains per	Parts per
	U. S. gallon.	million.
Silica	327	5.506
Oxide of iron and alumina	140	2.400
Carbonate of lime	13.058	233.865
Sulphate of lime	None	None
Carbonate of magnesia	433	7.423
Sodium and potassium sulphates	212	3.634
Sodium and potassium chlorides		40.802
Sodium and potassium carbonates		6.326
Loss, etc.	307	5.263
Total mineral solids	. 17.286	296.350
Organic matter	Trace	Trace
Total incrusting solids	. 13.958	239.294
Total non-incrusting solids	3.328	57.054

The following is a log of the well at Dania, obtained through the courtesy of Mr. G. A. Miller of the Florida East Coast Railway:

	Feet.
Sand	0- 6
Hard pan	6-8
Shell and rock	8-20
White rock	20-24
Shell, coarse sand and water	24-31
Rock	21 25
Sand and shell	25 40
Rock	33-40
	40-42

Sand and shell	42-52
Gravel	52-54
Hard rock	54-591/2

HOMESTEAD.

An experimental well was drilled at Homestead by the Florida East Coast Railway to a depth of 320 feet. The following is the analysis of the water from this well, at the depth of 16, 46, 66 and 320 feet. Analyses by the American Water Softener Company, Philadelphia, Pa.

No. 1, sample of water from the depth of 16 feet:

rior 1, bumple of water from the depth	or ro rece.	
	Grains per	Parts per
	U. S. gallon.	million.
Total solids	. 13.60	233.15
Calcium carbonate	. 9.85	168.85
Calcium sulphate	. 0.22	3.77
Calcium chloride	. 1.42	24.34
Calcium nitrate	. 0.48	8.22
Magnesium carbonate	. 0.91	15.59
Iron, alumina and silica	. 0.90	15.42
Incrusting solids	. 13.52	221.77

No. 2, sample of water from the depth of 45 feet. May 25, 1911:

*	Grains per	Parts per
	U. S. gallon.	million.
Total solids	13.50	185.24
Calcium carbonate	. 10.14	173.83
Calcium sulphate	0.22	3.77
Calcium chloride	. 1.32	22.62
Magnesium carbonate	0.45	7.71
Sodium chloride	0.66	11.31
Free carbon dioxide	0.90	15.42
Iron, alumina and silica	0.19	3.25
Incrusting solids	12.32	211.21
Non-incrusting solids	0.66	11.31

No. 3, sample of water from depth of 66 feet. June 29, 1911:

	Grains per	Parts per
	U. S. gallon.	million.
Total solids	14.00	240.01
Calcium carbonate	10.80	185.15
Calcium sulphate	0.39	6.68
Calcium chloride	0.77	13.20
Calcium nitrate	0.39	6.68
Magnesium carbonate	0.39	6.68
Free carbon dioxide	. 0.53	9.08
Iron, alumina and silica	. 0.89	15.25
Incrusting solids	. 13.63	233.67

No. 4, sample of water from the depth of 320 feet. August 4, 1911:

	Grains per	Parts per
	U. S. gallon.	million.
Total solids	. 57.40	984.05
Calcium carbonate	. 5.08	87.09
Magnesium carbonate	. 3.34	57,26
Sodium carbonate	. 14.35	246.01
Sodium sulphate	. 15.76	270.18
Sodium chloride	. 15.76	270.18
Free carbon dioxide	. 0.26	4.45
Iron, alumina and silica	. 0.25	4.28
Incrusting solids	. 8.76	150.18
Non-incrusting solids	. 45.87	786.39

The water at the depth of 320 feet being unsuited for boiler use, the well was plugged and a more shallow water is being used.

The following is a log of this well, supplied by Mr. G. A. Miller, of the Florida East Coast Railway:

Soft rock	0 - 10
Hard rock	20 - 30
Medium hard rock	30 - 40
Hard rock	40 - 50
Medium hard rock	50 - 551/2
Hard rock	551/ 591/
	0072- 0072

Sand	581/2	- 59	
Soft rock with sand pockets	59	- 62	
Loose rock and sand	62	- 66	
Sand	66	- 81	
Loose sand and rock	81	- 84	
Marl	84	- 841/	4
Sand	841/4	- 92	
Marl and shell	92	-115	
Gray clay with small amount of fine sand	115	-160	
Clay and marl	160	-167	
Marl containing a small quantity sand and shell. Sand			
increasing with depth	167	-197	
Marl or soft chalky rock	197	-204	
Tough slate colored clay	204	-217	
Marl containing sand, shell and gravel	217	-232	
Marl or soft chalk-like rock	232	-237	
Marl and sand	237	-240	
Slate colored clay	240	-268	
Clay	268	-294	
Marl and clay	294	-298	

MIAMI.

The public water supply at Miami is taken from seven wells, located on the north side of Miami River, about one and one-half miles west of the city. The principal supply of fresh water in these wells is obtained at a depth of about 85 feet, although some water is reported at 30 and at 80 feet. The water rises to within 14 inches of the surface and flows into a receiving basin. At 90 feet, in well number 7, recently drilled, salt water was reached. This well was plugged and fresh water admitted from above. The following notes were made from occasional samples from one of these wells. The samples were kept by the Florida East Coast Hotel Company:

The following is an analysis of the water from the Miami wells. Analysis by the American Water Softener Company, Philadelphia, Pa.:

	Grains per	Parts per
	U. S. gallon.	million.
Total solids	. 17.50	300.01
Calcium carbonate	. 12.68	217.38
Calcium sulphate	. 0.21	3.60
Calcium chloride	. 0.83	14.22
Magnesium carbonate	. 0.59	10.11
Sodium chloride	. 2,20	37.71
Free carbon dioxide		10.28
Iron, alumina and silica		3.08
Incrusting solids		156.98
Non-incrusting solids	. 2.20	37.71

MONROE COUNTY.

LOCATION AND SURFACE FEATURES.

Monroe County lies along the Gulf Coast, at the extreme southern end of Florida. The area of the land surface, including the numerous keys, is about 1,125 square miles.

WATER-BEARING FORMATIONS.

The Key Largo Coralline Limestone and the Key West Oolitic Limestone make up the surface formations along the keys. On the mainland the Lostman's River Limestone lies near the surface. The deep wells at Key West reach the Vicksburg Limestone.

ARTESIAN WELLS.

No flowing artesian wells have been reported from Monroe County. It is probable, however, that flowing wells could be obtained in the northern part of the county and along the Gulf Coast. Several wells have been drilled on the keys, along the line of the Florida East Coast Railway. None of these, however, have been successful in obtaining either flowing or fresh water.

LOCAL DETAILS.

KEY VACA.

Two deep wells have been drilled by the Florida East Coast Railway at Marathon, on Key Vaca. One of these wells reached a depth of 425 feet, the other 700 feet. The following is a combined record of these two wells by Samuel Sanford, who was in charge of the drilling. The log is republished from the Second Annual Report of this Survey, page 205:

	reet.
Reef rock	0-105
Hard to soft white limestone, with much white marl	105-148
Soft white limestone with shell casts	148-150
Medium hard white limestone, shell casts and shell frag-	
ments	150-155
Soft white limestone with quartz grains, proportion of	
quartz increasing with depth, shell fragments and	
casts	155-176
Medium fine quartz-sand containing numerous irregular	
nodules, with yellowish marly sand at 210 to 215 feet.	176-230
Quartz sand in a varying proportion of limy mud, sand	
grains, colorless mud, vellowish to dark green;	
streaks and beds of friable sandstone containing shell	
casts; bed of oyster shells at 240 feet	230-300
Quartz sands or beds of soft, friable sandstone, contain-	
taining shell casts; streaks of dark green, limy clay,	
306-310 feet; beds of shells, few determinable fossils,	
probably Miocene, 378-390	300-400
,	

Quartz sands as below 230 feet, beds of soft friable sand-	
stone with shell casts; gravel bed with much worn	
pebbles up to 40 mm. long; tough green, limy clay at	
407 to 410 feet	400-435
Quartz sands with little sandstone, tough, dark clay in	
occasional streaks	435-700

KEY WEST.

Two deep wells have been drilled at Key West. The first of these, drilled in 1895, is reported to have reached a depth of 2,000 feet. The water obtained from this well was too salty for drinking purposes, but is used for fire protection. The following is a log of this well, taken from the Second Annual Report of this Survey, page 206, abbreviated from the detailed description given by E. O. Hovey, of samples from this well:

	Fee	et.
Yellowish oolite	0-	25
White yellowish or light gray limestone, with oolitic		
lumps	50-	175
Fine white lime-sand rock	175-	200
White, porous oolitic and sandy limestone	200-	275
White, more or less solid limestone	300-	375
Friable soft gray lime-sand rock	400-	675
Yellowish to brownish lime-sand rock, Orbitoides, 800		
to 850 feet	700-1	075
Light gray, partly dense and partly porous limestone	1100-1	175
Gray lime-sand rock	1200-1	350
Yellowish gray lime-sand rock, with some porous lime-		
stone	1375-1	450
Lime-sand rock, varying in color and compactness, with		
strata of dense limestone	1475-1	975
Yellowish to light brownish-gray limestone, rather		
solid, with porous portions	1975-2	000

A second deep well was drilled at this locality by J. T. Brown for S. O. Johnson. This well is 1,010 feet deep and reached salty water. Occasional samples of the drillings from this well to a depth of 540 feet were forwarded to the Florida State Geological Survey. Below 540 feet only one sample was received, which was submitted as representing the material from 800 to 1,010 feet. The following partial log is made up from these occasional samples:

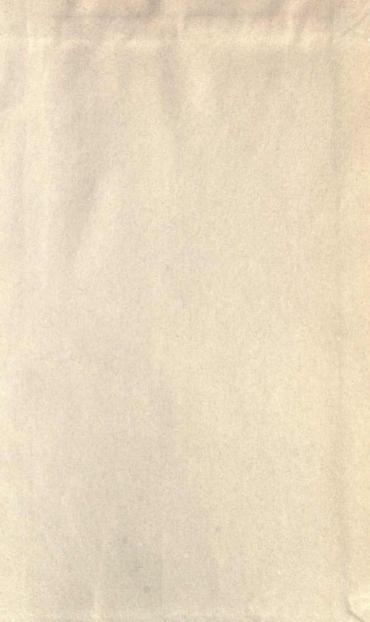
	Depth from w	hich th
	Character of rock. sample was	
		eet.
	Oolitic limestone with shell fragments. In color the oolitic	
	grains vary from light to pinkish. The sample contains	
	little or no quartz sand	30
	Soft limestone powdered very fine, not so conspicuously	
	oolitic as last sample	50
	Oolite, light and pinkish oolite grains	70
	Light colored oolitic limestone with fragments of shells	80
	Oolitic limestone with fragments of shell. Oolite grains vary	
	in color, from light to pinkish	100
	From 100 to 210 feet no fine material was brought up by the	
	drill. A salty sulphur water was reached at this depth,	
	and the fine material carried away apparently in the water.	
	The coarse pieces brought up in this distance were as	
	follows:	
P	iece of coral and limestone, consisting of fragments of shells	
	and other organisms, also pieces of dark-colored lime-	
	stones	
	Pieces of hard crystallized limestone and fragment of coral	150
	Rough white limestone pieces with shell fragments, also pieces	
	of limestone made up of a mass of shell fragments; also	
	colitic limestone	175
	Oolitic limestone with shells and shell fragments, also rough	
	white limestone and partially crystallized limestone	200
	Rough white limestone with shell fragments, partially crys-	
	tallized	
	Rough white limestone with shell fragments	260
	Oolite grains light and pinkish in color; also pieces of rough	0.00
	limestone	
	Mass of calcium crystals, stained brownish yellow	325
	Rough light-colored limestone pieces with fragments of shells	240
	and of corals	340
	Rough light-colored limestone pieces with fragments of shells,	
	corals, worm tubes, and light and pinkish oolite with	0.50
	admixture of greenish-gray calcarous material	350
	Greenish-gray calcareous sand, with occasional oolite grains	270
	imbedded, but no fossils and no siliceous sand	370
	Light-colored limestone with fossils and pieces of typical	200
	oolite	380

Gray calcareous sand with slight admixture of siliceous sand	390
Same gray calcareous sand with some pieces of impure light-	
colored limestone and with one calcite crystal	400
Same as above, gray calcareous sand, with some light-colored	
limestone	425
Light, rough limest with fossils and typical oolite	450
Gray calcareous sand and light limestone	475
Same gray calcareous sand with light-colored limestone	515
Same gray calcareous sand with fine siliceous sand	530
Same as above	540

The sample submitted, as representing the material from 800 to 1,010 feet, is limestone, apparently of the Vicksburg formation.







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