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T H E

AMERICAN JOURNAL OF SCIENCE

[FOURTH SERIES.]

ART. I.—*On the Energy of the Cathode Rays*; by W. G. CADY.

[An investigation carried out at the Physical Institute of the University at Berlin.]

IT has long been known that the cathode rays possess a considerable amount of energy; yet the relation between this and the other discharge phenomena has hitherto received but little attention. Ebert and Wiedemann,* E. Wiedemann,† and Ewers‡ measured the energy of cathode rays by means of calorimeters. For exact work, however, the thermo-element is to be greatly preferred. Among the recent investigations based upon the emission hypothesis, this method has been employed by J. J. Thomson§; O. Berg has also used a thermo-element in his work with the cathode rays.||

According to the emission hypothesis there exists the following relation between the energy of the cathode rays, potential of discharge, and amount of electricity transported by the rays. Suppose that we have a bundle of rays given off from a cathode of potential V , and striking a metallic conductor which is led to earth through a galvanometer; let the conductor be struck by N particles per second. We will at first assume that the entire charge of the particles, as well as all their energy in the form of heat, is given up to the conductor. Then if v [cm. sec⁻¹] be the velocity of the particles, and m

* H. Ebert and E. Wiedemann, Sitzungsber. der phys. med. Soc. zu Erlangen, Dec. 14, 1891.

† E. Wiedemann, Wied. Ann., lxvi, p. 61, 1898.

‡ P. Ewers, Wied. Ann., lxix, p. 167, 1899.

§ J. J. Thomson, Phil. Mag., xlv, p. 293, 1897.

|| O. Berg, Ber. d. naturforsch. Gesellsch. zu Freiburg i. Br. XI, vol. ii, p. 73, July, 1899.

[gr.] their mass, we have for the amount of heat given up to the conductor

$$Q = N \cdot \frac{1}{2}mv^2 \quad (1)$$

(ergs per second).

If the charge on each particle be ϵ , we shall have flowing to earth through the galvanometer a "cathode-current" of intensity

$$i = N \cdot \epsilon \quad (2)$$

Now the kinetic energy of each particle is

$$V \cdot \epsilon = \frac{1}{2}mv^2. \quad (3)$$

Hence

$$Q = N \cdot V \cdot \epsilon = i \cdot V,$$

or

$$\frac{iV}{Q} = 1. \quad (4)$$

According to Starke's investigations,* a part of the cathode rays are reflected. This can have no effect upon the above, provided that after reflection the velocity of the rays remains unchanged; and indeed the measurements made by Merritt† upon the deflection of reflected cathode rays seem to show this to be the case. We will, however, assume for the present that upon reflection the energy of each particle is diminished in the ratio $r' : 1$, while the charge remains unchanged. Then if the fraction r of the rays be reflected, we have

$$Q = N(1-r) \cdot \frac{1}{2}mv^2 + Nr(1-r') \cdot \frac{1}{2}mv^2 = \frac{1}{2}Nmv^2(1-rr') \quad (1a)$$

$$i = N\epsilon(1-r) \quad (2a)$$

$$V\epsilon = \frac{1}{2}mv^2 \quad (3a)$$

$$\frac{iV}{Q} = \frac{1-r}{1-rr'}. \quad (4a)$$

If $r' = 1$ (Merritt), we have still $iV/Q = 1$. If in the limiting case $r' = 0$, then $iV/Q = 1-r$.

These relations were investigated experimentally as described in the following paragraphs.

Apparatus.

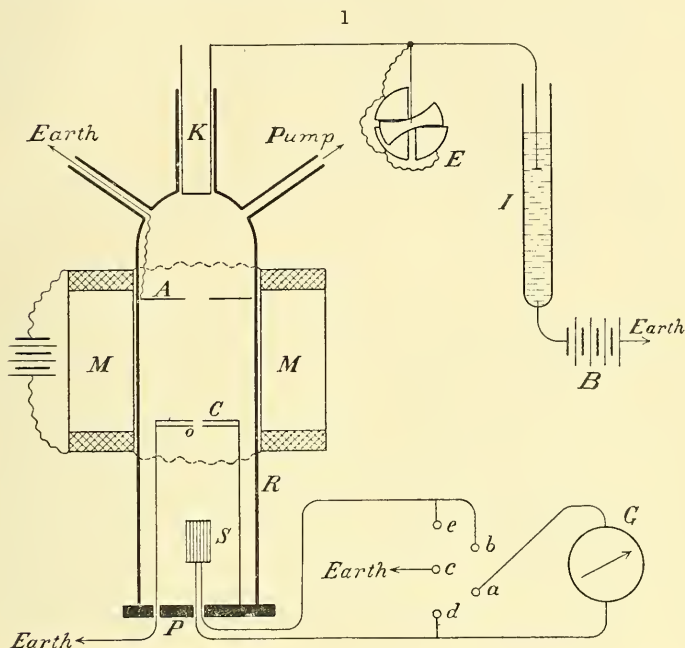
Fig. 1 shows the arrangement of apparatus at first used. For measuring the energy Q a thermo-element was here employed; it was later displaced by a bolometer (see below.) The thermo-element, S , was of the Melloni type, consisting of 49 bismuth-antimony couples each 25^{mm} long; the exposed surface of the junctions was 196^{mm}². It was placed inside the

* H. Starke, Wied. Ann., lxi, p. 49, 1899.

† E. Merritt, Phys. Rev., vii, p. 217, 1898.

Faraday cylinder *C*, which was connected to earth. The rays entered through the hole *o*, 6^{mm} in diameter. *R* is the glass discharge-tube, 25^{cm} long, cemented to the glass plate *P*.

The copper bottom of the brass tube *K* served as cathode; in order to reduce the heating as far as possible, a stream of water was kept constantly flowing through the tube. The anode *A* was a zinc diaphragm.



The discharge current was furnished by a secondary battery *B* of 2288 cells, supplemented at times by a rotary transformer, so that a maximum E. M. F. of 5400 volts could be reached. The current was regulated by the Hittorf cadmium-iodide resistance *I*, and the potential of the cathode measured by means of the quadrant-electrometer *E*. *G* is a Thomson galvanometer used in measuring the thermal and cathode-currents.

A serious source of error in observations lay in the evolution of gas from the cathode and other parts of the apparatus during discharge. This difficulty, so common in all gas-discharge work, was nearly always present, even after the apparatus had been in use for many hours. A partial remedy was found in running the air-pump slowly during observations, but even then the potential usually changed somewhat during the

minute or so necessary to take a set of readings. This is the chief cause of the irregularities in the numerical values given below.

Before the apparatus was put together, the thermo-element was calibrated as follows. The radiation from the blackened surface of a Leslie cube, filled with boiling water, passed through a series of diaphragms and fell upon the thermo-element. The constant of radiation for lampblack at 100° was taken according to Kurlbaum* as 0.0176 gr.-cal./sec. From the dimensions of the diaphragms could then be calculated, by use of the Stefan-Boltzmann law, the amount of heat Q_0 radiated from the hot surface at temperature T to the thermo-element at temperature t . Like the Leslie cube, the thermo-element was also covered with a layer of lampblack.

The galvanometer deflection s_0 was observed after the thermo-element had been exposed for one minute to radiation from the cube. The amount of heat per scale division is then $e = Q_0/s_0$, and the amount corresponding to the deflection s is $Q = s.e.$

During calibration the thermo-element was in air at atmospheric pressure; when in use with the cathode rays, it was in a partial vacuum, and moreover the layer of lampblack had then been removed. This would tend to make the rise in temperature of the junctions in the latter case greater than in the former, assuming the quantity of heat supplied to be the same: for the loss of heat by radiation and conduction to the air must have been greater during calibration. The effect of this error would be to make the observed energy of the cathode rays too great; but the error is smaller, the more the loss of heat by conduction and radiation to the air can be neglected in comparison to the loss by conduction through the metal of the thermo-element itself.

In measuring the cathode-current, the contact ac (fig. 1) was closed, thereby connecting the thermo-element to earth through the galvanometer. For the thermal circuit, ab was closed, and at the same time either ce or cd , since it was necessary to connect some point to earth. But when for example cd was closed, it was found that a small fraction of the cathode-current also flowed to earth through the galvanometer, along the path $SbaGdc$. Closing ce caused approximately the same deflection in the opposite direction. Hence the mean of the two readings could be taken as the deflection for the thermal current alone.

In order to eliminate the effect of heat radiated from the cathode, the rays were deflected before each observation by

* F. Kurlbaum, Wied. Ann., lxxv, p. 759, 1898.

means of the coils MM (fig. 1); the small remaining galvanometer deflection was then subtracted from the total deflection with rays falling.

Observations with Thermo-element.

The value $i \cdot V$ (eq. 4) in gram-calories is cathode-current in amperes \times difference of potential in volts $\times 0.24$. Pressures are always given in mm. of mercury.

The first measurements yielded values of $i V/Q$ much greater than 1, as shown in the following table:

TABLE I.

Pressure.	$i \cdot 10^7$ Amp.	V	$Q \cdot 10^5$ gr.-cal.	iV/Q .
0.067	152	2280	90.3	9.18
0.046	138	2470	141	5.80
0.044	113	2490	123	5.50
0.042	99.4	2600	127	4.89

The junctions of the thermo-element were here, as at the calibration, covered with a layer of lampblack. It may, however, be assumed that the long-wave heat radiations used in calibrating penetrate quite deeply into the interior of the lampblack, while the cathode rays are absorbed close to its surface. If the heat conductivity of the lampblack is small compared with that of the metal of the thermo-element, it is not impossible that for the same amount of heat, the rise in temperature of the junctions under the influence of the cathode rays may be smaller than at the calibration. The lampblack was therefore removed, and in fact the value of $i V/Q$ was thereby diminished.*

The following measurements were next made after removal of the lampblack:

TABLE II.

Pressure.	$i \cdot 10^7$ Amp.	V	$Q \cdot 10^5$ gr.-cal.	iV/Q .
0.046	180	3100	412	3.25
0.050	175	3100	397	3.28

TABLE III.

Pressure.	$i \cdot 10^7$ Amp.	V	$Q \cdot 10^5$ gr.-cal.	iV/Q .
0.043	122	3150	331	2.80
0.029	116	3500	416	2.34
0.026	100	3620	416	2.09
0.025	65.4	3820	309	1.94
0.022	61.7	4070	344	1.75

* This diminution no doubt comes in part as described above from the change in constant of the thermo-element after removal of the lampblack, also perhaps partly from the altered reflection of the cathode rays.

It is to be noted that in Table III as also in Table I, the measured value of iV/Q decreases with decreasing pressure and cathode-current.

In order to investigate this relation more closely, a series of observations was made, varying pressure and cathode-current within as wide limits as possible. The cathode-current could be altered either by changing the number of accumulator cells, or by varying the cadmium-iodide resistance. Owing to the evolution of occluded gases from the cathode and its gradual discoloration, causing sudden irregularities in potential, it was impossible to get a perfectly consistent set of readings. Still in general it was evident that the ratio iV/Q diminished with decreasing i , at first rapidly, then more slowly. At high pressures ($> 0.05^{\text{mm}}$), the ratio seemed to depend also upon the pressure; but when the pressure had sunk below 0.03^{mm} , the drift was hardly perceptible.

These results led to the supposition that beside the cathode rays, a fraction of the discharge current also was conducted from the cathode to the thermo-element; for as J. J. Thomson* has shown, the portion of gas transversed by the rays becomes a relatively good conductor, except at very low pressures. This supposition was confirmed by inserting a battery between c (fig. 1) and the earth, thereby charging thermo-element and Faraday cylinder to a positive or negative potential v .† This cannot affect the true value of the ratio iV/Q as long as V denotes the difference of potential between cathode and thermo-element; but the supposed disturbing current must grow stronger or weaker according as v is made positive or negative.

The following table shows how the apparent value of iV/Q was correspondingly increased or diminished according to the sign of v .

Pressure.	v	V	$i \cdot 10^7$ Amp.	$Q \cdot 10^5$ g.-cal.	iV/Q .
0.050	+70	4050	3.54	34.1	1.01
0.050	0	3980	2.64	30.2	0.84
0.050	-70	3910	2.29	30.5	0.71

The potential of the cathode as measured on the electrometer is here $V-v$; as in all measurements, i denotes the observed sum of cathode- and disturbing currents.

The variations in the value of Q result from the fact that the path of the rays was visibly altered while the thermo-element was charged.

* J. J. Thomson, Phil. Mag., xliv, p. 293, 1897.

† The cylinder was here connected with the point c . Owing to incomplete insulation between thermo-element and cylinder it was not practicable to connect the latter directly to earth.

The cathode-current is here only about $1/30$ as intense as in Tables II and III, hence the value of iV/Q is also much smaller. The extremely small value of iV/Q when the charge is negative is explained as follows: owing to the conductivity of the path of the rays, a weak current finds its way from the battery used in charging the thermo-element, along this path to the anode. This current is opposed to the cathode-current, thereby reducing the apparent value of the ratio iV/Q . When the thermo-element is positively charged, the ratio is on the other hand greatly increased.

Measurements similar to the above, made at a pressure of 0.012^{mm} , showed that charging the thermo-element no longer had any influence upon the value of the ratio. These observations on the conductivity of the cathode rays agree well with those of Starke* and Thomson.†

The attempt was also made to intercept the disturbing current by means of a grating of several fine wires laid over the opening of the Faraday cylinder. This was, however, of no practical assistance, as even at high pressures the effect was very slight.

The results that have been given show that the observed value of iV/Q was in general too large. The disturbing current was at times greater than the cathode-current itself, but it was impossible to measure directly its absolute value. It therefore remained to establish the lower limit of the ratio. For this purpose a series of 46 observations was selected, having a range of pressure from 0.01 to 0.05^{mm} , of potential from 2400 to 5400 volts, and of cathode-current from $40 \cdot 10^{-7}$ to $0.23 \cdot 10^{-7}$ amp. When arranged in order of decreasing cathode-current, the values of iV/Q showed great irregularities, since the order was without regard to pressure and other conditions. Still a distinct drift in the ratio was evident, especially when the observations were divided into five groups and the average values taken:

Pressure.	$i \cdot 10^7$ Amp.	iV/Q
0.041	19.6	1.32
0.035	7.41	0.97
0.026	4.29	0.85
0.033	2.02	0.84
0.023	0.49	0.83

From this it is evident that when the cathode-current decreases to the fortieth part of its original value, the ratio iV/Q is changed by only 38 per cent. This at least makes it seem

* H. Starke, Wied. Ann., lxxvi, p. 52, 1899.

† J. J. Thomson, Phil. Mag., xlv, p. 293, 1897.

probable that eq. (4a) is approximately correct, although the limiting value of the ratio is considerably smaller than unity. Still owing to the sources of error in the use of the thermo-element, no great dependence can be placed upon absolute values. It was therefore determined to measure the energy Q in a surer manner, by Kurlbaum's bolometric method.*

Observations with Bolometer.

A surface-bolometer forms one arm of a Wheatstone bridge; suppose the bridge balanced, and a current of intensity I_1 flowing through the bolometer. Let the resistance of the bolometer under these conditions be w . If now an amount of heat Q per second be supplied by the cathode rays to the bolometer, the galvanometer in the bridge will suffer a certain deflection. Suppose the same deflection caused by increasing I_1 to I_2 , the cathode rays no longer acting. Then from Kurlbaum's formula,

$$Q = w \cdot \frac{I_2^2}{I_1^2} (I_2^2 - I_1^2) \text{ watts,}$$

where I and w are expressed in amperes and ohms respectively.

If the four bridge resistances are equal, then $I = \frac{1}{2}c$, where c is the current in the battery branch, and

$$Q = \frac{1}{4}w \cdot \frac{c_2^2}{c_1^2} (c_2^2 - c_1^2) \text{ watts.}$$

As it was not practicable to make both deflections exactly equal, the following method was adopted:

1. galv. defl. a_1 with current c_1 , rays falling.
2. " " " a_2 " " " c_2 , " deflected.

We thus get the following equation:

$$Q = \frac{1}{4}w \frac{a_1}{a_2} \frac{c_2}{c_1} (c_2^2 - c_1^2)$$

The thermo-element formerly used (S , fig. 1) was replaced by a bolometer, whose terminals were led out through the plate P . The bolometer was made after Lummer and Kurlbaum's process,† with a few modifications. In the first observations it consisted of a platinum foil about 0.001^{mm} thick, mounted with mica insulation upon a brass frame. The form is shown in B (fig. 2). Care was taken that the entire bolometric resistance should be reached by the rays; also that all cathode rays entering the Faraday cylinder that did not strike

* F. Kurlbaum, Wied. Ann., lxxv, p. 746, 1898.

† O. Lummer and F. Kurlbaum, Wied. Ann., xlv, p. 204, 1892.

the bolometer, should give up their charge to earth through the metallic frame. Although the density of the rays is not perfectly uniform throughout the bundle, still it can be assumed that the distribution of temperature in the thin bolometric foil under the influence of the rays and of the electric current is practically identical.* The silver was not dissolved away from the platinum-silver foil until the bolometer had been mounted.

2

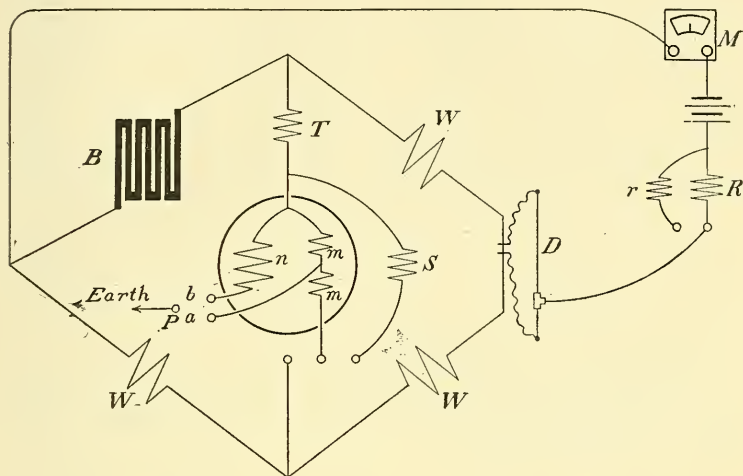


Fig. 2 shows the arrangement of the bridge. The current-intensity i in the battery branch was measured by means of the ampere-meter M . Shunting the resistance r parallel to R increased the current from i_1 to i_2 , and gave an accurate means for measuring the increase. A Thomson galvanometer of du Bois-Rubens type was employed. For the thermal measurements two 20-ohm coils mm were used; in order to get rid of the cathode-current, a point between the two coils was led to earth (contact a , fig. 2) and the resistance T so chosen that with battery disconnected the cathode-current caused no deflection of the galvanometer. Two 2000-ohm coils n served to measure the cathode-current, by closing the contact b , while a remained open.

S is a shunt for reducing the sensitiveness of the galvanometer when necessary. The bridge resistances were balanced by means of the sliding contacts at D . The manganin wire resistances W were as nearly equal to the bolometric resistance

* F. Kurlbaum, Wied. Ann., lxx, p. 755, 1898.

as possible, but small discrepancies were always taken into account. It was necessary to use great care in insulating the entire circuit from earth.

The bolometer was found to be much more sensitive than the thermo-element, although its effective surface was smaller. The source of error from gas-evolution was also less serious, owing to the rapidity with which observations could be made.

The process of etching and washing such small bolometers is somewhat complicated when a metallic frame is used, as the latter has to be protected by a coating of wax, which is afterwards dissolved away. Thus it is difficult to avoid straining the delicate platinum strips. This may become a very serious source of error, as it is impossible to know the extent of the injury. Moreover, a brownish deposit was nearly always formed upon the strips, probably from some action of the acid used to dissolve the silver, upon impurities in the wax. This may in some way have caused a further source of error.

Four different bolometers with brass frames were used. The results are briefly given below, though in at least two cases great irregularities occur. It was not until a frame of slate had been made that reliable observations were obtained (see below).

Bolometer I. Eleven strips, $11.5 \times 0.66^{\text{mm}}$. Resistance w at room temperature = 26.21 ohms.

The observations were all made at high pressures; the ratio iV/Q is therefore chiefly dependent upon the pressure and not upon the cathode-current.

Pressure.	$i \cdot 10^7$ Amp.	V	$Q \cdot 10^5$ Watt.	iV/Q .
0.116	18.0	2590	213	2.21
0.113	12.3	2610	140	2.28
0.064	22.3	4090	541	1.69
0.055	13.7	4210	376	1.53

Bolometer II. Nine strips, $11.5 \times 1.0^{\text{mm}}$. $w = 15.74$ ohms.

An influence machine was used instead of the battery as source of discharge. As the range of values of i is not great, only the average values are here given: $i = 1.40 \cdot 10^{-7}$ amp.; pres. = 0.020^{mm} ; $iV/Q = 0.81$.

Bolometer III. Seven strips, $11.5 \times 1.3^{\text{mm}}$. $w = 8.75$ ohms.

A battery of 30 cells was so arranged that it could be connected at will between P (fig. 2) and the earth. This raised the bolometer and of course the entire bridge to the potential of the battery. The Faraday cylinder was kept in contact with P . Readings were taken with the bolometer alternately at 0 and +60 volts.

From the table it is evident that as in the case of the thermo-element, charging the bolometer to a certain potential increases the ratio iV/Q to a greater extent, the stronger the cathode-current, that is, the greater the conductivity of the gas.

Pressure.	v	$i \cdot 10^7$	V	$Q \cdot 10^5$	iV/Q .
0.033	0	5.30	3130	106	1.57
	+60	7.20	3090	119	1.91
0.028	0	2.10	4320	80.8	1.12
	+60	3.23	4200	113	1.20
0.029	0	1.96	5150	75.5	1.33
	+60	4.91	4960	154	1.58
0.030	0	1.90	4230	67.5	1.19
	+60	2.80	4130	88.5	1.31
0.030	0	0.54	3910	20.8	1.02
	+60	0.67	3880	25.3	1.02
0.031	0	0.28	4220	11.9	0.98
	+60	0.56	3900	21.9	1.01

Bolometer IV. Seven strips, $11.5 \times 1.3^{\text{mm}}$. $w = 9.20$ ohms. Values of the ratio were irregular, but the averages from three groups of observations are here given :

Pressure.	$i \cdot 10^7$ Amp.	iV/Q .
0.070	9.43	0.81
0.040	11.48	0.76
0.023	0.18	0.72

As with Bol. I, the ratio at high pressures seems dependent upon the pressure alone. In the third group the cathode-current was diminished by increasing the cadmium iodide resistance.

These results are of only relative importance except as they indicate that the absolute value of iV/Q is somewhat less than 1.

A fifth bolometer was therefore made and mounted after the Kurlbaum process upon a slate frame. After the silver had been dissolved from the foil, the strips looked perfectly clean and seemed free from strains. In order to keep the cathode rays from the slate, a small brass diaphragm was placed close over the bolometer, with a square opening of the same size as the bolometer surface.

Bolometer V. Seven strips, $7 \times 0.73^{\text{mm}}$. $w = 8.72$ ohms. The following observations were taken :

Pressure.	$i \cdot 10^7$	V	$Q \cdot 10^5$	iV/Q .
0.038	6.10	2270	76.9	1.80
0.061	5.00	2690	82.3	1.58
0.035	4.80	3480	149	1.12
0.024	2.26	3770	85.5	1.00
0.024	2.09	3560	81.2	0.92
0.030	1.69	4460	92.6	0.81
0.022	1.36	3730	58.0	0.87
0.032	1.09	3740	50.4	0.81
0.032	1.06	3750	47.9	0.83
0.028	0.50	3900	23.6	0.83
0.026	0.255	4740	14.5	0.83
0.023	0.040	4800	2.29	0.84

This bolometer was apparently free from the faults of its predecessors: therefore only the usual errors of observation need be considered, especially that arising from the evolution of gas. This error can hardly have amounted to more than 3 per cent with weak cathode-current; hence it follows from the last table that the true value of the ratio iV/Q lies between 0.80 and 0.86.

Since the bolometer foil was only 0.001^{mm} thick, the question arises whether an appreciable part of the rays can have penetrated through to the other side. In this case it might easily happen that the particles on passing through would lose a part of their energy while retaining their charge. McClelland* and Wien† have shown that cathode rays after traversing an aluminium window carry a negative charge; but Lenard did not succeed by means of a thermo-element in detecting any heating effect.‡ Still if we assume according to Lenard that for solids the quotient of absorptive-power divided by density = 3200,§ we find from his formula that not more than 1 per cent of the rays can have passed through the bolometric foil used. This can have had practically no effect upon the ratio under discussion.

Conclusion.

It only remains to compare the limiting value of the ratio as experimentally determined, with the equation (4a)

$$\frac{iV}{Q} = \frac{i-r}{1-rr'},$$

where r denotes the reflected part of the rays, while upon reflection the rays lose their energy in the ratio $r':1$.

Starke|| has found that platinum reflects about 40 per cent of the rays. Assuming, therefore, that for the bolometer, $r = 0.4$, we find that when $iV/Q = 0.83$, $r' : 0.7$.

This result is in contradiction with the observations of Merritt already referred to. The latter are unfortunately not above criticism, as the author himself admits. Merritt compared the deflectibility of direct and reflected rays by observing the phosphorescence on the walls of the tube. It is hardly to be supposed that the spot from the reflected ray can have been even approximately as distinct as that from the

* J. A. McClelland, Proc. Roy. Soc., lxi, p. 227, 1897.

† W. Wien, Verh. Phys. Ges. Berlin, xvi, p. 165, 1897.

‡ Ph. Lenard, Wied. Ann., li, p. 239, 1894. This experiment was, however, performed in air at ordinary pressure, which rapidly absorbs the energy of the cathode rays.

§ Ph. Lenard, Wied. Ann., lvi, p. 274, 1895.

|| H. Starke, Wied., lvi, p. 58, 1899.

direct; moreover, since an induction coil was used as source of discharge, one would expect the formation of a "magnetic spectrum,"* which would greatly complicate the appearance of both spots. Thus it is easy to see that a systematic error may have crept into the results.

Whether such an error can have been great enough to give the value $r' = 1$ instead of $r' = 0.7$ is, of course, not certain; but it may be added that the energy of the cathode rays is proportional to the square of their deflectibility, hence an error of n per cent in the measurement of the deflection would correspond to an error of $2n$ per cent in the energy. In any case it is desirable to determine both the deflectibility and the energy of reflected cathode rays by as accurate a method as possible.

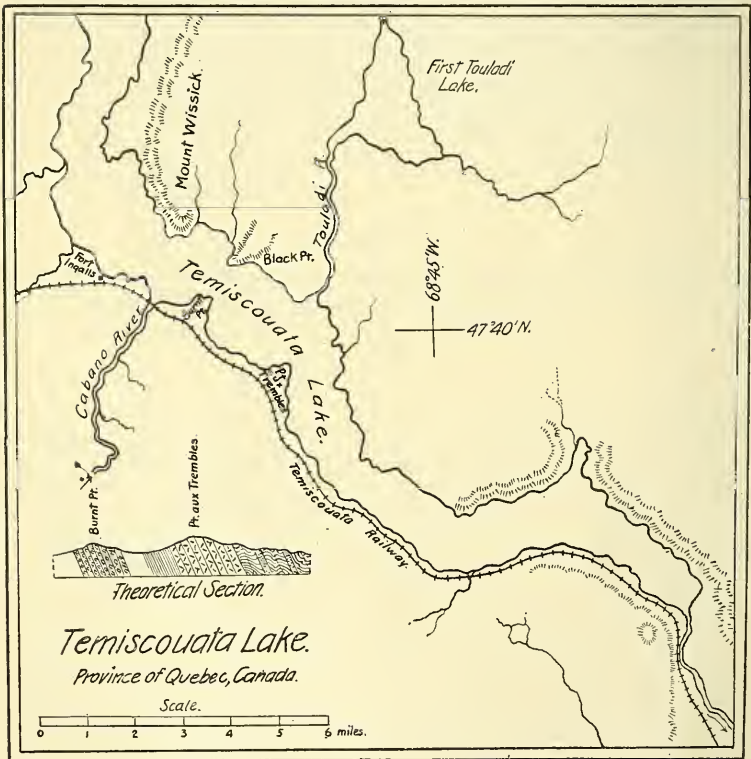
My thanks are due to Prof. Warburg for his kind advice and assistance during this investigation.

* K. Birkeland, *Comptes Rendus*, cxxiii, p. 492, 1896.

Providence, R. I., May 5, 1900.

ART. II. — *Volcanic Rocks from Temiscouata Lake, Quebec*; by HERBERT E. GREGORY.

TEMISCOUATA LAKE was explored by the early geologists of Canada, and Logan reported* the presence at Pointe aux Trembles of 'tough, green sandstone, with pebbles of metamorphic rock.' A survey of the region was made by Bailey and McInnes in 1886-87, and the presence of volcanic rocks noted. Their report in reference to these rocks reads as follows: "It is important to notice in connection with the Pointe aux Trembles sandstones the evidence which they appear to afford of contemporaneous volcanic activity. This is, perhaps,



partly indicated by the color of the rock, which varies from green to red and purple, but is more clearly seen in the abundance of epidote with which the rock is charged and in places

* Geology of Canada, 1863, p. 423.

its somewhat amygdaloidal aspect.”* During the summer of 1897, the region was studied by Prof. H. S. Williams, and the specimens then collected were turned over to the writer for examination. The outcrops visited extend along the lake for a distance of 20 miles, as shown on the accompanying map.

Mount Wissick.—The rocks examined from Mt. Wissick and the east shore of the lake are fossiliferous limestones, sandstones and shales, without admixture of volcanic materials. The shales and sandstones do not differ from similar rocks found elsewhere. The arenaceous limestone in places is separated into hexagonal prisms, probably from shrinkage, and appears as if composed of vertical columns with quite regular outline. The thin layers of limestone in the slates at the north of Mt. Wissick show under the microscope well-defined oolitic structure in which the little spheres are broken by numerous minute faults.

West Shore.—The section on the West Shore extends from Burnt Point, below Fort Ingalls, to the outlet of the lake.

At Burnt Point the rocks are coarse conglomerates made up of materials not much water worn, and show evidence of rapid deposition. Slates and limestones form the most abundant pebbles, and occasionally attain a diameter of 1–2 feet. The total thickness of the conglomerate at this point is about 1000', but such great thickness of the beds is reported to be quite local. Succeeding the Burnt Point conglomerate, to the south, occur thin-bedded shales and sandstones, which contain fossils of Niagara age.†

At Point aux Trembles, the rocks, both along the railroad and on the lake shore, appear at first sight to be greenish sandstones and coarse brown conglomerates. They are interstratified with the other beds of the region, and have practically the same dip and strike. A closer examination, however, shows them to be volcanic. The finer, more sandy beds, contain quantities of volcanic ash, and the coarser ones are conglomerates of typical andesitic fragments, with scarcely any foreign material. Parts of the rock contain very prominent rounded fragments of amygdaloidal andesite. As with the volcanics of northern Maine,‡ so here the gradation from the sandstones of the region to pure volcanic material can be traced, and like the Maine breccias and tuffs, these rocks show more or less rounding of their pebbles and sorting by water. It is believed that the volcanic vents were near some body of water, and that the material fell in or near the water and received a limited amount of wearing before final consolidation.

* Geol. Survey of Canada—Annual Report, 1887, p. 33M.

† Geol. Survey Canada, Annual Report, 1887, p. 33M.

‡ U. S. G. S. Bulletin 165.

From Point aux Trembles to the foot of the lake—a distance of about 10 miles—the rocks which outcrop are slates and impure sandstones with rare tufaceous material. In places the beds are much folded and traversed by faults of small displacement.

Petrography.

The ordinary stratified rocks of the region present no peculiarities which require detailed description. The volcanic series consists of fine tuff and coarse amygdaloidal conglomerate or breccia.

Fine Tuff.—The fine tuff appears in the hand specimen as a dense bluish-grey sandstone in beds a few feet in thickness. Under the microscope its true character is revealed, and it is found to consist of tiny fragments of andesite, broken crystals of plagioclase and olivine, and areas of devitrified glass with rare quartz grains and biotite shreds. The andesite has a hyalopilitic groundmass, which varies in amount of glass, so that some fragments appear to be practically all glass with only a few scattered laths of plagioclase. The larger crystals of andesine have albite twinning and are largely kaolinized. The olivines are represented in the slide by areas of serpentine with characteristic outlines and cleavages shown by dotted ore grains. Iron ore is sprinkled quite generally through the section.

Volcanic Conglomerate.—The volcanic conglomerate consists of subangular pebbles of andesite and amygdaloidal andesitic bombs, embedded in a finer matrix of red and green grains of the same material. The pebbles range in size from $\frac{1}{2}$ inch to 6 inches in diameter, and show in the hand specimen phenocrysts of plagioclase, and occasionally epidote and augite. The embedded bombs form a conspicuous feature of the rock. They are quite spherical and are distinctly different in appearance from the matrix. Amygdules, about the size of bird shot, make up fully one-half of the bomb and are filled with green balls of chlorite—rarely with calcite. Some of the larger amygdules are seen to be occupied by chalcedony inside the chlorite lining.

Microscopic Examination.—The microscope shows the general matrix of this coarse type to consist of crystals of feldspar, pyroxene and olivine, and fragments of andesite, devitrified glass, and jasper. Iron dust, a small amount of calcite, and a few shreds of biotite, are also present. The feldspar crystals consist of a few orthoclases, and many rather small andesines, with albite and carlsbad twins. The pyroxene crystals are represented only by their outline, in a base peppered with iron dust. The olivines are fragmentary crystals, and largely replaced by serpentine and iron. The bulk of the sec-

tions examined is formed of fragments of pyroxene-andesite of a type sparingly porphyritic with small stout phenocrysts set in a groundmass, having in some places a glassy base, in others being made up almost entirely of ragged feldspar microlites, arranged with flow structure. The fragments of devitrified glass present show occasional cusps, which represent expanded steam cavities. No close line, however, can be drawn between the fragments of glass and the fragments of andesite containing a few feldspar microlites in a glassy base. The dark-red iron dust is not generally distributed through the rock, but is present along the cracks of the olivine crystals, and it also gives color to the glassy fragments.

The Amygdaloidal Bombs.—The section cut from one of the amygdaloidal bombs reveals the composition and structure of an extrusive basic andesite. The minerals present as phenocrysts are plagioclase, pyroxene, olivine and iron ore, with infiltrated quartz and calcite in the amygdules. There is also a considerable development of minerals of the chlorite group, perhaps largely delessite. The plagioclases are short, rather square crystals with ragged ends and strongly-marked albite and pericline twinning. Many of them are bent and curved and irregularly broken. In composition they are between andesine and labradorite. Augite occurs as stout crystals always partly decayed and replaced by chlorite or epidote, or rarely by calcite. Olivine crystals are about as abundant as augite and show their characteristic parting along cracks now filled with iron. Serpentine has generally replaced the olivine. Iron ore occurs in a few ragged grains. The groundmass is of feldspar microlites usually untwinned and with little evidence of flowage. The structure was hyalopilitic, but the glass filling the interstices is devitrified and stained brown with iron. The steam cavities now filled and converted into amygdules, though very abundant, have exerted little influence on the arrangement of the feldspars in the groundmass—a fact which suggests that these cavities formed in part before any crystallization of feldspars had occurred.

All the amygdaloidal cavities are filled with secondary products which have a uniform arrangement. They are lined with a coating of a malachite-green mineral of the chlorite group, arranged in radial forms. Most of the smaller cavities are completely filled with this substance and appear on the rock surface as embedded green pellets. The larger cavities have the chlorite lining fractured, and fragments of it have floated in toward the center, leaving means of access for later infiltrations. Similar phenomena have been observed in certain

volcanic rocks in Maine.* The inside filling of the amygdules is silica or rarely calcite. The material next to the chlorite lining is chalcedony, with well-developed mammillary structure. The center of the cavity is a mass of quartz grains separated from the chalcedony by a definite boundary. It thus appears that the amygdules were filled by three or four successive periods of infiltration.

Summary.—While the mere description of these rocks and the determination of their volcanic origin is, in itself, chiefly of local interest, these facts have, however, important bearings on the general geological history of the Atlantic coastal region. That they are interbedded with Niagara sediments helps to determine the time when wide-spread volcanic activity gave rise to the numerous small areas of tuffs and lavas in the Maine-Quebec region, while this fact throws additional light on the physical conditions obtaining in that period. And finally, the recognition of so much distinctly contemporaneous volcanic material in well-bedded sediments affords another plea for the study of the sedimentaries by microscopic analysis.

* U. S. G. S. Bulletin 165, Plate XIV.

Yale University, New Haven, Conn., March, 1900.

ART. III.—*On the Interpretation of Mineral Analyses: a Criticism of recent Articles on the Constitution of Tourmaline*; by S. L. PENFIELD.

ABOUT eighteen months have elapsed since H. W. Foote and the author published in this Journal* a joint article on the chemical composition of tourmaline. Since that time two articles have appeared, presenting views differing from one another and from those of Foote and the author; one on the Constitution of Tourmaline, by Prof. F. W. Clarke† of Washington, the other *Über das Mischungsgesetz der Turmaline*, by Prof. G. Tschermak‡ of Vienna.

In order to discuss the views presented in these articles it is desirable to carefully consider some facts concerning chemical analyses. In the first place, a perfect chemical analysis cannot be made. There are, to be sure, a very few analytical processes for the determination of single constituents, which, when carefully executed, can be relied upon to give results varying less than 0.01 per cent from the theory; but when it comes to mineral analysis, necessitating the splitting up of a complex body and the determination of a number of constituents, such accuracy cannot be attained. In making a mineral analysis, one seldom feels confident that all determinations are correct, even within 0.25 per cent of the truth, although if duplicate analyses are made, it is expected that, for the majority of the constituents at least, the two determinations will agree within 0.10 or 0.20 per cent of one another. At times, of course, depending upon the difficulty of the analysis or the scarcity of available material, variations of 0.50 per cent, or even more, in duplicate determinations are not to be wondered at.

Secondly, analytical work may be of a high order, the results very accurate, and yet an analysis may not serve for the determination of a chemical formula because made on material more or less impure. The chemists of to-day have a decided advantage over those of a former generation, for the microscope enables them to study their material, select that which is best, and, if impurities cannot be avoided, to take their effect into consideration in discussing the analytical results. Then again the heavy solutions are invaluable for separating out material for analysis, and, what is considered of very great importance, for furnishing a guarantee of the purity of any given material; for if it can be stated that all of the mineral floats on a solution of a certain specific gravity and sinks when

* Vol. vii. p. 97, 1899.

† This Journal, IV, viii, p. 111, 1899.

‡ *Mineralog. und Petrogr. Mitth.*, xix, p. 155, 1899.

the specific gravity is lowered to a trifling extent, it gives one not only great confidence in the purity of the material, but, also, it enables the investigator to present data which others can make use of in judging the character of the work.

It has been the author's privilege during the past twenty-five years to make many analyses of minerals, and to superintend the making of many more in the Sheffield Mineralogical Laboratory; also to discuss the analytical results and derive therefrom the chemical formulas of minerals, and this occasion will be taken to call attention to certain features which are regarded as most important in mineralogical investigations. In the first place, the utmost pains should be taken to secure pure material, and, if the results are to be published, the character of the material should be described so that its degree of purity can be judged by others. Secondly, if an analysis presents any especially difficult features, the method of analysis should be carefully described, and it is in almost all cases well to give at least some brief outline of the analytical methods employed. Then, too, when material is abundant, it is advisable to make analyses in duplicate, and to give all of the determinations, together with the averages. Thus the investigator has from beginning to end the satisfaction of a control over all determinations, and, if agreements are close, others can form some estimate concerning the care with which the work was executed. There are those who apparently entertain the belief that closely agreeing duplicate determinations indicate great accuracy in analytical work, but that is not necessarily the case, for in some analytical methods there is a tendency for results to come too high, in others too low, and thus duplicate determinations, made under like conditions, either with faulty methods, or with good methods improperly executed, may be uniformly high or uniformly low, agreeing with one another, and yet varying considerably from the truth. Still two closely agreeing determinations carry with them a certain weight which cannot be ignored. Thirdly, with each analysis, the quotients obtained by dividing the several constituents by their molecular or atomic weights, as the case demands, should be given, and from the quotients thus obtained the ratio between the several constituents should be determined. The ratio ought not to be given simply rounded out to the nearest whole numbers, but, taking the quotient of the most characteristic or best determined constituent as unity, the ratio should be given to the second place of decimals. *It is safe to assume that the close approximation of a ratio to whole numbers constitutes the strongest argument that can be advanced in support of the excellence of an analysis and the correctness of the derived formula.* It will seldom happen that a ratio approximates to whole numbers merely as a matter of accident. Provided the

compound is a simple one, instead of giving the ratio, an excellent method is to give the calculated composition, which can then be compared directly with the results of the analysis. Lastly, for determining a formula one or two good analyses are of more value than many indifferent ones, hence it will often prove best to make new analyses on material of unquestioned purity. This may be done not wholly with the idea that the new analyses are better than those made by other investigators, but, knowing all about the quality of the material and the working of the analyses, it will be possible to exercise better judgment in summing up the results of the investigation, and to present with greater force the arguments needed in support of the proposed formula.

Turning now to the consideration of tourmaline, two new analyses were made by Foote and the author, upon material of ideal purity and with the use of most carefully studied methods. The results need not be repeated here, but it will be stated that, with the exception of a single water determination, all constituents were determined in duplicate; that in twenty out of a total of twenty-three instances, the discrepancy between duplicate determinations did not exceed 0.10 per cent; and that the maximum variation in the remaining three instances was 0.18 per cent. The single water determination which was not duplicated was controlled by a closely agreeing estimation of loss on ignition. In working out the ratios from these analyses, the method was adopted of calculating for the metals their equivalent of hydrogen, including fluorine with hydrogen, since tourmaline contains hydroxyl with which fluorine is isomorphous. Thus the ratio was found between SiO_2 , B_2O_3 , and Total Hydrogen, from which the empirical formula of the tourmaline acid was derived. For the sake of the present discussion the ratios will be repeated in two forms: with one-fourth of the SiO_2 as unity and also with one-twentieth of the Total Hydrogen as unity. This latter method has been here adopted, because a few relations can be brought out better in the discussion by so doing. The ratios of the two analyses are as follows:

	SiO_2	:	B_2O_3	:	Total H.		SiO_2	:	B_2O_3	:	Total H.
De Kalb	4.00	:	1.01	:	19.90		4.02	:	1.01	:	20.00
Haddam Neck.	4.00	:	1.02	:	19.98		4.00	:	1.02	:	20.00

These ratios approximate *very closely* to the whole numbers 4:1:20; such close approximations, in fact, are seldom obtained, and cannot in these two instances be regarded merely as matters of accident; they are the reward, rather, of careful analytical work on material of unquestionable purity. As soon as the ratios were worked out, it was seen at once that at least one important key to the solution of the tourmaline problem had

at last been found: *the empirical formula of the tourmaline acid must be $H_{20}B_2Si_4O_{21}$.*

And now, for the sake of the discussion, some space will be devoted to the ratios derived from the analyses of Riggs, and Jannasch and Kalb. These have already been given by Foote and the author* with $\frac{1}{4}$ SiO_2 as unity, and are now repeated, together with the ratios derived by taking $\frac{1}{27}$ Total Hydrogen as unity. They have moreover been arranged in series, commencing with the closest approximation to 4:20 between SiO_2 and Total Hydrogen, and proceeding to the maximum deviation from this ratio.

TOURMALINE RATIOS DERIVED FROM THE ANALYSES OF RIGGS.

No.	No. Dana.	$SiO_2 : B_2O_3 : Total\ H.$	$SiO_2 : B_2O_3 : Total\ H.$
1.	43.	4.00 : 0.94 : 20.03	3.99 : 0.94 : 20.00
2.	45.	4.00 : 0.95 : 20.03	3.99 : 0.95 : 20.00
3.	48.	4.00 : 1.01 : 20.06	3.99 : 1.00 : 20.00
4.	47.	4.00 : 0.98 : 20.08	3.98 : 0.97 : 20.00
5.	52.	4.00 : 0.94 : 20.11	3.98 : 0.93 : 20.00
6.	49.	4.00 : 1.01 : 20.12	3.97 : 1.00 : 20.00
7.	36.	4.00 : 0.90 : 20.2	3.96 : 0.89 : 20.00
8.	44.	4.00 : 0.88 : 20.2	3.96 : 0.87 : 20.00
9.	46.	4.00 : 0.96 : 20.2	3.96 : 0.95 : 20.00
10.	42.	4.00 : 0.97 : 19.8	4.04 : 0.98 : 20.00
11.	54.	4.00 : 0.98 : 19.8	4.04 : 0.99 : 20.00
12.	39.	4.00 : 0.94 : 19.7	4.06 : 0.95 : 20.00
13.	41.	4.00 : 0.92 : 19.7	4.06 : 0.93 : 20.00
14.	51.	4.00 : 0.91 : 19.6	4.08 : 0.93 : 20.00
15.	37.	4.00 : 0.93 : 20.5	3.90 : 0.91 : 20.00
16.	38.	4.00 : 0.92 : 19.5	4.10 : 0.93 : 20.00
17.	55.	4.00 : 1.01 : 20.6	3.88 : 0.98 : 20.00
18.	40.	4.00 : 0.96 : 19.3	4.14 : 1.00 : 20.00
19.	50.	4.00 : 0.98 : 19.2	4.16 : 1.02 : 20.00
20.	53.	4.00 : 0.97 : 18.9	4.23 : 1.00 : 20.00
Average		4.00 : 0.95 : 19.88	4.02 : 0.96 : 20.00

TOURMALINE RATIOS DERIVED FROM THE ANALYSES OF JANNASCH AND KALB.

No.	No. Dana.	$SiO_2 : B_2O_3 : Total\ H.$	$SiO_2 : B_2O_3 : Total\ H.$
1.	62.	4.00 : 0.80 : 20.00	4.00 : 0.80 : 20.00
2.	64.	4.00 : 0.84 : 20.01	4.00 : 0.84 : 20.00
3.	61.	4.00 : 0.95 : 20.2	3.96 : 0.94 : 20.00
4.	57.	4.00 : 0.99 : 19.8	4.04 : 1.00 : 20.00
5.	56.	4.00 : 0.96 : 19.7	4.06 : 0.97 : 20.00
6.	63.	4.00 : 0.98 : 19.7	4.06 : 0.99 : 20.00
7.	58.	4.00 : 0.95 : 20.4	3.92 : 0.93 : 20.00
8.	60.	4.00 : 0.88 : 20.4	3.92 : 0.86 : 20.00
9.	59.	4.00 : 0.92 : 18.8	4.25 : 0.98 : 20.00
Average		4.00 : 0.92 : 19.9	4.02 : 0.93 : 20.00

* Loc. cit., pp. 114 and 115.

Before entering upon the discussion of these ratios, let it be understood that the analysis of tourmaline is one of the difficult problems of analytical chemistry, and although Riggs made duplicate and often triplicate determinations of B_2O_3 and H_2O in almost all cases, and duplicated somewhat more than half of his determinations of SiO_2 and F, only single determinations of other constituents are recorded in his paper, while Jannasch and Kalb record only single determinations. Also it is to be borne in mind that although both Riggs and Jannasch and Kalb undoubtedly used carefully selected tourmaline fragments for analysis,* still there is nothing to indicate that slight amounts of foreign materials might not have been present in some of the specimens analyzed. Keeping these facts then well in mind, let us examine the ratios as presented in the foregoing tables. It is granted that the ratios are not exactly 4:1:20, and to get *exact ratios* from mineral analyses is not to be expected, but the close approximation to 4:1:20 in the case of the two analyses by Foote and the author, of sixteen out of the twenty analyses by Riggs, and of eight out of the nine analyses by Jannasch and Kalb, *constitutes an overwhelming amount of evidence in support of the empirical formula of the tourmaline acid, $H_{20}B_2Si_4O_{21}$.* It is safe to state that there does not exist a series of thirty silicate analyses of any one mineral yielding ratios which approximate so closely to whole numbers as the tourmaline analyses referred to above. That some analyses fail to yield a ratio as close to rational numbers as desired, reflects discredit neither upon the analyst nor upon the character of his work, for the material for analysis might not in all cases have been pure. Take, for example, No. 17 of the series of Riggs, brown tourmaline from Hamburg, N. J., occurring in calcite. The ratio of SiO_2 : Total Hydrogen=4:20·6. Evidently the bases are too high, and this particular analysis is peculiar in that it shows 5·09 per cent CaO, while the next highest percentage of CaO recorded is 3·70. The material might well have contained some calcite, either as small included nodules, or as an infiltration along cracks, and if the amount of calcite be assumed as 1·78 per cent, equivalent to 1 per cent CaO, the analysis would add up to 100·82, which is not too high for such a complicated substance, and the ratio of SiO_2 : Total Hydrogen would become 4·00:20·3 or 3·94:20·00. To assume that the Hamburg material probably contained some calcite seems far more reasonable than to speculate, as Prof. Clark does, upon some complex formula especially adapted to suit this particular analysis. Again, Nos. 18, 19, and 20 of Riggs, and 9 of Jannasch and Kalb indicate either

* Compare foot-note, this Journal, IV, vii, p. 115, 1899.

that the amount of base is low, SiO_2 being assumed as practically correct, or, what is far more likely, that the amount of SiO_2 is too high, as seen best when one-twentieth of the Total Hydrogen is taken as unity. Does the high silica ratio indicate that for these special cases a new type of tourmaline formula is needed, or is it not simpler to assume that the material from which these analyses were made might possibly have contained a little quartz or other silicate as impurity? It would take not over two per cent of quartz as an impurity to bring about the extreme amount of variation from the ratio 4:20 recorded in the foregoing tables.

Both Clarke and Tscherniak seem to place implicit confidence in *all* of the recent tourmaline analyses: they seem to regard them as perfect, and the material analyzed as necessarily pure; consequently they try to devise formulas or expressions (how complicated they are will be shown) to suit all of the analyses. Clarke states that a formula in order to be satisfactory "must adequately express the composition of the compound in question, covering all of its variations." It is evident, however, that a formula should not cover variations due to impurities in the material analyzed, nor possible inaccuracies in analytical work. Both Clarke and Tschermak have their well-known theories concerning the composition of mica, and, seeing in the lithia-, iron-, and magnesia-varieties of tourmaline certain analogies, respectively, to muscovite, biotite and phlogopite, they both endeavor to force the tourmaline formulas to conform to their ideas concerning the constitution of mica. True, as shown by the analyses of Riggs,* there occur at Auburn, Rumford and Hebron, Maine, pseudomorphs of muscovite after tourmaline; but the fact need not necessarily be taken to indicate that tourmaline is closely related to mica, nor that, by what may be designated as a sort of *molecular cleavage*, tourmaline is transformed into mica. The author is familiar with these Maine localities, and he does not believe that such alterations are common, nor does he believe that it is common to find similar alterations at other localities where tourmaline is found. Muscovite is evidently a very stable molecular compound; it occurs more or less pure as pseudomorphs after many minerals, and not, as it would seem, because each and every one of these minerals contains as a nucleus the muscovite molecule, but because muscovite has such a tendency to form under a variety of conditions that it develops, provided materials suitable for its formation are at hand. Thus it is supposed that an ancient mud flat becomes converted to a shale and eventually to a mica schist, as the result of indurating and metamorphic processes,

* This Journal, III, xxxv. p. 41, 1888.

and yet mud and shale are not supposed to contain the chemical nucleus of muscovite; but, rather, they contain constituents suitable for the formation of muscovite.

It is quite common at the Maine localities to find cavities in quartz and other minerals once occupied by tourmaline crystals, and often there still remain in these cavities remnants of the fresh, unaltered tourmaline. The tourmaline fragments appear as if etched, and much material has evidently been dissolved and carried away. Hence replacement pseudomorphs of muscovite after tourmaline might well result from the subsequent filling of such cavities by muscovite. Replacement pseudomorphs are well illustrated by specimens from the Maine localities in the Brush Collection. The specimens referred to exhibit cavities containing remnants of fresh unaltered tourmaline, and as secondary minerals we now find cookeite and quartz deposited on the walls of these cavities, with absolutely nothing to indicate that the material for the formation of the cookeite and quartz was derived from tourmaline, for it could equally well have been derived from lepidolite and other minerals; in fact the tourmaline originally in a cavity could scarcely have furnished sufficient lithia for the formation of the cookeite, provided that not a trace of the lithia was carried away. The fact, therefore, that pseudomorphs of muscovite after tourmaline have been described, cannot be taken and accepted as proof that tourmaline is closely related to the micas, and that the formula of tourmaline must show close analogies to the formulas of the minerals of the mica group.

Clarke's formula for the tourmaline acid is $H_{29}B_3Si_6O_{31}$, which requires the ratio of $SiO_2 : B_2O_3 : \text{Total Hydrogen}$ to be 6:1.5:29, or 4:1:19.33. It is true that three of the analyses of Riggs yield this ratio (Nos. 18 and 19, page 22, and also No. 20, if it is assumed, as is by no means proved,* that the titanium exists as Ti_2O_3). Assuming Ti_2O_3 as existing in No. 20, the ratio becomes 4.00:0.97:19.2.

These three analyses then, out of a total of twenty by Riggs, alone support Clarke's formula. Two of the three varieties (No. 19, black tourmaline from Pierrepont, N. Y., and No. 20, brown tourmaline from Gouverneur, N. Y.) have been analyzed by Rammelsberg, a fact ignored by Clarke. It may be claimed by some that Rammelsberg's analyses are not to be taken into consideration because they are old. It should be said, however, that Rammelsberg undoubtedly determined the bases and loss on ignition in his analyses with considerable accuracy, and his results need not be wholly ignored. Assuming that SiO_2 and B_2O_3 are present in tourmaline in the ratio 4:1, it becomes

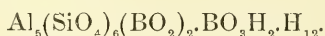
* See suggestions by Foote and the author, p. 117 of their article, loc. cit.

possible to determine both SiO_2 and B_2O_3 by difference, and this treatment of Rammelsberg's analyses was fully discussed by Foote and the author. In the present instance it is only necessary to repeat here the ratios between SiO_2 and Total Hydrogen derived from the two analyses, which are to be compared with Numbers 19 and 20 (page 22) of Riggs.

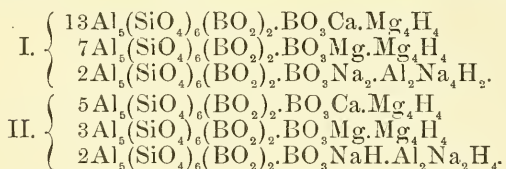
	SiO_2 : Total H.	SiO_2 : Total H.
Black tourmaline, Pierrepont, analyzed in 1889,	4.00 : 20.2	3.96 : 20.00
Brown tourmaline, Gouverneur, analyzed in 1850,	4.00 : 20.7	3.86 : 20.00

Thus Rammelsberg's results for these two varieties conform to the $\text{H}_{20}\text{B}_2\text{Si}_4\text{O}_{21}$ formula, rather than to the more complicated one proposed by Clarke. The analysis of the black tourmaline from Pierrepont is one of the last, if not the very last, analysis made by Rammelsberg, and he claimed for it a high degree of accuracy.* The ratio derived from it is certainly very close to 4:20.

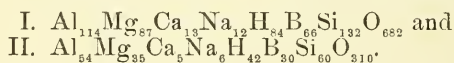
As an aluminium boro-silicic acid, Clarke writes his formula in linear form as follows :



The composition of the black tourmaline from Pierrepont (I) and the brown tourmaline from Gouverneur (II), as derived from Riggs' analyses, are then expressed by him as follows :



The above molecular mixtures when united become :



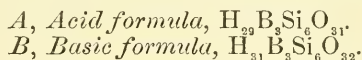
Thus, taking multiples of a molecule containing a large number of atoms, and admitting of many variations in the replacement of the hydrogen atoms, it becomes a matter not of chemical, but rather of arithmetical skill to so choose the numbers that the calculated percentage values shall agree with the results of the analyses: and all this labor when there is good evidence for believing that owing either to slight defects in the analyses or impurities in the materials the two analyses in question fail by a little only to give ratios like those of the remaining analyses.

* Abhand. K. Akad., Berlin, 1890.

Another difficulty also, which Clarke evidently recognizes, is that there is no appropriate place in his molecule for the attachment of fluorine, and hence he suggests that the fluorine may sometimes replace the group BO_2 , an equivalent, as stated by him, which is strongly indicated in the cappelinite group of minerals. Now from a chemical standpoint the group cited containing cappelinite, melanocerite, caryocerite and tritomite, seems poorly adapted for illustrating an important chemical principle, since the composition of all of the minerals of the group is very complicated, and only in the case of one mineral, tritomite, have direct determinations of B_2O_3 been made. Clarke's formula, it will be observed, is applicable directly to the three analyses of Riggs' (Nos. 18, 19 and 20, p. 22) which are low in bases, or high in silica: it will therefore be convenient to designate it as the *acid formula A*. In order to adapt his formula to the many analyses in which the ratio of SiO_2 : Total Hydrogen is approximately 4:20, Clarke presents a modification of his formula consisting in the substitution of a basic, bivalent, aluminium-hydroxide radical (AlOH) for two of the hydrogen atoms of his acid, and he always represents the (AlOH) radical as replacing the two hydrogen atoms attached to BO_3 . Fluorine in this formula is considered as combined with aluminium to form a bivalent radical (AlF) isomorphous with (AlOH) instead of replacing the BO_2 group. A hypothetical molecule containing the (AlOH) radical and employed for expressing the composition of the green tourmaline from Haddam Neck is as follows:

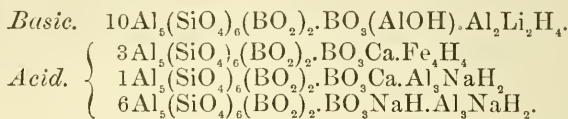


By substituting for the metals in the foregoing formula their equivalent of hydrogen, and simplifying, there results an expression which may be designated the *basic formula B*, so designated because it contains more hydrogen atoms than the *acid formula A*. The two are given together for comparison:



The empirical formula of the tourmaline acid, as derived by Foote and the author, $\text{H}_{20}\text{B}_2\text{Si}_4\text{O}_{21}$, is equivalent to $\text{H}_{30}\text{B}_3\text{Si}_6\text{O}_{31\frac{2}{3}}$, which is exactly midway between the *acid and basic formulas* of Clarke. Thus in order to find expressions that will yield calculated percentage values agreeing with the several analyses becomes again not a matter of chemical science, but rather an arithmetical problem, and one, too, which is bound to succeed; for, given the two formulas, some mixture of the molecules can be found to fit any analysis which falls within the limits

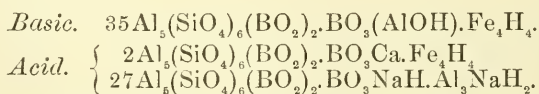
of Clarke's two types. Thus to express the composition of the green tourmaline from Haddam Neck, where the ratio of SiO_2 : Total Hydrogen is very close to 4:20, Clarke employs an equal number of *basic* and *acid* molecules as follows:



Prof. Clarke is very ingenious in his use of figures, and the agreement between the calculated percentages and the results of the analysis is indeed most satisfactory; but what is gained? The analysis gave a ratio of almost exactly 4:20 and a mixture of ten of the *basic* and ten of the *acid* types of Clarke's molecules must yield a ratio of exactly 4:20; hence it is a foregone conclusion that the theory must agree with the analysis. In order to bring about the agreement, however, a mixture of molecules is employed containing an aggregate of 1092 atoms.

If we are to accept the idea that formulas must in some way be found to suit all analyses, and, therefore, it is necessary to have pairs of formulas of acid and basic types, why not accept as the acid type for tourmaline the formula proposed by Rammeisberg, $\text{H}_{16}\text{B}_2\text{Si}_4\text{O}_{20}$? By substituting two bivalent (AlOH) groups for four hydrogen atoms, and then reducing to the simple hydrogen expression, we obtain for the basic type $\text{H}_{22}\text{B}_2\text{Si}_4\text{O}_{22}$, and by the proper replacement of the hydrogen atoms by metals and the combination of the two types of formulas, theoretical compositions could be calculated to a degree with all of the analyses. Indeed, these two formulas have some advantage over Clarke's, in that they are simpler, and priority, at least for one of them, can be claimed. Doubtless other pairs of formulas could be devised which would suit all variations presented by tourmaline analyses.

One other illustration presented by Clarke in support of his theory may here be cited, black tourmaline from Auburn, Maine, No. 2 of Riggs (page 22). The composition is expressed as follows:



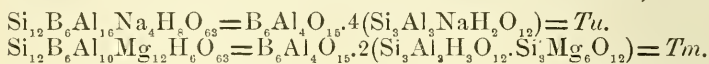
The aggregate number of atoms in the above mixture of molecules is 3499. Let it be observed (page 22) that as far as the ratio of SiO_2 : Total Hydrogen is concerned this analysis is one of the very best, the ratio of SiO_2 : Total Hydrogen being 4:20.03, which would require the basic and acid types of

Clarke's molecules to be present in nearly equal proportion. There are employed by Clarke, however, 35 *basic* and 29 *acid* molecules, from which the ratio of SiO_2 : Total Hydrogen = 4:20.06 is readily calculated: thus the ratio as derived from Clarke's molecular mixtures is farther removed from 4:20 than that derived from the analysis itself, and yet Prof. Clarke makes the following statement: "It will be noticed that the molecule A" (with him the basic one) "is in excess of the other two; a condition which fits the analyses, but which is incompatible with the formula proposed by Penfield and Foote." It may be stated, however, that the $\text{H}_{20}\text{B}_2\text{Si}_4\text{O}_{21}$ formula of tourmaline is not based upon hypothetical molecular mixtures, but, fortunately, upon actual analyses, and good ones, like the one by Riggs last cited, where the ratio is almost exactly 4:20. This example furnishes a good illustration of the fact that ratios are more serviceable in support of a formula than comparison between percentage values.

And, finally, Prof. Clarke, in the concluding pages of this article, presents his reasons for believing that the radical ($=\text{Al}-\text{BO}_2$) may be replaced "in part by the similar groups ($=\text{Al}-\text{OH}$) and ($=\text{Al}-\text{F}$)," or, in other words, that hydroxyl and fluorine are equivalent to, and isomorphous with, BO_2 . Clarke bases this conclusion upon the fact that, in many of the analyses of Riggs, and Jannasch and Kalb, the amount of B_2O_3 found is not sufficient to yield a ratio of SiO_2 : B_2O_3 =4:1 (see page 22), and the deficiency seems to him to be too great to be due to experimental errors. For explaining the occasional low determinations of B_2O_3 referred to, there is a far simpler way than the one proposed by Clarke: special pains were taken by Foote and the author to ascertain the conditions for accurately determining B_2O_3 in tourmaline, and it was found that by fusing the mineral with five times its weight of sodium carbonate and extracting with water, a little boron was still retained by the residue; a fusion of the residue with another portion of sodium carbonate was therefore made, and the boron was then determined by the well known Gooch method. It was further demonstrated that two fusions with sodium carbonate were sufficient for extracting all of the B_2O_3 . Professor Riggs has kindly informed the author that in making his boron determinations by means of the Gooch method, he fused only once, but used, however, ten times as much sodium carbonate as mineral. It is hence probable that a double fusion with five times the weight of sodium carbonate is better than a single fusion with ten times the weight. As may be seen on page 22 the ratio of SiO_2 : B_2O_3 is very close to 4:1 in the majority of Riggs' analyses. All who had occasion to determine boron prior to the description by Gooch of his admira-

ble method, unite in testifying that accurate and reliable estimations of B_2O_3 were exceedingly difficult, in fact almost impossible, to obtain. Jannasch and Kalb employed Bodewig's* modification of Marignac's method for estimating B_2O_3 , weighing boron as KBF_4 ; hence, considering the complex character of tourmaline, their determinations are certainly as close to the truth as could be expected (see page 22), especially when it is taken into consideration that some boron would be lost if a double fusion with alkali carbonate was not made.

Tschermak seeks to explain the composition of tourmaline as a mixture of two complex silicate molecules, as follows:



The radical $(Si_3Al_3NaH_2O_{12})$, in *Tu* is the generally accepted paragonite formula, and $(Si_3Al_3H_3O_{12} \cdot Si_3Mg_6O_{12})$ in *Tm* is Tschermak's typical meroxene formula, except that a part of the hydrogen in meroxene is replaced by potash. Thus tourmaline is supposed by Tschermak to contain mica molecules in combination with the boron radical $B_6Al_4O_{15}$.

In order to express the composition of the two varieties of tourmaline analyzed by Foote and the author, the following relations are employed: for Haddam Neck, *Tu*₄₄, *Tm*₇, and for DeKalb, *Tu*₁₂, *Tm*₃₇. The foregoing relations when written out in linear form are, respectively, $Si_{612}B_{306}Al_{774}Mg_{84}Na_{176}H_{394}O_{3213}$, and $Si_{600}B_{300}Al_{578}Mg_{444}Na_{52}H_{326}O_{3150}$. Thus with expressions, each containing about 5500 atoms, Tschermak shows that the calculated percentages agree in a satisfactory manner with the results of the analyses after the latter have been very much simplified by making numerous substitutions and recalculating to 100 per cent.

The two formulas of Tschermak, $Si_{12}B_6Al_{16}Na_4H_8O_{63}$ (*Tu*) and $Si_{12}B_6Al_{10}Mg_{12}H_6O_{63}$ (*Tm*) are both derived from an acid $H_{60}B_6Si_{12}O_{63}$, which is three times the empirical formula $H_{20}B_2Si_4O_{21}$ proposed by Foote and the author. Tschermak states that to refer tourmaline to the simpler acid $H_{20}B_2Si_4O_{21}$ is a decided step in the wrong direction, since it was shown ten years ago that the three-fold formula was correct. True, by taking suitable mixtures of Tschermak's *Tu* and *Tm* molecules, it is possible to calculate theoretical percentages which agree closely with the results of simplified analyses, but that fact does not necessarily prove the correctness of the formulas under consideration, and that the true constitution of tourmaline has been established. In his admirable *Lehrbuch der Mineralogie* Tschermak gives the composition of both pyrite

* Zeitschr. Kryst., viii, p. 211, 1883.

and marcasite as FeS_2 ; calcite and aragonite as CaCO_3 ; rutile, octahedrite and brookite as TiO_2 ; and cyanite, andalusite and sillimanite as Al_2SiO_6 . If, therefore, for these compounds the simplest formulas are employed, it seems certainly a great presumption to state that in the case of tourmaline the complex three-fold formula had been definitely proved.

Finally Tschermak, who, like Clarke, does not seem to take into consideration that some analyses may not be reliable, tries to explain the composition of the tourmalines from Pierrepont and Gouverneur (Nos. 19 and 20, p. 22), and Tamatawe (No. 9, page 22), by assuming the existence of a third molecule, $Tn = \text{Si}_{12}\text{B}_6\text{Al}_3\text{Mg}_{12}\text{H}_8\text{O}_{61} \approx \text{B}_6\text{Al}_4\text{O}_{16} \cdot 2(\text{Si}_3\text{Al}_2\text{H}_4\text{O}_{11} \cdot \text{Si}_3\text{Mg}_6\text{O}_{12})$, in which the portion $(\text{Si}_3\text{Al}_2\text{H}_4\text{O}_{11} \cdot \text{Si}_3\text{Mg}_6\text{O}_{12})$, is supposed to be analogous to phlogopite. It does not seem reasonable, however, to believe that the tourmalines from the localities in question are essentially different from those from other localities, and need, consequently, different formulas, especially since we have a recent analysis of the Pierrepont, and an early analysis of the Gouverneur varieties of tourmaline by Rammeisberg, both of which conform to the $\text{H}_{20}\text{B}_2\text{Si}_4\text{O}_{21}$ formula.

Summary.—As shown by the tabulation of ratios on pages 21 and 22 there exist a series of recently made and carefully executed tourmaline analyses which give ratios of $\text{SiO}_2 : \text{B}_2\text{O}_3 : \text{Total Hydrogen}$ approximating closely to $4 : 1 : 20$, from which the empirical formula of the tourmaline acid, $\text{H}_{20}\text{B}_2\text{Si}_4\text{O}_{21}$, is derived. That a few analyses do not yield ratios agreeing as closely as desired to $4 : 1 : 20$ is not to be wondered at, when the difficulties presented by the analysis are taken into consideration, together with the fact that the material analyzed might not in all cases have been perfectly pure and homogeneous. As far then as analytical evidence may be relied upon for establishing the formula of a mineral, it may be considered *as definitely proved that the empirical formula of the tourmaline acid is $\text{H}_{20}\text{B}_2\text{Si}_4\text{O}_{21}$* . The science of inorganic chemistry has not yet reached such a state of development that it can be proved, as stated by Tschermak, that the threefold formula, $\text{H}_{60}\text{B}_6\text{Si}_{12}\text{O}_{63}$, is the correct one. The empirical formula $\text{H}_{30}\text{B}_3\text{Si}_6\text{O}_{31}$, proposed by Clarke, can rest only on the analytical evidence supplied by a few analyses of Riggs and one by Jannasch and Kalb which yield ratios approximating to $4 : 1 : 19.33$ (page 22), and there are good reasons for believing that these ratios would not be obtained a second time if the analyses were repeated. Since tourmaline always yields sufficient water to form two hydroxyl radicals, it may be considered *as probably, if not absolutely, proved that the formula of the tourmaline acid should be $\text{H}_{18}(\text{OH})_2\text{B}_2\text{Si}_4\text{O}_{19}$* . Beyond this point it seems safe only to speculate and it cannot be con-

sidered that the ideas presented are capable of being definitely proved. All of the analyses indicate that at least half of the hydrogen atoms of the tourmaline acid, are replaced by aluminium, and this fact, coupled with the idea that it seems reasonable to unite the two hydroxyl radicals with the two boron atoms, led to the suggestion by Foote and the author* that the characteristic feature of all varieties of tourmaline is an *aluminium-borosilicic acid*, $H_2Al_3(B.OH)_2Si_4O_{19}$. In this acid the mass effect of the $(Al_3(B.OH)_2Si_4O_{19})$ is regarded as so overwhelming that it makes no difference how the nine remaining acid hydrogen atoms are replaced, whether largely by aluminium and to a trifling extent by bivalent metals and alkalis, or largely by magnesium and to a trifling extent by aluminium and alkalis, the result in all cases is tourmaline with its characteristic crystalline structure. That trivalent, bivalent and univalent metals, playing as it were the role of isomorphous constituents, may unite in replacing the nine hydrogen atoms of the tourmaline acid, is indeed a remarkable feature of isomorphism, but it furnishes an explanation of the composition of tourmaline, and one which can be comprehended, in part at least.

Looked at from the standpoint of an instructor, what explanation of the chemical composition of tourmaline can be given to a student provided the ideas of Tschermak prevail? Only this, that the composition is exceedingly complicated; that there are two molecules *Tu* and *Tm* (page 30), exhibiting certain analogies to minerals of the mica group, which mix in varying proportions, and that by taking appropriate multiples of the two molecules theoretical compositions can be calculated to agree with the results of analyses, provided the latter are very much simplified. By taking molecular mixtures containing in the aggregate several hundred and even thousand atoms, as done by Clarke (page 28) and Tschermak (page 30), it would seem as though chemists or perhaps arithmeticians might aspire to devise formulas for expressing the chemical composition of any sort of substance of which any kind of an analysis has ever been made.

* Loc. cit., p. 118.

ART. IV. — *Studies in the Cyperaceæ*; by THEO. HOLM.
 XIII. *Carex Willdenowii* and its allies. (With three figures
 in the text.)

TORREY, in his monograph of North American *Cyperaceæ*, was the first author to call attention to the peculiar habit of *Carex Willdenowii* Schk., which he suggested separating from *Carex* proper as a distinct genus, and to which the name "*Phyllostachys*" would be appropriate. The foliaceous bracts and the distinctly articulated base of the style were the characters, which Torrey considered as being of sufficient importance for the establishment of this new genus. Later authors have all agreed, however, in not adopting this genus, but merely to accept *Phyllostachys* as a section of *Carex*, and it seems very natural, since such articulation of the style is also observable in several other species of very remote relationship, besides that the foliaceous bracts are not characteristic of this species alone. A strange coincidence is, that another botanist made a similar suggestion, but many years later, concerning a *Carex*, which in many respects is analogous with *C. Willdenowii*. Duval-Jouve discovered in France, near Montpellier, a species of *Carex*, which he named *C. ædipostyla* on account of its articulated style-base, and in which he noticed the rhachis to be winged and the bracts to be foliaceous; this author suggested the establishment of a new section: "*Ædipostyla*," and he considered *C. moesta* Kunth and *C. phalaroides* Kunth as its nearest allies, since, also, these possess a "stylus basi bulboso-incrassatus." Moreover Duval-Jouve proposed *C. phyllostachys* Mey. as a member of this same section, on account of its small number of female flowers, its foliaceous bracts, its winged rhachis, the shape of its utricle and finally its articulated style-basis. But, on the other hand, Duval-Jouve does not seem to have known *C. Willdenowii* and he evidently did not know of Torrey's monograph, since he makes no allusion to this; he would, no doubt, have counted *C. Willdenowii* among his *Ædipostylæ* inasmuch as the utricle of this species differs less from that of his new species than that of *C. phalaroides* and *C. moesta*. We may thus consider *Phyllostachys* and *Ædipostyla* as analogous forms of our genus *Carex*, but it will be necessary to remove some of the species, which formerly have been considered as representatives of these sections.

In examining the section *Phyllostachys* as adopted by Carey,* this contains, as understood by him, only *C. Willdenowii* Schk., *C. Steudelii* Kth. and *C. Backii* Boott, all of

* For references consult the bibliography appended to this article.

which may surely be regarded as closely related to each other, thus forming an apparently natural section or group. The same view was, also, held by Tuckermann, who placed them as his *Phyllostachyæ* next his *Careyanæ*, in which we find *C. plantaginea*, *C. gynomane* and *C. depauperata*. It was Professor L. H. Bailey that first enlarged the section *Phyllostachys*, and this author includes not only the three species enumerated above, *C. Willdenowii*, etc., as a subsection "*Bracteata*" (*Bractioideæ* Bailey), but also *C. Geyeri* Boott and *C. multicaulis* Bail. as representing another subsection "*Phyllostachyæ*," however with no very clear distinction. As a matter of fact, the section *Phyllostachys* became confused by this enlargement, which is readily noticeable from the statement by Professor Bailey that the section is not only "to be regarded, probably, as an offshoot of the *Montanæ*, but also to be connected with the *Laxifloræ* by *C. multicaulis*, being related to *C. Hitchcockiana*." It is true that Drejer considered his *Carices dactylostachyæ* as derived from his *C. sphæridiophoræ*, but this author would not have included the true *Phyllostachys* in any of these, judging from the diagnosis, while he might have included *C. Geyeri* among the *Dactylostachyæ*. In a lately published paper on the *Carices* of South America, a German author, Rev. G. Kükenthal, follows Prof. Bailey's disposition of *Phyllostachys*, and includes another species, *C. Sellowiana* Schl., which, however, shows no marked affinity to either *C. Willdenowii* or to *C. Geyeri*. Rev. G. Kükenthal, moreover, considers *Phyllostachys* as a subsection of Drejer's *Sphæridiophoræ*, which does not seem to be in conformity with the original diagnosis of this section. While thus Torrey's *Phyllostachys*, even if reduced to a section, has been maintained through so many years, Duval-Jonve's *Ædipostyla* does not seem to have gained recognition at all. In the Catalogue of European species of *Carex*, Dr. H. Christ has only enumerated *C. ædipostyla* as one of the *Digitatæ* Fries and placed it between *C. Halleriana* Asso and *C. pilulifera* L., while this same author has counted *C. depauperata* Good and *C. olbiensis* Jord. among the *Phyllostachyæ* of Torrey. If we consider the *Digitatæ* as understood by Dr. Christ, these contain representatives that are so little related to each other, that we must combine the *Sphæridiophoræ*, the *Dactylostachyæ* and the *Lamprochloenæ* with each other in order to get them all together. It may be, however, that *C. ædipostyla* is in some respects related to, for instance, *C. digitata* L., as suggested by Dr. Christ, but we cannot find any characters by which we might connect this with *C. pilulifera* or with *C. supina* Wahlbg. Future research must decide the proper place for *C. ædipostyla*; in the present paper we intend only

to discuss the position of the old section *Phyllostachys*, as defined by Torrey.

We have seen from the above that the section has been considered as an offshoot of the *Montanæ* Fr., which would mean to indicate some relationship between the *Sphæridiophoræ* on the one side and *Phyllostachys* on the other. We have, also, seen that it has been regarded as a subsection of the *Sphæridiophoræ*, which would even indicate a still closer relationship to these. The question is now whether our conception of the *Sphæridiophoræ* allows any such combination, if we take this in absolute accordance with Drejer, the author, who established the section. As long as we quote these sections as those of Drejer, and not "ex parte," we are supposed to follow the diagnosis strictly with no modifications whatsoever.

Let us read what Drejer says about the *Sphæridiophoræ*, on page 9 in his *Symbolæ Caricologicæ*: "Spica mascula unica. Spicæ foemineæ una vel plures, rotundatæ, sibi et masculæ approximatae, sessiles. Bracteæ membranaceæ nervo dorsali excurrente herbaceo, vel herbaceæ ad basin membranaceæ. Perigynium membranaceum, pube hirsutie vel tomento vestitum, caryopsin arcte includens, rostro brevius longioreve apice scarioso-hyalino subbilobo. Stigmata terna, caryopsis trigona. Centrum habet hic grex in *C. pilulifera*, *ericetorum* cet., ex quibus character sumendus est. Ortum ducit inter monostachyas, nominatim in *C. Wormskjoldiana* Hornem," etc. The central species are all of ordinary habit; none possess foliaceous bracts replacing the *squamæ*; in none of these are there androgynous spikes on long axillary, basal peduncles; in none of these is the rachis winged or zigzagged, nor is the male portion of the inflorescence so inconspicuous as in *Phyllostachys*; furthermore the utricle is not extended into a long, straight and rough beak in any of these. It seems difficult to understand, how any author could see such close relationship to exist between these sections, so distinct do they really appear. And even if we, agreeing with Drejer, see no objection for admitting species with almost simple or very little decompound inflorescence among such, which possess a number of separate spikes, androgynous or with separate sexes, we nevertheless cannot detect any connecting forms by which we might unite the *Sphæridiophoræ* or *Dactylostachyæ* with *Phyllostachys*, inasmuch as we cannot even consider *C. Geyeri* and *C. multicaulis* as members of this section, the *Phyllostachys* of Torrey.

It is now interesting to notice, that Drejer did not only decline to accept *Phyllostachys* as a genus, but that he did not even consider any of its species with enough significance to constitute even a special section of *Carex*, although he was well acquainted with these, at least with *C. phyllostachys*, *C. Willde-*

nowii and *C. Steudelii*. However, since authors of recent years have seen no difficulty in retaining this as a section, it will be necessary to analyze it again in order to decide the actual importance of its morphological characters: the structure of the inflorescence.

Carex Willdenowii possesses a terminal, androgynous spike with male flowers at the top, and a few lateral similarly androgynous, which are developed on long peduncles from the axils of the basal leaves. The plant thus shows a somewhat peculiar habit, which is not frequently met with in the *Cavices genuinæ*, but which, to some extent, suggests certain species of *Vigneastra* Tuckerm. If we consider the simple androgynous spike by itself, this is just one of the fundamental characters of the old group "*Monostachyæ*" with its sections *Psyllophoræ* and *Capitatae*; but if we, furthermore, include the basal, long-peduncled spikes as being merely a part of the main inflorescence, the plant becomes better referable to the group *Heterostachyæ*. In these the lateral peduncles are provided with a clado-prophyllon at their base, an organ which is, also, present in *C. Willdenowii*. The basal position of these lateral spikes, or to be exact "spicate inflorescences," is very characteristic in *C. Willdenowii*, while these are situated higher up on the culm in most of the *Heterostachyæ*, and are more or less remote from each other, as we remember from *C. sylvatica* Huds., *C. laxiflora* Lam., etc. This distinction, drawn from the basal position of the lateral inflorescences in *C. Willdenowii* and its nearest allies *C. Steudelii* and *C. Backii*, is not, however, applicable only to these species, but to several others, which show no affinities or immediate relationship to the members of *Phyllostachys*. It is to be observed in *C. basilaris* Jord., *C. phalaroides* Kth., *C. pedunculata* Muhl., *C. Linkii* Schk. and in *C. adpostyla* Duv.-Jouve, which may belong to the sections *Trachychlænæ* and *Dactylostachyæ* of Drejer. While thus our species of *Phyllostachys* may be arranged very naturally among the *Heterostachyæ*, still another disposition may be made when we examine the section *Vigneastra* of Tuckermann. As described by this author, these species possess decompound, ramified inflorescences, which are always androgynous, male at the top, and may have two or three stigmata. If we, carefully, compare these species with *Phyllostachys*, we soon discover that each branch of *C. cladostachya* Wahlbg., for instance, corresponds to a complete flower-bearing stem of *C. Backii* with its terminal and axillary inflorescences.

The clado-prophylla P in *C. cladostachya* (fig. I) correspond to those same organs, P, in *C. Backii* (fig. II); the bracts L correspond to the green leaves L in *C. Backii*, of which only the scars are indicated; the androgynous spikes A, B, C and D to those of

C. Backii. Finally the terminal androgynous spike A is equivalent with the same, A, in this species, and the utriculi, painted black in the figures, correspond to those of *C. Backii*. We might even go still farther and confine ourselves to the uppermost lateral branch in *C. Backii* (B in fig. II), which showed the



EXPLANATION OF FIGURES.

FIG. I. *Carex cladostachya* Wahlbg., lateral branch of inflorescence, enlarged. A = the terminal, B, C and D = the lateral spikes, all androgynous with the male flowers at the top; P = clado-prophyllon; L = leaves, the lowest one removed. Utriculus is painted black in all the spikes.

FIG. II. *Carex Backii* Boott, the complete inflorescence of a single shoot, enlarged. Letters as above; the leaves (L) have been removed. All the spikes are androgynous with the male flowers at the top; the uppermost lateral spike (B) is developed upon the rhacheola of an abnormal utriculus, of which the female flower is merely rudimentary. Utriculus is painted black.

FIG. III. *Carex adipostyla* Duv.-Jouve; the terminal inflorescence, showing the zigzagged rhachis and the leaf-like bracts; copied from Duval-Jouve, but reduced one-half.

common case of a prolonged, flower-bearing rhacheola. In this specimen, the little branch B constitutes the rhacheola bearing four female, one-flowered spikelets and one terminal, male spike, while the utricle at the base has acquired a corresponding development as the clado-prophyllon P in *C. cladostachya*: thus the ramified rhacheola demonstrates exactly the same structure as the branches B, C and D in *C. cladostachya*. Considered from this point of view the section *Phyllostachys* Torr. is morphologically inseparable from any of the other sections of *Eucareæ*, of *Vigneæ* or *Vigneastræ*. Utriculus shows in these sections or subgenera, as they have been considered by some authors, the same plasticity of developing either as a membranaceous clado-prophyllon at the base of peduncles or as a normal utricle, surrounding the female flower. The extension of the rhacheola may sometimes result in the suppression of the female flower, but not always. Both Duval-Jouve, Gay, Roeser and Schulz enumerate not a few cases where the female flower had developed normally; besides that the rhacheola protruded through the orifice of the same utricle, bearing one or two female flowers with bracts, utricles and pistils. The clado-prophyllon at the base of the lateral peduncles has, also, been observed to contain a pistil, perfect or rudimentary. But a case like that which we have observed in *C. Backii* differs from most others by the development of a complete male spike besides the female ones at the base; it is evidently an exceptional occurrence, and it illustrates exactly one of the minor inflorescences in the genus *Schænoxiphium*. As a matter of fact, this genus, as well as *Uncinia* and *Kobresia*, show a very striking resemblance to *Careæ*, when considered from a morphological viewpoint.

As regards the development of the lower bracts in *Phyllostachys* into green leaves instead of membranaceous scales, this character does not seem to be of much importance as far as concerns the systematic position of the species. Similar green bracts are, also, known from the other sections, for instance in specimens of *Careæ scirpoidea* Michx. of the *Sphæridiophoræ*, where a single or sometimes even several female spikelets may be seen subtended by leaf-like bracts, a case which we noticed in the mountains of Greenland. Moreover in *Careæ Hilairei* Boott, the lower bracts of the androgynous spike are developed into long leaves, as figured by Boott (Plate 468); this species may probably be referred to the *Dactylostachyæ*. Furthermore *C. phyllostachys* Mey., *C. multicaulis* Bail. and *C. ædipostyla* Duv.-Jouve show similar leafy bracts, and have, therefore, been regarded as members of *Phyllostachys*, although they do not seem to possess the most important character of this section and are, perhaps, better placed in a section parallel

with the *Dactylostachyæ*. The zigzagged rachis is another character, which has been claimed as important to *Phyllostachys* Torr., yet this is, also, noticeable in *C. arctata* Boott, in *C. phyllostachys* Mey. and *C. ædipostyla* Duv.-Jouve and several others. The winged margins of the rachis is, on the other hand, a rare feature in *Carex*, and seems only to have been observed in the species of *Phyllostachys* Torr., in *C. ædipostyla* and *C. phyllostachys*. Finally to be mentioned is the utricle, which, as already stated by Drejer, is one of the most essential organs for classifying the *Carices* in sections, and he distinguishes between three categories within *Eucarex*: "Prima est, ubi margines tam arcte circa caryopsin concresecunt, ut verum rostrum non formetur; perigynium tum aut apice perforatum est, aut brevissime rostellatum ore integerrimo vel subemarginato. (*C. polytrichoides*, *baldensis*, *pallescens*, *atrata*, *vulgaris* cett.) Secunda est, ubi rostrum quidem formatur, sed ab inferiore perigynii parte non aut vix distinguitur, estque apice hyalinum, bilobum aut irregulariter bifidum. (*C. pilulifera*, *ericetorum* et aff., *frigida* et aff. cet.) Tertia denique et perfectissima est, ubi rostrum verum et distinctum, apice distincte bifidum aut bicuspidatum formatur. (*C. distans* et aff., *C. vesicaria* et aff. cet.) These sections are again divided into "greges": *Carices melananthæ*, etc., which are distinguished by the consistency of utriculus, whether membranaceous or spongy, glabrous or hairy, and furthermore by the disposition of the sexes, though of lesser importance.

If we now apply the classification, based upon the structure of utriculus, it is readily to be seen that *Phyllostachys* as it stands at present, including *Bracteata* and *Phyllostachyæ*, is no very natural section: that the *Bracteata* are quite distinct from the latter. We have already in the preceding given our views about the position of Professor Bailey's *Phyllostachyæ* as more properly to be arranged parallel with the *Dactylostachyæ*, while *Phyllostachys* proper shows certain analogies with Drejer's *Hymenochlænæ*. In any case the distinction that has hitherto been drawn between *Phyllostachys* and the other sections is by no means tenable, nor have any accordances been proved to exist between these species and Professor Bailey's *Phyllostachyæ*: *C. Geyeri*, etc. Moreover, when Drejer did not exclude the distribution of the sexes as being of some importance to the classification, he, nevertheless, was well aware of the fact that "*formæ hebetatæ*" do exist in most, if not in all, the "greges," which he proposed. These "*formæ hebetatæ*" constitute such species as may naturally be looked upon as old types of the respective sections or "greges," for instance *C. scirpoidea* Michx. as being the type of *Sphæridiophoræ*, etc. By combining the morphological

features of the extreme forms, the simpler with the higher developed, we cannot avoid noticing that they occasionally unite in some characters, which may be understood as exhibiting their descent through modifications from a common, or fundamental type. We may in this manner take *Phyllostachys* of Torrey to be a lesser developed form of *Hymenochlænæ*, and these species "sensu strictiori" do represent several analogies in habit and structure. But we have, also, demonstrated in the preceding, that *C. Backii*, *C. Willdenowii* and *C. Steudelii* show characters that are, also, common to *Vigneastræ*, which, however, does not necessarily indicate anything more than this section possesses characters that may be compared with those of the "*Carices genuinæ*." *Vigneastræ*, as far as these peculiar species are understood, represent some certain transition between *Vignæ* and *Eucareæ*. And if it were not for the structure of utriculus and the three stigmata in *C. Willdenowii*, we might have had just as good reason for placing this among the *Vignæ* as a highly developed type, combining these with *Vigneastræ* of Tuckermann. And in regard to the systematic position of *Vigneastræ*, if this is to be retained as a subgenus, it may be most naturally placed between *Vignæ* and *Eucareæ*, as suggested by Rev. G. Kükenthal, instead of as a section between *Microrhynchæ* and *Hymenochlænæ* as proposed by Professor Bailey. In returning to our *Phyllostachys*, the species by which it is represented in accordance with Torrey, may seem naturally to be referred to the *Hymenochlænæ*, but as extreme and poorly developed forms; they may occupy a position almost as far from the central types: *C. sylvatica* Huds. and *C. cherokeensis* Schw. as *C. nepalensis* Sprgl. and *C. longipes* Don:

Hymenochlænæ.

- I. Spicis androgynis, apice masculis, paucifloris. (*C. Willdenowii*, *C. Steudelii* et *C. Backii*.)
- II. Spicis simplicibus sexu distincto plus minus densifloris. (*C. sylvatica* Huds., *C. cherokeensis* Schw. cet.)
- III. Spicis androgynis, apice masculis, plus minus densifloris. (*C. nepalensis* Sprgl., *C. longipes* Don).

In the preceding pages we have attempted to demonstrate the systematic position of these species, formerly constituting the genus *Phyllostachys* Torr.; we shall, now, briefly discuss the internal structure of these in order to demonstrate the differences between them and *C. Geyeri*, *C. adipostyla* and *C. multicaulis*. On the other hand, we freely admit, that divergencies in anatomical structure may not necessarily indicate distant relationship, especially not if we adopt the classifi-

cation of Drejer, where the species are arranged not in accordance with such artificial characters as are applied for the separation of *mono-*, *homo-* and *heterostachyæ*, or of *di-* and *tri-stigmaticæ*, but in accordance with purely natural affinities. It becomes, thus, evident that each of Drejer's "*greges*" or sections, as we may call them, is really an aggregate of a number of apparently very diverse forms, beginning with the "*formæ hebetatæ*" and passing gradually over into "*formæ centrales*," from which again a number of allied species, but somewhat differently developed, extend towards the limits of the section as "*formæ desciscentes*." An anatomical investigation of some of these parallel forms, the "*hebetatæ*" for instance, will no doubt reveal certain analogies with the corresponding "*formæ hebetatæ*" of other sections, while the central types may be very distinct. In this way the anatomical characters become obscure, and it will be necessary to make many comparisons between the various forms of each section, before we shall know where to draw the distinction between the characters of the section "as its own," and those, which may be regarded as inherited from old types through its "*formæ hebetatæ*." Our anatomical study of *Carex Willdenowii* and its allies must, therefore, necessarily be incomplete, but may, perhaps, be useful in further research, at least as a contribution to comparative studies of other species of the genus. We shall not, however, confine ourselves to these species, but we will compare their structure with that of *C. ædipostyla*, *C. Geyeri* and *C. multicaulis*. The structure of utriculus is, no doubt, the most important, yet, as will be shown, some interesting characters may, also, be obtained from the root, the stem and the leaves. In regard to the localities, from where our material was gathered, we might state that both *C. Willdenowii* and *C. Steudelii* were collected in the woods near Brookland, D. C.; that the former especially prefers shaded places in rich woods, while the latter inhabits drier ground in thickets or open woods; *C. Backii* was received from Mr. James M. Macoun in Ottawa, as collected on dry, grassy and rocky places in open woods and thickets near Ottawa; *C. Geyeri* from the Rocky Mountains in Montana and British Columbia, in somewhat damp ground; *C. multicaulis* from Wyoming; *C. ædipostyla* from thickets near Montpellier in southern France.

The root.

The epidermis and hypoderm show a uniform structure in all three species: *Carex Willdenowii*, *C. Steudelii* and *C. Backii*. The cortex is thin-walled throughout in the last species, but

shows a distinctly thickened outer zone in the two other species. The inner cortex shows the usual tangential collapsing. Endodermis is exceedingly thick-walled in *C. Willdenowii* and *C. Backii*, but not so in the third species. The pericambium is more or less thick-walled in all three species and it is, also, interrupted by all the proto-hadrome vessels. In *C. Backii* we noticed that the average number of pericambium-cells between each two proto-hadrome vessels was only two or three, while there were four in the others. The conjunctive tissue was quite thick-walled in *C. Steudelii* and *C. Willdenowii*, but much less so in *C. Backii*. The root-structure is thus very uniform in these species, and agrees very much with that of *C. ædipostyla*, except that the pericambium is much more thick-walled in this species, besides that there are here mostly five pericambium-cells between each two proto-hadrome vessels. Very different is the structure exhibited by *C. Geyeri* and *C. multicaulis*. The cortex in the former consists exclusively of very thick-walled parenchyma, which is more or less collapsed though without leaving any lacunes; in the other species, *C. multicaulis*, it is only the outermost five or six strata that are thick-walled, while the inner are quite thin-walled, but not collapsed, although the root was apparently of the same age as the others. The endodermis is exceedingly thickened, especially on the inner walls, and very porose. The pericambium, which is not very thick-walled, is in both species only interrupted by some of the proto-hadrome vessels, but in different ways: in *C. multicaulis* every other proto-hadrome vessel bordered on endodermis, in almost regular alternation, while in *C. Geyeri* there was only one vessel out of thirteen which did not border on the endodermis; in some other specimens we noticed that fourteen of twenty-three proto-hadrome vessels were situated inside the pericambium.

The rhizome

is cespitose and very short in our species with the exception of *C. Geyeri*, in which there is a creeping rhizome, but of very short internodes and sympodial as in most of the other *Carices*. The cortical parenchyma is quite broad and somewhat thickened all through in *C. Steudelii*, but not so in *C. Willdenowii*. The endodermis is thickened like an O—endodermis in both species. The stereome is only to be observed as accompanying the mestome-bundles, which it surrounds completely, while it does not occur in the cortex as independent strata or groups. The mestome-bundles are irregularly scattered in several, about four bands, and are nearly all perihadromatic; there is, furthermore, a solid, central pith of slightly thickened cells. In *C.*

Geyeri the cortical parenchyma shows about eight thick-walled strata near the periphery, while the inner ones are very thin-walled; otherwise the structure of the stereome, the endodermis, etc., is like that of *C. Willdenowii* and *C. Steudelii*.

The stem.

The above-ground stem is somewhat rough from prickle-like projections from the epidermis and triangular in *C. Willdenowii*, *C. Steudelii* and *C. Backii*. However, when we examine the structure closer, it is readily seen that the outline is so sharply triangular that the term "three-winged" may be well applicable to these species. The cuticle is smooth and thin in all three species; the epidermis shows a distinct thickening of the outer walls, with the exception of the cone-cells; stomata are quite abundant. The cortical parenchyma is very poorly developed in *C. Steudelii*, consisting only of one stratum short, roundish cells, which border on wide lacunes. In *C. Willdenowii* this same parenchyma is better developed and shows several layers with relatively narrow lacunes; in *C. Backii* there are distinct palisades, but the lacunes are very wide in this. The stereome is thickwalled in *C. Willdenowii* and *C. Backii*, but not so in the third species; it accompanies the mestome-bundles either as hypodermal groups or separated from the epidermis by the cortex; it is, also, present on the hadrome side of the mestome-bundles; it reaches its highest development in the three wings, where it, furthermore, occurs as isolated groups, not being in contact with any mestome-bundles. These, the mestome-bundles, are very regularly arranged in one band near the periphery and represent larger oval in alternation with smaller, almost orbicular bundles, when viewed in transverse sections; the parenchyma-sheath is thinwalled, and the mestome-sheath shows its inner cell-walls to be distinctly thickened. A thinwalled pith occupies the center of the stem; it is hollow in *C. Backii*, solid in the other species.

In *C. ædipostyla* the stem is triangular, but without being winged, and the outer cell-walls of epidermis are more heavily thickened than in the former species; the cortex is more open, there being very distinct intercellular spaces besides lacunes of quite considerable width. The stereome is thickwalled, but does not occur separate from the mestome-bundles, which are here represented in two concentric bands, instead of but one as in the former species; the mestome-sheath shows the cells to be thickened all around. A partly hollow pith occupies the center of the stem. In *C. Geyeri* and *C. multicaulis* the stem is longer than in the former species and shows, also, a

much more solid structure; the cortex constitutes a compact palisade-tissue with relatively small lacunes, especially in the latter species. The stereome occurs in larger groups and is very thickwalled, but invariably accompanying the mestome-bundles; the sharp angles of the pentagonal stem of *C. Geyeri* have no isolated stereome, but merely cortical parenchyma; the terete stem of the other species possesses a larger number of mestome-bundles and, consequently, it contains, also, more stereome. The mestome-bundles constitute a single band and the mestome-sheath is very heavily thickened. The solid pith is thickwalled in *C. Geyeri*, but thinwalled in the other.

The leaf

is narrow, but flat in *C. Willdenowii* and *C. Steudelii*, relatively broad in *C. Backii*. The cuticle is thin and smooth in the former species, but quite thick in the latter. Epidermis exhibits several modifications; it is perfectly glabrous and smooth on the upper face in *C. Steudelii* and *C. Willdenowii*, and is differentiated into a row of bulliform cells just above the midrib; above the stereome the epidermis, viewed "en face" consists of narrower and shorter cells than those which cover the mesophyll. If we examine the lower surface of the leaf we notice a corresponding structure, but there are no bulliform cells, and stomata are observable outside the mesophyll; these are but slightly projecting and possess a wide but shallow air-chamber. If we consider the leaf of these species in cross-sections, we notice that the outer wall of epidermis is slightly thickened and somewhat projecting in the shape of very minute papillæ on either side of the midrib, but only on the lower surface. In *Carex Backii* the epidermis-cells are very short outside the stereome, much shorter than we observed in the former two species; their outer wall is very thick and we find here small wart-like projections from the epidermis, which form several rows on each side of the bulliform-cells, besides that there are four such projections around each stoma. The mesophyll is differentiated into a distinct palisade-tissue on the upper face of the leaf of *C. Steudelii* and *C. Willdenowii*, while it consists of roundish cells with large intercellular spaces on the lower face, almost as a typical pneumatic tissue; large lacunes traverse the mesophyll in both species. In *C. Backii*, on the other hand, the mesophyll is nearly homogeneous throughout the blade, very open and without palisade-cells. The stereome is thickwalled in *C. Willdenowii* and *C. Steudelii*, but quite thinwalled in *C. Backii*. It accompanies the mestome-bundles as hypodermal groups on either face, at least in the larger bundles, and there is, also, an iso-

lated group in each of the two margins. The mestome-bundles are arranged in one plane, larger alternating with smaller ones; the parenchyma-sheath is thin-walled and colorless; the mestome sheath shows distinctly thickened inner cell-walls. The bracts, which subtend the one-flowered, female spiculæ in *C. Steudelii*, *C. Backii* and *C. Willdenowii* are leaf-like, at least the lower ones, and show a structure corresponding to that of the proper leaves, only that they are relatively narrower and possess more stereome than these. The bracts of the uppermost female spikelets are, on the contrary, scale-like, awned, membranaceous and with hyaline margins like the scales of the *Hymenochlænæ* in general.

In comparing the leaf of *C. Steudelii*, *C. Backii* and *C. Willdenowii* with that of *C. adipostyla* we notice almost the same structure; but the outer wall of epidermis is thicker in this species, and without any development of papillæ, as we noticed in *C. Backii* for instance; the mesophyll is very open and is not differentiated into any distinct palisade-tissue. The stereome is quite thick-walled, and the mestome-sheath has its inner cell-walls considerably more thickened than we observed in the former species. The leaves of *C. Geyeri* and *C. multicaulis* are longer and much narrower than in any of the former species; they are slightly conduplicate with a distinct keel in the former, but in the latter the midrib is not projecting. The outer cell-wall of epidermis is quite thick, especially so in the bulliform cells, which are here relatively small and occur as a single group above the midrib, when viewed in transverse sections. No papillæ are visible around the stomata, but in *C. multicaulis* there are a few rows of such projections on each side of the bulliform cells, similar to those observed in *C. Backii*. The stomata are level with epidermis and confined to the lower surface of the blade. The mesophyll is a dense palisade-tissue with a lacune between each two mestome-bundles in *C. Geyeri*, but is more open in the other species. The stereome is extremely thick-walled in both species and accompanies the mestome-bundles, besides that it, also, occurs as isolated groups in each of the two margins in *C. multicaulis*, but not in *C. Geyeri*. The mestome-bundles themselves show the same structure as we have found in the other species, described above, but the mestome-sheath is more thick-walled than in these.

Utriculus.

This little organ furnishes the most essential characters in distinguishing the sections and even the species of *Carex*. It is almost bottle-shaped in *C. Steudelii* with a long, straight and rough beak, and is readily distinguished from that of *C. Willde-*

nowii, where the beak is shorter and less pronounced. However the internal structure is the same for these two species. The outer wall of epidermis on the dorsal face is somewhat thickened, on the ventral face it is much less so; there is but little mesophyll displayed in a few strata and no stereome, not even by the two mestome-bundles, in which, furthermore, the mestome-sheath is quite thin-walled. In *C. Backii* the utricle shows a similar structure with the exception of a few stereome-cells on the leptome-side of the two mestome-bundles. A much firmer structure is exhibited by *C. ædipostyla* in which the epidermis is quite thick-walled on either face, besides that the mesophyll is more compact and occurs in several layers; moreover we find in this species very thick-walled stereome, not only on either face of the mestome-bundles, but also as several isolated groups between these. There are more than two mestome-bundles and these possess a very thick-walled mestome-sheath. The utricle of *C. Geyeri* and *C. multicaulis* has a very short beak like *C. ædipostyla*; it is perfectly glabrous and smooth. The epidermis is heavily thickened on both faces; the mesophyll is poorly developed, and the stereome, which is quite thick-walled, is in *C. Geyeri* confined to the mestome-bundles, while in the other species it occurs as several isolated groups between them.

It would seem from the above as if *C. Steudelii*, *C. Willdenowii* and *C. Backii* are closely related to each other, but not to *C. ædipostyla*, *C. Geyeri* or *C. multicaulis*, hence the section *Phyllostachys*, as defined by recent authors, does not constitute a natural section. If we, moreover, consider those species, which by Torrey were the fundamental ones for his genus, it appears as if they are not sufficiently characteristic among themselves to necessitate the establishment of even a section. And the same is the case with *C. ædipostyla*. In regard to *C. Geyeri* and *C. multicaulis*, these do not possess very pronounced characters either, but similarly to the others may be arranged in some of the larger sections as "formæ hebetatæ." The species of Torrey's *Phyllostachys* appear, on the other hand, as inseparable from the *Hymenochlænæ*, of which they probably represent one of the earlier types. The earliest is unknown in this section, but was perhaps monostachyous; but we might, also, suppose that the *Hymenochlænæ* developed parallel with one of the other sections from one common "forma hebetata," which then branched out into two or more distinct sections.

Brookland, D. C., February, 1900.

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ART. V. — *The Titration of Mercury by Sodium Thio-sulphate*; by JOHN T. NORTON, JR.

[Contributions from the Kent Chemical Laboratory of Yale University—XCV.]

ACCORDING to J. J. Scherer* mercurous nitrate, mercuric nitrate and mercuric chloride may be estimated by direct titration with sodium thiosulphate, Hg_2S , 2HgS , $\text{Hg}(\text{NO}_3)_2$, and $2\text{HgS} \cdot \text{HgCl}_2$ being the precipitates obtained in each case. I have been unable to obtain access to Scherer's original publication, but Sutton† gives the following very general directions for this process:

(a) "*Mercurous salts*.—The solution containing the metal as a protosalt only is diluted, gently heated and the thiosulphate delivered in from the burette at intervals, meanwhile well shaking until the last drop produces no brown color. The sulphide settles freely and allows the end of the reaction to be easily seen. One cm^3 of the $\frac{1}{20}$ normal solution of thiosulphate = 0.02 grams Hg or 0.0208 HgO.

(b) "*Mercuric nitrate*.—The solution is considerably diluted, put into a stoppered flask, nitric acid added and the thiosulphate cautiously added from the burette, vigorously shaken meanwhile, until the last drop produces no further precipitate. Scherer recommends that when the greater part of the metal is precipitated the mixture should be diluted to a definite volume, the precipitate allowed to settle and a measured quantity of the clear liquid taken for titration; the analysis may then be checked by a second titration of the clear liquid if needful. One cm^3 of $\frac{1}{20}$ normal thiosulphate = 0.15 of Hg. or 0.162 of HgO.

(c) "*Mercuric chloride*.—With mercuric chloride the end of the process is not so easily seen. The very dilute solution is acidified with hydrochloric acid, heated nearly to boiling, and the thiosulphate cautiously added so long as a white precipitate is seen to form; any great excess of the precipitant produces a dirty-looking color. Filtration is necessary to distinguish the exact ending of the reaction. One cm^3 of $\frac{1}{20}$ th normal thiosulphate = 0.15 Hg or 0.162 HgO."

Fresenius‡ gives practically the same directions, but omits all mention of that portion of the process dealing with mercurous nitrate.

In view, therefore, of the scant information available on the subject and of the apparent difficulty of working the process accurately according to the directions given, an attempt was

* His Lehrbuch der Chemie, i, 513.

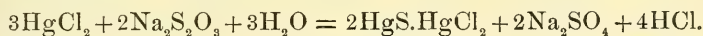
† Volumetric Analysis, p. 220.

‡ Quantitative Analysis.

made to ascertain whether the careful regulation of temperature, dilution and amount of acid present might not produce beneficial results.

That portion of the process dealing with mercuric chloride was first taken up. The mercuric chloride used was pulverized, dried at 100° and its purity proved by several determinations as mercuric sulphide. The sodium thiosulphate was made up of approximately $\frac{1}{20}$ th normal strength and standardized on decinormal iodine, which in turn was titrated against decinormal arsenious acid made from pure resublimed arsenious oxide.

For the action of sodium thiosulphate upon the mercuric chloride Scherer gives the equation,



According to my experience, the action results in the formation of a dense white precipitate which refuses to settle either by shaking or standing, thus making it impossible to fix the end reaction by reading the first drop of thiosulphate which produces no further white precipitate in the solution containing the mercuric chloride. Recourse must be had therefore to filtering. By far the quickest and neatest method is to use the asbestos filter deposited on a large perforated platinum cone.* This cone is set in a glass funnel by means of a rubber connector and the funnel is passed through the stopper of a large side-necked Erlenmeyer connected with an exhaust pump. A little asbestos fiber shaken in the liquid to be filtered was found to be very beneficial in preventing the precipitate from running through the filter. In all the following experiments the thiosulphate was run into the solution containing the mercuric chloride in excess, the whole shaken up with asbestos fiber, filtered and the excess of thiosulphate determined by $\frac{1}{20}$ th normal iodine. This method of procedure seems to be far preferable to attempting to catch the end of the reaction by running in the thiosulphate until the last drop produces no precipitate. In the experiments shown in Table I no attention was paid to the temperature of the solution and the thiosulphate was run in until the liquid turned brown. In every case the solution was allowed to stand until there was no further visible change of color.

A glance at the table shows that the results are most irregular. In Table II is seen the result of regulating the temperature and the length of standing after the addition of the sodium thiosulphate.

* Amer. Chem. Jour., i, 321.

TABLE I.

	HgCl ₂ taken calc'd as Hg. grams.	Na ₂ S ₂ O ₃ in excess. cm ³ .	Volume at beginning. cm ³ .	HgCl ₂ found calc'd as Hg. grams.	Error. grams.
1.	0.0446	46.28	200	0.0343	0.0103—
2.	0.0354	46.28	400	0.0326	0.0028—
3.	0.0356	44.97	400	0.0225	0.0131—
4.	0.0345	44.5	100	0.0308	0.0037—
5.	0.0354	44.43	50	0.0326	0.0028—
6.	0.0382	22.59	50	0.0354	0.0028—
7.	0.0375	8.58	50	0.0385	0.0010+
8.	0.0371	1.84	50	0.0304	0.0067—
9.	0.0731	2.28	50	0.0774	0.0043+
10.	0.1486	9.34	50	0.1489	0.0003+

TABLE II.

	HgCl ₂ taken as Hg. grams.	Volume at beginning. cm ³ .	Tem- perature. C.	Stand- ing minutes.	Na ₂ S ₂ O ₃ in excess. cm ³ .	HgCl ₂ as Hg. found.	Error. grams.
1.	0.0738	50	36°	40	16.68	0.0494	0.0244—
2.	0.0741	50	70	15	15.42	0.0738	0.0003—
3.	0.0741	75	70	12	16.07	0.0733	0.0008—
4.	0.0744	50	70	10	14.6	0.0755	0.0011+
5.	0.0764	50	72	7	6.77	0.0771	0.0007+
6.	0.0762	50	75	10	8.54	0.0799	0.0037+
7.	0.0756	50	73	15	9.99	0.0815	0.0059+
8.	0.0774	50	68	15	10.84	0.0767	0.0007—
9.	0.0745	75	69	7	6.62	0.0805	0.0060+
10.	0.0736	50	68	5	15.82	0.0714	0.0022—

These results, although better than those of Table I, are still very uncertain. On the supposition that the change from white to black, which takes place in the solution after the addition of an excess of sodium thiosulphate more or less quickly according to the temperature, was due to an increased amount of HgS in the compound 2HgS.HgCl₂, the next step was to ascertain whether this could be avoided by stopping the addition of the thiosulphate at the first indication of a change of color in the white precipitate, diluting the solution with a large amount of cold water and immediately throwing it on the filter. The following table (III) shows the result of the experiments.

In the case of quantities of mercuric chloride up to 0.1 gram the results shown in Table III are very satisfactory, but when larger amounts of mercuric chloride are used the errors again become prominent. In Table IV, the effect of lowering the temperature to 60° C. and of increasing the dilution to 100^{cm}³ is shown.

TABLE III.

	HgCl ₂ taken as Hg. grams.	Volume at beginning. cm ³ .	Tem- perature. C.	Na ₂ S ₂ O ₃ in excess. cm ³ .	HgCl ₂ found as Hg. grams.	Error. grams.
1.	0.0749	50	70°	4.15	0.0751	0.0002 +
2.	0.0749	50	75	0.72	0.0728	0.0021 -
3.	0.0756	50	72	1.46	0.0759	0.0003 +
4.	0.0753	50	70	2.57	0.0750	0.0003 -
5.	0.0390	50	70	3.43	0.0395	0.0005 +
6.	0.0388	50	72	8.19	0.0390	0.0002 +
7.	0.0380	50	76	2.03	0.0393	0.0013 +
8.	0.1494	50	78	4.99	0.1498	0.0004 +
9.	0.1489	150	78	4.38	0.1512	0.0023 +
10.	0.1480	50	70	0.52	0.1438	0.0042 -
11.	0.1498	50	78	1.47	0.1540	0.0042 +
12.	0.1484	50	71	2.09	0.1517	0.0033 +
13.	0.1480	75	72	1.59	0.1509	0.0029 +

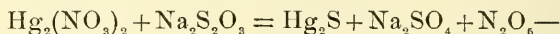
TABLE IV.

	HgCl ₂ taken as Hg. grams.	Volume at beginning. cm ³ .	Tem- perature. C.	Na ₂ S ₂ O ₃ in excess.	HgCl ₂ found as Hg. grams.	Error. grams.
1.	0.0759	100	60°	3.06	0.0766	0.0007 +
2.	0.0384	"	"	2.81	0.0387	0.0003 +
3.	0.1498	"	"	1.1	0.1500	0.0008 +
4.	0.1503	"	"	1.63	0.1506	0.0003 +
5.	0.1479	"	"	2.41	0.1480	0.0001 +
6.	0.1489	"	"	2.12	0.1503	0.0014 +
7.	0.2244	"	"	2.63	0.2259	0.0015 +
8.	0.1490	"	"	2.33	0.1484	0.0006 -
9.	0.0758	"	"	2.	0.0762	0.0004 +
10.	0.0383	"	"	2.58	0.0379	0.0004 -

From this table it is plain that Scherer's process for the estimation of mercury in the form of mercuric chloride is capable of yielding accurate results if carried out under certain fixed conditions. These conditions, which must be closely adhered to, are as follows: The solution containing the mercury in the form of mercuric chloride is placed in a liter flask, diluted to 100^{cm}³ and heated to a temperature of 60° C. The sodium thiosulphate in $\frac{1}{20}$ th normal solution is run in from a burette until the white precipitate formed begins to take on a brownish tinge. The solution is then diluted with cold water, some asbestos fiber added to coagulate the precipitate and the whole is quickly thrown on the filter. After careful washing of the precipitate, the filtrate is diluted to a definite volume, 3 grams of potassium iodide added and the excess thiosulphate

titrated with iodine and starch solution. The duration of the process need not exceed 15 minutes. It is worthy of note that there is no necessity of using any hydrochloric acid in addition to that formed in the reaction. This certainly eliminates one probable source of error—the interaction of hydrochloric acid and sodium thiosulphate.

In dealing with the estimation of mercury in the form of mercurous nitrate the same method of procedure was employed as in the case of mercuric chloride. A solution of mercurous nitrate was prepared by dissolving as much as possible of 20 grams of the salt in about 200^{cm}³ of water, filtering off the clear liquid and diluting to a definite volume. The standard of the solution was determined by precipitation as metallic mercury by means of the electric current. Contrary to the statement made in Sutton, the brown precipitate of Hg₂S—formed as shown in the equation,



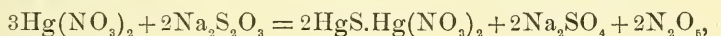
does not settle and leave a clear supernatant liquid, but the solution remains cloudy and it is impossible to see any end reaction. Although the conditions of dilution, temperature and amount of acid present were carefully considered, no arrangement or adjustment of these conditions was found under which satisfactory results could be obtained. Table V gives the result of experiments.

TABLE V.

	Hg ₂ (NO ₃) ₂ taken as Hg. grams.	HNO ₃ 1:3. cm ³ .	Volume at beginning. cm ³ .	Tem- perature. C.	Na ₂ S ₂ O ₃ in excess. cm ³ .	Hg ₂ (NO ₃) ₂ found as Hg. grams.	Error. grams.
1.	0.0148	none	50	50°	4.28	0.0129	0.0019—
2.	0.0148	“	75	60	4.45	0.0117	0.0031—
3.	0.2976	“	300	60	12.67	0.2760	0.0216—
4.	0.1488	“	100	40	2.19	0.1386	0.0102—
5.	0.1488	“	100	50	6.92	0.1378	0.0110—
6.	0.1488	“	200	50	0.78	0.1388	0.0100—
7.	0.0744	“	100	65	0.73	0.0636	0.0108—
8.	0.0744	1	100	55	0.49	0.0733	0.0011—
9.	0.0744	2	100	40	1.16	0.0660	0.0084—
10.	0.0744	1	100	55	1.35	0.0685	0.0059—
11.	0.0744	1	100	55	0.82	0.0686	0.0058—
12.	0.0744	½	200	55	1.71	0.0685	0.0059—
13.	0.0744	4	100	40	1.77	0.0649	0.0095—
14.	0.0744	5½	100	40	1.34	0.0645	0.0099—
15.	0.0744	1	100	45	1.71	0.0654	0.0090—
16.	0.0744	10	100	45	1.55	0.0669	0.0075—
17.	0.1488	1	100	40	0.73	0.1391	0.0097—

The errors in experiments 1, 3, 4, 5, 6, 10, 11, 12 and 17 are, proportionally to the amount of material handled, practically the same and this fact caused me to make a careful revision of all standards; but no mistake could be found. The reaction upon which the process depends requires the formation of Hg_2S , but this mercurous sulphide breaks down immediately into mercuric sulphide and mercury. The latter is probably acted upon by the free nitric acid present to form mercuric nitrate, which in turn is transformed into the compound $2\text{HgS} \cdot \text{Hg}(\text{NO}_3)_2$ by the action of the thiosulphate. At any rate, with an error so large, whatever its source may be, the process is plainly impracticable.

The third step in Scherer's process deals with the action of sodium thiosulphate on mercuric nitrate. In the following experiments a solution of mercuric nitrate was prepared either by dissolving as far as possible 20 grams of mercuric nitrate in about 200cm^3 of cold water, filtering off the supernatant liquid and diluting to a definite volume (exps. 1-8), or by dissolving the salt in strong nitric acid and diluting (exps. 9-19). The standard of the solution was obtained by precipitation as metallic mercury by means of the electric current. The yellow precipitate, formed according to Scherer's reaction,



on adding the sodium thiosulphate settles much better than in the case of either mercuric chloride or mercurous nitrate; but, as the supernatant liquid takes on a permanent yellow color towards the end of the reaction, it is impossible to see when the thiosulphate produces no further precipitation. On this account, therefore, the same method of procedure was adopted as in the case of the mercuric chloride and mercurous nitrate, i. e., filtration and titration of the excess of sodium thiosulphate with iodine and starch solution. The result of the experiments is shown in the following table.

These results seem to show the impossibility of obtaining accurate results according to Scherer's method for the determination of mercuric nitrate by direct titration with sodium thiosulphate. The constant plus error cannot be accounted for on the hypothesis that the nitric acid present decomposes the sodium thiosulphate, for in that case the error would lie in the other direction. It is more probable that the constitution of the compound $2\text{HgS} \cdot \text{Hg}(\text{NO}_3)_2$ is not definite enough to make it the basis for an analytical process.

TABLE VI.

	Hg(NO ₃) ₂ taken as Hg. grams.	Volume at begin- ning. cm ³ .	HNO ₃ 1 : 3 cm ³ .	Tem- perature. C.	Na ₂ S ₂ O ₃ in excess. cm ³ .	Hg(NO ₃) ₂ found as Hg. grams.	Error. grams.
1.	0·1167	200	none	60°	0·46	0·1384	0·0217 +
2.	0·1167	100	1	60	3·63	0·1348	0·0181 +
3.	0·1167	100	2	60	0·23	0·1375	0·0208 +
4.	0·1167	100	none	60	0·17	0·1360	0·0193 +
5.	0·1167	300	"	21	0·12	0·1232	0·0065 +
6.	0·1167	300	"	21	0·15	0·1375	0·0208 +
7.	0·1167	200	"	21	11·8	0·1461	0·0294 +
8.	0·1167	200	20	21	15·97	0·1403	0·0236 +
9.	0·1278	200	none	21	5·23	0·1647	0·0369 +
10.	0·1278	200	"	"	1·35	0·1653	0·0375 +
11.	0·1278	100	"	"	1·43	0·1662	0·0384 +
12.	0·0752	200	"	"	2·09	0·0996	0·0244 +
13.	0·0255	100	"	"	2·01	0·0280	0·0025 +
14.	0·0255	200	5	"	3·44	0·0264	0·0009 +
15.	0·0255	200	10	"	0·85	0·0334	0·0079 +
16.	0·0255	300	none	"	1·96	0·0264	0·0009 +
17.	0·0255	200	"	"	1·9	0·0287	0·0032 +
18.	0·0255	200	"	"	1·76	0·0290	0·0035 +
19.	0·0639	200	"	60	1·53	0·0831	0·0192 +

In conclusion I wish to thank Prof. F. A. Gooch for his kind advice and assistance.

ART. VI.—*Selenium Interference Rings*; by A. C. LONGDEN.

SOME months ago, while at work on thin films deposited by cathode discharge, at the request of Mr. C. C. Trowbridge of Columbia University, I attempted to deposit for him a thin conducting film of selenium. The attempt was not successful; because, although the selenium was in the conducting form in the cathode, the film was always in the amorphous form, which is not an electrical conductor. A number of films were deposited under all sorts of conditions, but they were invariably amorphous and therefore non-conducting.

During the deposition of selenium, although no conducting films were produced, a very interesting and beautiful phenomenon was observed. The selenium cathode was small in comparison with the size of the glass plates upon which the films were deposited, and therefore the distribution of selenium was not uniform, the film being considerably thicker in the center than at the edges. The deposited material is sufficiently transparent to transmit large quantities of light through thicknesses of several wave-lengths; so that when light falls upon a film of varying thickness, the beams reflected from the upper and lower surfaces present interference phenomena similar to Newton's rings. The interference bands which were at first produced were so irregular that they could hardly be called rings, because of the lack of uniformity in the upper surfaces of the films; but after arranging the cathode in the form of a small globule, films increasing uniformly in thickness toward the center were obtained.

These films, when viewed by reflected light, present concentric systems of circular interference fringes of great beauty and remarkable brilliancy. The transparent selenium film is deposited upon plane glass, and, having a convex upper surface, it constitutes a very thin plano-convex lens. The interfering beams of light, however, are not reflected from the surfaces of an air film, between the convex lens and a second plane, but from the two surfaces of the selenium lens itself. As this lens is thickest in the center, instead of being thinnest in the center as in Newton's air film, the selenium rings increase instead of decrease in width and brilliancy, counting from the center outward. For the same reason the order of the colors is red, yellow, green, blue, violet; instead of violet, blue, green, yellow, red; and, instead of having a black center when viewed by reflected light, as is the case with Newton's rings, the color of the center depends upon the order of the ring, counting from the margin inward.

It will be seen by considering the geometrical conditions which govern radiation from a point to a plane, that if our cathode is a small enough globule, the radius of curvature of the film will depend upon the distance of the cathode from the glass plane upon which the film is deposited. Accordingly, this distance determines the diameter of the rings, and the width of any particular ring. By placing the cathode at a suitable distance from the glass, it is very easy to obtain exceedingly brilliant rings as much as 8 or 10 millimeters in width and 4 or 5 centimeters in diameter.

These rings are so brilliant that they are as easily projected upon the screen as ordinary lantern slides; and without the usual precautions in regard to size and illumination demanded in projecting Newton's rings with a glass lens and plate. They should, however, be projected by reflected light.

I am not aware that Newton's rings of this character, namely, interference rings from the reflecting surfaces of a very thin plano-convex lens, have heretofore been produced by any process whatsoever. The physical conditions connected with selenium rings are of a simpler character than those which exist in the case of ordinary Newton's rings, and they may, therefore, afford physicists new opportunities for the study of this subject.

Physical Laboratory of Columbia University,
April 20, 1900.

ART. VII.—*Carboniferous Boulders from India*; by B. K. EMERSON.

IN the spring of 1894, I was in the office of the Geological Survey of India and Dr. Fritz Nötling unpacked for me and explained a fine series of the boulders of the Carboniferous Glacial period of the Salt range in northwestern India.

They had been obtained partly from the talus and partly from the conglomerate itself, and he was kind enough to give me samples of the different forms. Later, Dr. King, the Director of the Survey, observing the unconscious but silent admiration with which I examined the finest boulder of the series, packed it up and sent it after me to my lodgings. This is the largest boulder in the figure and is 10 inches long.



It is a dark red rock halfway between a granite and a quartz porphyry and is perfectly scratched on four sides and rough and battered on the ends.

It is in color, shape, and kind and perfection of striation, so like the great boulder, about 9 feet long, of red sandstone which stands in front of the Geological Museum at Amherst and which was taken from one of the streets at Amherst during the lowering of the latter, that I may cite a figure of the latter for comparison from my monograph of old Hampshire Co., Mon. xxix, U. S. G. S., pl. xxxiii, p. 192.

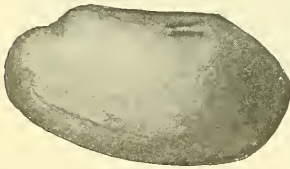
The boulder with the cord around it was broken in dislodging it from the ledge and shows remains of the conglomerate cemented onto its upper surface. The boulder on the small box is scratched on several small facets separated by sharp crests and has also fragments of the conglomerate attached.

The boulder in front was not quite in focus but a dim line runs across the middle of the front from the upper left hand corner to the lower right hand corner and the portion below that line is distinctly striated, while the part above that line is weathered several millimeters deep by long exposure to the weather as it projected on the cliff while the lower portion was protected within the solid mass of the rock. Dr. Nötling explained to me the conditions of the occurrence and there can be no doubt about the glacial character and the Carboniferous age of the deposit. I send this short article because doubt has been expressed in a recent standard text-book as to the reality of the Carboniferous Glacial period.

ART. VIII.—*A new Bivalve from the Connecticut River Trias*; by B. K. EMERSON.

A FEW days ago my former assistant, Mr. Chas. S. Merrick, of Wilbraham sent me a large slab of sandstone of a buff color, much stained by malachite, containing several indistinct casts of a unio-like bivalve.

The shell seems to be one of those fresh or brackish water forms common in the Trias allied to the Unionidæ. It may be compared with the *Anoplophora lettica* Quenstedt,* and may be called *Anoplophora Wilbrahamensis*. It is a distinctly unio-shaped shell 38^{mm} long, 18^{mm} high, 7^{mm} thick. The exterior is smooth with fine lines of growth. The mantel impression is quite deep as preserved on the central portion of the length. The beak rises very slightly and the hinge line is long and straight and there is in the fig-



ure what seems to be the impression of a long, posterior tooth, and above this a slight groove at the place of attachment of the ligament.

The large slab contains 14 imperfect casts, all of which may well belong to one species.

I do not know of any other bivalve shell from the Trias in Massachusetts.

* Pal., plate lxiii, fig. 28.

ART. IX.—*The Statement of Rock Analyses*; by HENRY S. WASHINGTON.

OF late years the importance of chemical analyses in the study of igneous rocks is generally recognized, and their publication and use is becoming more and more common.

Apart from improvements in methods and facilities, as well as increase in the number of workers, the main factor in this increase in the number of analyses published is the growing appreciation of their vital importance for a thorough comprehension of the rocks of the globe.

This importance of rock analyses to-day lies in their disclosing the bearing of the chemical composition of the rock masses on the deeper theoretical problems of petrology. Many of these, such as the differentiation of magmas, the genetic relationships of various rock bodies or parts of a complex to each other, or the true character and meaning of the so-called "petrographic provinces," are only to be solved by means of thoroughly good and trustworthy analyses.

There is an increasing tendency among some petrographers to regard the magma as one of the main objects of study, of which the solid rocks are simply the consolidated and accessible portions, many of their characters being to a great extent fortuitous and dependent on extraneous conditions of solidification and the like.

The chemical composition of the magma either persists as such in the rocks formed from it (as in Brögger's aschistic dikes), or is to be inferred if the magma has undergone changes such as those induced by differentiation (as in Brögger's diaschistic dikes), absorption of country rock or other causes. This chemical composition is, as far as we can tell at present, the single original character of the magma which is open to study after it has undergone the physical and physico-chemical changes which result in the formation of igneous rocks, whether aschistic or diaschistic.

Hence for obtaining a knowledge of these vast and highly important portions of the earth's crust, of their original characters, the conditions under which they exist and the changes which they may undergo, rock analyses are absolutely essential. It is true that physics also comes to our aid in certain directions, but this does not lessen the importance of the chemical investigation.

The practical use of analyses involves their collation or comparison one with another, either those of the rocks of one given region, or those of different regions with each other, or

of rocks belonging to any given group or kind. Scattered as they are through the literature, it is often a matter of some practical difficulty, or at least inconvenience, to bring several together under the eye so as to observe readily the features which it is desired to see. This inconvenience is increased by the fact that there is no uniformly adopted sequence in which the various constituents are stated.

In some cases, as in the publications of the U. S. Geological Survey, the analytical order is followed. The constituents are put roughly in the order in which they are determined in the course of the analysis.

Again a roughly chemical order is followed, either by putting all the acid radicals at the top of the column and the basic radicals below, as is usually done in Tschermak's *Mittheilungen*, or by putting SiO_2 first, followed by TiO_2 , then the metallic oxides in the order R_2O_3 , RO , R_2O , with H_2O , P_2O_5 , Cl , etc. last of all, as is the custom of Roth, Zirkel, Rosenbusch, and many others, including the writer. But even here there are variations in the order followed, as in some cases CaO is put before MgO , and K_2O before Na_2O .

A third method is that adopted of late years by Pirsson. This, which may be called the petrographic order, consists in putting the nine most important oxides first, beginning with SiO_2 and ending with H_2O , while the subordinate and more rarely occurring constituents follow after in the same general order.

A last method shows a lack of any system, the constituents being put down almost at haphazard, and with little apparent attempt at order or natural association. This method is fortunately rarely met with, and is to be dismissed at once as quite unworthy of consideration.

In consequence of this confusion not only is the inconvenience of and time required for copying and comparing analyses greatly increased, but there is introduced a positive danger of error due to unavoidable slips, which, unless the results are carefully and laboriously checked, is apt to lead to wrong conclusions.

The benefits then of a uniform system are evident, and it is probable that every petrographer has realized this to a greater or less extent. After some experience in the copying and use of rock analyses and full consideration of the subject, as well as discussion of the matter with other petrographers, I would propose that the third method mentioned above, which was, I think, inaugurated by Prof. Pirsson, be generally adopted, this presenting the greatest advantages to the petrographer, with a minimum of disadvantage.

A rock analysis, it must be remembered, is primarily intended for, and almost exclusively used by, petrographers. Therefore for him an arrangement on analytical or strictly chemical lines is neither advantageous nor appropriate. What he needs especially is an arrangement which shall bring the essential chemical features—both the percentage figures and the molecular ratios—prominently and compactly before the eye, so that the general chemical character and the relations of the various constituents may be seen at a glance. It is also of importance that the arrangement be such as to facilitate comparison of one analysis with another.

To the petrographer the eight oxides, SiO_2 , Al_2O_3 , Fe_2O_3 , FeO , MgO , CaO , Na_2O , and K_2O , which in practically every case are present in preponderating amount, are, and must always remain, of prime importance. They are the oxides which, by their relative amounts, determine the chemical character of the rock. H_2O and CO_2 , which are also found sometimes in notable quantity, are chiefly of value as a measure of the freshness of the rock. The other constituents, while of great interest, are present in minute quantities and in general, especially as compared with the main oxides, influence the character of the rock only to a very limited extent, either entering into the composition of accessory and subordinate minerals, or replacing to a very small extent the more important oxides in the essential minerals.

Hillebrand's plea* for their determination is well founded, and it is of course a desideratum that all analyses should be complete as to the rarer constituents. This, however, is not always possible, and it will probably remain true as it is at present, that many otherwise good and serviceable analyses are incomplete in this respect, or only show determinations of a few of the rarer constituents, notably TiO_2 , MnO , P_2O_5 , and Cl .

By putting the eight main oxides together, then, the petrographer is able to see at a glance the character of the rock under study, and the molecular ratios of these oxides, which are the only ones of practical importance, may be written after them and so easily compared *inter se*. Furthermore, whether an analysis is complete or incomplete, these oxides, which are determined in every case, are always in the same relative position, so that the eye finds them without trouble, thus immensely facilitating comparison and study.

The question naturally arises as to the advisability of stating the analytical results, not in oxides, but in terms of the elements, oxygen being given as a separate constituent. This is

* Hillebrand, Bull. 148, U. S. G. S., p. 15, 1897.

done by some petrographers in the reduction of analyses for certain lines of work, though in these cases oxygen is omitted. This last seems to me a somewhat unjustifiable procedure, in view of the fact that oxygen makes up by far the greatest percentage in all the known terrestrial rocks.

It is true that we do not yet know whether the oxides exist as such in the magma or not, though the theories of physical chemistry seem to favor the view that they do. But apart from such considerations, the possible benefits to be obtained are so doubtful and meager that it seems scarcely justifiable at present to make any such sweeping change. Experience has shown the convenience and practicability of the present method for the study of rocks, and as it is sanctioned by long and universal usage, it had better be retained, and the results stated in oxides.

The statement of the analysis then will be divided into two divisions, the main portion, in which are placed the principal oxides, and the subordinate in which occur all the rest.

The order in which the constituents may be put is the next consideration. That which is here proposed is as follows :

SiO₂, Al₂O₃, Fe₂O₃, FeO, MgO, CaO, Na₂O, K₂O, H₂O (ignit.), H₂O (110°), CO₂, TiO₂, ZrO₂, P₂O₅, SO₂, Cl, F, S (FeS₂), Cr₂O₃, NiO, CoO, MnO, BaO, SrO, Li₂O.

As regards the main portion the sequence usually adopted seems eminently proper. This is SiO₂, Al₂O₃, Fe₂O₃, FeO, MgO, CaO, Na₂O and K₂O. In this we start out with the chief acid radical and the constituent present in largest amount, and go through successively lower orders of oxides to the most positive radicals, the alkalis. At the same time they are presented in a way which brings the oxides together in their natural petrographic and mineralogic relations. The sesquioxides are together; ferrous iron follows ferric, MgO is next to FeO, as the two go hand in hand in the ferro-magnesian minerals; CaO is intermediate as is proper, since it is a constituent both of these and of the feldspars, and it is next to Na₂O, as it is associated with it in plagioclase.

H₂O should follow K₂O, as it is a highly important and regularly determined constituent. I follow Hillebrand* in urging the separation of ignition and hygroscopic water, putting the former first, as being an essential ingredient. Next to these should come CO₂, as this, with H₂O, is a measure of the freshness of the specimen, and this character can, therefore, be told at a glance. Together also they constitute the "loss on ignition" so often given, and in this case can be connected by a bracket.

* Hillebrand, Bull. 148, U. S. G. S., p. 29, 1897.

After these it seems well to put the acid radicals which have been determined, following the main principle of the other division; TiO_2 and ZrO_2 should come first, then P_2O_5 and V_2O_5 . SO_3 is put next instead of at the head of the lesser acid radicals, as it, like Cl which follows, is a constituent of the sodalite group. F comes after Cl, and then S or FeS_2 may be put.

Then come the subordinate metallic oxides, in the order R_2O_3 , RO, R_2O . These would include Cr_2O_3 , (Ce_2O_3) , NiO, (CoO) , MnO, (CuO) , BaO, SrO, Li_2O . Finally in rare cases C or N may be put at the last.

While the above order of arrangement is undoubtedly open to criticism in certain respects, yet it seems to the writer, as well as to other petrographers to whom it has been submitted, to be on the whole the most logical and convenient. It may be added that practical experience of it in the collection of a large number of analyses sustains this view.

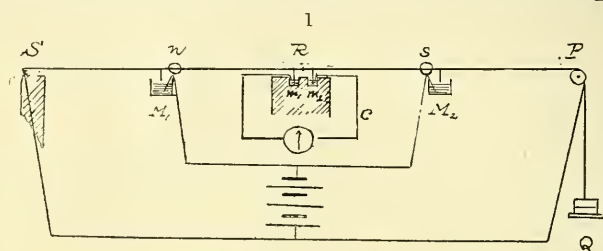
I would also urge the advisability of printing the important molecular ratios with each new analysis, wherever this is possible. The reader will be saved the trouble of calculating them for himself, and will thus immediately have a fuller comprehension of the character of the rock.

In this connection I would suggest, as a practical and convenient method for using analyses, the use of a card catalogue, the cards being printed with the oxides in their proper order. Experience has shown the utility of such a plan, and the time given to its preparation will be more than compensated for by that saved in reference and collation.

ART. X.—*A String Alternator*; by K. HONDA and S. SHIMIZU.

THE transverse vibration of a stretched iron wire was first used by M. Wien* as an interrupter for an induction coil. The vibration was produced by the mutual attraction between a magnet and the wire. Pupin,† however, replaced the driving force by the motion of a wire carrying an electric current in the magnetic field. A similar arrangement was also used by L. Arons.‡

Pupin's interrupter may be modified to serve as an alternator in the following way. As shown in the annexed figure, a



copper wire is horizontally stretched; one of the ends is fixed to a stout support S, and the other, passing over a small metallic pulley P, is attached to a weight Q. The wire is electrically insulated at the middle point R, so that no current can flow through the point but through the circuit C. Close to the middle point and just under it, two mercury cups m_1 and m_2 are placed; two short vertical wires fixed to the horizontal one are dipped into each of them. Again, at quarter distances from S and P, two short wires about 2^{cm} long are vertically fixed to the main wire; two cups M_1 and M_2 are placed under them for mercury contact. One pole of an accumulator is connected to these cups, and the other to S and P; n and s are two poles of a strong magnet (or better an electromagnet) horizontally placed close to the wire; the magnet maintains the wire in constant vibration. Evidently, there is a definite relation between the direction of the current and the poles of the magnet for continuing the vibration. This direction of the current is easily found by trials. A tap given to the wire so as to produce vibrations with a single node at the middle point sets the wire in constant vibration. By this arrangement, the mercury contact is made in turn at M_1 and M_2 , so that the current passing through the portion of the circuit C changes

* M. Wien, *Wied. Ann.*, xlv, 683, 1891.

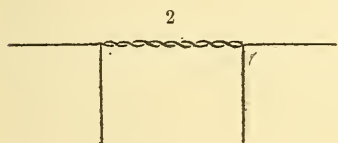
† M. I. Pupin, *this Journal*, III, xlv, 325, 1893.

‡ L. Arons, *Wied. Ann.*, lxvi, 1177, 1898; lxvii, 682, 1899.

its direction once in a complete vibration. Thus, suitably changing the length (by sliding two wedge-shaped blocks) or the tension (by changing the suspended weight) of the string, an alternating current of desired frequency is obtained.

Now, in the above method, only a part of the current employed can be turned to the alternating current; in case where a strong alternating current is required, it is sufficient to use two equal sets of accumulators; two poles of the one set being connected to M_1 and P, and those of the other to M_2 and S, so that the currents through the portion of the circuit C are directed in opposite sense. A sensitive galvanometer, inserted in the circuit, shows whether the mean strength of the alternating current in both directions is equal or not. The small difference in the current-strength can easily be effaced by adjusting the spark gap of the mercury contact.

To cut off the electrical connection at the middle point of the wire, the following method proved to be the most satisfactory. A silk-covered wire of suitable length is cut into two parts. The ends are twisted together like a rope for the length of 1^{cm} and bent at right angles to the wire as shown in the



annexed cut. The twisted portion of the wire is firmly bound together by a silk thread, and the remaining portion of the cover is scraped off. Though different specimens of wire were tried, copper and alumi-

num wires gave the most satisfactory result.

The simplicity of the present arrangement for obtaining an alternating current of definite frequency is one of its chief merits. The actual frequency, if the tension of the string be not small, was found not to differ by more than 1 per cent from the value calculated by the formula

$$n = \frac{1}{l} \sqrt{\frac{T}{\rho}},$$

where l is the length, T the tension and ρ the linear density of the wire. It was not difficult to obtain vibrations ranging from 30 to 1000 per second.

It may not be out of place to remark that in using the present arrangement, the self-induction of the circuit is to be reduced as small as possible, inasmuch as it diminishes the strength of the current. Hence in case the self-induction of the circuit be not small, a large number of accumulators must be used; otherwise the string would not maintain its vibrations. Measuring the mean strength of the current by an ampere-balance, we found but a small fluctuation, if the current does not exceed 1.5 amperes.

Physical Laboratory of the Imperial University, Tokyo, Japan.

AM. JOUR. SCI.—FOURTH SERIES, VOL. X, No. 55.—JULY, 1900.

ART. XI.—*The Action of Light on Magnetism*; by JOSEPH HALL HART.

THE electro-magnetic theory of light assumes the existence of electro-magnetic waves of the required frequency in the ether and of their identity with ordinary light-waves. In consequence of the alternate electric displacements there should be an alternating magnetic field and we might expect the demagnetization of iron by light as a result, if the intensity of the light were gradually diminished. We have, however, little trustworthy evidence of the action of light on magnetism. The contradictory results obtained and the smallness of the effect have prevented the attainment of experimental results sufficient for definite conclusions on the subject. In particular, the action of light on the magnetism of iron bars presents a field in which, if ordinary care is taken, positive results should be obtained.

Several experimenters in the early part of the present century tried to magnetize iron and steel by the action of light. Morichini,* Christie† and Mrs. Somerville‡ obtained positive results from a series of experiments, while P. Riess and L. Moser,§ and John W. Draper|| obtained negative results from precisely similar experiments, and the thing has been generally regarded as impossible. Under ordinary circumstances there can be little doubt that this is the case, but, according to Bidwell,¶ in a paper before the Physical Society in 1889, if a certain initial condition is fulfilled, we can find evidence of the action of radiation upon the magnetism of iron, and the fulfillment or non-fulfillment of this condition explains the diverse results of previous experimenters.

The condition is that the susceptibility of the bar AB to be operated upon shall be greater (or less) for a magnetic force in the direction AB than for an equal one in the direction BA. Iron bars having this property were obtained by the following method: A piece of soft iron rod, which may conveniently be 10 or 12^{cm} long and from 0.5 to 1^{cm} in diameter, is raised to a bright yellow heat and slowly cooled. When cold, it is placed inside a solenoid, through which is passed a battery current of sufficient strength to produce a field of about 350 or 400 C. G. S. units. The iron when removed from the coil is found to be permanently magnetized, and its north pole is marked

* See Christie. † Phil. Trans., 1826, p. 219.

‡ Phil. Trans., 1826, p. 132.

§ Annales de Chimie, xlii, 304.

|| Journ. of Franklin Institute, Feb. 1835.

¶ Proc. Phys. Soc., London, April, 1889, p. 455.

for the sake of distinction with red sealing-wax varnish. The bar is then supported horizontally and in an east-and-west direction behind a small reflecting magnetometer, and over it is slipped a coil, which is shunted with a rheostat, the resistance of which can be gradually increased from 0 to 26 ohms. The coil can be connected by a key with a single battery cell, which is so arranged as to produce a demagnetizing force inside the coil. The resistance of the rheostat is slowly raised, so that more and more current passes through the coil, the battery key being alternately lifted and depressed until the magnetometer indicates that the iron bar as a whole is perfectly demagnetized. The strength of the demagnetizing force required varies according to circumstances: it is generally about one thirtieth or one twenty-fifth of the original magnetizing force.

A piece of iron thus treated possesses certain remarkable properties. The magnetization induced by a force acting in such a direction as to make the marked end a north pole, is greater than that caused by an equal force in an opposite direction. If such a bar be held horizontally east and west and tapped with a mallet, the marked end at once becomes a north pole. Application of heat or instantaneous application of flame causes a similar effect. If subjected to the action of a series of equal, alternating feeble magnetic forces, the marked end generally becomes a north pole, even if the last of the alternating forces may have tended to induce the opposite polarity. These were well-known effects, but in addition, Bidwell found that such a rod was remarkably sensitive to the action of light. When placed behind the magnetometer and illuminated by an oxy-hydrogen lamp about 70^{cm} distant, there occurred an immediate deflection of from 10 to 200 scale divisions, the magnitude of the effect varying in different specimens of iron. The magnetometer mirror was one meter distant from the scale and each division=0.64^{mm}. In these experiments, as the action of the light was continued, the deflection slowly increased. When the light was shut off, the magnetometer instantly went back over a range equal to that of the first sudden deflection, then continued to move slowly in the backward direction toward zero. With a thick rod the slow movement was barely perceptible. He believed the first quick movement to be due to the direct action of radiation, and the subsequent slow movement to the gradually rising temperature of the bar. In several cases, the magnetism was of the opposite kind, and sometimes became north when certain portions of the rod were illuminated and south when the light acted upon other portions. This was probably due to irregular annealing and a consequent local reversal of the

direction of maximum susceptibility: it indicated, probably that the light effect was local, and confined to the illuminated surface. The effect varied directly as the intensity of the illumination; it was strongest for red light and independent of direction of vibration of incident light; and blackening the bar made the action much slower.

I first repeated and verified Bidwell's experiments. I procured eight pieces of the softest iron obtainable, of the following dimensions: each was 20^{cm} long; four were cylindrical rods, with diameters 13^{mm}, 9·5^{mm}, 6·5^{mm}, 3·5^{mm} respectively: they were designated rods A, B, C and D; four were of rectangular cross-section, 12×3·5^{mm}, 12×2·5^{mm}, 12×1·5^{mm}, 12×0·75^{mm} respectively: they were designated, rods E, F, G and H. They were treated in the same manner as in Bidwell's experiments; they were carefully annealed, being heated in iron filings almost to the point of fusion and then cooled in ashes for about 25 hours. They were first strongly magnetized and then carefully and totally demagnetized. As a result of numerous trials and careful manipulation, I was able to get distinct effects from a gas jet at a distance of 6 inches. The following is a characteristic set of experiments: the light was perpendicular to the axis of the bar and only 3^{cm} of the north end of each rod was exposed. The deflection represents increase of magnetism as shown with magnetometer.

Rod.	Area exposed to light.	Cross-section of rod.	Deflection.
A	6·15 ^{sq} cm	1·33 ^{sq} cm	22·5
B	4·48	0·71	15·0
C	3·06	0·33	7·1
D	1·65	0·10	1·5
E	3·60	0·42	10·0
F	3·60	0·30	7·5
G	3·60	0·18	4·5
H	3·60	0·09	1·5

The effect was proportional to the area exposed, but not directly. It depended also on the intensity of the illumination. The action of the light does not appear to be here confined to the illuminated surface, as the effect is approximately the same, no matter which 3^{cm} of length of the bar is illuminated, and hence distance of illuminated surface from magnetometer has little or no effect. This is contrary to Bidwell's results and is probably due to the very regular annealing. The heating effect was also present. Light of long wave-length had greater effect. I did not attempt Bidwell's experiments in regards to polarized light and its action on the magnetism of the iron, since no positive results were obtained by him.

I then extended these experiments of Bidwell in an attempt to learn more in regard to the nature of this action of light on magnetism. Every cause of molecular disturbance favors the magnetization of a bar subjected to a magnetizing force, and also favors its demagnetization after it has been withdrawn from the field. Vibration has a specially marked effect upon iron. Ewing* has shown that if a bar of this metal be kept from the slightest vibration one can obtain residual magnetizations much greater than those shown in steel bars, but the least vibration causes the acquired magnetism to vanish almost completely. If the action of the light is of the same nature as that of a mechanical shock on a magnet and not dependent on the special structure of the magnetic field in Bidwell's experiment, we ought to obtain effects from light in experiments like those of Ewing, if they are sufficiently delicate. I have been able to get very good results from a series of experiments with this object in view. The bars were magnetized in a magnetic field of about 400 C. G. S. units and the current was steadily and rapidly shut off by means of a continuous current rheostat. The slightest vibration caused a sudden fall in the residual magnetism. Extreme care was necessary. Only a small portion of the bar was exposed to the light, as it was found that the magnetizing coil could not be removed without causing the sudden fall in the magnetic strength of the bar. Electric disturbances due to the surface-road and electric light wires had the same effect, and it was necessary to work during the quiet portions of the day. The following is the average result of a number of experiments. Only 3^{cm} of the bar was exposed, that is, there was the same surface exposed as in previous experiments, and the light was that of an ordinary gas jet at a distance of six inches; other conditions were the same. The direction of the deflection indicated in every case a decrease of magnetism.

Rod.	Area exposed to light.	Deflection.
A	6·15 sq cm	44·0
B	4·48	26·1
C	3·06	12·5
D	1·65	6·4
E	3·60	13·1
F	3·60	8·4
G	3·60	4·3
H	3·60	0·9

These results differ from those of Bidwell in the fact that the deflections here are permanent, even after light has been

* Proc. Roy. Soc., 1890, p. 342.

removed. In my own repetition of Bidwell's experiment, the magnetometer image did not always return to the same spot after the light had been flashed upon the iron and removed. There was in general a partial return, the law or cause of which I was unable to discover. I then attempted a similar experiment obviously suggested by the previous one, namely, that of increasing by light the magnetic strength of a bar which was in a steady field. I here also had great difficulty on account of external vibrations and electrical disturbances. I steadily increased the current up to different strengths and then flashed the light on the exposed portion of the bar and noted increase with the magnetometer. The first series here given is when the field is of about 10–15 C. G. S. units and the second, of about 350 C. G. S. units, 3^{cm} was exposed and the light was a gas jet at a distance of 6 inches.

Rod.	Area exposed.	Deflection—I.	Deflection—II.
A	6.15 ^{sq cm}	none	none
B	4.48	none	none
C	3.06	42	none
D	1.65	6	10
E	3.60	none	none
F	3.60	36	none
G	3.60	23	15
H	3.60	7	14

It is a noteworthy fact that it is only the larger bars from which no deflection was obtained. The conditions were precisely the same for the series and therefore for the larger bars there is either an increased sensitiveness to vibrations coming from without, or the action is dependent on the mass in some manner not understood. It is probably only a small portion of the iron which is so extremely sensitive as to be influenced by light. A slight but sudden increase of the field up to the maximum strength used in this experiment caused a rise in the magnetism of one bar amounting to 125 divisions, while a slight and sudden fall in the strength would also cause a slight increase, when the sensitive state was obtained.

All these results can be effectively accounted for by the assumption of a regular closed magnetic circuit (in iron bars carefully annealed, magnetized, and immediately demagnetized as above described) in which the lines of force return through a thin layer at the surface. The circuit is wholly or partially closed in every case except that of complete saturation, and the thickness of this surface layer depends on the magnetic state of the bar. In Bidwell's experiments, we may suppose the larger portion of the lines of force to return through the surface layer, which is here of considerable thickness. Here

the action of the light is to break or destroy a number of those lines of force. On removal of the acting force, or source of energy, there exists a tendency to return to the original form, owing to action of magnetic lines in the immediate vicinity. In the next two experiments, the instability is of the opposite kind and there is no tendency to return to the original form. In both cases, where the iron was demagnetized by light, and where the iron was magnetized by light, the action is similar to that of a blow under the same circumstances.

While this satisfactorily explains the mechanism of the results of the action of light on a magnetized iron bar, that of the direct action of the light is in doubt. Analogous effects can be obtained from purely mechanical actions and I have obtained them also from the action of electric waves. Whether these quantitatively remarkable results are due to the magnetic or purely mechanical action of light, is unknown. I determined therefore to make some experiment in regard to the action of polarized light on the magnetism of iron. Bidwell has shown that polarized light in the ordinary experiments has the same effect as ordinary light. In order to make a more sensitive action, I obtained from Professor A. W. Wright a number of films of iron deposited on thin glass plates by electrical discharge in exhausted tubes. The method of production and general properties of these films are given in this Journal for January and September, 1877. The films of iron had a brilliant luster and a high degree of transparency and polarized the light transmitted by them powerfully. The films were approximately 2^{cm} long and 0.5^{cm} in width and the thickness was very small, about a quarter of the wave length of red light. The amount of iron in one of these magnets was extremely small; the magnetic moment of the largest one was so small that when suspended by a single silk fiber, the finest obtainable, the magnetic action of the earth's field was scarcely noticeable. But the magnet possesses relatively strong permanent magnetism for the amount of iron involved, in fact, the iron appears permanently saturated. A remarkable feature of this magnetism is its strong resistance to a reversing force. If we assume the reversal of any ordinary magnet by an opposite magnetizing force to be due to the cumulative action resulting when a few unstable molecular magnets are reversed and these produce other unstable groups which are in their turn reversed and so on, then this phenomenon can be explained by the almost total absence of unstable groupings of molecular magnets in the mass of the iron. In a magnetic field of 400 C. G. S. units it required from three to five seconds to reverse it, and in a field of 25 C. G. S. units it was im-

possible to reverse it without tapping. Under these circumstances it was found to be impossible to demagnetize this magnet totally by any of the ordinary methods. The magnetometer used in the previous experiments, while extremely delicate, was not of the slightest use in these, and therefore it was necessary to measure the time of vibration of the magnet itself, with and without light polarized in different directions falling on its surface. The system was almost aperiodic, the magnet coming completely to rest in five half vibrations. No difference in time amounting to 0.2 seconds was observable out of a total of 40 seconds approximately. Owing to the unexpected magnetic behavior of the iron, it was not thought that the light polarized so as to have its vibrations perpendicular to the magnetic axis of the magnet would demagnetize it completely, but that a slight fall from completely saturated state might occur, and so the experiments were continued. The magnet was kept suspended in a field of 600 to 800 C. G. S. units for some time; the field was gradually diminished to 200 C. G. S. units and then torsion for given deflection was measured with and without the application of polarized light. Five complete revolutions of the torsion head caused a deflection of 55° of the magnet without the application of light, while 4.8 revolutions caused the same deflection when light so polarized that the vibrations were perpendicular to the magnetic axis was applied to the surface. There is a possible error of 20° for each of these readings. The same experiment was repeated with light parallel to the magnetic axis of the magnet and a difference in the torsion amounting to 40° was observed. The quantitative difference in the two experiments is very small, but the results of a great number of the observations were taken and were fairly consistent. This apparently shows a difference depending on direction of the plane of polarized light analogous to the demagnetizing effect of an alternating current.

Owing to the anomalous behavior of iron films when placed in a magnetic field and their strong resistance to a reversal of the magnetization, it was thought advantageous to attempt the experiment with a nickel film, since the resistance to demagnetization would probably be much less than for iron films. Although the film which was used contained nearly fifty times the amount of metal of the heaviest of the iron films and was exposed in a field of nearly 2000 C. G. S. units, there existed no appreciable magnetic moment in the nickel and the experiment was given up.

From the results of this investigation the following conclusions may be drawn:

1. An effect of light on magnetism is, in general, easily obtained and depends only on a certain condition of sensibility artificially produced.

2. This condition is, apparently, that a peculiar and unstable distribution of the magnetic lines of force shall exist in the magnet.

3. The action is, in general, a purely mechanical one, similar to the effect of a blow.

4. Evidence of the magnetic action of polarized light has probably been obtained by the use of iron films on glass.

In conclusion, I wish to thank Prof. A. W. Wright for his assistance throughout the investigation. He suggested the subject and general plan of research, and afforded me ample facilities for performing the experiments.

Sloane Physical Laboratory, June 15, 1900.

SCIENTIFIC INTELLIGENCE.

I. CHEMISTRY AND PHYSICS.

1. *On Radio-active Barium.*—In a preliminary notice BÉLA V. LENGYEL calls attention to the fact that the bodies thus far known which emit Becquerel's rays are of five different kinds, viz., the compounds of uranium, of thorium, polonium, radium, and Debierne's body which is analogous to titanium. Uranium and thorium are well-defined chemical substances, while the three others are merely hypothetical elements. Radium is the best known of these, but from the statements made in regard to it the author believes that it can hardly be regarded as an existing element. In favor of its existence the two principal facts from a chemical point of view are the higher atomic weight found by Madam Curie for the radio-active barium, and the single foreign line found by Demarcey in the spectrum of such material. In regard to the first point, the author thinks that a body which is present in sufficient amount to raise the atomic weight eight units above barium should not fail to reveal itself in the course of various chemical decompositions. He believes, moreover, that the single spectrum-line noticed by Demarcey does not point to the presence of an element closely analogous to barium, for, as is well known, the spectra of barium, strontium, and calcium consist of many sharp lines and less clearly defined bands, similarly grouped in the three cases, and an element almost identical with barium should possess an analogous spectrum. He states also that it should be remembered that these hypothetical elements are always found adhering to other well-known chemical elements. Curie found polonium with bismuth, Giesel found it with lead; radium is attached to barium, Debierne's element goes with titanium. All these radio-active substances come from pitchblende, being separated from it by analytical processes. It is difficult to assume that elements exist which differ from others that are well known in nothing except their radio-activity.

Such considerations as the above led the author to investigate experimentally the question whether radio-active bodies contain new elements or not. He ignited uranium nitrate with two or three per cent of barium nitrate and finally fused the resulting oxides in the electric arc. The mass was dissolved in nitric acid, the solution was evaporated, whereupon a large part of the barium separated as nitrate, then from the decanted liquid a precipitate of barium sulphate was produced which showed radio-activity. The author believes that he has prepared radium synthetically, but does not consider that these preliminary experiments are sufficient to decide with certainty whether this is a definite chemical element or not.

There appears to be a weak point in this supposed synthesis of radium from the possibility that the uranium nitrate used may

have contained the "element" in question, although v. Lengyel does not mention such a possibility. At all events, the result of the experiment makes it seem possible that the radio-activity of uranium may be due to the same cause as that of radium, and perhaps also of polonium and the others.—*Berichte deutsch. chem. Gesellsch.*, xxxiii, 1237.

H. L. W.

2. *Mercuric Antimonide, and Stibonium Compounds.*—PARTHIEL and MANNHEIM have prepared mercuric antimonide, Sb_2Hg_3 , by passing hydrogen antimonide over finely powdered mercuric chloride which was mixed with sand and fragments of glass. The reaction took place slowly, the white mercuric chloride becoming black. This mercuric compound is interesting because like potassium antimonide it gives stibonium compounds when treated with alkyl iodides. The authors have studied its behavior with ethyl and propyl iodides, heated under pressure. They thus obtained the double salts $Sb(C_2H_5)_4I.HgI_2$ and $Sb(C_3H_7)_4I.HgI_2$. By treatment with silver oxide these bodies yielded tetraethyl stibonium hydroxide, $Sb(C_2H_5)_4OH$, and the corresponding propyl compound, from which various salts were prepared.—*Chem. Centralbl.*, 1900, 1, 1091.

H. L. W.

3. *The Atomic Weight of Iron.*—This important constant has not been revised since about fifty years ago; meanwhile analytical methods have been much improved, and the sources of error involved in atomic weight determinations have been more carefully studied. In view of these facts, RICHARDS and BAXTER have made a preliminary series of seven determinations by reducing ferric oxide to metallic iron. Their results show that the previously accepted atomic weight, 56, is somewhat too high and that in all probability the value is very near 55.88.—*Zeitschr. anorg. Chem.*, xxiii, 255.

H. L. W.

4. *The Physical Properties of Cæsium.*—ECKARDT and GRAEFE have made an elaborate examination of the physical properties of metallic cæsium which was prepared by heating pure cæsium carbonate with metallic magnesium in an iron tube in a current of dry hydrogen. The summary of their results is as follows:

Sp. gr. of liquid Cs at 40°	1.827
Sp. gr. of liquid Cs at 27°	1.836
Sp. gr. of solid Cs at 26°	1.886
Solidifying point of Cs	26.37°
Conductivity at 27° (Ag=100)	3.63
Specific heat04817
Atomic heat	6.406
Heat of fusion (for 1 g.)	3.73 cal.
Contraction upon solidification	2.627 vol. per cent.
Coefficient of expansion0003948

The specific gravity found by the authors corresponds with the value 1.88 previously determined by Setterberg, but it does not agree with the value 2.40003 recently given by Mencke. The latter is probably incorrect, not only in view of the agreement of

the other investigators, but because it gives an atomic volume for caesium, which is not in harmony with the periodic system.—*Zeitschr. anorg. Chem.*, xxiii, 378. H. L. W.

5. *On Compounds of Telluric Acid with Iodates.*—WEINLAND and PRAUSE have prepared the salt $I_2O_5 \cdot 2TeO_3 \cdot K_2O \cdot 6H_2O$, as well as an exactly corresponding rubidium compound, and an ammonium salt which varies from the foregoing in containing $8H_2O$. The potassium salt of a different acid, $I_2O_5 \cdot TeO_3 \cdot K_2O \cdot 3H_2O$ was also obtained. These results are interesting inasmuch as they show that TeO_3 forms complex acids with iodates similarly to the oxides SO_3 , CrO_3 , MoO_3 , and WO_3 .—*Berichte deutsch. chem. Gesellsch.*, xxxiii, 1015. H. L. W.

6. *A Simple Method of Decomposing Chromite.*—For getting this very refractory mineral into a soluble condition FIEBER recommends heating .5 grams of the finest possible material with six parts of sodium-potassium carbonate for ten minutes in a platinum crucible, then after cooling adding six parts of borax and fusing at a gradually increasing heat, finally over the blast-lamp for three-quarters of an hour. This generally effects the decomposition. If it fails, it is only necessary to add another portion of sodium-potassium carbonate and fuse again.—*Chem. Zeitung*, xxiv, 333. H. L. W.

7. *The Theory of Electrolytic Dissociation and Some of its Applications*; by HARRY C. JONES, 8vo, pp. xii-289, New York, 1900 (The Macmillan Company).—The title of this valuable addition to chemical literature hardly gives a proper idea of the scope of the work. The first chapter, 68 pages, is devoted to an exposition of "The Earlier Physical Chemistry," which serves as an excellent introduction to the main subject of the book. The four chapters are designed to answer the questions: What was physical chemistry before the theory of electrolytic dissociation arose? How did the theory arise? Is it true? What is its scientific use? The author has given very satisfactory answers to these questions, and the book will undoubtedly be useful to those who wish to gain a clear notion of some of the chief points of modern physical chemistry. H. L. W.

8. *Optical Activity and Chemical Composition*; by Dr. H. LANDOLT, translated by JOHN McCRAE, Ph.D., 12mo, pp. xi-158, London, 1899 (Whittaker & Co., from The Macmillan Company, New York).—This work forms a chapter in Graham-Otto's "Lehrbuch der Chemie." The topics dealt with are the general principles of optical activity, the connection between the rotary power and the chemical composition of carbon compounds, and the connection between the degree of rotation and chemical constitution. H. L. W.

9. *Variation of the Electric Intensity and Conductivity along the Electric Discharge in Rarefied Gases.*—Skinner has shown that very near the anode the electric intensity is very small or zero when the positive column is not striated. He also mentions that with a striated positive column the potential-difference

between the anode and an exploring wire near it, was a minimum when the exploring wire was at a short distance from the anode, so that the apparent electric intensity near the anode was negative. These results are confirmed by HAROLD A. WILSON, who measured the differences of potential between two exploring wires kept at a fixed distance apart in the discharge, which could be brought into any desired portion of the discharge by moving the electrodes between which the discharge took place. This discharge was produced by a storage battery of 600 cells. The gases employed were air, nitrogen, and hydrogen. The positive drop close to the anode was very apparent from the experiments. Very close to the positive electrode there is apparently a very intense ionization. In the negative glow and Faraday dark space there is also ionization in excess. In every kind of electric discharge ionization appears to occur on the metallic electrode when it is red hot and most easily there. The negative ions formed at the surface of the cathode constitute the cathode rays, and produce the ionization which the intensity curves show occurs in the negative glow.—*Phil. Mag.*, June, 1900, pp. 505-516. J. T.

10. *Mechanical possibilities of flight.*—Lord RAYLEIGH in the Wilde lecture published in the Manchester Memoirs 1899, No. 5, discusses this proposition and shows that in order for a man to support himself by a vertical screw by working at the power an average man can maintain for eight hours a day, he would require a screw ninety meters in diameter; and in this estimate no account is taken of the weight of the mechanism or of frictional losses.—*Nature*, May 31, 1900. J. T.

11. *Magnetic Screening.*—The great extension of electric circuits for lighting and for power makes it impossible in most cases to use the ordinary form of mirror galvanometer. H. DU BOIS and A. P. WILLS discuss the subject of the protection of galvanometers from outside changes in the magnetic field produced by the commercial employments of strong currents. The cases of iron cylindrical shells and of spherical shells is treated and the results are used by H. Du Bois and H. Rubens in the construction of a protected or armored galvanometer (Panzergalvanometer). Cast-steel cylinders having a radius ratio of air space 1.5 to 1.6 and spherical cast-steel shells 1.3 to 1.4 give in general this order of protection: one shell, 10; two shells, 100; three shells, 1000.—*Ann. der Physik*, pp. 78-95. J. T.

12. *The Sun's Corona.*—Mathias Cantor in the *Ann. der Physik* for March, 1900, believes from his experiments that a rarefied gas through which an electrified discharge is passing shows no perceptible absorption corresponding to its emission spectrum, and Professor Fitzgerald in *Nature*, May 3, 1900, remarks that this fact confirms the suggestion that the sun's corona is an electrical discharge around the sun since the bright spectrum line of the corona is not represented by a dark line in the solar spectrum. E. PRINGSHEIM criticises severely the results of Cantor and believes that he has not taken suitable precautions in his experiments.

He points out that the discharge through the gases was an intermittent one and the gases were only a portion of the time in that condition when, according to Kirchhoff's laws, an absorption was to be expected. For the remaining portion of the time the light of the arc lamp employed could go through the gas completely unabsorbed. Furthermore since the spectral lines in a Geissler tube represent homogeneous light, an absorption could only be expected, according to Kirchhoff's laws, in a very narrow spectral region. In order to show it, a spectrum of a large dispersion must be employed and a narrow slit, in order that no light of the neighboring portion of the arc spectrum can intrude upon the homogeneous portion under examination. These precautions were not taken by Cantor.—*Ann. der Physik.*, No. 5, 1800, p. 199–200.

J. T.

13. *On some Properties of Light-struck Photographic Plates*; by FRANCIS E. NIPHER. (Abstract by the author.)—A paper with the above title recently published* by the Academy of Science of St. Louis, deals among other things with some features which seem to be new in photography.

In a general way it is probably true that any plate, upon which a camera impression has been made, may be developed either as a positive, or as a negative. Suppose the object to be a white design on a black ground. Let the camera be replaced by a printing frame, containing the plate and an opaque stencil with some design punched through it. Place this plate at any distance from a known source of light. Give any exposure to the plate. If the plate be now developed at a considerable distance from the lamp, a negative image will appear. The lamp should be vertically over the developing bath. If the bath in a similar exposure is taken nearer to the lamp, the negative will be a poorer one. At a still nearer distance, nothing will develop. In the parlance of the photographer, the plate will fog. This is called a zero plate. If the plate be developed at a point still nearer the lamp, the picture will appear as a positive. The conditions of zero plate as a function of exposure is now being studied. A given exposure is a time integral of the effect due to any given illumination, at the point of exposure.

The paper referred to gives half-tone reproductions of photographs thus developed in the light as positives. In one an attempt was made to "fog" a plate into a zero condition before it was put into the camera for exposure to a street scene. The plate was fogged by X-rays for two hours, and was exposed in the camera for ten minutes. It was developed in a hydrochinone bath, within one foot of a sixteen-candle lamp. The result is a superb positive of the street scene in which the moving objects leave no trace. The plate was an instantaneous isochromatic plate by Cramer.

Another picture reproduced was a picture on the same kind of plate, obtained by an exposure of one minute in the camera. The

* No. 6, vol. x, Transactions.

street was strongly lighted, and the plate was converted into a zero condition, and the positive was then produced, during the single minute and while in the camera.

It is also found that the camera exposure may be shortened, by fogging the plate in open lamp-light before it is put into the camera. If the plate is to be developed within 20 cm. from a sixteen-candle lamp, these fast plates will be put into a zero condition by holding them for 90 seconds at a distance of one meter from the lamp. When exposed in the camera such a plate yields a perfect positive picture. If traces of fogging appear, the plate should be moved nearer the lamp. The precise conditions which yield the best results have not yet been determined, nor is it yet known how short the exposure may be made. Some of the plates turn out badly until experience has been gained.

It is evident that if a similar change can be produced in the operation of printing, so that a positive print may be obtained from a positive plate, the dark-room may perhaps be dispensed with for purposes of developing.

In some cases the most sensitive plates have been exposed for from three to four hours to brilliantly lighted street scenes, with the diaphragm fully open. There is not the least trouble in developing such plates, as positives, in a strongly lighted room. In general, the greater the exposure, the darker the developing room must be in order to get a zero or a negative result.

Large pin-hole images of the sun may be developed in this way, and there appear to be many ways in which the process may yield good results. One of the incidental features of value is that a little of what photographers most fear does not interfere with the results.

II. GEOLOGY AND MINERALOGY.

1. *A Monograph of Christmas Island (Indian Ocean): Physical Features and Geology*; by CHARLES W. ANDREWS. With descriptions of the Fauna and Flora by numerous contributors. Pages x, 337; plates xxii. London, 1900 (British Museum of Natural History).—The unique character of Christmas Island, in its position, history, and life, gives peculiar interest to this account of the results obtained from the ten months' vigorous explorations made by Mr. Andrews, of the Geological Department of the British Museum.

Christmas Island has an area of 43 square miles, and rises in places to a height of 1000 feet; it is covered with a dense tropical vegetation. It is situated in the eastern part of the Indian Ocean, 190 miles to the south of Java, 900 miles northwest of the coast of Australia, and 550 miles east of the atolls of Cocos and North Keeling. The submarine slopes about it are so steep that a depth of 1000 fathoms is found within two or three miles of the coast, while to the north, a depth of 3200 fathoms was found (Maclear Deep), and to the south and southwest, of 3000 fathoms

(Wharton Deep). The island is described as forming the summit of a submarine peak, the base of which rises from a low saddle which separates the two abysses named, and on the western end of which the Cocos-Keeling Islands are situated. Its peculiarly isolated position, hence, is most striking. Its history is also unique, since, although known to navigators since the middle of the 17th century, no one seems to have penetrated into the interior until 1887, and, as remarked by Dr. Murray, down to a few years ago it was probably the only existing tropical island of any large extent that had never been inhabited by man, savage or civilized. Its animal and vegetable life, therefore, are thus far almost unchanged by the conditions introduced by human life.

Geologically, the island consists largely of elevated Tertiary limestones with extensive series of eruptives; briefly, it may be considered as an ancient atoll raised to a considerable height above the level of the sea. The "central nucleus" is made up of compact yellow limestone in places very hard and showing no traces of bedding or jointing. This is referred to the Eocene (or Oligocene) and is accompanied by basalts and trachytes both beneath and between the beds. The total thickness of these older Tertiary and accompanying volcanic rocks is estimated to be 600 feet. Forming the mass of the island is the Miocene Orbitoidal limestone, separated from the older rocks by basalts and basic tuffs. The higher elevations are dolomitic limestones containing 34 to 41 p. c. of magnesium carbonate; these show traces of coral structure and imperfect remains of Foraminifera. Thick beds of phosphate of lime, in part limestone beds altered by overlying guano, in part phosphatized volcanic tuffs, occur on some of the elevated points and have proved to be of economic value. The Tertiary limestone, especially the Miocene, forms abrupt vertical cliffs, sometimes 250 feet in height, along a large part of the coast line. A series of terraces is also noted around the shore, and outlying the whole is the fringing coral reef. The author remarks upon the remarkable development of elevated Tertiary rocks and the difficulty in explaining their deposit over an area so isolated. He adds that the great thickness of reef limestone, required by the Darwinian theory of atoll formation, is not found, and although there may be some evidence that subsidence did occur in the history of the island, it is clear that it was not for any long period nor of any great extent. It is interesting to recall in this connection the similar observations recently made by Agassiz on the elevated Tertiary limestones of the Fijis and other islands of the Pacific.

The life of the island is fully described from the collections made and it is shown to be to a remarkable extent endemic. Thus of the 319 species of animals recorded, about 45 per cent are described as peculiar to it, although this percentage may perhaps be reduced when the fauna of Java and other neighboring islands is more minutely known. For example, of the mammals, all are peculiar species except one; of six reptiles, four are pecu-

liar; of fourteen species of land shells, six are peculiar, etc. The author gives an interesting summary of the conditions, as to winds and ocean currents, which have been instrumental in the introduction of the fauna and flora.

2. *On the Lower Silurian (Trenton) Fauna of Baffin Land*; by CHARLES SCHUCHERT. Proc. U. S. Nat. Mus., vol. xxii, pp. 143-177 (with plates xii-xiv), 1900.—The author reports upon several collections placed in his hands for study. The following excellent summary is quoted from page 175 of the report:

“The only Lower Silurian horizons known in northeastern Arctic America are of Trenton and Utica age. The latter zone appears only on the north shore of Frobisher Bay, but the Trenton is found in various places from the north shore of Hudson Strait to latitude 81° north. The Lower Silurian is thickest on Akpatok Island, where it is from 400 to 500 feet in depth. Dr. Bell, however, estimates the entire thickness of these strata in this region to be not less than 900 feet.

“In Baffin Land, and apparently elsewhere in Arctic America, the Lower Silurian strata rests unconformably on old crystalline rocks. To the north of Baffin Land, the former are overlain by beds of Niagara or Wenlock age.

“The Trenton faunas, occurring in various places around the insular Archæan nucleus of North America, have much in common, and this indicates that the conditions at that time were very similar, while the sea was in communication throughout. As yet, however, the distribution of the strata, together with their faunas, are well known only to the south and southeast of the Archæan nucleus, yet that of the west (Manitoba) and of the northeast (Baffin Land) show direct communication.

“The Baffin Land fauna had an early introduction of Upper Silurian genera in the corals *Halysites*, *Lyellia*, and *Plasmopora*. In Manitoba similar conditions occur in the presence of *Halysites*, *Favosites*, and *Diphyphyllum*. Other Upper Silurian types do not appear to be present.

“The Trenton fauna of Silliman’s Fossil Mount, at the head of Frobisher Bay, has seventy-two species, of which twenty-eight are restricted to it. This fauna shows an intimate relationship with that of the Galena of Minnesota, Iowa and Wisconsin. Fifty-seven per cent of the species of Baffin Land also occur in the Galena of the regions just mentioned.

“The Trenton fauna of Baffin Land shows that the corals, brachiopods, gastropods, and trilobites have wide distribution, and are therefore less sensitive to differing habitats apt to occur in widely separated regions. On the other hand, the cephalopods, and particularly the pelecypods, indicate a shorter geographical range. The almost complete absence of Bryozoa in the Baffin Land Trenton contrasts strongly with the great development of these animals in Minnesota and elsewhere in the United States.”

3. *A Preliminary Report on the Geology of Louisiana*; by GILBERT D. HARRIS and A. C. VEATCH, Part V, Geology and

Agriculture, plates 1-62, figures 1-7, map, pp. 4-354, 1899. (Geological Survey of Louisiana, Wm. C. Stubbs, Director.) The author recognizes in Section II the following formations:

Cretaceous series—Ripley stage.

Eocene series—Midway, Lignitic, Lower Claiborne, and Jackson stages.

Oligocene—Vicksburg and Grand Gulf.

Lafayette—

Quaternary—(including the Columbia of McGee and Biloxi sands of Johnson) Basal gravel, Port Hudson, Loess and Yellow Loam, Alluvium and recent coastal formations.

In Section III—special reports—nine miscellaneous reports are given by various authors, the last on “Wood-destroying Fungi,” by Professor Atkinson.

4. *Geology of Panama*.—In a study of the geology of the isthmus of Panama recently published by Messrs. BERTRAND and ZURCHER,* Bertrand draws the following important conclusions regarding the probability of earthquake shocks in the Panama region:

“There are no volcanoes near Panama, all eruptions having ceased since the Miocene; this is the first and most important point of all.

Since the earthquake of 1621, which is in reality disputed, there have been in the region only very feeble shocks, a part of which were due to the echo of far distant earthquakes.

The depression made use of by the Panama canal project is not a transverse fracture.

The sinking of the Pacific coast, and especially the subsidence of the bay of Panama, of which there are numerous indications, are not phenomena now going on but finished, so far at least as regards the present geological epoch. There is there no special reason for crustal movement.

Finally, the plan of the lines of folding and the distribution of volcanic and seismic activity following these lines, shows that Panama is situated in a sort of dead angle, in a tranquil zone, at an equal distance north and south from the lines of disturbance.

Thus all considerations, whether statistic, volcanic or seismic, lead to the same conclusion, that Panama is the most stable and least menaced region of Central America.”

L. V. P.

5. *Enrichment of Mineral of Veins by later Metallic Sulphides*; by WALTER HARVEY WEED. Bulletin of the Geological Society of America, vol. II, pp. 179-206.—Secondary sulphide enrichment is certainly one of the most interesting subjects of economic geology, and, as noted by Mr. Weed in the opening pages of his article, is one upon which surprisingly little has been written, considering its scientific and practical importance. Among prominent writers on ore deposits, De Launay, Posepny, Emmons and

* *Études géologiques sur l'isthme Panama and Phénomènes volcaniques et les tremblements de terre de l'Amérique Centrale.* Paris, 1900.

Kemp have mentioned the subject, but the present paper is the first to treat it separately and comprehensively.

After some introductory remarks and definitions a discussion follows of the chemical reactions which take place during the leaching of the mineral contents of the gossan zone; then the chemical processes concerned with mineral deposition in the enrichment zone and of the alterations of various vein minerals, are considered. A more extended description and discussion of the manner of occurrence of secondary enrichments is then entered upon, dealing with the deposits of copper, silver and zinc in prominent localities, here and abroad, which have been carefully studied.

To the material taken from various outside sources Mr. Weed has added much valuable information, gained as a result of his own extensive observations and studies of the ore deposits of the West.

The conclusions reached are worthy of careful notice and are quoted here in full. Conclusions: "From what has been shown it is concluded that later enrichment of mineral veins is as important as the formation of the veins themselves, particularly from the economic standpoint. The enrichment is usually due to downward-moving surface waters, leaching the upper part of the vein and precipitating copper, silver, et cetera, by reaction with the unaltered ore below. In many cases the enrichment proceeds along barren fractures and makes bonanzas. In others it forms films, pay streaks, or ore shoots in the body of leaner original ore. In still other cases the leaching, transportation, and redeposition are performed by deep-seated uprising waters acting upon the vein.

As a consequence of this, veins do not increase in richness in depths below the zone of enrichment.

The practical bearing of the phenomena described and the deduction drawn from them will, I think, be apparent to every mining engineer and geologist. If my views be correct, the future of many ore deposits is to be judged in the light of these facts, and the value of the mine must not be based on the presumption that the ore will continue in unabated richness in depth."

A forthcoming paper, which is promised on the secondary sulphide enrichments of Butte, Mont., will be awaited with no little interest.

C. H. W.

6. *Mineralogical Notes*.—Number 7 (vol. i) of the Geological Series of Publications by the Field Columbian Museum (pp. 221–240), by Dr. O. C. FARRINGTON, contains an account of several new mineral occurrences and also a description of the fine calcite crystals from Joplin, Missouri. The rare species *inesite* is noted from a mine near Villa Corona, Durango, Mexico. It occurs in tufts of radiating crystals of flesh-red color. These correspond in angle with the results of Scheibe and show the new forms *k* (11·0·12) and *s* (946). The following analysis was made:

	SiO ₂	MnO	FeO	CaO	MgO	H ₂ O (cryst.)	H ₂ O (const.)
G.=2.965	44.89	36.53	2.48	8.24	tr.	5.99	2.21=100.34

From this, the composition $H_2(Mn, Ca)_6Si_6O_{16} + 3H_2O$ is deduced, which varies somewhat widely from the results hitherto obtained.

Caledonite in distinct crystals is noted from the Stevenson-Bennett mine, Organ Mountains, New Mexico; also *gay-lussite* from Sweet Water Valley, Wyoming. The use of *dolomite* as money by the Indians in Lake County, California, is noted. The tokens are shaped in cylindrical forms and burned, which brings out reddish streaks in the oxidation of the iron, and then polished and perforated. In this form they are highly valued by the natives.

The various types of the beautiful Joplin crystals of calcite from Joplin are described and well figured. The observations form an important addition to the literature of the species.

7. *Tabellen zur Bestimmung der Mineralien mittels äusserer Kennzeichen*. Herausgegeben von Dr. ALBIN WEISBACH. Fünfte Auflage. Pp. 106. Leipzig, 1900 (Arthur Felix).—The Determinative Tables of Prof. Weisbach, first issued in 1866, are so well known by those interested in mineralogy that their value hardly needs to be remarked upon here. Based upon external characters alone, they give the student a simple means of determining species, applicable in ordinary cases, and teach him to use his powers of observation with accuracy and discrimination.

8. *Repertorium der Mineralogischen und Krystallographischen Literatur vom Anfang d. J. 1891 bis Anfang d. J. 1897, und Generalregister der Zeitschrift für Krystallographie und Mineralogie*, Band xxi-xxx. Herausgegeben und bearbeitet von E. WEINSCHENK und F. GRÜNLING. II. Theil (Generalregister von F. Grünling). Pp. 394. Leipzig, 1900 (Wilhelm Engelmann).—The First Part of this Index has already been noticed in the present volume of this Journal (p. 229). Part II, now issued, completes the work, which is invaluable for every one interested in the progress of Mineralogy. The fact that this general index extends to nearly four hundred pages shows strikingly the immense amount of material in the *Zeitschrift* to which references are made.

III. BOTANY AND ZOOLOGY.

1. *Les Carex de l'Asie orientale*; by A. FRANCHET, Nouv. archiv. du muséum d'hist. nat., Ser. 3, vols. 8-10, Paris, 1896-1898.—The present comprehensive work deals with the vegetation of *Carex* in eastern Asia, of which 274 species are described and a number of new ones are figured; the geographical distribution is given to each species. Inasmuch as the number of species of *Carex* is in the neighborhood of some 800, the author has rendered excellent service to the study of the genus by presenting figures of the new with their general habit and other details. Eastern Asia, and especially Japan, appears to be ex-

ceedingly rich in representatives of this genus, in spite of the fact that many large areas of, for instance, China and even Japan are yet unexplored, and the extreme north-east is almost unknown, save the collections brought home by Eschscholtz and Kjellman. In looking over the vast material treated in the present work, one notices readily that a number of very peculiar types are characteristic of that part of the world; there are, indeed, several which are utterly unlike the ordinary types of *Carex*: *C. podogyna*, for instance, possesses a utricle borne on a pilose stipe, 4–8^{mm} in length; in *C. Nambuensis* there are several lateral androgynous spikes on long, filiform peduncles; *C. pachygyna* has inflated, bladeless sheaths subtending globose, pistillate inflorescences; the leaves of *C. capilliformis* are numerous, long and capillary; in *C. hakkodensis* and *C. rhizopoda* we meet with types of very much the same habit and structure as the singular *C. lejocarpa* and *C. circinata*; *C. gentilis* shows the habit of the *Indicæ* Tuckm., while *C. moupinensis* reminds one of a *Scirpus* or *Rhynchospora* rather than a *Carex*. Only two diœcious species are known from eastern Asia: *C. Redowskiana* C. A. Mey. (*C. gynocrates* Wormskj.) and *C. grallatoria*. Among the monœcious we meet with the circumpolar *C. rupestris*, and with *C. pyrenaica*, of which the latter shows a most singular geographical distribution: Mountains of Middle- and South-Europe, western Asia, New Zealand, Rocky Mountains of Colorado and Alaska. *C. pauciflora* and *C. microglochis* are, also, represented in this Flora; *C. incurva* is reported from West China, and is known besides from so remote localities as the Arctic region, the Magellan strait, Himalaya, the coast and higher mountains of Europe. The cosmopolitan *C. vulgaris* is of course, included, besides some species which are very abundant in Europe, as for instance: *C. vesicaria*, *C. filiformis*, *C. Pseudocyperus*, *C. Buxbaumii*, etc.

It appears, altogether, as if the genus in eastern Asia possesses a number of species in common with other parts of the globe, and many that are not known from elsewhere. But in these special types, special to eastern Asia, are several which actually represent forms analogous to those which inhabit other parts of the world. There is a series of *Microrhynchæ*, of *Melananthæ* and even of *Dactylostachyæ*, which correspond very well with such types as exist in Europe, Asia and America. It is this part of the work, the geographical distribution in connection with the characterization of the various types, so excellently done by Franchet, which makes his treatise of the East Asiatic *Carices* the most important and instructive in this line of studies. T. H.

2. *Systematische Anatomie der Dicotyledonen*; by H. SOLE-REDER. Stuttgart, 1899—"Ergo species tot sunt, quot diversæ formæ seu structuræ plantarum, rejectis istis, quas locus vel casus parum differentes (Varietates) exhibuit, hodiernum occurrunt"—these words of Linnæus may be well understood as an indication of the importance of structural characters to systematic work in

Botany, but fully a century elapsed before this thought ripened and became realized as "the anatomical method" nowadays almost universally adopted in the scientific world. French botanists, prominent among whom were Mirbel and Chatin, were the first to take up the idea, that affinities and divergences in the vegetable kingdom might be sought in the internal structure, and the rapid increase in the number of genera and species made it almost necessary to invent other characters than those supplied by the study of morphology alone. In large genera it had, already, become difficult to distinguish the species of closely allied types satisfactorily, and although the very beginning of the anatomical method was purely for the sake of applying the internal characters in the service of systematic work, the study of plant-anatomy soon broadened into other lines, where the anatomy became a branch of its own, making the first and principal foundation for development of physiological research. It is the most rational development of science, when we look back at the work of the earlier botanists, when they began to discriminate organs of plants and compose the systems, and then gradually became aware of natural groups of plants, until genera and species became adopted as a means of expressing in brief the mutual affinities. Then followed the doctrine of morphology, first as a mere guide in systematic research, the terminology; then it gradually developed into the study of analogies and homologies amongst the plant-organs themselves, while contemporarily anatomical research was found to be useful for controlling the validity of morphological identities. Soon anatomy was applied as an aid in systematic work, and it is this branch of Botany which Dr. Solereder has illustrated in his present book.

In looking through the pages of this elaborate work, one gets an idea of the history of the anatomical method from its beginning to its present stage, and it is, really, an enormous quantity of labor that has already been bestowed upon this line of Botany. It seems as if the first, or at least the most effective, impulse was given in the works of Radlkofer and Vesque, and since then a number of other botanists have taken the subject up with strenuous efforts to prove the validity of anatomical characters. And so far have we reached now in the last decennia that no systematic work seems complete unless this method is considered. Engler's and Prantl's systematic treatise of the natural orders is an excellent example of what great importance is attributed to anatomical characters, and the results of such studies are, really, twofold: they bring to light a number of structural details, more or less applicable to demonstrate the affinities between many genera and species, and, moreover, we obtain by these same results an accumulation of data useful to the understanding of the internal life of the plants, the structure of the tissues, etc. From these may again be drawn conclusions as to the functions of the structural elements, their necessity to plant-life. The study of anatomy became thus

greatly encouraged when botanists discovered that such observations might, also, be useful to classification, and it is no exaggeration to say, that the anatomical method has brought out a keener interest in the study of anatomy, than if plant-structures were worked up for the mere sake of learning how plants are built.

Never before have botanists cared to give so many detailed accounts of the internal structure of the reproductive and vegetative plant-organs as they do now, and moreover, this same method has given rise to another still more modern branch of Botany, generally known as "ecology," which has attracted so much attention lately. From the knowledge of the structure, considered by itself, investigations have been broadened into the study of its application to systematic botany, while at the same time these same anatomical characters have thrown light upon the connection between these and the conditions under which plants live. But investigations of that kind require long time and steady attention, and it is very complimentary to European botanists that so much work has already been accomplished in this line: "the anatomical method." The large number of papers published upon the subject are, however, very scattered, and there has been great need for a book in which all the facts hitherto known, were brought together in systematic form. It must be said that the author has been very successful in his compilation of this enormous material to which he has himself contributed extensively. Besides giving a skillful and thoroughly scientific representation of the anatomical characters of dicotyledonous plants, the author has rendered excellent service to further studies by appending a bibliography to each of the orders treated. It would be impossible to review all the results which are laid down in this book, inasmuch as it is not a book to be simply read, but to be studied. A very detailed account is given of each order, including a large number of genera and species, and much can be learned about the anatomy of the root, the stem and the leaves, as these have been treated in their various modifications. The anatomical characters are represented as generic or specific, and it is very interesting to see the great variation that exists in a number of species, even of the same genus, when we consider, for instance, the mere structure of epidermis with its stomata and hairs, which furnish so many and such prominent distinctions. Very important characters are, also, derived from the modifications that are frequently met with in the mesophyll of the leaf or the bark of the stem with its reservoirs, as cells or ducts, the contents: crystals or liquids; moreover, the very varied development of the mestome in root, stem and leaf, the structure of the pith, etc.

But when we look at the material, upon which these investigations are based, it is at once noticed, that only a relatively small number of North American plants have, so far, been studied from this viewpoint. This ought to give an impulse to similar studies in this country, inasmuch as the North American vegeta-

tion contains so many types which no doubt are of great interest anatomically, and might even serve to solve various difficult problems in systematic botany. It is sad to see, however, that anatomical work is so little appreciated in this country, while physiological research seems to be "fashionable." And it is very astonishing that American investigators seem constantly to overlook the importance of morphological and anatomical research as the principal foundation of physiological work.

Systematic Botany in this country had prominent leaders in Nuttall, Elliott, Torrey and Gray, but we see no reason why this branch of Botany should not be extended still further, in the same scope as abroad. Europe has the advantage, however, of having had a systematic epoch, a morphological and an anatomical, followed by a physiological, while in this country physiological research was taken up before the systematic had been more than justly commenced. While recommending Dr. Solereder's book to students in this country, we hope that American investigators will take so much interest in the work that the second edition may, also, contain results gained by anatomical studies of North American genera.

T. H.

3. *Éléments de Paléobotanique*, par R. ZEILLER, Ingénieur en Chef des Mines, Professeur à l'École Nationale Supérieure des Mines. Pp. 421, 8vo. Paris, 1900. (Georges Carré et C. Naud Éditeurs, 3, Rue Racine.)—The present handsome volume has been prepared from the botanical standpoint, and forms an important addition to the small but growing group of text-books on the subject which Lesquereux twenty-five years since considered yet in its "infancy." The work of an accomplished systematist, the various groups of fossil plants and their principal forms and relationships are treated with charming clearness and precision. This text is not only a most timely one to the special student, but will be indispensable to the general reader since it displays so clearly the progress which has been made in Paleobotany, and the light which fossil plants shed upon problems of development and descent. The more important structural details are succinctly treated, and a chapter on Floral succession is added. The 210 figures inserted in the text add much to convenience of use, and the volume is a highly satisfactory one in every respect.

G. R. W.

4. *Lehrbuch der Pflanzenpalaeontologie mit besonderer Rücksicht auf die Bedürfnisse der Geologen*; von Dr. H. POTONIÉ. Pp. 402, 8vo; with 3 plates and 355 text figures. Berlin, 1899. (Ferd. Dümmlers.)—Differing essentially from Zeiller's text-book, that of Potonié approaches the subject of fossil plants with reference first to the needs of the stratigrapher; thus these two works are in a large measure supplementary to each other. The treatment is in fact more distinctly geological than in the case of any of the text-books on Paleobotany which have yet appeared.

The opening chapter on Vermeintliche und Zweifelhafte Fossilien is very interesting, and the closing one—Charakterisirung

der Fossilen Floren—is a useful résumé. The profuse introduction of illustrations throughout the body of the work is a most commendable and labor-saving feature in which it excels. The ferns are given extended treatment in consonance with their stratigraphical importance, while the generalized group of ancient plants now known as the *Cycadofilices* receive a careful structural description. Though giving, in accordance with the general plan of the work as outlined, the greater prominence to earlier forms, the author has yet found space to include many of the more fundamental details of structure characterizing the representative groups of fossil plants. The colored frontispiece showing a Carboniferous landscape is new and effective. Both this and the preceding text-book must prove widely useful. G. R. W.

5. *The Birds of Eastern North America. Part II. Land Birds.* Key to the Families and Species; by CHARLES B. CORY, Curator of Department of Ornithology in the Field Columbian Museum. Pp. i-ix, 131-387. Chicago, 1899 (Special edition printed for the Field Columbian Museum).—The second part (this Journal, viii, 398) of Mr. Cory's classified catalogue of North American Birds has recently been issued. It is liberally illustrated and will be found very useful by the many who are interested in this subject.

6. *Zoological Results based on material from New Britain, New Guinea, Loyalty Island and elsewhere, collected during the years 1895-1897*; by ARTHUR WILLEY. Part iv, pp. 357-530, plates xxxiv-liii; May, 1900. Cambridge (University Press).—The earlier parts of this important series of papers have already been mentioned in this Journal (vii, 79, 322; viii, 398). The present part contains ten papers (Nos. 18 to 27) by different authors. Among these may be mentioned as of especial interest the description by J. J. Lister of *Astrosclera willeyana*, the type of a new family of Sponges.

7. *Das Tierreich. Eine Zusammenstellung und Kennzeichnung der rezenten Tierformen.* Herausgegeben von der Deutschen Zoologischen Gesellschaft. Generalredakteur, FRANZ EILHARD SCHULZE. 9 Lieferung. *Aves*, Redakteur, A. REICHENOW. *Trochilidæ* bearbeitet von ERNST HARTERT. Pp. i-ix, 1-254, with 34 text figures and alphabetical Index. Berlin, 1900. (R. Friedländer u. Sohn.) Earlier numbers of this great work have been noticed in these pages. The present part includes the family *Trochilidæ* of the Birds and has been prepared by Mr. Ernst Hartert of the Zoological Museum at Tring, England.

IV. MISCELLANEOUS SCIENTIFIC INTELLIGENCE.

1. *The Solar Eclipse.*—The total solar eclipse of May 28th was observed under exceptionally favorable weather conditions both in this country and Europe. Nowhere on the line of totality, did clouds interfere with the observations. In consequence a great quantity of material was obtained, but as yet no results have been announced.

In addition to photographic and spectroscopic work, the bolometer was used by at least two parties with special reference to variations of heat radiation in the rifts and streamers of the corona near the sun's limb. The corona showed a conformity to the types of the last three periods of minimum sun spots so close as to be very striking even to the naked eye. The equatorial streamers were of moderate extent, and the curved polar rays sharply defined to very near the photosphere.

W. B.

2. *Lessons in Elementary Physiology*; by THOMAS H. HUXLEY, LL.D., F.R.S., edited for the use of American Schools and Colleges by FREDERIC S. LEE, Ph.D. Pp. 577, 8vo. New York, 1900 (The Macmillan Company).—Huxley's *Physiology* is so well and favorably known to teachers and students of biology that an extended review of the aim and methods of the book is scarcely necessary in this place. Despite the marked advances in physiology since the preceding revision of the *Lessons* in 1885, many teachers have felt reluctant to abandon the book, owing to the many points of excellence which it retained. Indeed, it can fairly be stated that few text-books of science equal Huxley's *Lessons* in the clearness of exposition and a type of analytic treatment which has contributed largely to the development of a proper scientific attitude. The American publishers are therefore to be complimented in having secured the assistance of Professor Lee in a new revision which is extremely satisfactory. The histological portions have appropriately been transferred from their former separate chapter to the descriptions of the various tissues to which they apply. The descriptions of the chemistry of the blood and lymph have been revised and extended; the chapter on digestion has been corrected and greatly improved; statistics of nutrition have been added and the metric system has been introduced. The parts dealing with the nervous system and innervation have been rewritten almost entirely and have been adapted to the recent progress in this department of physiology. Many new illustrations and diagrams have been inserted; and though the book has grown by nearly 200 pages, the spirit of the earlier work is retained. We cannot refrain from quoting from Dr. Lee's Preface the concluding sentence, which describes the attitude of more than one physiologist towards Huxley's *Lessons*: "The present writer has performed his task with a long-standing feeling of affection for the pages which introduced him to the study of Physiology, and first gave him a clear insight into the nature of scientific conceptions and scientific reasoning."

L. B. M.

3. *Ostwald's Klassiker der Exakten Wissenschaften*. Leipzig, 1899 (Wilhelm Engelmann).—Number 109 of this valuable series has recently been issued, it contains a memoir entitled "Ueber die Mathematische Theorie der elektrodynamischen Induction von Riccardo Felici." Translated by Dr. B. DESSAU (Bologna) and edited by E. WIEDEMANN (Erlangen).

THE

AMERICAN JOURNAL OF SCIENCE

[FOURTH SERIES.]

ART. XII.—*On Rowland's new Method for measuring Electric Absorption, and Losses of Energy due to Hysteresis and Foucault Currents, and on the Detection of short Circuits in Coils*; by LOUIS M. POTTS.

THE following investigation has had as its object the testing of methods devised by Professor Rowland* for the measurement, in the first place, of electric absorption; further, of the energy losses due to hysteresis and Foucault currents; and, finally, for the detection of short circuits in coils.

I. ELECTRIC ABSORPTION.

Historical.—It has long been known that a Leyden jar, which has been charged and then discharged, will show another charge after standing a short time. If this is discharged, after a short time the jar will show another charge; this may be repeated indefinitely. These “after-charges” are known as residual charges and are due to the phenomenon now called electric absorption. Faraday† made some experiments on this phenomenon in Leyden jars, and seems to have attributed it to a conduction of the charge into the interior of the dielectric, and after discharge creeping back again to the coatings and manifesting itself as the residual charge. Kohlrausch‡ was the first to make any elaborate investigation of the subject. He charged the condenser and then measured the potential at certain intervals with an electrometer. In this way he obtained the relation between the potential and time. He advanced the idea that the phenomenon was due to an electric polarity of

* See this Journal for July, 1899, pp. 35-57.

† Faraday, *Experimental Researches*.

‡ Pogg. *Annalen*, vol. xci, pp. 59-82, 179-214.

the particles of the dielectric, produced by the electric force between the plates of the condenser. Rowland and Nichols* have shown that certain homogeneous crystals show no electric absorption. H. Hertz† has shown that pure benzine possesses no electric absorption, while impure does.‡ The great sensitiveness of this phenomenon to change of temperature has been noted. The energy loss in condensers due to it has also been studied.§

Theory of Electric Absorption.

The theory of electric absorption has been developed by Clausius,|| Riemann,¶ Maxwell** and Rowland.†† The following development is that of Maxwell applied by Prof. Rowland to the case of a dielectric acted upon by an e. m. f. varying harmonically.

A dielectric such as paraffin paper is made of a substance of a certain dielectric capacity and specific resistance having imbedded in it particles of a different dielectric capacity and different specific resistance. Now we can very closely approximate to this case by considering a plane plate condenser, in which the dielectric is made up of a number of layers of different substances. An ordinary condenser is merely a great number of very small condensers like this, joined in multiple.

The theory of electric absorption as extended by Prof. Rowland shows that a condenser possessing electric absorption should act as a capacity in series with a certain resistance. The value of each depends on the period of the current. If b_1 , b_2 , etc., are constants and T the period of the current, the resistance is of the form

$$R = b_1 T^2 - b_2 T^4 + b_3 T^6, \text{ etc.}$$

and if a_1 , a_2 , etc., are constants the capacity is of the form

$$\frac{1}{c} = a_1 - a_2 T^2 + a_3 T^4, \text{ etc.}$$

General Theory.—The arrangement adopted is essentially a Wheatstone bridge, in which the fixed coils of an electro-dynamometer were placed in one arm of the bridge and the hang-

* Phil. Mag., p. 414, 1881.

† Wied. Annalen, p. 281, 1883.

‡ Phil. Trans., p. 599, 167; Proc. Roy. Soc., p. 468, 1875.

§ Physical Review, 1899, p. 79.

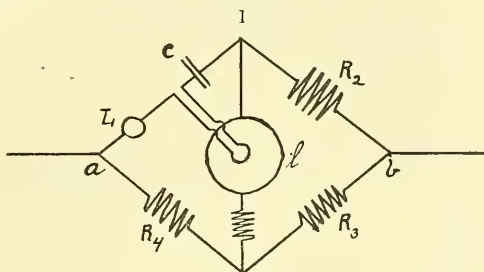
|| Théorie Mécanique de la Chaleur; deuxième partie.

¶ Riemann, Mathematische Werke, p. 48.

** Electr. and Mag., vol. i, p. 452. †† This Journal, Dec. 1897, p. 429.

ing coil in the cross-connection, in place of the galvanometer in the direct-current method of use. The adjustment of the bridge thus used depends upon the fact that there will be no deflection of the electro-dynamometer if the phase difference of the current in the fixed coils and those in the hanging coil is 90° .

Fig. 1 is the arrangement used. Let R_1, R_2, R_3 and R_4 be the resistances of the different arms and r that of the cross-connection. Let C_n be the current in the arm n of the bridge, and C_5 in the cross-connection.



If we apply at a and b a direct electromotive force E , we shall have the following expressions for the currents :

$$C'_5 = E \frac{R_2 R_4 - R_3 R_1}{\Delta} \quad (1)$$

$$C'_1 = E \frac{R_4(R_2 + R_3) + r(R_3 + R_4)}{\Delta} \quad (2)$$

where

$$\Delta = r(R_3 + R_4)(R_1 + R_2) + R_1 R_2 (R_3 + R_4) + R_3 R_4 (R_1 + R_2)$$

If in place of using a direct current we apply to the terminals a, b , a simple alternating e.m.f. $E e^{ibt}$, we shall get the corresponding quantities by the following substitutions :

$$C'_1 = C_1 e^{i(bt + \phi_1)} \quad C'_5 = C_5 e^{i(bt + \phi_5)}$$

and if we place in the arm 1 a coil of self-induction L and a capacity c , we must substitute for R_1, R'_1 , where

$$R'_1 = R_1 + ibL - \frac{i}{bc}$$

and, if l is the self-induction of the hanging coils of the electro-dynamometer, we must replace r by r' , where

$$r' = r + ibl$$

Making these substitutions and dividing (1) by (2), we have

$$\frac{C_5}{C_1} \epsilon^{i(\phi_3 - \phi_1)} = \frac{R_2 R_4 - R_1 R_3 - i \left(bL - \frac{1}{bc} \right) R_3}{R_4 (R_2 + R_3) + r (R_3 + R_4) + ibl (R_3 + R_4)} \quad (3)$$

Now taking only the real part of the above quantities we have

$$\frac{C_5}{C_1} \cos \phi = \frac{(R_2 R_4 - R_1 R_3) \{ R_4 (R_2 + R_3) + r (R_3 + R_4) \} + \left(\frac{l}{c} - b^2 l L \right) R_3 (R_3 + R_4)}{\{ R_4 (R_2 + R_3) + r (R_3 + R_4) \}^2 + b^2 l^2 (R_3 + R_4)^2} \quad (4)$$

If $\phi = 90^\circ$, $\cos \phi = 0$ or

$$(R_2 R_3 - R_1 R_3) \{ R_4 (R_2 + R_3) + r (R_3 + R_4) \} + \left(\frac{l}{c} - b^2 l L \right) R_3 (R_3 + R_4) = 0$$

$$\text{or} \quad R_1 = \frac{R_2 R_4}{R_3} + \left(\frac{l}{c} - b^2 l L \right) \frac{R_3 + R_4}{R_4 (R_2 + R_3) + r (R_3 + R_4)} \quad (5)$$

This then is the condition satisfied when there is no deflection of the electro-dynamometer.

The first term of the above is the same as the expression for R_1 , when the adjustment is conditioned by no deflection of a galvanometer in the cross-connection, and a direct current is used. The second is a correction term, always very small, at most one per cent and often entirely negligible. In a circuit carrying an alternating current the effective values of R_1 , etc., are not usually the same as their actual ohmic values, but are larger. They include that part of the impedance against which work must be done to maintain the current. Let R_2 , R_3 and R_4 be as nearly as possible pure ohmic resistances, i. e. let their actual values be equal to their effective ones. And let the arm 1 contain iron, a condenser possessing electric absorption, or any piece of apparatus using energy which is not expended in heating the conductors of that arm. The value of R_1 calculated by the above formula will be the effective resistance, and this, less the actual value as measured by a direct current, will be the increased resistance due to the hysteresis of the iron, the energy-loss by electric absorption, etc.

In the measurement of certain quantities (*e. g.* hysteresis loss in iron) by this method, it is necessary to insert in arm 1 of the bridge a large coil. Such a coil acts not as a pure self-induction, but on account of the numerous turns of the wire so close to one another as a self-induction in *multiple* with a

capacity. On this account formula (5) is not strictly accurate but should include a term involving the capacity of the coil. Let the capacity which in parallel with the self-induction, L , will have the same effect as the given coil be k and let R'_1 be the resistance of the coil and R''_1 the resistance of the remainder of arm 1. Substitute in (3) in place of

$$R_1 + ibL - \frac{i}{bc}$$

$$R''_1 + \frac{R'_1 + ibL}{(1 - b^2kL) + ibkR'_1}$$

and also put

$$R_4(R_3 + R_2) + r(R_4 + R_3) = A$$

Then

$$\frac{C_5}{C_1} \varepsilon^{i(\phi_5 - \phi_1)} =$$

$$\frac{(R_2R''_1 - R_2R_4)(1 - b^2kL) + R'_1R_3 + i\{bkR'_1R''_1R_3 + bLR_3 - A(1 - b^2kL) - b^2kLR'_1(R_3 + R_4) + i\{AbkR'_1 + (1 - b^2kL) \frac{bkR_1R_2R_4}{bL(R_2 + R_4)}\}}}{(6)}$$

As before, the condition for no deflection is that the real part of (6) vanish, or

$$[(R_2R_4 - R_3R''_1)(1 - b^2kL) - R'_1R_3][A(1 - b^2kL) - b^2kLR'_1(R_3 + R_4)] = [AbkR'_1 + (1 - b^2kL)bL(R_3 + R_4)][bLR_3 - bkR'_1(R_2R_4 - R_3R''_1)]$$

Expanding this it becomes,

$$A(1 - b^2kL)^2(R_2R_4 - R_3R''_1) - AR'_1R_3(1 - b^2kL) + b^2kLR''_1(R_3 + R_4)R_3 - (R_2R_4 - R_3R''_1)(1 - b^2kL)b^2kLR'_1(R_3 + R_4) = Ab^2kLR'_1R_3 - Ab^2kR'^2_1(R_2R_4 - R''_1R_3) (1 - b^2kL)(R_3 + R_4)R_3b^2L - (1 - b^2kL)(R_3 + R_4)(R_2R_4 - R''_1R_3)b^2kLR'_1$$

Now since k is in all cases small and l is also small, the terms above which involve k^2 and kl may be dropped, whence we get,

$$R_1 = \frac{A(1 - b^2kL)^2(R_2R_4 - R_3R''_1) - b^2LR_3(R_3 + R_4)R_3(1 - b^2kL)}{AR_3}$$

$$= (1 - b^2cL)^2 \frac{R_2R_4 - R_3R''_1}{R_3} - b^2L \frac{R_3 + R_4}{A} (1 - b^2kL)$$

Now since the last term is very small and $1 - b^2kL$ is nearly 1, it may be dropped; and we have for the final formula,

$$R'_1 = \frac{R_2R_4 - R_3R''_1}{R_3} + \left(\frac{l}{c} - b^2L\right) \frac{R_3 + R_4}{R_4(R_3 + R_2) + r(R_3 + R_4)} - 2b^2kL \frac{R_2R_4 - R_3R''_1}{R_3} \quad (7)$$

In the measurement of electric absorption and hysteresis loss, it is necessary to know the period of the current used. In this investigation a speed counter and chronograph were used; the speed counter was placed at the end of the dynamo shaft, and directly connected to it, a contact was so arranged that for every one hundred revolutions of the dynamo armature a circuit was closed and a record made on a chronograph sheet. On a table beside the electro-dynamometer was a key, which also could be used to make a record on the chronograph sheet. As soon as the bridge had been balanced, this key was pressed. And by the measurement of this sheet the period of the current at the time of the observations was quite accurately determined. The error from this source was usually not more than one part in 1000, never more than 1 in 100.

After each adjustment, the resistance of each arm was determined by the use of a "Post-office Box" (when the current through any part of the apparatus was not large, it was not necessary to measure its resistance after each adjustment). The resistance of that part which was affected the most by heating was measured first, and in this way the actual value at the time of adjustment was quite closely determined. In cases where the heating was large, the error from this cause might reach several parts in 1000 in R' , and consequently a considerably greater amount in the value of the electric absorption resistance.

Errors due to induction and electrostatic action of the different portions of the apparatus were carefully guarded against by the arrangement. And induction was tested for by reversal of the relative directions of the currents in different portions of the apparatus. Usually no effect was noticed or at most it was very small. The errors introduced by the self-induction of the electro-dynamometer coils and also that caused by the electrostatic action of the turns of a large coil on one another, were determined and corrected for when sufficiently large in amount.

Apparatus.

*Electro-dynamometer.**—The self-induction of the fixed coils was .0165 henry, and of the hanging coil .0007 henry.

Dynamos.—The current used in this investigation was furnished by one of three dynamos. The Westinghouse alternator in power house of the University furnished a current of period .0075, i. e. 133 complete periods per sec. This was used for only a few observations. In most of the work two small dynamos constructed in the University workshop were

* See this Journal, July, 1897, p. 35.

employed. Both were directly connected to small electric motors. Both had armatures of the pancake type. The one had four coils in the armature and four poles and thus produced a current of two complete periods for each revolution of the armature. The other was larger but of similar construction, having six coils in the armature and six poles, and gave three complete periods for one revolution. If the load on the dynamo was not changed, these dynamos would run at a very nearly constant speed. With the second dynamo, the number of complete periods per second could be varied from 6 to 70. However, at the lower speed the electro-dynamometer was difficult to balance, since the hanging coil would vibrate with the current and blur the image of the scale. The voltage furnished could be controlled very well by changing the strength of the field. In any one series of observations the same dynamo was used, as the results using different dynamos would not be comparable, on account of the different harmonics introduced. The small dynamos which were almost exclusively used gave, however, very good sine curves.

Resistances.—The high resistances and those which were required to carry very small currents were made of fine german-silver wire wound on thin sheets of fiber. The self-induction and electrostatic action of these was practically zero. The lower resistances, and those required to carry larger currents were made of a special resistance wire, which had a very slight negative temperature coefficient, and would bear considerable heating with a very small change of resistance. This wire was wound on slates. Each slate contained sufficient number 30 wire to have nearly 2000 ohms resistance. These were conveniently subdivided for adjustment. For the final adjustment an ordinary resistance box was used, but never more than fifty ohms were used in this box, and then the total resistance of that arm was at least 1000 ohms.

Self-inductance.—Two coils were used.

A. External diameter 35.46^{cm}; internal diameter 23.8^{cm}, 3700 turns No. 20 B. and S. Self-inductance 5.30 henrys. Resistance 188 ohms.

C. Same dimensions as A except depth. Self-inductance 1.30 henrys, 1747 turns No. 22 B. and S., single cotton covered copper wire. Resistance about 78 ohms.

Condensers.—2 and 3. Paper condensers made by Marshall of 2 and 3 micro-farads capacity.

Willyoung.—8 micro-farad wax condenser, made by Willyoung & Co. and divided into sections of one micro-farad each.

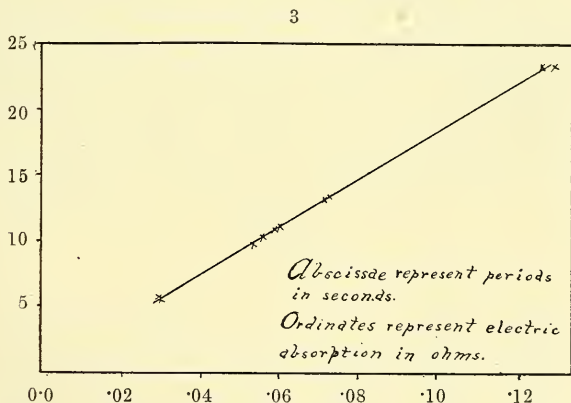
Mica condensers.— $\frac{1}{8}$ M. F. standard condenser made by Elliot Bros.

r' equal to the resistance of the hanging coil of the electro-dynamometer. This avoided any sudden change of current and hence a change of speed of dynamo. The key R_2 was the chronograph key described above and was closed just at the time the bridge was balanced.

TABLE I.
 $T = .0075$ $r = 6535$

Condenser.	R_1	R_2	R_3	R_4	Calculated R'_1	A $R'_1 - R_1$
3 -----	34.04	2022	4600	99.7	43.86	9.82
3 -----	34.45	2020	9090	203.4	45.22	10.77
2 -----	34.45	2020	2475	99.7	81.46	47.01
2, 3[S] ---	34.45	2020	8920	407.0	92.23	57.78
2, 3[P] ---	34.45	2020	2205	51.7	47.37	12.92
2, 3[P] ---	34.48	1991	4130	99.7	48.09	13.61
2, 3[S] ---	34.19	1992	2031	99.6	97.83	63.64
3 -----	34.45	1992	3325	99.6	59.72	25.27
2 -----	34.49	1992	2414	99.6	82.28	47.79

A few measurements were first made using the paper condensers 2 and 3. Table I shows the results. In this and the following tables, R'_1 denotes the effective value of R_1 calculated by formula (5) and $A(=R'_1 - R_1)$ the resistance due to the electric absorption. T is the period of the current. The



results with these condensers was very unsatisfactory, as the heating was so great that it was difficult to make accurate determinations.

The wax condenser made by Willyoung was next used. This condenser had been made in a vacuum under pressure, and showed quite small heating by the current. By taking a series of measurements during an afternoon, the results were not affected to any great extent by the changes in temperature

of the room, as these were comparatively slow. Of course there was still some slight heating by the current; but this in extreme cases did not amount to enough to occasion an error of more than one or two per cent.

Table II gives results for the absorption of all eight sections of this condenser in parallel, for different periods of the current. In Fig. 3 the above results are plotted; the ordinates

TABLE II.

Date, etc.	$r=6020$		$R_2=1060$		$R_4=303.1$		
	R_3	R_1	Calculated. R_1'	A $(R_1'-R_1)$	T.	$\frac{A}{T}$	
2-23-99	1	7192	34.68	44.81	10.13	.0557	181.7
-22-	2	7251	34.70	44.45	9.75	.0544	179.2
	3	5564	34.67	57.91	23.24	.1269	183.2
	4	5556	34.66	58.00	23.34	.1291	180.8
	5	6694	34.73	48.15	13.42	.0735	182.6
	6	6718	34.77	47.97	13.20	.0717	183.6
	7	7016	34.85	45.93	11.08	.0602	184.1
	8	7044	34.88	45.75	10.87	.0592	183.6
	9	7921	35.11	40.68	5.57	.0301	185.2
	10	7910	35.10	40.74	5.64	.0301	187.4

represent A , the resistance, which in series with the condenser, would be equivalent to the electric absorption; and the abscissæ represent the period of the current in seconds. The curve proves to be a straight line, or A/T is a constant within the limits of error of the experiment. On account of this very simple relation connecting A and T , this condenser was exceedingly convenient for a test of the method.

The first test applied was to change the relations of the resistances in the different arms of the bridge. Table III

TABLE III.

Date, etc.	$r=6536$				R_1' Calculated.	A $(R_1'-R_1)$	T	$\frac{A}{T}$		
	R_2	R_3	R_4	R_1						
3-16-99	1	509.2	5561	507.2	34.61	46.57	11.96	.0640	186.7	
	29	2	509.2	5581	507.2	34.66	46.40	11.74	.0628	187.0
	3	2071.	4598	99.56	34.65	44.97	10.32	.0558	184.9	
	4	2071.	4622	99.56	34.65	44.74	10.09	.0541	186.5	
	5	1009.5	6750	302.9	34.71	45.42	10.71	.0572	187.2	
	6	1009.5	6785	302.9	34.68	45.19	10.51	.0563	186.7	
	7	302.1	6732	1009.2	34.83	45.41	10.58	.0563	187.5	
	8	302.1	6728	1009.2	34.84	45.38	10.54	.0563	186.8	
	9	99.46	4511	2044.5	35.00	45.19	10.19	.0542	188.0	
	10	99.46	4511	2044.5	35.04	45.19	10.16	.0536	189.7	

gives the results for R_2R_4 nearly constant and R_2/R_4 varied. The variation in R_2/R_4 was about 200 per cent and A/T is practically constant. The slight increase of A/T in the last two measurements is due to the larger current in arm 1 and a consequent heating of the condenser and also to the fact that after the coil of the electro-dynamometer had been heated slightly, it would be cooled a small amount before its resistance could be measured. Table IV shows the results when R_4

TABLE IV.

Date, etc.	$R_4 = 1505.8$		$r = 4810$		T	A ($R'_1 - R_1$)	$\frac{A}{T}$	Mean.
	R_2	R_3	R_1	Calculated, R'_1				
4-19 1	2070	76,600	34.68	40.70	.0312	6.02	192.9	194.8
118 5	2070	74,200	34.80	42.02	.0367	7.22	196.7	
2	1009	36,120	34.76	42.07	.0375	7.31	194.9	196.0
4	1009	35,970	34.80	42.25	.0378	7.45	197.1	
3	99.48	3,542	34.75	42.30	.0386	7.55	195.6	195.6

is kept constant and R_2 and R_3 are varied. These also show A/T constant. It appears then from the above facts that A is a quantity independent of the relative values of the resistances in the different branches of the bridge.

In the next test the period was kept constant and the electromotive force acting on the condenser was varied about 300 per cent. The results of this test are given in Table V. The values for the higher electromotive forces are slightly greater, owing to the two heating effects mentioned above. Aside from

TABLE V.

Date, etc.	$R_2 = 1008.9$		$R_4 = 303.8$		A $R'_1 - R_1$	T	$\frac{A}{T}$	Volts e.m.f. acting on condenser $\equiv V$
	R_3	r	R_1	Calculated, R'_1				
3-10-99								
1	6649	24,410	35.45	46.14	10.69	.0523	204.3	240.
27								
2	6593	24,410	35.34	46.53	11.19	.0562	199.3	240.
3	6797	12,490	35.13	45.17	10.04	.0508	197.6	152.
4	6777	12,490	35.15	45.31	10.16	.0511	198.9	152.
5	6606	6,348	35.03	46.56	11.53	.0584	197.3	77.
6	6473	6,348	34.96	47.57	12.55	.06418	195.5	77.
3-14								
6	6591	4,807	34.73	46.60	11.87	.0590	201.2	76.
28								
7	6803	6,371	34.79	45.14	10.35	.0515	201.1	142.
8	6732	6,371	34.78	45.61	10.83	.0540	200.6	142.
9	6555	18,830	35.01	46.74	11.73	.0567	203.5	243.

this, A/T is constant, i. e., A is independent of the current flowing through the condenser.

The variation in A/T due to changes of temperature was obtained as follows: in an opening made in the side of the box containing the condenser, a mercury thermometer was placed. The temperature indicated by the thermometer was of course not that of the inside of the condenser but that of the outer edge. The condenser, however, was kept within a degree or two at least of the desired temperature for some six or more hours before being used.

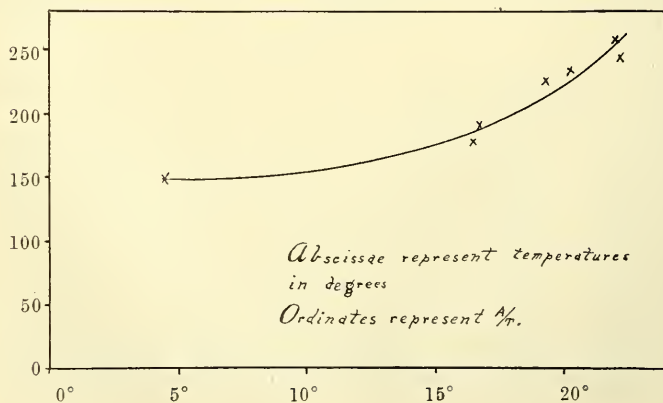
This method gave sufficiently accurate results, as there was no occasion for an accurate determination of the temperature.

TABLE VI.

Date, etc.	Temp. cent. t	$r=4810$		R_4	R_1	Calcu- lated. R'_1	A $R'_1 - R_1$	T	$\frac{A}{T}$
		R_2	R_3						
11-16-99	19.3	709.0	2557	203.62	35.73	56.48	20.75	.0920	226.0
102									
1	22.0	707.8	2195	202.6	34.90	65.35	30.45	.1167	258.3
11-2-99									
7	22.2	707.8	2614	202.6	34.97	54.88	19.91	.0816	242.7
1	16.5	709.9	3271	203.0	34.40	44.08	9.68	.0543	178.6
12-8									
2	16.7	709.6	2862	203.0	34.44	50.04	15.60	.0816	191.3
106									
	20.2	2985.	2985	203.3	34.93	48.27	13.34	.0569	234.5
1	4.5	302.4	2696	407.3	34.55	45.87	11.32	.0752	150.9

Table VI gives the results of this investigation, and in fig. 4 they are plotted with temperatures as abscissas and A/T as

4



ordinates. It appears that at ordinary laboratory temperatures a variation of $\cdot 1^{\circ}$ C. will cause a change of about 1 per cent in the value of A/T and consequently in A for a given T . From this it appears that the slight variation noted above in the constancy of A/T would be easily accounted for by the changes of temperature due to the current or the gradual changes due to changes in the temperature of the room, since a series of observations usually occupied three hours or longer.

Table VII gives the results for the two $\frac{1}{3}$ microfarad condensers described above connected in parallel; and in fig. 5 the results are plotted as before.

TABLE VII.
 $R_4 = 303\cdot 3$ $r = 4811$

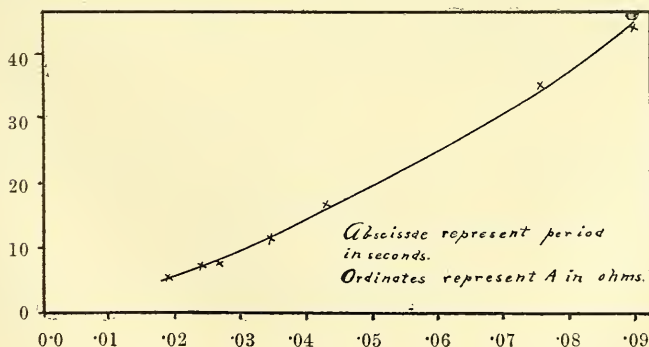
Date, etc.		R_2	R_3	R_1	Calculated, R'_1	A ($R'_1 - R_1$)	T
11-29-99	1	409·7	2984·	34·07	41·66	7·59	·0271
105	2	409·7	2723·	34·07	45·59	11·52	·0345
	3	409·7	1807·	34·07	68·79	34·72	·0752
	4	409·7	1604·	34·07	77·49	43·42	·0893
	5	409·7	2454·	34·07	50·66	16·59	·0428
	6	409·7	3010·	34·07	41·30	7·23	·0239
	7	410·6	3157·	34·04	39·38	5·34	·0187

It was necessary to use the two in parallel in order to get sufficient current through the fixed coils of the electro-dynamometer. This was especially true for long periods, as then the impedance of the condenser increased and at the same time the available electromotive force from the dynamo decreased.

Capacity of a Condenser which shows Electric Absorption.

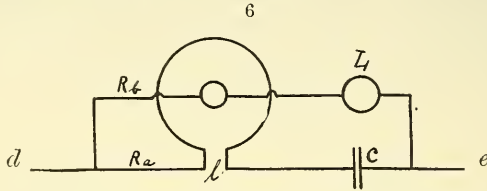
From the theory of electric absorption as based on the heterogeneous nature of the dielectric it appears that there should be a variable value of the capacity of such a condenser depend-

5



ing upon the period of the current flowing through the condenser. The method chosen for the measurement of capacity was one described in this Journal, December, 1897.

Method.—This consisted in the use of a divided circuit. One branch, a , contained a resistance R_a , either the fixed or hanging coils of the electro-dynamometer, and a condenser whose capacity c was to be studied. The other arm, b , contained a resistance R_b , a coil with which the capacity c is com-



pared, and either the hanging or fixed coils of the electro-dynamometer. Let L be the coefficient of self-induction of the coil plus that of the coils of the electro-dynamometer in that arm, and let l be the self-induction of the coils of the electro-dynamometer in the arm b . Let an electromotive force

$$E = E_0 e^{i\omega t}$$

be applied to the terminals de . Now representing the maximum values of the current by C_a and C_b and the phases by ϕ_a and ϕ_b we have for the branch a

$$C_a e^{i(\omega t + \phi_a)} = E_0 \left(R_a + i\omega l - \frac{i}{\omega c} \right)$$

and for branch b

$$C_b e^{i(\omega t + \phi_b)} = E_0 (R_b + i\omega L)$$

If we adjust the resistances until there is no deflection of the electro-dynamometer, we shall have the difference of phase $(\phi_b - \phi_a) = 90^\circ$ or $\cos(\phi_b - \phi_a) = 0$. Further,

$$\frac{C_b}{C_a} e^{i(\phi_b - \phi_a)} = \frac{R_b + i\omega L}{R_a + i\omega l - \frac{i}{\omega c}} \quad (12)$$

Hence since $\cos(\phi_b - \phi_a) = 0$ we must have the real part of this equal to zero, or

$$(1 - b^2 c l) (b^2 c L) = b^2 c^2 R_a R_b$$

or

$$\frac{L}{c} = R_a R_b \left(\frac{1}{1 - b^2 c l} \right) \quad (13)$$

As in the case considered above, a coil does not act as a self-induction alone but as a self-induction in *parallel* with a capacity due to the electrostatic action of the turns of the coil on one another. For this reason the above formula is not exact for any actual case; but there must be substituted above in place of $R_b + i\delta L$

$$R'_b + \frac{\frac{R''_b + i\delta L}{i\delta c'}}{R''_b + i\delta L - \frac{i}{\delta c'}} \tag{14}$$

Substituting this in (12) we have

$$\begin{aligned} \frac{C_b}{C_a} e^{i(\phi_b - \phi_a)} &= \frac{R'_b + \frac{R''_b + i\delta L}{i\delta c' R''_b + 1 - b^2 c' L}}{R_a + i\delta l - \frac{i}{\delta c}} \\ &= \frac{R'_b(1 - b^2 c' L)\delta c + R''_b \delta c + i[(b^2 c c' R'_b R''_b + b^2 c L)]}{R_a(1 - b^2 c' L)\delta c + (1 - b^2 c l)\delta c R''_b - i[(1 - b^2 c' L)(1 - b^2 c l) - b^2 c c'] [R''_b R_a]} \end{aligned}$$

As before, the condition for no deflection is that the real part of this equals zero. Hence

$$\begin{aligned} b^2 c [R'_b(1 - b^2 c' L) + R''_b] [R_a(1 - b^2 c' L)c + R''_b(1 - b^2 c l)c'] \\ = b^2 c [R'_b R''_b c' + L] [(1 - b^2 c' L)(1 - b^2 c l) - b^2 c c' R''_b R_a] \end{aligned}$$

Expanding,

$$\begin{aligned} R'_b R_a (1 - b^2 c' L)^2 c + R''_b (1 - b^2 c l) (1 - b^2 c' L) R'_b c' \\ + R''_b R_a (1 - b^2 c' L)c + R''_b{}^2 (1 - b^2 c l)c' = \\ \frac{R'_b R''_b c' (1 - b^2 c' L) (1 - b^2 c l) + L(1 - b^2 c' L) (1 - b^2 c l)}{-b^2 c c' R''_b R_a (L + R'_b R''_b c')} \end{aligned}$$

Since c' is small, we can drop the term in c'^2 . We have, on dividing by $(1 - b^2 c' L) (1 - b^2 c l)c$ and rearranging terms,

$$\begin{aligned} \frac{L}{C} &= \frac{R'_b R_a (1 - b^2 c' L)}{1 - b^2 c l} + \frac{R''_b R_a}{(1 - b^2 c l)} + R''_b{}^2 \frac{e'}{e} \frac{1}{(b^2 - b^2 c' L)} \\ &+ \frac{b^2 c' R''_b R_a L}{(1 - b^2 c l) (1 - b^2 c' L)} \end{aligned}$$

or

$$\begin{aligned} \frac{L}{C} &= \frac{(R'_b + R''_b) R_a}{(1 - b^2 c l)} + R''_b{}^2 \frac{c'}{c} \frac{1}{(1 - b^2 c' L)} \\ &- \frac{b^2 c' L R_a R'_b}{(1 - b^2 c l)} + \frac{b^2 c' L R_a R''_b}{(1 - b^2 c l) (1 - b^2 c' L)} \end{aligned}$$

Now, since in any case the last three terms are small, and $(1-b^2cl)$ and $(1-b^2c'L)$ are nearly one, they may be dropped from the last terms and we have

$$\frac{L}{C} = \frac{(R'_b + R''_b)R_a}{(1-b^2cl)} + \frac{c'}{c}R''_b{}^2 + (R''_b - R'_b)b^2c'L \quad (15)$$

Investigation—The arrangement is the same as in fig. 2. In position *a* and with a resistance in arm 4, electric absorption can be found. In position *b* and with a coil *L* in arm 4 in place of the resistance, the capacity of the condenser can be compared with the standard coil *L*. In formula (13) R_a includes not only the ohmic resistance of branch *a*, but also the added resistance due to electric absorption.

A preliminary investigation was carried out to find whether the correction due to the capacity of the coil *C'* were appreciable and, if so, to ascertain its amount.

The method chosen was as follows: An arrangement was made as in fig. 2, except that in place of the condenser *C* the coil *L* with which the capacity of the condenser is to be compared is placed in arm 1 of the bridge. The arrangement was first balanced with a direct current; and the value of R_1 as measured and as calculated from R_2, R_3, R_4 were the same. In Table VIII are given the results for R'_1 as calculated for three

TABLE VIII.

Date	$R_2=1010$		$R_4=912.5$	$r=4810$	Coil A=5.3 henry
	R_3	R_1	R'_1	T	
	By formula 13				
4.20	1 4080	224.6	225.8	.0188	1.0
	2 4092	224.8	225.2	.0488	.4
	3 4074	225.0	226.2	.0182	1.2
	4 4034	225.0	228.4	.0182	3.4 with .01 microfarad in parallel with coil.

If formula (15) is used in calculating R'_1 it is 226.2. Hence $D=1.2$ as without condenser.

periods of the current. In all three cases R'_1 is greater than R_1 by an amount *D*. A third observation was made with a condenser of .01 microfarad capacity shunted across the terminals of the coil. The corrected formula (7) was tested in this way. The result gives the same value of *D*, with or without the condenser, thus verifying the formula. By assuming the quantity *D* as entirely due to the electrostatic action of the coil, which if not absolutely true, the formula will at least give a value of c' the equivalent capacity of the coil, which may be used as the limiting value. The value of c' is .006 microfarad. In formula (15) the last term will be very small as

compared to the first for any values of b used in this work. The second term will in the most unfavorable circumstances amount to only 1 part in 10,000, so that the corrections due to the electrostatic action of the coil may be entirely neglected.

The next point investigated was the correction due to the electric absorption. In these observations the absorption was determined, the capacity then measured, and the electric absorption again determined. In Table IX are given the

TABLE IX.

Date, etc.	$r=4811$		$R_2=707.8$		$R_4=202.6$	$L=5.302$ in 2 and 5.318 in rest.	
	R_3	R_1	Calculated. R'_1	$\frac{A}{(R'_1-R_1)}$	T	$\frac{A}{T}$	Mean. $\frac{A}{T}$
N-16	1 2195.	34.90	65.35	30.45	.117	261.	253.
	7 2614.	34.97	54.88	19.91	.0816	244.	
	R_a	R_b	T	A	Corrected for electric absorption. R_a R_b		$10^6 \times \frac{L}{R_a R_b} = C$
	2 238.2	2604	.121	30.6	268.8	2604	7.574 M.F.
	3 228.8	3053	.103	26.1	—	3079.	7.549
	4 328.2	2119	.103	26.1	—	2145.	7.554
	5 431.5	1602	.102	25.8	—	1628.	7.569
	6 531.5	1302	.0908	23.0	—	1325.	7.552

results with the condenser in one arm and then changing to the other, and also changing the resistance in series with the condenser. The results are corrected for electric absorption and the change in L due to the change of c and the coil from one arm to the other, caused by the coils of the electro-dynamometer having different coefficients of induction. The greatest difference between two determinations under these different conditions is 3 parts in 1,000.

The change of capacity with the period of the current was now tried. Table X shows the results of the investigation. The error of each observation has a limit of about 1 part in 1,000, if the observations are compared among themselves, while the actual error, as compared with the true ratio $\frac{L}{C}$ may be in error two or three times this; but we are not particularly concerned here with the actual value, but merely the change with change of period. In fig. 7 the results are plotted on two scales. The results for the capacity seem to agree very well

consumption of energy will manifest itself by an apparent increase in the resistance and consequently a greater i^2R loss. This fact is made use of in the following method for the detection of short circuits in coils and was suggested by Professor Rowland.*

The method is as follows: the connections are as in fig. 2, with the exception that a coil of wire is introduced in arm 1 in place of the condenser. If the resistances are now balanced until there is no deflection of the electro-dynamometer, and a mass of metal approaches the coil, there will be a deflection of the electro-dynamometer; owing to the increase in the effective resistance of the arm 1. If a coil of wire whose ends are not connected be laid on top of the coil in arm 4 there will be no deflection, while if the ends are connected or if there is a short circuit in the coil there will be a deflection.

TABLE XI.

No. of turns.	$R_2=1010$		$R_4=1507$	$R_1=225.7$	$T=.015$	
		Dia. of wire.	R_3	Uncorrected. R'_1	D	R_c
0		—	6711	226.8	0	
2.	14	1.59	5882	258.8	9.15 cm.	32.0
3.	22	.62	6497	234.3	2.21	7.5
5.	25	.44	6520	233.4	1.90	6.6

Table XI shows the sensitiveness of the method. The coil A of 5.3 henrys was used. The bridge was balanced and then small coils of wire the same size as the inner diameter of the coil were placed on the coil and the deflection noted; and the apparent increase of resistance was determined in the same manner as the electric absorption in the case when the coil L was replaced by a condenser. Column D gives the deflection after the coils were placed on the large coil in arm 1 and R_c is the apparent increase in resistance of arm 1. It appears from this table that with a coil of the size used, a short circuit in another coil of same size could be detected, even though the coil were of quite fine wire and only one turn was crossed. Other conditions being the same, the sensitiveness varies directly as the cross section of the wire in the coil to be tested, if the resistance of the contact between the two ends of wire is neglected. In cases where small coils are to be tested, the sensitiveness may be increased by filling the center of the coil with iron. And of course as short a period of current as available should be used.

* This Journal, December, 1897.

III. HYSTERESIS.

The arrangement used for the determination of losses of energy due to hysteresis and Foucault currents is the same as for the determination of electric absorption, with several additional elements.

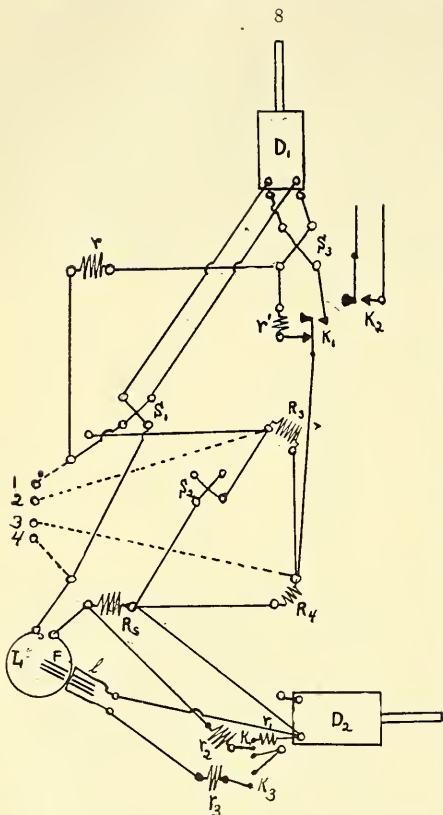


Fig. 8 shows the arrangement used. All the arms are the same as in fig. 2 except arm 1. Arm 1 contains a coil L . In this coil is placed the iron to be tested, F . l is a small coil of wire surrounding the iron, and used to determine the induction through the iron. The two coils of the electro-dynamometer D_2 are connected in series. The key K_3 serves to put the coil l , the resistance r_3 and the coils of the electro-dynamometer in series. By noting the deflection produced by the current induced in l , the induction may be calculated, the electro-dynamometer having been previously calibrated.

R_3 is a small resistance in arm 1. A small current is shunted off from the terminals of R_3 and can be sent through the resistance r_2 and the electro-dynamometer by the key K_4 ; and thus the total current in arm 1 may be determined.

Measurement of the current.—The calibration of the electro-dynamometer showed the current to be quite accurately proportional to the square root of the deflection (\sqrt{D}). A deflection of 1 cm. corresponds to a current of .00213 ampere. We shall then have the total current in arm 1

$$C_1 = .00213 \sqrt{D} \frac{r_2 + R_3}{R_1}$$

In place of measuring the current by an electro-dynamometer, in some cases a small ammeter might be used, but as the current enters as a square in the energy loss, it must be accurately determined.

Determination of the induction: For the determination of the induction several different forms of the secondary coil were tried. Coils of the same size as the internal diameter of the coils A and C were first used. These were found unsuitable, however, as a considerable current was induced in them, even when they contained no iron, and their coefficient of self-induction was not negligible. The best form was found to be a coil of fifty turns, just large enough to contain the iron used. No measurable current was induced in this, when it contained no iron, and its self-induction was negligible. The e.m.f. around the circuit when K_1 is closed will be E_1 where

$$E_1 = \sqrt{D} \times .00213 \times R \text{ volts}$$

Now let N be the number of turns in the coil l and S the average cross section of the iron surrounding the coil L and we have the induction per sq. cm. B

$$B = \frac{\sqrt{D} \times .00213 \times R}{N \times 4.44 \times r} \times 10^8 \text{ c. g. s. units}$$

where N is the number of complete periods of the current.

Energy loss.—The energy loss due to hysteresis is ordinarily expressed as a certain loss per c.c. of iron per cycle. Energy loss = $i^2 R t$.

If C_1 is the current in arm 1, H the total added resistance due to the hysteresis and Foucault currents, v the volume of the iron used, we have the energy loss per cycle,

$$C_1^2, \frac{H}{V} T$$

since t will equal T , the period of the current.

Experiment.—The iron used was ordinary transformer iron. The plates were L-shape, and could be fitted about the coil C very nicely. In the first place the uniformity of B was tested for different quantities of iron. A slot was cut in the center of one side of an L plate, thus dividing the side into two parts which were made as nearly equal as possible. A coil of wire of one hundred turns was wound on each of these parts of the iron. By noting the deflection produced by the two coils in succession when the iron was placed in coil A, the relative value of B close to the coil A and farther away could be tested. By placing in the coil different numbers of the plates and

placing the test plate at different points, the uniformity of B was tested.

Table XII shows the results of this test. It appears from this table that with either 8 or 10 pieces of iron B is practi-

TABLE XII.

No. of pieces of iron.	Test plate on outside of bundle.		Per cent var. B.	Test plate next to outside plate.			Two plates outside of test plate.		
	Lower coil.	Upper coil.		Lower coil.	Upper coil.		Lower coil.	Upper coil.	
2	6·51	6·03	3·	—	—	—	—	—	—
4	4·14	3·82	3·	—	—	—	—	—	—
6	3·60	3·48	1·9	3·60	3·62	·7	—	—	—
8	2·65	2·55	2·	2·71	2·71	0·	—	—	—
10	—	—	—	—	—	—	2·75	2·85	2·

cally uniform. With fewer pieces the induction was greatest next the coil, and when more were used the magnetism was not as great in the central pieces.

Table XIII shows the results of a series of measurements, and in fig. 9 a curve is plotted showing the relation between B and the energy loss per cycle.

TABLE XIII.

Weight of iron=340 g. Volume=43·3 cc. Area section=981 sq. cm. D_1 =deflection dynamometer by current C. D_2 =same for secondary current.

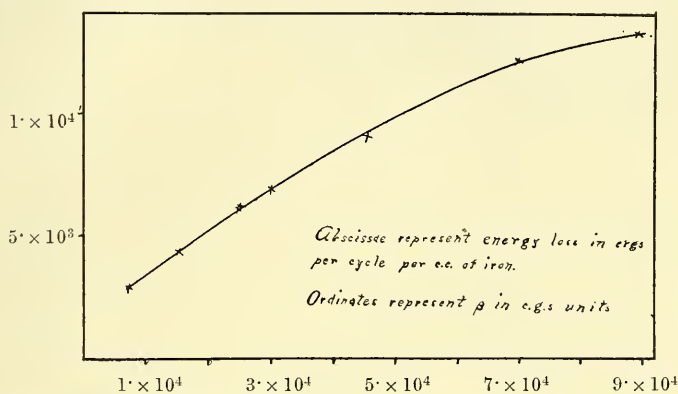
3-21-00 -115-	R_2	R_3	R_4	R_1	R'_1	H	T	D_1	D_2	r_1	r_2
1	913·2	2945·	1011·	148·3	313·5	165·2	·0277	3·62	3·18	166·7	59·72
2	505·9	2140·	1508·	148·3	356·5	207·8	·0280	5·85	7·40	166·7	59·72
3	208·6	1106·	2016·	148·3	380·2	231·9	·0275	8·78	14·70	166·7	59·72
5	1769·	2408·	507·5	148·3	372·8	224·5	·0276	9·00	3·00	166·7	129·72
6	1563·	2100·	507·5	148·3	377·7	229·4	·0273	10·61	3·86	166·7	129·72
8	1060·	1431·	507·5	153·0	375·9	222·9	·0267	4·00	6·95	377·0	129·72
9	503·1	1671·	1010·5	153·0	304·2	151·2	·0254	9·60	14·10	377·0	129·72
11	206·9	1639·	2014·	153·1	254·1	101·1	·0255	18·10	14·36	377·0	149·8

	B	C_1	H per c.c.	C_1^2HT
1	$2·88 \times 10^3$	·0262	3·815	$7·10 \times 10^3$
2	$4·32 \times 10^3$	·0337	4·799	$1·524 \times 10^4$
3	$6·17 \times 10^3$	·0412	5·356	$2·51 \times 10^4$
5	$6·07 \times 10^3$	·0418	5·185	$2·50 \times 10^4$
6	$6·86 \times 10^3$	·0454	5·298	$2·98 \times 10^4$
8	$8·92 \times 10^3$	·0576	5·148	$4·56 \times 10^4$
9	$1·207 \times 10^4$	·0891	3·482	$7·00 \times 10^4$
11	$1·318 \times 10^4$	·1226	2·335	$8·93 \times 10^4$

SUMMARY.

From the above it appears that the method described is a perfectly good method for the measurement of electric absorption. In all cases tried the electric absorption has acted as a resistance in series with a capacity. This resistance is independent of the current. The temperature has a decided effect. The value of this absorption increases very rapidly with rising temperature. The theory as given above appears to be verified

9



by the results in as far as a condenser possessing electric absorption may be considered as a capacity in series with a resistance, both of which depend upon the period of the current. While the variation is in the proper direction in each case, its rate of change with the period does not agree with the theoretical formula, especially in the case of the wax and paper condenser.

It also appears that the method given is a good one for the determination of the capacity of a condenser, which shows electric absorption. If electric absorption is corrected for, the capacity of such a condenser is a quantity which can readily be determined, and may be compared with a self-induction standard to 1 part in 1,000. It also appears that no correction is necessary for the electrostatic action of the turns of the standard coil on one another, at least if the relative size of coil and condenser is properly chosen, and if the resistance of the coil is not too large as compared with the total resistance in the branches of the circuit.

The method given for the detection of short circuits in coils proves to be an exceedingly sensitive one.

The method described for the measurement of losses due to hysteresis and Foucault currents gives very good results. Its

chief advantage lies in the fact that a very small quantity of the material is necessary. In the experiment only 360 grams were used. Much less than this amount could be used, with nearly as great accuracy as in the case given. For the value of the resistance H could be increased by decreasing the period of the current used.

In conclusion I wish to express my sincere appreciation of much assistance received from Professors Rowland and Ames during three years spent in study at the Johns Hopkins University. The investigation was suggested by Professor Rowland, and the methods used were those devised by him and described in the articles noted above.

Johns Hopkins University,
May 1, 1900.

ART. XIII.—*Some New Jurassic Vertebrates*; by WILBUR C. KNIGHT. Paper No. III.

FOR several years the Geological Department of the University of Wyoming has been directing its entire energy and what little money it could procure toward building up a collection of Jurassic vertebrates. On account of splendid success in the field the collection has already assumed considerable proportions and from time to time new animals are being discovered. The bulk of the material at the present time is Dinosaurus, but the marine beds of the Rocky mountain Jurassic have not been neglected and a few marine reptiles have been discovered. Among these are two new species of the order of Sauropterygia, which belong to two genera, neither of which have been reported from the Rocky mountain Jurassic; one of them being a Plesiosaur and the other a Cimoliosaur. With this addition to the fauna of the Eastern Rocky mountain region, there are now known two species of Ichthyosaurs* and three species of the Plesiosaur† type. The two skeletons which form the basis for this paper are not as complete as one would like to have them for generic and specific determinations, but there are sufficient remains in each case to give a very good idea of the species.

Plesiosaurus shirleyensis sp. nov.

The remains of this species consists of vertebra from all parts of the column; there being many from the cervical region; numerous teeth and fragments of teeth; a distorted portion of the lower jaw and many phalanges. Teeth large and numerous, incurved, elliptical in cross-section, interior surface of the teeth covered with numerous very fine angular striæ, exterior surface nearly smooth and showing faint marks of striæ. The greater portion of one side of the lower jaw measures .280^m in length and has a depth of .035^m. The vertebra are slightly biconcave, and all wider than long; but in the dorsals and posterior cervicals the length and breadth are nearly equal. The neural spines were found attached to the centra of the cervical vertebra, but the suture was obliterated. On the anterior caudles the neural spines are of considerable height.

Nothing of importance is known of the dorsal vertebrae, excepting that they are slightly biconcave and circular in

* *Baptanodon discus* Marsh; *Baptanodon nateus* Marsh.

† *Megalneusaurus rex* Knight. See this Journal, vol. v, p. 378.

transverse sections. Anterior caudles are flattened beneath and have two large circular facets for the articulation of the chevrons; neural arches firmly attached to centra. The basal phalanges have compressed, angular shafts, their shape depending largely upon their position in the paddle. Those in the central region approach a quadrangular section; while those on the exterior are V-shaped in cross-section with the opening of the V to the interior of the paddle. Terminations of phalanges rugose and strongly biconvex.

Measurements.

		M.
Length of fragment of lower jaw280
Depth of fragment of lower jaw035
Length of teeth about049
Transverse diameter of teeth at base of striations..		.006
Length of striated portion014
Dorsal vertebra.		
Length048
Width055
Height to neural platform045
Posterior cervical vertebra.		
Length050
Width055
Height to neural platform045
Height of process from neural platform090
Caudle vertebra (anterior).		
Length029
Width046
Height of neural platform034
Phalanges.	M.	
Basal.	large.	medium.
Length049	.045
Width, distal end022	.020
Width, proximal end023	.020
Width of shaft015	.012
Intermediate.		
Length035	.030
Width, distal end020	.018
Width, proximal end022	.019
Width of shaft013	.013
Terminal (nearly).		
Length025	.021
Width, distal end012	.012
Width, proximal end013	.013
Width of shaft008	.007

This animal had a long neck and large paddles; but did not attain a length of over fourteen to sixteen feet. It was dis-

covered in the Shirley* stage of the Jurassic rocks of Albany County, Wyoming, and was associated with *Baptanodon discus*, *Ostrea strigilecula*, *Camptonectes bellistriatus* and *Astarta Packardi*. Type specimen marked letter H in the collection of the University of Wyoming.

Cimoliosaurus laramiensis sp. nov.

This specimen consists of numerous vertebra and nearly a complete front limb. The humerus conforms to the general shape of Cimoliosaurs, but whether it belongs to the division where ulna and radius articulate with the humerus or to the other group where the pisiform is present, cannot be satisfactorily determined. There are two very distinct facets at the distal end of the humerus; the remainder of the margin is rounded and shows no sign of a pisiform. The facet for the radius occupies over one-half of the width of the expanded end, and is nearly flat. The humerus is a short, heavy bone, broadly expanded distally, and the post-axial border slightly recurved. Trochanteric ridge very prominent; shaft transversely elliptical with a rugose prominence on the anterior border, $\cdot 120^m$ below the head; apparently for the attachment of a large muscle; head spherical. Radius V-shape with the angle truncated and toward the ulna. The articulations of the ulna slightly convex; the proximal end very wide, distal narrow, and the ulna and radius evidently uniting without a central opening; exterior margin curved and very thin. Ulna not known. There are six carpal bones, all angular, but their relative position has not been determined.

Metacarpals and basal phalanges only slightly compressed and with flattened shafts, the upper surface slightly rounded, with a depression on either side, making a transverse section as figure in plate B, No. 1; they are also biconvex with pitted terminations. The terminal phalanges are much more compressed and without the depressions on the surface. All the vertebra wider than long and moderately biconcave, with neural arches firmly attached. Dorsal vertebra with a forward overhanging of the centra as is usually found in Cimoliosaurs, and circular in transverse section. Cervical vertebra numerous and anterior ones very small.

Caudle vertebra large, elliptical in transverse section. Anterior ones with large transverse processes, and large angular chevron facets.

*This is a new name applied to the Rocky Mountain marine Jurassic. See paper presented by me to the Geological Society of America, December 30, 1899.

Measurements.

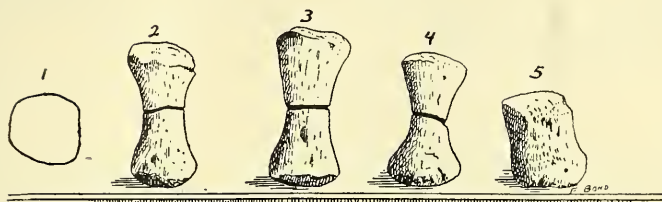
	M.
Length of humerus312
Width of humerus, distal end136
Thickness, maximum, of distal end041
Height of head100
Width of head068
Width of shaft humerus midway066
Thickness of shaft of humerus midway051
Radius.	
Length095
Width of	
} proximal articulation085
} distal articulation064
Thickness on interior edge033
A. Phalanx.	
Basal.	
Length053
Width of base030
Width of shaft023
Width at distal end027
Vertebra.	
Dorsal.	
Length037
Width045
Height of centrum040
Concavity of ends of centrum006
Anterior caudle.	
Length032
Width050
Height of centrum035
Concavity of ends of centrum006
Cervicals.	
Anterior.	
Length022
Width030
Height032
Posterior.	
Length026
Width041
Height032

This was a small animal, probably not more than twelve feet in length. It was found associated with *Baptanodon* remains in Shirley stage of the Freezeout Hills, Carbon County, Wyoming. This specimen is marked letter T in the collection of the University of Wyoming.

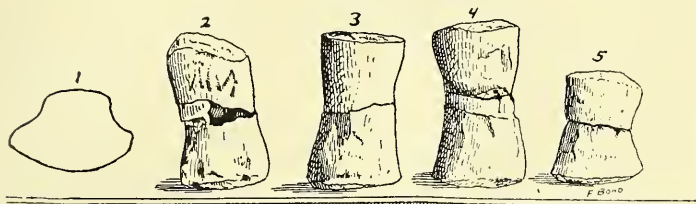
Geological Department,

University of Wyoming, April 6, 1900.

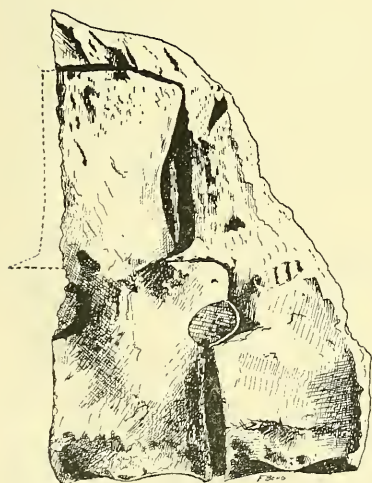
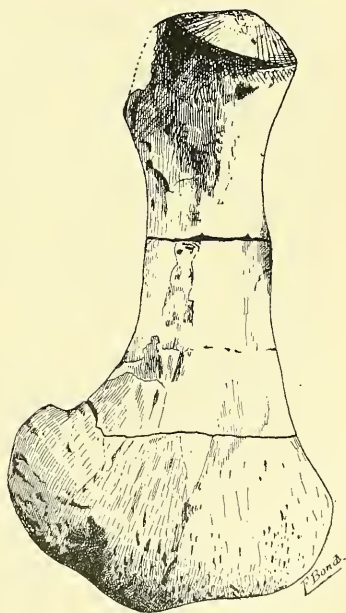
A



B



D



C

EXPLANATION OF FIGURES.

A and C.—*Plesiosaurus shirleyensis*.

A, No. 1.—Transverse section of a basal phalanx. ($\times \frac{2}{3}$)

Nos. 2, 3, 4, 5, basal phalanges. ($\times \frac{7}{16}$)

C.—Side view of cervical vertebra. ($\times \frac{2}{3}$)

B and D.—*Cimoliosaurus laramienseis*.

B, No. 1, transverse section of a basal phalanx. ($\times \frac{2}{3}$)

Nos. 2, 3, 4, 5, basal phalanges. ($\times \frac{7}{16}$)

D.—Humerus. ($\times \frac{1}{2}$)

ART. XIV.—*On Carnotite and Associated Vanadiferous Minerals in Western Colorado*; by W. F. HILLEBRAND and F. LESLIE RANSOME.

INTRODUCTION.

W. F. HILLEBRAND.

THE rather wide-spread occurrence in western Colorado of considerable quantities of a canary-yellow ore of uranium has been known for a few years past. It has been marketed to the extent of a few tons, but its mineralogical nature was unknown or incorrectly surmised until about a year ago.

In the spring of 1899 a specimen of this ore first reached my hands through Dr. Geo. P. Merrill of the U. S. National Museum, and was speedily recognized to contain a mineral or minerals new to science. Within a week appeared the announcement by MM. C. Friedel and E. Cumenge* of a new mineral, carnotite, a hydrous vanadate of uranium and potassium, obtained through Mr. Poulot of Denver, from Roc Creek, Montrose Co., Colo. Mr. Poulot had already identified vanadium in it. It was at once seen, despite certain differences in composition, that the two were identical.

According to the French authors the mineral is of simple composition, as above expressed, with only a little iron and mere traces of Al, Ba, Cu, Pb, and also according to M. and Mme. Curie, of the radio-active substances radium and polonium. The empirical formula $2U_2O_5, V_2O_5, K_2O, 3H_2O$, was assigned to it, some doubt attaching to the water.

Since then I have been able to examine carnotite from several localities in western Colorado, finding in each case the same lack of agreement with the analyses of Friedel and Cumenge.

While engaged in this work there were brought to my attention certain more or less greenish sandstones from the vicinity of Placerville on the San Miguel River, San Miguel Co., Colo., which were said to be highly vanadiferous and of considerable extent, and in which a zone a few inches thick was rather strongly impregnated with a yellow mineral resembling and probably identical with carnotite. This latter appeared also scattered through the sandstone at other points in sporadic small patches, sometimes only visible by aid of a lens.

This occurrence led to the thought that the carnotite bodies farther west might also be associated with existing or depend-

* Bull. Soc. Chim. de Paris (3), xxi, 328, 1899; Bull. Soc. Franc. Min., xxii, 26, 1899; Comptes Rend., cxxviii, 532, 1899; Chemical News, lxxx, 16, 1899. The papers as published in French differ slightly.

ent on preëxistent vanadiferous sandstones. For the carnotite of Montrose and Mesa Counties, as mentioned by the French authors, occurs mixed in all proportions with quartz-sand grains, the remnants beyond doubt of former sandstone bodies, and it was soon recognized that in the carnotite bodies the vanadium existed in two conditions, the larger part by far as pentavalent vanadium in the easily soluble carnotite, and a smaller and sometimes hardly distinguishable portion as trivalent vanadium in a much less soluble silicate which was free from uranium.

It was hoped that by a field reconnaissance, observations bearing on these points would be obtained and material assembled which would on analysis help to solve the nature and explain the association of these two entirely different classes of mineral substances.

The field and microscopical observations of Messrs. Ransome and Spencer are embodied in the pages immediately following these introductory remarks. Unfortunately their collections, except from Placerville, came to hand so late that the chemical work on the more western occurrences has been confined to the carnotite bodies alone, of which material was already in my hands. Greenish sandstones have been observed by Messrs. Ransome and Spencer in those regions, but whether any of them are highly vanadiferous, or what their connection with the carnotite may be, remains yet unknown. The only two examined did not owe their color to vanadium. Meanwhile the chemical results thus far obtained, which it is advisable to put on record at once, will be found in the concluding section of this paper.

In this place it is my pleasure to record my appreciation of the readiness with which the following gentlemen have supplied me with material for study: Messrs. Poulot and Voillequé of Denver, Mr. A. B. Frenzel of Placerville, and Mr. J. R. Duling of Paradox. These gentlemen have, also, not hesitated to furnish me with all information at their disposal as to occurrence, etc., of these interesting ore bodies.

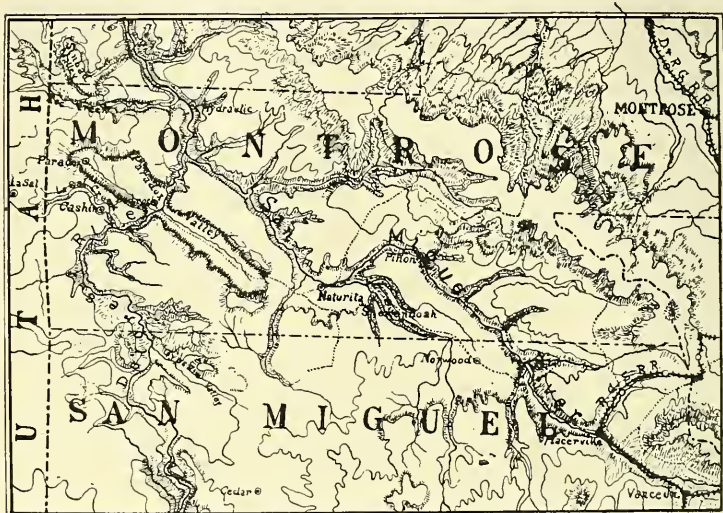
OCURRENCE OF THE URANIUM AND VANADIUM ORES.

F. L. RANSOME.

General.—In the autumn of 1899, a brief reconnaissance trip was undertaken into the western portions of San Miguel, Montrose and Mesa Counties, near the Utah-Colorado line. I was accompanied by Dr. A. C. Spencer, whose knowledge of the stratigraphy of the region was of great assistance. The primary object of the expedition in accordance with a suggestion from Mr. S. F. Emmons, was to investigate the

copper deposits of La Sal Creek, Paradox and Sinbad Valleys, and the vicinity, which had been responsible for some mining excitement a few months previously. A memorandum from Dr. W. F. Hillebrand, received just before starting, indicated that it would be well also to examine certain prospects on which some preliminary work had been done, looking toward the extraction of ores of uranium and vanadium. It is to the latter that these present notes are confined.

1



The most convenient way of reaching Paradox Valley was found to be the stage road, which, starting from Placerville, a settlement and station on the Rio Grande Southern R.R., runs by way of Norwood, Shenandoah and Naturita to Paradox, a distance of 60 or 70 miles.

The region can also be reached from the west by way of Moab, in Utah. The Placerville route and the general geographical relations are roughly indicated in the accompanying sketch map, fig. 1.

The topography of the region west of Placerville is that characteristic of the "mesa country" of western Colorado and southeastern Utah. Broad stretches of plateau are intersected by steep-walled cañons and cliff-encircled valleys. The underlying rocks comprise the "red beds" of the Dolores forma-

tion* (Triassic), the La Plata formation (Jurassic), the McElmo formation (Jurassic), the Dakota sandstone (Cretaceous), and the Mancos shales (Cretaceous). Carboniferous rocks occur in Sinbad Valley, and a series of gypsum-bearing shales of unknown age in the bottoms of Sinbad and Paradox Valleys, but as the ore-deposits to be described all occur in the beds of the La Plata and McElmo formations, these older rocks need not be again referred to. The sediments making up these various formations lie usually nearly horizontal, but they are sometimes flexed and frequently faulted.

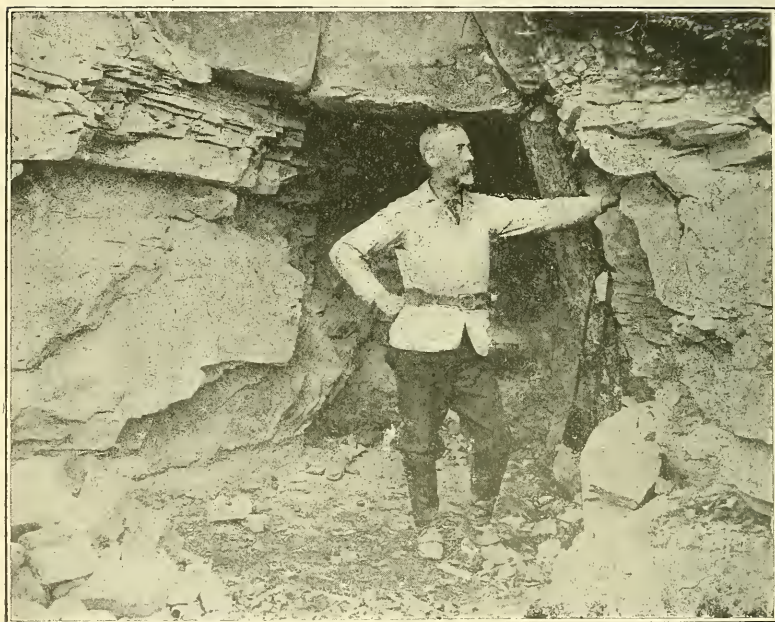
In all of the prospects examined, the ore of uranium occurs in the form of the recently described bright yellow carnotite. In one case this is intimately associated with a dull olive-green mineral which according to Dr. Hillebrand is either identical with, or very closely allied to, the vanadium-mica roscoelite. Deposits of one or both of these minerals occur widely scattered over San Miguel and Montrose Counties, Colorado, and in the Blue Mountain (Sierra Abajo) district of southeastern Utah; but a portion only of the known deposits were personally examined.

The Placerville Deposits.—These are essentially vanadium deposits, and occur 4500 feet nearly northeast of the railway station near Placerville, and about 1000 feet above the San Miguel River. The lower 900 feet of the San Miguel cañon is cut in the typical “red beds” of the Dolores formation. Above the Dolores comes the La Plata sandstone in typical development as described in the Telluride folios, viz., two heavy beds of light-colored sandstone, separated by a much thinner bed of dark limestone. The roscoelite occurs as an impregnation in the lower bed of the La Plata sandstone, about 100 feet above the base, and just beneath the bed of dark limestone. All the beds at this point are practically horizontal. The dark green vanadiferous sandstone occurs in a nearly continuous band, approximately parallel to the bedding planes, and varying in thickness from a few inches up to five or six feet. This band extends along the sandstone cliffs for an estimated distance of about 2000 feet. The roscoelite occurs more or less thoroughly impregnating portions of the fine-grained sandstone which constitutes the mass of the bed. It sometimes makes up more than 20 per cent of the vanadiferous facies. Normally the sandstone of this lower bed of the La Plata is light buff in color, with patches, mottlings, and stripes of pale pink, the latter color being apparently due to small amounts of ferric oxide. But when richly impregnated with roscoelite, this light-

* The formation names used in these notes are those adopted by Cross in the text of the Telluride Folio, of the U. S. Geol. Survey. The reader is referred to this folio for fuller descriptions.

colored sandstone becomes dark olive-green, the contrast being most marked when the cliffs are wet. Although fairly continuous for the distance stated, this zone of vanadiferous sandstone is by no means regular. It varies much in thickness, and in one place splits into two or more branches. At some points two or more distinct streaks of roscoelite-bearing sand were found at different horizons in the main sandstone bed. The carnotite is not nearly so abundant as the roscoelite at this locality. It occurs as minute yellow specks in the sandstone, and particularly as thin horizontal seams or streaks near the bottom of the vanadiferous zone.

2



The work thus far done on the several claims which have been located along these deposits is of the most superficial character. On the Canary claim a tunnel of about 18 feet has been run by Mr. A. B. Frenzel which exposes a typical section of the impregnated zone (fig. 2). The roof of the tunnel is formed by the underside of the bed of dark limestone previously referred to. This limestone is underlain by a few inches of sandy limestone which passes into the light-buff La Plata sandstone without break. The latter contains abundant calcite

as a cement or matrix for the sand grains. At a distance of a foot below the limestone, the sandstone shows, on close inspection, numerous specks of carnotite and gives a qualitative reaction for vanadium (Hillebrand). Examined in thin section under the microscope, this portion of the bed shows a fine-grained homogeneous sandstone, in which well rounded grains of quartz are held together by a rather abundant matrix of calcite. The latter is crystallized as a fine granular aggregate and includes small indeterminable particles of various kinds. Many of the quartz grains as seen in section are entirely surrounded by calcite, and so isolated from adjacent grains. A pale lemon-yellow substance occurs sporadically in the section as an interstitial material between the quartz grains. This is probably the carnotite which was seen as small yellow specks in the hand specimen. It is very minutely crystalline, recalling the habit of some obscure chloritic aggregates in decomposed igneous rocks, and is too indistinct for successful optical study. Small crystal grains of zircon are scattered through the sandstone, and are readily concentrated in the residue after treating with hydrofluoric acid.

In the next two feet below where it exhibits the foregoing facies, the sandstone becomes pinkish in color, due to the presence of ferric oxide, and the lens fails to reveal any roscoelite or carnotite. But just below the pink coloration, the sandstone begins to show yellow and green specks. The latter especially become more numerous and larger, until at from 3 to 4 feet below the limestone the sandstone has a decided green tint. This deepens, going downward, until near the floor of the tunnel the sandstone is a deep uniform olive-green, rich in roscoelite (over 20 per cent) and showing many small yellow specks of carnotite. This is regarded as the first-class ore, and the chief value of the deposit is supposed to be in its vanadium, the uranium being very subordinate. In this respect the Placerville deposit differs from those on La Sal and Roc Creeks, presently to be described, where the mineral sought for and occurring most abundantly is carnotite.

Thin sections of the green-spotted sandstone, when examined microscopically, resemble those of the light-buff sandstone just above it, as far as character of quartz grains and relative abundance of matrix are concerned. But the character of the matrix or cement is different. Calcite is much less abundant and its place is largely taken by roscoelite. This is grass-green in thin section, and might readily be mistaken for indistinct wisps and areas of chlorite. It sometimes forms a distinct envelope around the quartz grains, showing an indistinct foliated or fibrous structure normal to the surface of the latter.

The uniformly dark green sandstone richest in roscoelite does not effervesce with acids and shows no calcite cement in thin section. The usual quartz grains are held together by roscoelite. This is crystalline, but the highest available powers show only an indistinct and minute foliation, such as may be observed in some very finely crystallized chlorites.

Close to the floor of the tunnel is a fairly regular, nearly horizontal streak of carnotite, varying in width, but usually less than an inch, and showing noticeable diminution in thickness in the face of the tunnel. This small seam is not solid carnotite, but is merely a zone in the sandstone impregnated with the bright-yellow uranium mineral. It is not nearly so continuous as the main vanadiferous belt, and was seen only at three or four places along the 2000 feet or so of outcrop of the latter. When this seam is closely examined, it is seen that narrow bands rich in carnotite alternate with green bands carrying mostly roscoelite. There is also usually present a seam generally about an eighth of an inch in thickness (though often thicker) which is almost wholly quartz. The microscope shows it to be a true quartzite, in which the original rounded detrital grains of quartz have been cemented by fresh quartz in optical continuity with the older granules. A similar quartzite occurs in the green vanadiferous sandstone above the carnotite, where it forms concretionary knots and nodules. It was not noted in the buff sandstone however, where the cementing material is calcite.

Immediately below the carnotite seam there is a parting or "floor" in the sandstone, probably originally a very thin layer of shale, which forms the working floor of the tunnel. The sandstone below this floor is plainly impregnated for a short distance with roscoelite, but the thin shale seam is regarded as the practical bottom of the deposit.

The questions of the origin and actual extent of this deposit are closely related, and of much interest. Their discussion will be deferred until the other deposits visited have been described.

Some distance below the Placerville vanadium deposit, sandstone, presumably belonging to the Dolores formation, was observed to be colored green, of a somewhat brighter hue than the vanadiferous sandstone higher up the slope. As there are some copper prospects near by, from which ore has been taken, this was supposed to be a copper stain. Qualitative tests by Dr. Hillebrand show however that the color is due to a compound of chromium.

Similar green sandstones occur on the western side of Sinbad Valley in what is apparently the La Plata formation, and were originally supposed to be impregnated with roscoelite. Dr. Hillebrand's investigations, however, show that they too owe

their color to some chromium mineral. These occurrences are interesting as showing that a green color in sandstones may result from various causes, and that even a bright green tint cannot be taken as an infallible indication of copper.

The La Sal Creek deposits.—These occur in the extreme western portion of Montrose County, southwest of Paradox, and about six miles up La Sal Creek from Cashin P. O. They are reached by trails from Paradox Valley and from Cashin. The deposits are on the south side of La Sal Creek and about 700 feet above the stream. They occur for a distance estimated at more than a quarter of a mile, along the sandstone cliffs which descend from the mesa into the cañon of La Sal Creek, and only a few feet below the level of the mesa surface.

In the absence of continuous stratigraphic work, it is impossible to correlate certainly and finally the rocks on La Sal Creek with the divisions established by Cross and Spencer in the Telluride quadrangle to the eastward. It seems probable, however, that the La Plata sandstone attains a much greater thickness in this portion of western Colorado than it does between Telluride and Placerville. The limestone bed, so characteristic a feature of the formation near Placerville and further east, is not uniformly present in this western region, and the La Plata sandstone (Variegated Beds, in part, of the Hayden Survey) is not always readily differentiated from the underlying Dolores formation (Red Beds). For a vertical distance of about 400 feet above the bed of La Sal Creek the rock is a heavy-bedded, rather fine-grained, light-colored sandstone, which is considered by Dr. Spencer to be the La Plata. Above this come thinner-bedded sandstones, with some conglomerates and shales, which are included in the McElmo formation. It is in this upper series that the uranium deposits occur. All the beds are here approximately horizontal.

As revealed by numerous small openings near the crest of the bluff, the carnotite, which is the material here sought, is found chiefly in a massive bed of nearly white sandstone. Some of the ore, however, lies between the sandstone and a lower bed of light-gray shale. Although the prospecting openings all lie at about the same level along the cliffs, the deposit is not nearly so regular as the vanadiferous band near Placerville.

The carnotite of La Sal Creek occurs as irregular, bunchy "pockets" in the sandstone, or along the contact of the sandstone with the underlying shale. These have all the appearance of being impregnation deposits, the solutions carrying the uranium compounds having deposited the ore wherever they found ready passage through the rock—usually along bedding planes. No roscoelite was detected with the carnotite.

The most remarkable and interesting fact in regard to the La Sal Creek deposits is their very superficial character. The ore bodies are usually flat-lying streaks, a few inches thick, which grade above and below into the common light-buff sandstone, and which die out and disappear when followed into the hillside. In tunnels run but a few feet underground the yellow impregnation of carnotite can be seen to gradually die out, to be succeeded by light-colored sandstone, showing no apparent trace of the mineral. It is doubtful whether any appreciable quantity of carnotite occurs as much as 20 feet from the surface, on any of the locations, although this distance is given from memory and not from measurements on the ground. As before stated, the impregnation has usually taken place along bedding planes; it has also proceeded along surfaces of minor and superficial movement in the rocks. In one case it was observed that a portion of the overlying sandstone had moved upon the underlying shales, the disturbance being apparently a superficial one, of a kind commonly enough observed where massive beds rest on yielding shales on a steep hillside. In other words, the movement appeared to be directly related to the present topography. The deposition of carnotite was here plainly subsequent to the movement and had taken advantage of the small openings and dislocations in the shale afforded by this very recent disturbance. It was reported that some of the best nests of ore had been found in "slide rock," i. e., rock which had slipped to some extent down the slope; but I was unable to verify this statement further than is indicated in the preceding description. There can be little doubt but that the deposits of carnotite on La Sal Creek are not only very superficial in character but very recent in age.

The Roc Creek Deposit.—The principal claim is the Copper Prince, owned by J. R. Duling, on the north side of Roc Creek, 3 or 4 miles above its mouth, and near the foot of the Miller trail to Paradox. It is reached by this trail from Paradox Valley, or by trail from Hydraulic on the Dolores River. This deposit is in the La Plata sandstone according to Dr. Spencer's observations, and has been more extensively exploited than any other seen. The sandstone, which at this point is nearly horizontal, is cut by an east and west fault, the fault plane dipping north at about 70°. The amount and character of the throw could not be determined. It is probably less than 50 feet. The carnotite occurs in the hanging wall of the fissure as small irregular branches in a loose mass of crushed sandstone and also as an impregnation of some of the firmer portions of the bed. No roscoelite was seen. Several small tunnels have been run in on the hanging wall side of the fissure (which itself carries no vein or ore-body), but they had

been securely closed by planking, and no examination of their inner ends could be made. I am thus unable to tell what becomes of the deposit as it is followed in from the surface. A few hundred pounds of picked carnotite ore has been shipped from this claim and is reported to have sold for \$1.25 a pound in Denver. The character of this deposit is similar to others examined, except that in this case a well-defined fault has provided a zone of crushed and porous rock in the hanging wall, along which impregnation could take place. A few hundred feet farther west the crushed sandstone adjoining the fault has been impregnated with cupriferous solutions and is spotted with stains of the blue and green carbonates of copper.

Other Deposits.—It is known that several carnotite claims have been taken up in Gypsum Valley in what is known as the Disappointment District. The impregnated sandstone is said to cap a hill and to constitute an extensive deposit, but it was not visited. Carnotite is also reported from the Blue Mountain District, but I have no personal knowledge of these occurrences. It seems highly probable that the material will be found widely distributed in the Mesozoic sandstones of western Colorado and eastern Utah, although perhaps nowhere in very extensive bodies.

Origin of Deposits.—That the deposits of carnotite and roscoelite were formed subsequently to the deposition of the sandstones is evident from the facts presented in the preceding pages. It is equally plain that the minerals could not have resulted from the alteration, in place, of other compounds of vanadium and uranium originally deposited with the sands. The shape and position of the deposits indicate clearly that the ores have been deposited in their present position only after transportation from a greater or less distance. Moreover, the recency of the deposits and the fact that they are sometimes directly connected with faults and dislocations in the sandstones shows that the vanadium and uranium compounds could not have been the original cementing material of the quartz grains, but have in all probability locally replaced the calcite which acts as matrix to the ordinary light-colored sandstone in which the ore-bodies occur. The deposits of roscoelite appear to be comparable to the impregnations of the sandstones with cupriferous solutions observed in many places in this region, particularly on La Sal Creek near Cashin and in Sinbad Valley, whereby the sandstone becomes colored bright green with the carbonate of copper. In these cases, however, the copper appears to have been previously deposited in part as chalcocite. An analogy might also be drawn with the green chromiferous sandstone near Placerville already referred to. In all these cases the actual sources of the materials which have

been deposited by solutions in their present position are not known. Dr. Hillebrand has shown,* however, that vanadium in small amounts is widely distributed in sandstones, limestones, and igneous rocks. It is perhaps present in very small amounts in the bulk of the sandstone, and the deposits described may simply represent a concentration of this material under certain favorable conditions of solution and re-deposition.

In the absence of exploitation it is manifestly impossible to predict the probable shape and size of ore-bodies formed in this manner. The roscoelite as seen near Placerville appears, however, to be much more persistent than the carnotite. There is no apparent reason why a mass of sandstone, impregnated with roscoelite, which is continuously exposed for several hundred feet along a cliff, should not extend for a considerable distance inward from the cliff face. The carnotite, on the other hand, appears to be a much more superficial occurrence and, in fact, to have a not yet fully understood connection with the present surface of the ground. This would indicate that the carnotite results from a local concentration of material already existent in the sandstone, and the deposition of this material in the form of carnotite under conditions determined by proximity to the surface, and probably partly dependent upon a semi-arid climate.

CHEMICAL ANALYSES AND DISCUSSION.

W. F. HILLEBRAND.

I. *The Green Coloring Constituent of the Placerville Sandstone.*

As noted by Mr. Ransome, the greener the sandstone, the greater has been the replacement of the calcareous cementing material by the crystalline vanadiferous mineral. Fortunately the greenest available specimen was entirely free from carbonate, and this was used for analysis, 10 grams being taken for the main portion.

The green substance was but slightly attacked by cold acids and not rapidly by hot ones; nevertheless after 12–24 hours digestion on the water-bath with diluted nitric acid, the greenish color had been transferred to the liquid, the residue of sand was nearly white, and a rather voluminous separation of non-gelatinous silica had taken place. The filtration and washing of the insoluble matter was accomplished without difficulty. The free silica in it was extracted by two or three digestions of 15 minutes each on the water-bath with sodium carbonate solution of 5 per cent strength. The flocculent matter had entirely disappeared and the sandy residue was collected in a

* This Journal, IV, vol. vi, pp. 209–216, 1898.

Gooch crucible and washed with hot sodium carbonate solution, followed by very dilute nitric acid and finally by alcohol to prevent turbid filtration. Drying of the sand was effected by allowing the pump to draw air through the crucible until no further loss in weight resulted.

The water in the sand was then determined at 105°, at 300° and above 300°. Similar water determinations having been made on the unattacked substance, the differences gave the water expelled at those temperatures from the soluble part.

The sodium carbonate solution was acidified, evaporated, and filtered, and the filtrate again evaporated, to recover the last of the silica. To this was added a very small portion which the nitric acid extract of the sandstone held in solution. The sum represents the total silica of the decomposed portion of the sandstone.

The further analysis of the nitric solution presented difficulty because of the vanadium it contained. It was carried out in a variety of ways, of which the following seemed to be most satisfactory.

Barium was first precipitated by sulphuric acid and subsequently separated from traces of lead, calcium and vanadium. Hydrogen sulphide then threw out further traces of lead and copper with much sulphur from reduction of V_2O_5 to V_2O_4 . After evaporation of the filtrate to reoxidize vanadium and iron, the alumina, iron and trace of uranium with much vanadium were separated from calcium, magnesium and alkalis by three precipitations by ammonia solution. The combined filtrates were evaporated, ignited, the residue transferred to a porcelain boat with nitric acid, evaporated therein to complete dryness, and exposed in a glass tube to a current of dry hydrochloric acid gas, as recommended by Smith and Hibbs* for the expulsion of vanadium from alkaline vanadates. The removal of vanadium being incomplete, the contents of the boat were again evaporated with nitric acid and again distilled, and these operations were repeated till no further evidence of a brown distillate appeared. The boat now contained only magnesium and alkalis with a little calcium. The distillates, containing all the vanadium that was in the boat, had been collected in U-tubes charged with water, and were eventually obtained in sulphuric solution for further treatment by evaporating with sulphuric acid in porcelain.

The precipitate of Al_2O_3 , etc., was dissolved in nitric acid, evaporated almost to dryness in platinum and boiled with sodium hydroxide solution. These operations were repeated on the precipitate after filtration. The iron and titanium thus

**Jour. Am. Chem. Soc.*, xvi, 578, 1894.

purified were redissolved, precipitated by ammonia, weighed, redissolved by acid potassium sulphate, the ferric iron was reduced by hydrogen sulphide and titrated with permanganate after boiling out the reducing agent. The titanium was then found colorimetrically by hydrogen peroxide.

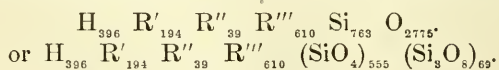
The combined alkaline filtrates were heated with excess of ammonium nitrate to separate alumina from most of the vanadium. The precipitate was redissolved in nitric acid, reprecipitated by ammonia, then ignited and weighed and redissolved by acid potassium sulphate. The sulphuric solution was reduced by hydrogen sulphide, boiled and filtered from platinum sulphide, again boiled in a current of carbon dioxide and titrated hot with permanganate. Sulphur dioxide gas was then introduced, boiled out in a current of carbon dioxide, and the titration repeated, this second result being taken as representing the V_2O_4 equivalent to the V_2O_5 that still contaminated the Al_2O_3 . Deducting the V_2O_5 thus found from the combined weight of Al_2O_3 and V_2O_5 gave the Al_2O_3 of the soluble constituent of the sandstone.

The combined filtrates from the alumina were evaporated, ignited to remove ammoniacal salts, the residue converted to sulphates, and united with the earlier solution of vanadium sulphate. The vanadium in it was then determined precisely in the manner already described, and when added to that found with the alumina and calculated to V_2O_5 , gave the total in the sandstone.

A check was made by dissolving the sandstone in hydrofluoric and sulphuric acids in a current of carbon dioxide and titrating the V_2O_5 with permanganate. The figure thus obtained was a trifle below that found as above detailed, and this is to be attributed to the presence of a very little vanadium as V_2O_5 in carnotite, and probably as a lead vanadate. All the uranium and most of the trifling amount of lead could be extracted by cold dilute nitric acid, thus showing them to be foreign to the green substance.

It is preferable to regard iron as in the ferric state rather than as ferrous, otherwise an equivalent amount of vanadium would have to be considered as V_2O_5 , for which there is no occasion. It is quite possible that the iron is in part, if not wholly, foreign to the green matter, for it would be surprising to find an exposed sandstone free from ferric oxide.

If the iron is disregarded as probably extraneous, the ratios given below afford the following empirical formula :



Simplified this is $H_{40} R'_{20} R''_4 R'''_{61} X_{62.4}$.

Considering that the mineral is probably not perfectly pure and the uncertainties affecting the amount of water to be considered, etc., the above affords a close approach to the typical phengitic muscovite formula H_2KR''', X_3 in which X comprises the groups SiO_4 and Si_3O_8 .

Analysis of the vanadiferous sandstone.

		Sandstone.		Silicate.		
		Per cent.	Ratios.	Per cent.	Roscoelite.	
Soluble in Nitric acid.	{ SiO_2 -----	12.56	.763	46.06	45.17	
	{ TiO_2 -----	.02			.78	
	{ V_2O_5 -----	3.50*	.0846	} .305	12.84	24.01
	{ Al_2O_3 -----	6.15	.2206		22.55	11.54
	{ Fe_2O_3 -----	.20	.0046		.73 FeO	1.60
	{ CaO -----	.12	.0078	} .039	.44	
	{ BaO -----	.37	.0088		1.35	
	{ MgO -----	.25	.0228		.92	1.64
	{ K_2O -----	2.41	.0938	} .097	8.84	10.37
	{ Na_2O -----	.06	.0035		.22	.06
	{ H_2O at 105° --	.54	.1100		1.98	.40
	{ H_2O 105° - 300°	.14	.0283		.51	.17
	{ H_2O above 300°	.97	.1978	.198	3.56	4.12
	{ UO_3 -----	.05				
	{ PbO -----	.06			100.00	99.86
{ V_2O_5 -----	.05					
Insoluble in Nitric acid.	{ Quartz, etc. --	72.24†				
	{ H_2O at 105° --	.04				
	{ H_2O above 105°	.20				
		99.93				

Traces of Li, Cu, Mo, Bi. No Cl, SO_3 , or P_2O_5 .

An attempt to satisfactorily distribute the constituents among different chemical molecules would be futile, since, for instance, we do not know whether to ascribe the barium to a biotite or muscovite molecule, but the close approach to the muscovite ratio leaves little room to doubt the character of the body. Notwithstanding the chloritic aspect of the mineral under the microscope, the analysis shows that it cannot be a chlorite, and we have undoubtedly to do with a body closely related to the mica roscoelite, wherein, however, the percentage proportions of Al_2O_3 and V_2O_5 are reversed, thus affording a further striking example of the mutual replaceability of these two oxides.

* Oxidation by permanganate of the HFl and H_2SO_4 solution of the sandstone indicated 3.43 per cent V_2O_5 (mean of 3.50, 3.32, 3.48, 3.43).

† Containing about 65 per cent SiO_2 , 7 per cent Al_2O_3 (Fe_2O_3 , TiO_2) and about 0.3 per cent of zircon and other minerals not decomposed by repeated evaporations with HFl and H_2SO_4 .

For comparison, the latest analysis of roscoelite from Placerville, Calif., is given a place in the above table. Peculiar, though unimportant, is the coincidence that the only two known localities for this mineral should bear the name Placerville.

II. *Green Sandstone Colored by Chromium.*

Other sandstones much brighter green than the vanadiferous one were observed and collected by Dr. Ransome, both at Placerville and 60 miles distant in Sinbad Valley. The color suggested a salt of copper as its cause, but analysis showed it to be due to chromium. Time has failed for an examination as to the nature of this coloring body. It is very difficultly soluble, thus presenting greater hindrance to analysis than was the case with the vanadium compound. It would be interesting to find it to be a micaceous mineral analogous to the one just described. Under the microscope it presents a chloritic appearance (Ransome). If opportunity offers, the problem of its nature may yet be attacked.

In still another greenish sandstone from the west bank of the Dolores River, near the mouth of La Sal Creek, analysis failed to show either vanadium or chromium.

III. *The Carnotite Ores.*

Optical Examination.—Dr. Geo. P. Merrill has kindly submitted the following notes.

“The carnotite powder appears under the microscope in the form of exceedingly minute dust-like particles, without crystal outlines and acting so faintly on polarized light as to at first seem almost amorphous. Much of the matter appears merely as a fine brownish clay, stained yellow by an amorphous pigment, but occasionally a well-defined fragment of a light yellow, translucent mineral is met with which doubtless represents the vanadium compound in its condition of ideal purity.

Working over a considerable amount of the powder, I have found occasional clusters of this yellow mineral in the form of flattened radiating crystals with pyramidal terminations which are without evident pleochroism, polarize only in dull colors, and give extinctions always parallel to the axis of elongation. These are so minute (not over 0.25^{mm} in length) and so thin that I have never been able to find a crystal so oriented as to give an opportunity of determining its exact character, and I can only say that the general shape is such as to suggest a hexagonal mineral, though this is by no means certain.”

Chemical Examination of the Carnotite Ores.—The chemical problems involved in the analysis of the carnotite ores were peculiarly intricate. Ideal material was quite unobtainable

and mechanical separation impossible. The first specimen received happened to be of a higher grade than any of the subsequent ones, containing about 5 per cent of sand grains and showing only very faint delicate reddish tracings indicative of some foreign iron mineral. In mass it was of a beautiful canary-yellow color and easily broken down by pressure. Other specimens were more coherent, the degree depending altogether on the extent to which the sandstone had been impregnated and altered, and they sometimes showed more of the peculiar reddish admixture.

A fact only suspected in making the first analysis was confirmed by subsequent work on lower grade ores, namely, that the vanadium existed in two conditions, in entirely distinct minerals; the greater part by far as pentavalent vanadium in the easily soluble carnotite, and a smaller, much less soluble portion, almost vanishing in the purest ores, in the trivalent state as a constituent of a silicate free from uranium. This observation explained the statement of Messrs. Poulot and Voillequé that they had found the low grade ores to be relatively richer in vanadium as compared with uranium than the high grade ores.

While the carnotite dissolves at once in cold dilute nitric acid, unfortunately the vanadiferous silicate is not quite insoluble, hence arose an important difficulty in the way of arriving at the true composition of the carnotite. It is true that in one case the silicate has been analyzed (p. 143), but it would be unsafe to correct the carnotite analysis on that basis. Moreover, the analyses were not all made on the same plan; some are less complete than others; they cannot in some respects be rigidly compared with each other.

As to the carrying out of the analysis, various procedures were tried and no one found which gave altogether satisfactory results as to each constituent, though many could be determined with the usual degree of accuracy. The difficulty was mainly due to vanadium and the small amount of phosphoric acid usually present. These two constituents were likely to be found in different precipitates and could not be separated from them at one stroke. Their complete removal from other bodies was at times impossible and the weight of the latter had then to be corrected for these residual amounts.

Two entirely different lines of attack were open. One, that of Friedel and Cumenge, described in their paper on carnotite, had already been used with apparent success. It depends on rendering the vanadium insoluble in water by evaporating the nitric acid solution to dryness. Water extracts the alkalies and uranium without dissolving the vanadium, iron, or aluminum. Repetition of this process yielded Friedel and Cumenge satis-

factory results. They recommended washing with water containing ammonium nitrate.

The other method contemplated the quantitative removal of vanadium by dry hydrochloric acid gas as already described (p. 131). This had been shown to be perfect for alkaline pyrovanadates by Smith and Hibbs and it was hoped might succeed even with such complex mixtures as the present. The distillations were made on the crude ore and also on the nitric acid solution after filtration from the insoluble matter and evaporation to thorough dryness. The action is immediate in the cold, copious red-brown vapors coming off and condensing in part as a dark red liquid in the tube. But repeated distillations and the application of heat were required to effect complete removal of the vanadium accompanied by the arsenic and molybdenum. After each distillation the contents of the boat had to be evaporated with nitric acid. The blue and green colorations apparent on adding this acid showed that the hydrochloric acid gas had reduced a good deal of the vanadium to a lower state of valence, and in this condition it was incapable of forming the volatile body.

A somewhat more satisfactory separation sometimes resulted when this distillation method was combined with that of Friedel and Cumenge by subjecting both the evaporated uranium nitrate solution and the residue insoluble in water to the action of hydrochloric acid gas.

The distillates obtained by either way were evaporated with sulphuric acid, the arsenic and molybdenum were separated by hydrogen sulphide, and the vanadium was then titrated by permanganate at a temperature near boiling and again after reduction by sulphur dioxide gas. In one or two cases when the temperature of distillation had been high and it was feared some iron had passed over, the distillates were evaporated in porcelain with nitric acid, transferred to a platinum crucible, evaporated therein with sulphuric acid and fused with sodium carbonate. The aqueous extract was then treated as above for arsenic, molybdenum and vanadium.

The separation of Friedel and Cumenge, while perhaps adequate for technical purposes in ores free from phosphorus,* does not in my hands give perfect satisfaction even then. It is impossible to prevent a little of the vanadium, also of the iron and aluminum, from going with the uranium, and on the other hand a little uranium may stay with the vanadium.

* With even only half a per cent of P_2O_5 many times that amount of UO_3 is rendered insoluble in water after evaporation to dryness with nitric acid. The compound formed is of a lemon-yellow color, which is masked by the separated vanadic acid until this has been removed by ammonia. This insoluble body was treated as follows in order to arrive at the P_2O_5 and UO_3 it contained. After solution in nitric acid the phosphorus was precipitated by ammonium molybdate, and from the filtrate the uranium by three precipitations by ammonia.

Again when the vanadic acid is extracted from the residue by ammonia, as prescribed by Friedel and Cumenge, a little of the other constituents of the residue accompany the vanadium into solution.

If the combination of the two methods is used, it is better after removal of the vanadium by distillation to further treat the two residues separately at first instead of to unite them at once. Full details of this treatment are unnecessary and would unduly extend this paper. Suffice it to say that from the uranium portion the little iron and aluminum present are separated by ammonium sulphide and carbonate, and after removal of the latter, and acidification, the uranium can be thrown down by ammonium sulphide, and then by at least two precipitations by ammonia, or far better, by precipitating the neutral hydrochloric solution by freshly precipitated and alkali-free mercuric oxide at boiling heat, as prescribed by Alibegoff.* As pointed out by von Foullon and also by Alibegoff, contrary to certain still widely disseminated statements, ammonium sulphide does not afford a good separation of uranium from calcium. This is perhaps especially true if the solution contains any phosphorus. Again, contrary to another statement, it seems perfectly possible to separate uranium completely from alkalis by a few ammonia precipitations.

The finally ignited and weighed U_3O_8 was redissolved in nitric acid, filtered if necessary (SiO_2 , Al_2O_3 , Fe_2O_3), and divided into two parts of which the one was tested for the very little P_2O_5 usually present, and the other for vanadium by conversion into sulphate, reduction by sulphur dioxide gas, and titration by very dilute permanganate solution.

Table of Analyses of Carnotite Ores.

I. Copper Prince Claim, Roc Creek, Montrose Co., Colo.

According to the donor of the specimen, Mr. J. R. Duling, I-a is from the same lot as that from which Mr. Poulot obtained the specimens afterwards analyzed by Friedel and Cumenge.

II. Yellow Boy Claim, La Sal Creek, Montrose Co., Colo.

III. Yellow Bird Claim, La Sal Creek, Montrose Co., Colo.

The last two claims belong to the same group and the ore is from the same "blanket" (Voillequé).

I-b and II-b, the first analyses made, are not strictly comparable with the other analyses since they represent the effect of warm dilute hydrochloric acid, whose greater action is shown by the nearly complete solution of the ferruginous admixture. Cold, dilute nitric acid was used for the other analyses.

* Ann. Chem. u. Phar., vol. cxxxiii, 133, 1886; Zeit. für anal. Chemie, xxvi, 632, 1887.

	I.			II.		III.
	a	b	c	a	b	
Insol.	7.10*	8.34†	19.00‡	10.33§		¶
UO ₃	54.89	52.25	47.42	54.00	52.28	20.51
V ₂ O ₅	18.49	18.35	15.76	18.05	17.50	7.20
P ₂ O ₅80	.35	.40	.05	tr.	none
As ₂ O ₅	tr.	.25	none	none	none	none
Al ₂ O ₃09	?	.08	.29	?	.08?
Fe ₂ O ₃21	1.77	.72	.42	3.36	.25
CaO	3.34	2.85	2.57	1.86	1.85	1.64
SrO02	?	?	tr.	tr.	?
BaO90	.72	.65	2.83	3.21	.29
MgO22	.20	.24	.14	.17	.07
K ₂ O	6.52	6.73	6.57	5.46	5.11	1.51
Na ₂ O14	.09	.07	.13	.02?	.01
Li ₂ O	tr.	?	?	tr.	?	?
H ₂ O 105°	2.43	2.59	1.85	3.16	4.52	} total H ₂ O in ore 1.85
H ₂ O 350°	2.11	3.06	2.79	2.21	3.49	
H ₂ O + 350°	none	none	none	none		.19 (+ 300°)
PbO13	.25	.18	.07		.09
CuO15	.20	.22	tr.		tr.
SO ₃	none	.12	.18	none		none
MoO ₃18	.23	.18	.05		.04
SiO ₂15	.06	.13	.20		.07
TiO ₂03	.10	?	?		.06?
CO ₂56	.33	none	none		
	98.46	98.84	99.01	99.25		

* Containing .54H₂O, .09V₂O₅.

† The insoluble matter had the following composition: SiO₂ 5.18, V₂O₅ .21, P₂O₅ .09, K₂O .26, Na₂O .04, H₂O (105°) .56, (350°) .32 (+ 350°) .48, Al₂O₃, TiO₂, Cr₂O₃, etc., by diff. 1.20. The SO₃ of this analysis is not combined with BaO, for it is wholly extracted by dilute acids. Once in solution the acid used is sufficient to prevent its immediate precipitation as BaSO₄ by union with some of the barium present. The same holds true for the SO₃ of I-c.

‡ This material was obtained by floating off the finer matter, allowing it to settle, collecting on a Gooch filter, and drying it in a current of air drawn through the crucible. The insoluble matter held in addition to 16.41 quartz and silicates including .39 V₂O₅ and a little UO₃, H₂O (105°) .83, (300°) .73, (+ 300°) 1.03; total H₂O 2.59.

§ Containing .16 V₂O₅ and 1.90 H₂O.

|| The insoluble matter contained besides quartz and silicates .25 V₂O₅, .21K₂O, .05 Na₂O.

¶ This analysis was made purposely on a relatively poor ore, furnished by Messrs Poulot and Voillequé, with the object of determining, if possible, the composition of the vanadiferous silicate which it contained (see p. 143). The data for calculating the H₂O values of both analyses are as follows:

	Ore.	After extraction of carnotite by cold, dilute nitric acid.	After extraction of residue by hot nitric acid, sod. carb., etc.
	(a)	(b)	(c)
H ₂ O 105°	3.53	1.68	.02
H ₂ O 300°	2.11	.47	.03
H ₂ O + 300°83	.64	.02
	6.47	2.79	.07

a-b furnishes the values for the carnotite as shown in analysis III; b-c gives those for the less soluble silicate (p. 143).

Discussion of the Carnotite Analyses.—It will be noted that a somewhat marked deficiency appears in most of the analyses, the cause of which is quite unknown. Great care was exercised in most cases and especially in those which show the greatest loss. It seems hardly possible that any serious constant loss of a known constituent should have occurred, but the only alternative demands the presence of an element or elements unnoticed and which cannot have been weighed with the known constituents. The researches of M. and Mme. Curie have shown that these ores contain traces of radio-active elements, precipitated the one by hydrogen sulphide, the other by sulphuric acid. Their presence, however, in quantity sufficient to account for the observed losses in the above analyses, especially when 10 grams of ore were operated on, could not possibly have escaped observation. To whatever cause it may be due, this loss alone suffices to render somewhat uncertain any calculations based on the analytical figures, though if the loss is to be ascribed to uranium or vanadium the ratios would not be sufficiently affected to obscure any simple relations that might exist.*

Another difficulty is the impossibility of knowing what constituents to exclude and what to include in deducing molecular ratios. It is certain that most if not all of the iron is foreign to the yellow body. It is probable that phosphorus is likewise so, since its extraction by dilute acids does not keep pace with that of the uranium and vanadium. It may possibly be in combination with the iron, in part at least. The alumina doubtless is derived from the vanadiferous silicate which seems to exist in all the ores and which is not quite insoluble in cold dilute acids. If so, a small portion of the vanadium, potassium, magnesium, and water are to be attributed to this mineral, but a general correction based on the analysis of this compound (see p. 143) would not be justifiable. Its application leads to nothing definite, even in the case of the particular ore No. III, from which the silicate analyzed was derived.

In the following tables are given first the recalculated analyses and then the molecular ratios. All those constituents have been excluded which pretty certainly do not belong to the carnotite, but small portions of some of those retained are unquestionably extraneous. In two cases (*I-a* and *I-b*) a certain proportion of lime has been arbitrarily excluded equivalent to

* According to Dr. Harry C. Jones of Johns Hopkins University, who very kindly undertook to examine a specimen of the Copper Prince ore for rare gaseous elements, helium is not present. Faint hydrogen lines were observed, the source of which was ascribed to water vapor. Other lines, due probably to hydrocarbons, were fairly strong, but the specimen had been long enough exposed in our laboratory and elsewhere to have accumulated enough dust to account for them.

the acid anhydrides CO_2 , MoO_3 , and SO_3 , less what is needed to offset PbO and CuO .

Carnotite Analyses Recalculated.

	I-a	I-b	I-c	II-a	II-b	III
V_2O_5 -----	20.72	21.09	20.12	20.54	19.85	20.62
P_2O_5 -----	.90	.40	.51	.06		
As_2O_5 -----		.29				
UO_3 -----	61.53	60.06	60.55	61.44	59.31	58.75
CaO -----	3.03	2.77	3.28	2.11	2.10	4.70
$\text{BaO}(\text{SrO})$ -	1.03	.83	.83	3.22	3.64	.83
MgO -----	.25	.23	.31	.16	.19	.20
K_2O -----	7.31	7.73	8.39	6.21	5.80	4.33
Na_2O -----	.15	.10	.09	.15	.02	.03
$\text{H}_2\text{O}-105^\circ$	2.72	2.98	2.36	3.59	5.13	5.30
$\text{H}_2\text{O}+105^\circ$	2.36	3.52	3.56	2.52	3.96	5.24
	100.00	100.00	100.00	100.00	100.00	100.00

	I-a	I-b	Ratios. I-c	II-a	II-b	III
V_2O_5 -----	1133	1208	1101	1124	1081	1128
P_2O_5 -----	63	28	36	4		
As_2O_5 -----		12				
UO_3 -----	2140	2088	2105	2136	2062	2043
CaO -----	541	495	586	377	375	839
$\text{BaO}(\text{SrO})$ -	68	54	54	210	238	54
MgO -----	62	57	78	40	47	50
K_2O -----	776	826	891	659	616	460
Na_2O -----	33	22	20	33	4	6
$\text{H}_2\text{O}-105^\circ$	1511	1656	1311	1994	2850	2944
$\text{H}_2\text{O}+105^\circ$	1311	1956	1978	1400	2200	2911

These ratios lead to the following empirical formulas, in which only the water given off above 105° is considered. Since the water is wholly removable below 350° it is regarded as water of crystallization and not of constitution.

	R'	R''	R ^v	U	O	H ₂ O
I-a ..	1618	671	2392	2140	13880	1311
I-b ..	1696	606	2496	2088	13958	1956
I-c ..	1822	718	2274	2105	13629	1978
II-a .	1384	627	2256	2136	13367	1400
II-b .	1240	660	2162	2062	12871	2200
III ..	932	943	2256	2043	13178	2911

If it be assumed that the bivalent elements offer the most accurate determinations, the above ratios may be reduced to the following simpler terms on that basis :

	R'	R''	R ^v	U	O	H ₂ O
I-a	2.41	1	3.56	3.19	20.68	1.95
I-b	2.80	1	4.12	3.45	23.03	3.23
I-c	2.54	1	3.17	2.93	18.98	2.75
II-a	2.21	1	3.60	3.41	21.32	2.23
II-b	1.88	1	3.27	3.12	19.50	3.33
III99	1	2.39	2.16	13.98	3.09

The results, however, show a great lack of agreement and wide variation. It is plain that no probable formula can be calculated for the yellow body. The variations are of such a nature as to indicate in the plainest manner that it is a mixture of several substances.

Such a detailed discussion as the foregoing would hardly have been justified in view of the negative conclusions arrived at, but for the fact that Messrs. Friedel and Cumenge in their paper announced a simple formula for the body examined by them and gave it the specific name carnotite. Their published analyses are as given below, from which they have excluded considerable sand and traces of barium, aluminum, lead, copper, and radio-active bodies as present in excessively small quantities. They make no mention of calcium, and admit that their values for water are open to doubt. The formula deduced by them is $2U_2O_3, * V_2O_5, K_2O, 3H_2O$.

Analyses of Carnotite by Friedel and Cumenge.

	Found.		Calculated.
$U_2O_3^*$ -----	64.70	62.46	63.54
V_2O_5 -----	20.31	19.95	20.12
K_2O -----	10.97	11.09	10.37
H_2O -----		5.29	5.95
Fe_2O_3 -----	.96	.65	
			99.98

It appears from these analyses that Messrs. F. and C. by great good fortune obtained a variety of samples of the pure potassium compound, free from calcium and without appreciable admixture of barium. This is very remarkable in view of the fact that all the ores from different localities examined by me show large admixture of calcium or barium salts or both, even that which is certified to have come from the same lot as that from which their material was taken. The French authors give a brief outline of their methods of analysis. That one which afforded them the best results would involve the weighing of any calcium present as sulphate along with the potassium, on the assumption that the presence of that element had been overlooked. It is much to be desired that a re-analysis of their material should be made, if there is any of it still available, in order to clear up the doubt connected with the first analysis.

In the light of the evidence herein set forth, the existence of a distinct mineral species having the composition claimed for carnotite can by no means be considered as established.

* Old notation equivalent to the modern UO_3 .

Average quality of marketed ore.—As these carnotite ore bodies are being exploited for the market, it is of some interest to know the average quality of each commercial lot. A carefully prepared sample representing several tons of ore was received from one of the commercial houses of Denver and was found to carry 11.49 per cent of uranium counted as U_3O_8 , and 6.40 per cent of vanadium counted as V_2O_5 . Over one-sixth of the vanadium existed, however, in the trivalent state, not as a constituent of the yellow body, but doubtless of a silicate like the one whose composition is given on p. 143.

Commercial assay.—The commercial assay of these ores has presented difficulties to the technical chemist, the results being sometimes very discordant.

As to uranium this is not surprising. The methods that have probably been commonly employed will give varying results according to the contents of the ore in phosphorus and alkaline earths. Possibly the old Patera process, described in most text-books on analytical chemistry, might be made to serve, with modifications called for by the large amount of vanadium present.

The assay for vanadium presents little difficulty and does not require much time. The ore is fused with sodium carbonate, leached with water and the fusion repeated on the residue. The combined filtrates are acidified by sulphuric acid, arsenic and molybdenum are precipitated in the hot solution by hydrogen sulphide, whereby the V_2O_5 is reduced to V_2O_4 . After filtration and expulsion of hydrogen sulphide by boiling, the vanadium is titrated in hot solution by permanganate. It is then reduced by sulphur dioxide gas, and after boiling this out the titration is repeated. The results are exact, and they are not affected by the uranium that may be present.

IV. *Composition of the Vanadiferous Silicate in the Carnotite Ore.*

As already mentioned the ores contain a vanadiferous silicate free from uranium. To the end of ascertaining its composition if possible, ore No. III (p. 128), from the Yellow Bird claim, low in carnotite but relatively rich in vanadium, was treated as follows:

The carnotite from 10 grams was extracted by cold, dilute nitric acid, and the well washed residue, consisting of coarse sand and an utterly amorphous mud, by 4 per cent sodium carbonate solution to get rid of the small amount of silica presumably set free but not dissolved by the acid. This amounted to 0.35 per cent in duplicate determinations, and together with 0.06 per cent in the acid solution or 0.41 per cent in all, may

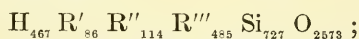
serve as a maximum figure by which to gauge the action of the cold acid on the silicate or silicates in the ore. The residue was then digested for several hours with warm nitric acid of about 1·2 sp. gr. until as shown by a companion test its action had ceased. It was then filtered, washed and digested with 5 per cent sodium carbonate solution to dissolve the copious deposit of silica. The final residue was collected in a Gooch crucible, washed with sodium carbonate, followed by dilute nitric acid to remove all alkali, then by alcohol to prevent turbidity in the filtrate, and dried by suction of the pump. In it the water was determined at different temperatures, also its general composition.* From the sodium carbonate filtrate the silica was obtained by two evaporations and filtrations, also the trifling amount held by the acid solution. This latter was then fully analyzed and the complete results follow.

Composition of vanadiferous silicate in carnotite ore.

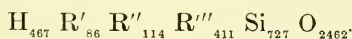
	Per cent in ore.	Per cent calculated to 100.	Ratios.
SiO ₂	6·48	43·94	·7275
Al ₂ O ₃	2·445 (2·45 and 2·44)	16·58	·1622
V ₂ O ₅	·965 (·93 and 1·00)	6·54	·0434
Fe ₂ O ₃	·875 (·87 and ·88)	5·93	·0370
CaO	·035	·24	·0043
MgO	·654 (·652 and ·657)	4·43	·1099
K ₂ O	·546	3·70	·0393
Na ₂ O	·03	·20	·0037
H ₂ O 105° C.	1·66†	11·26	·9222
H ₂ O 300° C.	·44†	2·98	·1655
H ₂ O above 300° C. .	·62†	4·20	·2333
	14·75	100·00	

Also traces of titanium, manganese, and lithium.

On the improbable assumption that the iron is to be wholly included, and regarding only the water given off above 300°, the following ratios result:



which become, if the iron is excluded,



These figures, while strongly suggesting definite ratios between certain of the constituents, do not under the circumstances warrant the deduction of a formula, nor do they lead

* SiO₂ 48·89, Al₂O₃, Fe₂O₃, zircon etc., ·44, TiO₂ ·08, MgO ·01, K₂O ·11, Na₂O ·03, H₂O ·07; total 49·63.

† Determined on a separate portion of the same powdered sample. See p. 138 footnote for data.

to the same conclusion as in the case of the green cementing material of the Placerville sandstone. The phengite-muscovite ratio of that is not apparent here. Yet it is not at all unlikely that a mineral like the one from Placerville is present, but contaminated with some other. In fact it would be surprising to find in such thoroughly altered sandstones anything but a mixture. The fact of the existence of such vanadiferous transition products is itself highly interesting, and these tedious analyses cannot therefore be considered as made in vain. The mud-like amorphous character of this material precludes any hope of aid from the microscope in solving the question of its homogeneity.

Summary.

The body called carnotite is probably a mixture of minerals of which analysis fails to reveal the exact nature. Instead of being the pure uranyl-potassium vanadate, it is to a large extent made up of calcium and barium compounds. Intimately mixed with and entirely obscured by it is an amorphous substance—a silicate or mixture of silicates—containing vanadium in the trivalent state, probably replacing aluminum.

The deposits of carnotite, though distributed over a wide area of country, are, for the most part, if not altogether, very superficial in character and of recent origin.

The green coloring and cementing material of certain sandstones near Placerville, Col., is a crypto-crystalline aluminovanadio-potassium silicate resembling rooseelite, but with the percentage proportions of Al_2O_3 and V_2O_5 reversed. It constitutes over 25 per cent of the sandstone at times, and contains nearly 13 per cent of V_2O_5 , the latter amounting in the maximum case observed to 3.5 per cent of the sandstone.

As yet these highly vanadiferous sandstones have been found only at Placerville, where it is intended to work them for vanadium. Carnotite is associated with them in only trifling amount.

Other sandstones noticed owe their bright green color to chromium. In yet another case where the color was dull green, this was not due to either chromium or vanadium.

ART. XV.—*Restoration of Stylonurus Lacoanus, a Giant Arthropod from the Upper Devonian of the United States*; by CHARLES E. BEECHER. (With Plate I.)

IN the animal kingdom, the attribute of bigness has come to be regarded as one of the prerogatives of the vertebrates. On this account, invertebrates seldom receive credit for having a size of more than a fraction of a cubit, and are looked upon as objects to be held in the hand or viewed under a lens. As a matter of common experience, and probably also of congratulation, large invertebrates *are* rare, and some whole classes cannot furnish a single individual measuring more than a few inches in greatest diameter.

In a list of arthropod giants, the subject of the present note must be included, and will take equal rank with the Giant Spider-Crab of Japan (*Macrocheira Kaempferi*) and the great "Seraphim" of the Scotch quarrymen (*Pterygotus anglicus*). The former can safely claim to be the largest representative of the Brachyurans that has ever existed, and to the latter may be accorded the same distinction among the Merostomes.

The living species of the Merostomata comprise only the American and Moluccan Horse-Shoe Crabs, *Limulus polyphemus* and *L. moluccanus*. The latter sometimes attains a length of three feet, and measures eighteen inches across the carapace. To find other species in this order worthy of comparison with the huge Brachyuran of Japan, it is necessary to go back to the Paleozoic forms, and among these the larger species of *Pterygotus*, and the *Stylonurus* here noticed, fill all the requirements. It should be borne in mind, however, that these statements are based upon comparative lengths and breadths. If bulk alone were considered, the common lobster (*Homarus americanus* and *H. vulgaris*) should be mentioned, though in length and extent of limbs it would be considerably smaller.

Concerning the size of the Scotch "Seraphim," Dr. Henry Woodward¹⁰ states that "From our present knowledge of the almost perfect remains of *Pterygotus anglicus*, and on the evidence of the numerous detached portions of this extinct genus, we are justified in concluding that it attained a length of six feet, and a breadth of nearly two feet at the widest part of its body." This huge Merostome has been found in the Lower Old Red Sandstone of Scotland, at a horizon nearly equivalent to the one furnishing the remains of *Stylonurus* in America. Thus what seem to be the two largest species of this class were contemporaries, though not associates.

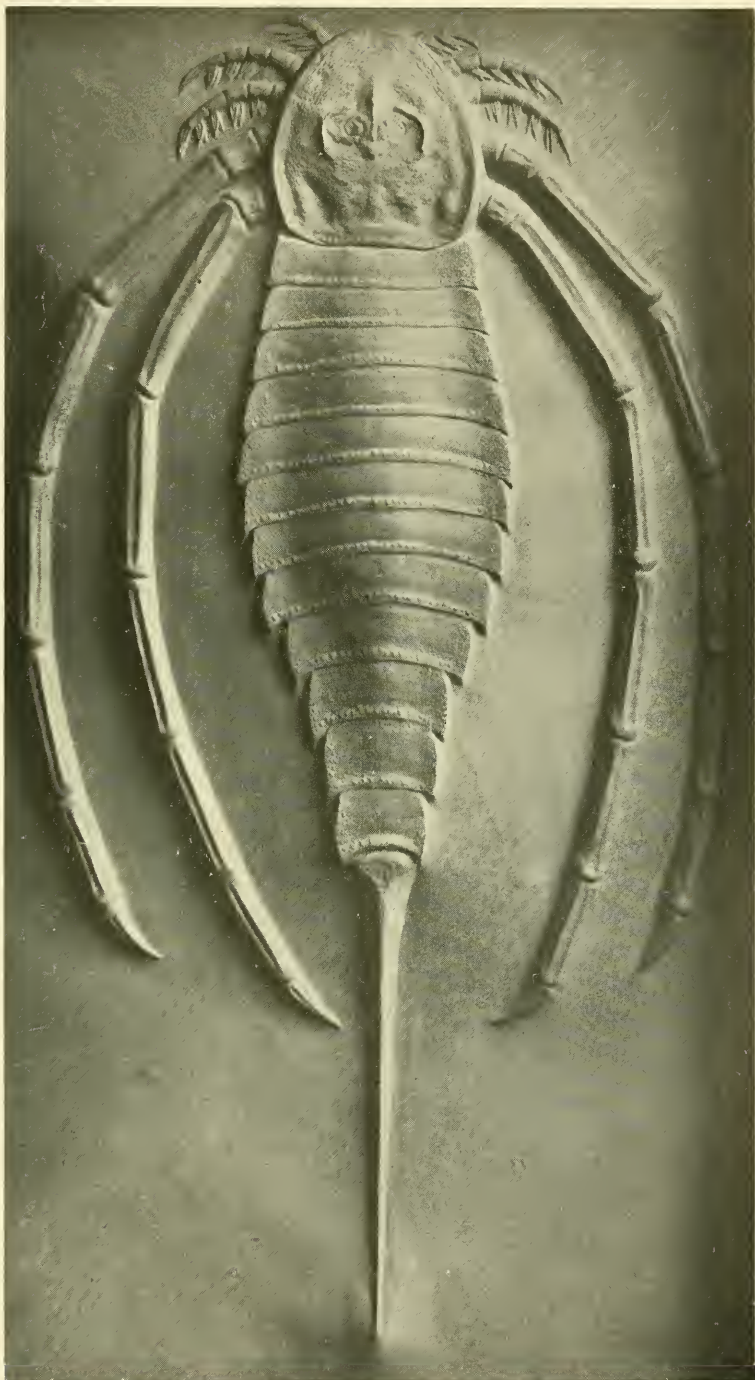
Historical.

The first specimen found in America that can be referred to the genus *Stylonurus* was collected by the writer about 1870, and loaned to Professor James Hall. It remained in his hands unnoticed until 1884, when he described it as *Eurypterus Beecheri*.⁴ The specimen preserves the abdomen and portions of two of the large posterior limbs. No species of *Eurypterus* known possessed such greatly elongated limb joints, and there seems to be no good reason for not referring it to *Stylonurus*, in which this is a normal character. The specimen of *Stylonurus Beecheri* is uncompressed, and apparently retains the proportions of form and convexity as in life. On this account, it was of considerable importance in the restoration of the larger species.

In 1882, Hall was furnished with a plaster cast of the carapace of a large arthropod, by Dr. Cook, then State Geologist of New Jersey. The original specimen was from the Catskill group at Andes, Delaware County, New York, and had been sent to the museum at Rutgers College, New Brunswick, New Jersey. Prof. D. S. Martin⁹ made the first reference to this species in some remarks on "A New Eurypterid from the Catskill Group," before the New York Academy of Sciences, October 16, 1882, an abstract of this note appearing in the transactions of the same society some time after June, 1883. In this abstract, the species is neither described nor figured, and Hall is not mentioned in any connection. Martin states that he saw the specimen (=cast sent to Hall) in the State Museum at Albany, and it bore the name *Stylomurus excelsior* (evidently a misprint for *Stylonurus*).

The next reference to this form in point of time, and the first publication of a generic and specific name accompanied with a description and accurate illustration, was given by E. W. Clappole,¹ in a paper read before the American Philo-osophical Society, September 21, 1883, under the title "Note on a large Crustacean from the Catskill Group of Pennsylvania." It is stated on the signature containing this paper that it was printed November 2, 1883. Clappole's description was based upon a second specimen found in Wyoming County, Pennsylvania, which preserves about three-fourths of the cephalothorax, and belonged to the collection of R. D. Lacoë of Pittston. This was given the name *Dolichocephala Lacoëna*, and rightly classified with the Merostomata. It therefore appears that, up to this time, the name *Stylonurus excelsior* was simply *nomen nudum*, and as such cannot be recognized as valid.

In 1884, Hall⁹ published his description and figure of the



STYLONURUS (length nearly five feet).

New York specimen, in the thirty-sixth Annual Report of the New York State Museum, in a paper entitled "Description of a New Species of *Stylonurus* from the Catskill Group." It is here that the New York specimen was first figured and a description given, and the date of publication of this paper is the one to be considered in deciding the claims of *Stylonurus excelsior* as Hall's species.

At the Philadelphia meeting of the American Association for the Advancement of Science, September, 1884 (Proc. A. A. S., vol. xxxiii, published 1885), Hall^e presented a note on *Stylonurus excelsior*, merely referring to its occurrence, and citing Martin's abstract with page and month of publication. This citation is repeated by Hall in each of his notices of this species, for only by thus establishing the species could he have any claim to priority. As already mentioned, Martin's paper does not attempt any description of this form, and Hall is not mentioned. Hall further says: "The carapace is described and figured in the 36th Report of the N. Y. State Museum of Natural History," without reference to plate, page, or year, and it is therefore quite possible that this description was not published until after the meeting of the Association. In any case, it appeared some months later than Claypole's paper, and the name *Dolichocephala Lacoana* has priority over *Stylonurus excelsior*, and must be recognized.

Claypole failed to point out the affinities of this form with *Stylonurus*, and proposed a new generic term for his species. Although there are differences that may prove of generic value when more complete specimens of the American species have been studied, yet at the present time there seem to be no strong reasons why the specimen in question should not be considered as belonging to *Stylonurus*, and it is upon this ground that the present restoration is attempted.

Material available for a Restoration.

Restorations of extinct organisms are largely exhibits of mental architecture, based upon the personal interpretation of a certain number of real things. Some statement, therefore, should be given of the character and amount of the material that has been collated to furnish a restoration of *Stylonurus Lacoanus*.

(1) The specimen of the cephalothorax described by Hall shows the complete outline and upper surface of this part, and a cast from the original was taken to represent this portion in the restoration. (2) The type of *S. Lacoanus* Claypole includes a large part of the cephalothorax of an individual nearly the same in size as the preceding. (3) Dr. J. M. Clarke

discovered that this specimen also preserved considerable evidence as to the nature of the appendages, and he succeeded in developing what appear to be the chelate antennæ, the first pair of gnathopods, and the mandibular bases of at least three others. (4) The length and number of joints in the limbs are taken from the English species *S. Logani* and *S. Powriei*, of which quite complete, though smaller, individuals have been described by Woodward.¹⁰ (5) The outline and proportions of the abdomen follow closely those of the English forms and of *S. Beecheri*, the latter giving the natural convexity. (6) A portion of a large abdominal segment found by the writer in the Chemung group at Warren, Pennsylvania, and apparently belonging to a nearly related species, has an ornamentation closely approaching that on the cephalothorax of the type, and was used to elaborate the sculpture over the abdomen of the restoration. (7) The form and character of the telson spine correspond to *S. Logani* and also to some large fragments found by F. A. Randall at Warren and Ackley, Pennsylvania, in the Chemung group, and probably belonging with the abdominal somite already mentioned.

With the data furnished by the foregoing material, the restoration was undertaken. The first model in relief was constructed in clay, and from it a plaster mold was taken. A number of casts have been made since, and a photograph of one of them is represented in the accompanying plate (Plate I).

In this connection, it may be suggested that the type specimen of *Stylonurus* (?) (*Echinocaris*?) *Wrightianus* (Dawson sp.) represents two proximal joints of one of the large crawling feet of a form related to *Stylonurus*, and not two somites of the abdomen as indicated by Hall.⁵ Any reference is at present somewhat uncertain, owing to lack of positive knowledge, and the fact that the specimen in question was first described as a plant (*Equisetides*²), then referred to the Phyllocarida (*Echinocaris*⁸), and lastly appeared as a possible Merosome, shows how this form may be interpreted by different observers. No one can doubt its arthropod nature, on account of the characteristic surface markings. Its elliptical or ovoid section without any flattening of the epimera, the very considerable overlapping of the joints, and the configuration of the suture, are more strongly indicative of the nature and requirements of a limb than of abdominal segments.

The Relief Model.

In this restoration, the animal is represented as lying on a slab, with the entire dorsal surface exposed. The cephalothorax has an axial length of 25^{cm} and a width of 22^{cm}.

The chelate antennæ were doubtless carried in a folded position, as in most related genera, and seldom were visible from the dorsal side. They are, therefore, not shown. The three pairs of short gnathopods, serving partly as swimming organs, are seen extending outward from the antero-lateral margins of the cephalothorax. Several of their distal joints are each provided with a pair of flat, ridged, spinous processes, and a similar spine at the termination of the limb.

The two pairs of great crawling feet extend outward and backward from the postero-lateral margins of the cephalothorax. The anterior pair expose 109^{cm} of length, and the posterior pair about 108^{cm}. The elements of the limbs are represented as grooved, as this character seems necessary to give the needed strength to long slender joints, and also because a similar conformation is present in *S. Beecheri*.

The abdomen measures 30^{cm} in greatest width at the fifth segment, and 66^{cm} in length exclusive of the telson. The posterior abdominal segments are represented without detachable epimera, as this feature is not as yet known to be constant for the genus, although present in some species.

The telson spine agrees proportionally in length with the same member in *S. Logani* and *S. Powriei* as described by Woodward,¹⁰ and was given a slight upward curvature as in *Limulus*. It measures 54^{cm} in length and 7.5^{cm} across at the proximal end.

Altogether the animal as restored has a length of nearly five feet (147^{cm}), and with the legs extended it would measure about eight feet (242^{cm}) across.

It is not intended to claim any high degree of accuracy for this restoration, but merely to represent in some graphic form an animal approximating in size and character an individual of the species *Stylonurus Lacoanus*. Its size alone was the chief incentive for attempting a reconstruction, and some sacrifice of exact detail may well be allowed, in order to make any presentation of this magnificent arthropod.

Yale University Museum, New Haven, Conn., September 5, 1899.

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ART. XVI.—*The Iodometric Estimation of Arsenic Acid;*
by F. A. GOOCH and JULIA C. MORRIS.

[Contributions from the Kent Laboratory of Yale University—XCVI.]

THE interaction of a soluble arseniate and a soluble iodide in a suitably acidulated solution results, as is well known, in the reduction of the arsenic acid (more or less completely according to conditions of temperature and proportions of reagents and solvents) with the corresponding liberation of two atoms of iodine for every molecule of arsenic acid (H_3O_3AsO) reduced. Inasmuch, however, as the reaction of this process is reversible, it is necessary, in order that the reduction may be complete, to nullify the oxidizing action of the iodine liberated. Theoretically this end may be accomplished in either of two ways, by volatilizing the free iodine bodily or by destroying the oxidizing power of the iodine by converting it to hydriodic acid. The former method was followed in a process devised for the estimation of arsenic acid and elaborated in this laboratory.* This method, as originally put forward, consisted in adding to the solution of the arseniate potassium iodide in excess of the amount theoretically indicated, with 10^{cm^3} of sulphuric acid of half strength, and so arranging the dilution that the total volume of the liquid should be about 100^{cm^3} , boiling until the volume decreased to 40^{cm^3} , bleaching by the cautious addition of sulphurous acid the trace of free iodine still held by the hydriodic acids, diluting, cooling, neutralizing with acid potassium carbonate, and titrating with iodine, after adding the starch indicator. This process, depending upon the removal by volatilization of all but the last traces of liberated iodine and the conversion of this minute residue by sulphurous acid, involves no secondary reactions of a sort likely to influence the main effect. It is exact and fairly rapid.

The method of Williamson,† brought forward more recently, depends upon the conversion of the liberated iodine to hydriodic acid. The interaction at ordinary temperatures of a suitably strong acid, hydrochloric or sulphuric acid, upon the mixture of the arseniate and iodide sets free iodine, and the liberated iodine is converted to hydriodic acid by the action of sodium thiosulphate, the end point being the disappearance of the iodine color.

According to Williamson's directions, 25^{cm^3} portions of the solution of the arseniate are treated with potassium iodide and mixed with an equal volume of hydrochloric acid of sp. gr.

* Gooch and Browning, this Journal, xl (1890), 66.

† Jour. Soc. Dyers and Colorists, 1896, 86-89.

1.16. The precaution is recommended that the strength of the solution of the arseniate shall not exceed the decinormal value, in order that the dilution consequent upon titration by the thiosulphate may not be too great—the reducing action brought about by the action of the strong acid upon the arseniate and iodide being reversible upon the dilution of liquid with water. This procedure thus limits the process to the determination of about 0.18 gram of arsenic acid in 25^{cm}³ of the solution to be treated with an equal volume of hydrochloric acid of sp. gr. 1.16. Obviously, however, the process should, so far as the reduction is concerned, be applicable to larger amounts of arsenic provided the strength of the acid is kept up proportionately. It is essential that the liquid at the end of the titration should contain approximately ten per cent of its mass of absolute hydrochloric acid or about one-third of its volume of the aqueous acid of sp. gr. 1.16.

The arsenic acid is measured either by the amount of standard thiosulphate required to bleach the iodine or by the amount of iodine required to reoxidize the reduced arsenious acid, after neutralizing with acid potassium carbonate. If the former alternative is followed, the end-reaction must be the disappearance of the yellow color of the iodine, since in solutions so strongly acid it is impossible to place dependence upon the starch indicator; in using the latter alternative, the starch indicator is, of course, permissible and preferable.

In the direct titration of the iodine by thiosulphate two sources of error present themselves as possibilities; first, the excessive liberation of iodine by the action of air upon the strongly acidulated iodide; and second, the liability of the thiosulphate,* if present even in momentary or local excess during the process of titration, to break down under the action of strong acid, thus changing its capacity to convert iodine to hydriodic acid. The latter contingency should be remote in proportion to the caution used in adding the thiosulphate and in keeping the liquid well stirred; the former must of necessity vary with the acidity of the solution containing the iodide, the time of exposure to atmospheric action, and the degree of contact with the air incidental to stirring. We have thought it desirable, therefore, to see how far each of these possibilities is likely to interfere in the practical conduct of an ordinary analysis.

The effects likely to result simply from the strong acidification of the solution containing potassium iodide and their variation for conditions of dilution representing the beginning and the end of a titration on the lines laid down are shown in

* Norton, this Journal, vii, 287.

the following table. The solution of potassium iodide was diluted as indicated before the addition of the acid and the iodine set free was titrated by thiosulphate.

Sp. gr. HCl taken.	KI taken.	Total volume.	Na ₂ S ₂ O ₃ added at once. In terms of H ₃ O ₃ AsO.	Na ₂ S ₂ O ₃ added after 5 minutes. In terms of H ₃ O ₃ AsO.	Na ₂ S ₂ O ₃ added after stirring 5 minutes. In terms of H ₃ O ₃ AsO.
cm ³ .	gm.	cm ³ .	gm.	gm.	gm.
25	2	50	0·0013		
25	2	75	0·0004		
25	2	50		0·0035	
25	2	75		0·0019	
25	2	50			0·0042
25	2	75			0·0021
50	2	100	0·0017		
50	2	150	0·0004		
50	2	100		0·0035	
50	2	150		0·0019	
50	2	100			0·0035
50	2	150			0·0014

The proportionate strength of acid and the time before titration are, obviously, the essential factors. The absolute amount of acid present and the stirring seem to make little difference.

As to the action of the hydrochloric acid on small amounts of the thiosulphate, we have the evidence of the experiments detailed in the following statements, in which 1, 2, and 5^{cm³} of nearly $\frac{n}{10}$ thiosulphate are exposed to the action of 25^{cm³} hydrochloric acid, sp. gr. 1·16, without dilution or diluted with

HCl Sp. gr. 1·16.	Volume before titration.	Na ₂ S ₂ O ₃ nearly $\frac{n}{10}$ in terms of H ₃ O ₃ AsO.	Iodine to color without dilution in terms of H ₃ O ₃ AsO.	Error of titration without dilution in terms of H ₃ O ₃ AsO.	Iodine to color after diluting to 75 ^{cm³} in terms of H ₃ O ₃ AsO.	Error of titration after dilution in terms of H ₃ O ₃ AsO.	
cm ³ .	cm ³ .	cm ³ . gram.	gram.	gram.	gram.	gram.	
25	26	1	0·0071	0·0062	—0·0009	0·0071	0·0000
25	50	1	0·0071	0·0071	0·0000	0·0071	0·0000
25*	50	1	0·0071	0·0079	+0·0008	0·0079	+0·0008
25	50	2	0·0141	0·0146	+0·0005	0·0146	+0·0005
25*	50	2	0·0141	0·0157	+0·0016	0·0157	+0·0016
25	30	5	0·0353	0·0336	—0·0017	0·0374	+0·0024
25	50	5	0·0353	0·0359	+0·0006	0·0359	+0·0006
25*	50	5	0·0353	0·0411	+0·0058	0·0411	+0·0058

* In these experiments the acid stood in contact with the thiosulphate 5 minutes before titration.

an equal volume of water, were titrated with nearly $\frac{n}{10}$ iodine.

The condition of acidity when the volume of 50cm^3 contains 25cm^3 of hydrochloric acid, sp. gr. 1.16, is that of the beginning of titration of Williamson's process. In order that the effect of error due to such action upon the determination of arsenic acid may appear immediately, the thiosulphate and iodine used are expressed in terms of that acid.

These two sources of error, the one due to a liberation of iodine and the other due to decomposition of the thiosulphate, would naturally tend to overcome one another, but the completeness of such neutralization would naturally be largely a matter of chance in the varying conditions of actual analysis. The experiments of the following table, however, in which $\frac{n}{10}$ thiosulphate, to the amount of 1, 2, and 5cm^3 , was added to the liquid, 50cm^3 and 75cm^3 , containing 25cm^3 acid, and titrated with iodine at once, and after five minutes, were made to test the matter for the conditions of dilution at the beginning and at the end of a titration.

HCl Sp. gr. 1.16.	KI.	Volume.		Na ₂ S ₂ O ₃ nearly $\frac{n}{10}$ in terms of H ₃ O ₃ AsO.		Iodine in terms of H ₃ O ₃ AsO, at once. gram.	Iodine in terms of H ₃ O ₃ AsO, after 5 min. gram.	Error in terms of H ₃ O ₃ AsO. gram.
		cm ³ .	gm.	cm ³ .	cm ³ .			
25	2	50	1	0.0071	0.0057		-0.0014	
25	2	75	1	0.0071	0.0071		0.0000	
25	2	50	2	0.0141	0.0131		-0.0010	
25	2	75	2	0.0141	0.0143		+0.0002	
25	2	50	5	0.0353	0.0322		-0.0021	
25	2	75	5	0.0353	0.0357		+0.0004	
25	2	50	1	0.0071		0.0028	-0.0043	
25	2	75	1	0.0071		0.0067	-0.0004	
25	2	50	2	0.0141		0.0116	-0.0025	
25	2	75	2	0.0141		0.0139	-0.0002	
25	2	50	5	0.0353		0.0314	-0.0041	
25	2	75	5	0.0353		0.0361	+0.0008	

It is clear that under the conditions covered by the experiments of the two preceding tables the decomposition of the thiosulphate is likely to occur in greater or less degree, and that when the acid of sp. gr. 1.16 is not much diluted, the products of decomposition are not oxidized by the iodine completely. The latter observation is quite in harmony with the fact that sulphur dioxide bleaches iodine in strong hydrochloric acid only slowly and incompletely. In such cases dilution

favors further action of the iodine, but results obtained by titration with iodine in the acid solution diluted with an equal amount of water are unmodified by further dilution.

In the following tables are recorded actual determinations of arsenic according to Williamson's process. To each 25^{cm}³ of the arseniate were added 1, 2, or 3 grams of potassium iodide and 25^{cm}³ hydrochloric acid, sp. gr. 1.16. The iodine was bleached by nearly decinormal thiosulphate without addition of the starch indicator, which loses all delicacy in the presence of strong acid. The time occupied by each titration was about five minutes. The standards of the arseniate were determined by the vaporization process,* the purity of reagents employed in that process having been proved by trying the process in the estimation of a solution of arsenic acid made by oxidizing pure decinormal arsenious acid by iodine.

HCl cm ³ .	KI. gram.	Volume at beginning of titration. cm ³ .	Volume at end of titration. cm ³ .	H ₂ KAsO ₄ in terms of H ₃ O ₃ AsO. gram.	H ₃ O ₃ AsO found. gram.	Error. gram.
25	2	50	51	0.0062	0.0085	+0.0023
25	2	50	52	0.0125	0.0156	+0.0031
25	2	50	55	0.0312	0.0350	+0.0038
25	2	50	55	0.0624	0.0666	+0.0042
25	2	50	73	0.1559	0.1588	+0.0029
25	2	50	73	0.1559	0.1587	+0.0028
25	2	50	73	0.1559	0.1591	+0.0032
25	2	50	73	0.1559	0.1595	+0.0036
25	3	50	73	0.1559	0.1595	+0.0036
25	1	50	73	0.1559	0.1581	+0.0022
25	2	50	73	0.1559	0.1581	+0.0022
25	2	50	73	0.1559	0.1588	+0.0029

The range of error in these results is from +0.0023 gram to +0.0042 gram with a mean of +0.0031 gram—not very different from what might be expected from the effect of the interaction of the strong hydrochloric acid and the iodide alone. The counter-effect due to the decomposition of the thiosulphate is not large, yet it is probably real, as will appear in the sequel. In the following series of determinations, made with new solutions and new standards throughout, the arsenic acid was determined, first, by titrating the iodine set free by 25^{cm}³ of hydrochloric acid, sp. gr. 1.16 and 3 grams potassium iodide, the solution having a total volume of 50^{cm}³ at beginning and of 75^{cm}³ at the end of titration and, secondly, the arsenious acid produced in the first reaction was titrated, after being neutralized with acid potassium carbonate by iodine in the presence of the starch indicator.

* Gooch and Browning, loc. cit.

H ₂ KAsO ₄ taken in terms of H ₃ O ₃ AsO. gram.	H ₃ O ₃ AsO found by the thiosulphate. gram.	Error. gram.	H ₃ O ₃ AsO found by titration of H ₃ O ₃ As with iodine. gram.	Error. gram.
0·1767	0·1798	+0·0031	0·1776	+0·0009
0·1767	0·1798	+0·0031	0·1777	+0·0010
0·1767	0·1795	+0·0028	0·1785	+0·0018
0·1767	0·1793	+0·0026	0·1785	+0·0018
0·1767	0·1794	+0·0027	0·1780	+0·0013
0·1767	0·1798	+0·0031	0·1785	+0·0018

The average error of the first operation is 0·0029 gram, not far from that of the previous series; the error of the second operation, the titration of the arsenious acid, amounts on the average to 0·0014 gram. In the second operation the error due to over-use of the thiosulphate by iodine set free outside the main reaction is obviously eliminated. The tetrathionate present after neutralization with acid potassium carbonate is unaffected by iodine, as we have found by titrating $25^{\text{cm}^3} \frac{n}{10}$ iodine mixed with 25^{cm^3} hydrochloric acid, sp. gr. 1·16, by the thiosulphate, neutralizing with acid potassium carbonate,* adding starch and getting the starch blue with a single drop of $\frac{n}{10}$ iodine. The average error of this process, therefore, 0·0014, is probably due to the products of decomposition of the thiosulphate in the first operation.

From the foregoing experiments it is clear that an arbitrary correction of about 0·0030 gram must be deducted from the indications of Williamson's process of direct titration by thiosulphate, made with the greatest care under the conditions mentioned; and that a correction varying from one-half that amount (0·0015 gram) to nothing (according to the amount of arsenious acid present) when the determination is made by iodine after neutralization with acid potassium carbonate. After making these arbitrary corrections in the results of the preceding table, the individual variations fall within reasonable limits.

On the other hand, the vaporization process, in which the arseniate is reduced by boiling with sulphuric acid and potassium iodide in the manner described,† gives indications reason-

*It is worthy of note, that, as we have found by experience, it is not possible to substitute an alkaline hydroxide for the carbonate in the early stages of the process of neutralization, on account of the decomposing effect of the former reagent upon the tetrathionate. This effect is in proportion to the heating of the solution, but is never wholly absent even when ice is intermixed with the liquid and the greatest care taken to prevent a rise of temperature.

† Loc. cit.

ably regular and accurate without the application of an arbitrary correction. This process, moreover, may be shortened by restricting the volume at which heating begins so that the boiling need not be extended beyond five or six minutes. According to this slight modification, the solution of the arseniate is heated in an Erlenmeyer flask with potassium iodide to an amount about 0.5 gram in excess of the amount theoretically required and 10^{cm³} of sulphuric acid of half strength in a total volume of between 50^{cm³} and 75^{cm³}. The liquid is boiled till the iodine vapors are no longer visible in the flask above the liquid, the iodine color in the still hot liquid is bleached by the cautious addition of sulphurous acid, the whole is diluted with cold water, and cooled quickly. The solution is nearly neutralized with potassium hydroxide and the neutralization is completed with acid potassium carbonate. The reduced acid is titrated with iodine after adding the starch indicator. By this procedure the results of the following table were obtained.

Volume. cm ³ .	H ₃ O ₃ AsO taken. gram.	H ₃ O ₃ AsO found. gram.	Error. gram.
35	0.1559	0.1559	0.0000
35	0.1559	0.1560	+0.0001
40	0.1559	0.1559	0.0000
65	0.1559	0.1559	0.0000
50	0.2495	0.2499	+0.0004
50	0.2557	0.2449	-0.0008
60	0.3119	0.3117	-0.0002
60	0.3119	0.3120	+0.0001
75	0.3119	0.3124	+0.0005
75	0.3119	0.3132	+0.0013
75	0.3119	0.3121	+0.0002
75	0.3119	0.3115	-0.0004
75	0.3119	0.3124	+0.0005

ART. XVII.—*Further Notes on Preglacial Drainage in Michigan*; by E. H. MUDGE.

THE most complete study and discussion of the surface geology of Michigan that has recently been published is contained in No. 30 of the Water Supply and Irrigation Papers of the U. S. Geological Survey, under the title, "Water Resources of the Lower Peninsula of Michigan," by Alfred C. Lane, State geologist. Besides much information of economical importance, this pamphlet contains a series of very useful maps, an attempt to work out more fully than has heretofore been done the system of Preglacial drainage, and other matters of more special scientific interest. Prof. Lane has probably given more study to the records of deep wells and other similar data than any other man in the state, and concludes that the theory of Dr. Spencer, heretofore adopted by the writer, that in Preglacial times a large stream flowed across the state from west to east at about the locality of the Grand-Saginaw valley, is incorrect. He believes that the principal Preglacial stream had its rise in the vicinity of Saginaw bay, flowing west and north-west across the peninsula to the vicinity of Manistee, on the Lake Michigan shore, a route far to the north of the Grand-Saginaw valley.

This conclusion is based, first, on the fact that a limestone ridge is believed to extend clear across Saginaw bay at a level so high as to render a large drainage channel in that direction extremely improbable; second, on the condition of the rock surface in the vicinity of Manistee, which so far as known is extremely broken and uneven, characteristic of a locality well down the course of a stream rather than near its head; and third, deep borings at scattered intervals along the line find the rock-surface at lower levels than is known elsewhere in the state.

The writer has no preconceived theories which he is bound to maintain, and concedes that Prof. Lane's theory rests on coherent even if somewhat meager evidences. However, the element of speculation is necessarily large, and the problem can scarcely yet be considered definitely settled. I desire to present some new facts bearing on the question, which render it necessary to modify quite materially Prof. Lane's contour map of the rock surface (Plate VI of his pamphlet). As a matter of course, in the present state of knowledge, large areas of the rock surface being totally unknown, such a map must contain a large element of guesswork, and Prof. Lane, I am sure, will welcome any corrections of the same. His map

is constructed with contour intervals of 100 feet, the contour lines being arranged in harmony with the theory of a deep central valley in the location above indicated. The valley corresponds to the 300-foot contour (above sea level), the lines rising to the north and south. There is but a brief space to the south between the 300-foot and 400-foot contours; thence to the 500-foot contour the distance is from five to ten miles, and fifteen or twenty miles farther to the 600-foot line. The 600-foot line enters Ionia County near the northwest corner and traverses the county in a southeasterly direction, crossing Grand river a short distance east of Ionia City, at the locality of the outcrop of Carboniferous sandstone, known locally as the Ionia sandstone quarry. This outcrop in its bearing upon another problem has been described more fully in a previous paper.* It is the only outcrop in this part of the state, is located on the flat bottom of the river valley, and has an altitude of about 650 feet, or fifty feet higher than the contour line, which presumably is intended to represent the average altitude of the rock surface in this vicinity.

A comparison of this outcrop with some neighboring localities reveals some surprising irregularities. To the east the rock surface is quite unknown, but at Lyons, three miles east, a well has been sunk to a point a little below the 600-foot level, and another at Pewamo, six miles farther east, nearly to the 500-foot level, without penetrating rock. To the west, however, more definite results have been attained. In the city of Ionia, three miles down the river from the sandstone outcrop, two deep wells were put down two years ago, penetrating rock at about the 550-foot level. As Ionia is located between the 600-foot and 700-foot contour lines on the map, it is seen that the 600-foot line must be removed several miles to the southwest at this point.

But this is not all. Nine miles still farther down the valley the village of Saranac now has in process of construction a well which at last reports was down to a depth of more than 250 feet, the rock surface having been found at a depth of 248 feet. The location of the well has an elevation of about 650 feet, hence the rock surface is at about the 400-foot level. It follows that not only must the 600-foot contour line be shifted so as to include this locality, but also the 500-foot line, which appears on the map 25 miles to the north, while the 400-foot line would pass through Saranac instead of bordering the central valley 30 to 40 miles away. That is, the rock surface at Saranac is only about 100 feet above the bottom of the central valley.

* This Journal, November, 1895.

Probably the most natural conclusion to be derived from the above facts is that a branch of the main valley has its rise somewhere to the south of the Saranac. Prof. Lane's contours indicate a branch valley coming in from the south along the east line of Ionia and Montcalm Counties. As no deep borings are known along this line, it is presumed that such branch valley is simply a probability, based upon little known conditions. It would not require a very great change in the direction of this supposed branch valley to bring it in line with the Saranac boring.

It is not likely that the Saranac well indicates an east and west valley, because the rocks to the east rise rapidly, as shown by the Ionia wells and the sandstone outcrop above mentioned, while to the west in the vicinity of Grand Rapids the rock surface is above the 600-foot level. Neither is it probable that the valley opened to the south. The only other theory is that the main valley had become so extensive in Preglacial times as to include the Saranac location, but this presumes an amount of base-leveling out of harmony with the general situation. It may, therefore, be taken as at least fairly well settled that an important branch of the drainage system, probably not inferior to the eastern division, exists beneath the drift not far from the line indicated. Future borings along this line in Ionia and Montcalm Counties will be watched with interest.

Grand Rapids, Mich.

SCIENTIFIC INTELLIGENCE.

I. CHEMISTRY AND PHYSICS.

1. *The various forms of Sulphur.*—A paper by R. Brauns in a recent number of the *Jahrbuch für Mineralogie** discusses with much fullness the various forms assumed by sulphur on crystallizing from a state of fusion. Various new points of interest and importance are brought out, and the author concludes that there are eight modifications of sulphur to be recognized, as follows: (1) orthorhombic; (2) monoclinic-prismatic, Mitscherlich; (3) crypto-fibrous-radiated, with concentric structure, monoclinic? (No. 3 of Muthmann); (4) fibrous-radiated, monoclinic (soufre nacré of Gernez); (5) fibrous-radiated, orthorhombic; (6) triclinic; (7) monoclinic, in six-sided plates (Muthmann); (8) rhombohedral (Engel).

In addition to these he mentions, also, two forms of fluid fused sulphur and two or three forms of sulphur vapor, differing in their vapor density. The conclusion is, therefore, that sulphur can assume twelve states differing in the degree of energy characterizing them.

Another investigation on a similar subject, published just before this, is by O. Bütschli, Professor of Zoology at Heidelberg.† The author has been led, through his investigation of the various forms of microscopic structure belonging to different organisms, to extend his study also to inorganic substances. He here discusses the behavior of drops of sulphur produced from sublimation: he shows that a vaporization of the solid sulphur could be noted at 58° , and that at this temperature, as well as at the point of fusion, minute superfused drops are yielded which remain fluid for a long time, and solidify spontaneously or through pressure in the form of doubly refracting spherulites. The microscopic structure of these is given with minuteness, and a series of plates—one of them colored—shows the various forms observed.

2. *The Diurnal Variation of Atmospheric Electricity.*—For a period of seven years, from 1892 on, observations have been conducted on the summit of the Eiffel Tower in Paris, during the months from May to October, having as their object the determination of the electrical condition of the atmosphere; continuous observations have been also made at the *Bureau central* in Paris for a slightly longer period. A discussion of these by A.-B. CHAUVEAU shows that they lead to the conclusion that there exist in our temperate regions two very different types of diurnal variation near the surface of the earth. One corresponds to the

* Beilage-Band, xiii, pp. 39–89.

† Untersuchungen über Mikrostrukturen des erstarrten Schwefels nebst Bemerkungen über Sublimation, Überschmelzung und Übersättigung des Schwefels und einiger anderer Körper. Pages 96, with 6 text figures and 4 plates. Leipzig, 1900 (Wm. Engelmann).

warm, the other to the cold season. During the summer there is a pronounced minimum in the hot hours of the day, which is more pronounced the nearer the point of observation is to the surface of the earth, and the more it is free from the influence of trees or buildings. There is another minimum in the night, the importance of which varies in an inverse direction; the existence, however, of a double daily oscillation is well marked. During the winter the afternoon minimum disappears more or less completely, while that of the night increases; there is thus a single oscillation with a maximum by day and a minimum at night; the hour of the latter corresponds to that of the second minimum of summer, about 4 P. M. Observations at other localities show very much the same phenomena.

It is concluded, in the second place, that the diurnal variation at the summit of the Eiffel Tower during the summer differs entirely from that observed simultaneously at the *Bureau central* a few hundred meters distant, but offers a close analogy in its single oscillation to the variation observed in winter. Similar results are obtained at the summit of a wooden pylon used to support anemometers at the observatory at Trappes.

The final conclusions reached are to this effect: that the influence of the soil and the maximum in summer (due, according to the suggestion of Peltier, to the water vapor emanating from the surface, and negative like that) is the cause of the perturbation in the diurnal variation; that the general law of this variation is very simple, having a maximum in the day and a minimum between 3:30 and 4:30 A. M.—*Séances Soc. Franç. de Physique*, 1899, p. 91.

3. *Lehrbuch der Photochromie (Photographie der natürlichen Farben)* von WILHELM ZENKER. Neu herausgegeben von Prof. Dr. B. SCHWALBE. Mit einem Bildniss des Verfassers und einer Spectraltafel. Pages 157. Braunschweig, 1900 (Fr. Vieweg u. Sohn.—This volume is for the most part a re-publication of a work by Zenker (1829–1899) issued by the author, in 1868, but thus far little known. The sketch of the author's life by G. Krech, with which the volume opens, shows that he was a remarkable man of original talents, whose contributions to science covered a wide range of subjects. Early interested in the work of E. Becquerel (1848) on color photography, he devoted much time and study to the subject, and his results are given in the volume before us. The chief point which he regarded as being established was this: that silver chloride, particularly the violet silver sub-chloride, takes the same colors as the light rays which act upon it, and shows them both by reflected and transmitted light. This result was explained physically as due to the formation of stationary light-waves which caused the separation of particles of metallic silver in the silver chloride. It will be recognized at once that this is the line in which such brilliant work has since been done by Lippmann, and the relation of the experiments and theoretical discussion of Zenker to the labors of those who have followed him is given in the closing pages of the volume (133–157) by E. Tonn.

4. *On the Joint Transmission of Direct and Alternating Currents.*—In a paper presented before the recent meeting of the American Association, Dr. FREDERICK BEDELL discussed the simultaneous transmission of direct and alternating currents by the same conductor. He shows that each current acts as though it had the whole conductor to itself and the other current were absent. Further there results a saving in weight of copper and copper losses, which saving, as shown by calculation, may amount to as much as fifty per cent.

II. GEOLOGY AND MINERALOGY.

1. *Geology of the Narragansett Basin*; by N. S. SHALER, J. B. WOODWORTH and A. F. FOERSTE. Monograph XXXIII, United States Geological Survey, pls. i-xxxix, figs. 1-30, pp. v-xx, 1-402, Washington, 1899.—The investigations reported in this Monograph were begun as early as 1865 by Professor Shaler, at first in connection with his university classes in geology. Ten years ago the completion of the work for publication as a monograph was undertaken in connection with the United States Geological Survey, and has been completed with the assistance of Mr. Foerste, working in the southern section, and Mr. Woodworth, in the northern portion of the field. The study of the construction of the basin has led to one important conclusion, which requires special notice, which may be best expressed by quoting the words of Professor Shaler: "The judgment as to the nature of the mountain-building work rests in part upon observations—in the main unpublished—which I have made in other somewhat similar basins that lie along the Atlantic coast from Newfoundland to North Carolina. The general proposition that the basins are characteristically old river-valleys which have been depressed below the sea level, filled with sediments—the sedimentation increasing the depth of the depression—and afterwards corrugated by the mountain-building forces, will derive its verification in part, if at all, from the study of other troughs of the Atlantic coast. It may, however, fairly be claimed that the facts set forth in this memoir show that this succession of actions has taken place in the Narragansett field."

Mr. J. B. Woodworth has written a report upon the northern and eastern portion of the basin (pp. 99-214). In his report the pre-Carboniferous rocks recognized are classified as follows:

- Algonkian Period
 - Blackstone series
 - Cumberland quartzites
 - Ashton schists
 - Smithfield limestones.

Rocks of Cambrian age are recognized in drift pebbles, but not in place in the northern basin.

The rocks of the Carboniferous system are classified under the following names: from below upward:

Pondville group	100 ft.
Millers River conglomerates	
Wamsutta group	1,000 ft.
Wamsutta conglomerates	
Attleboro sandstone	
Wamsutta slates and shales.	

These rocks are regarded as below the Coal Measures which begin with the following :

Rhode Island Coal Measures	10,000 ft.
Cranston beds	
Pawtucket shales	
Sockanasset sandstones	
Mansfield beds	
Tenmile River beds	
Seekonk sandstones	
Westville shales.	
Dighton group	1,000-1,500 ft.
Seekonk conglomerate	
Rocky Woods conglomerate.	

The classification of the rocks of the southern field is made and reported on by Mr. Foerste (pp. 223-393). His basal conglomerate and arkose are the equivalent of Mr. Woodworth's Pondville group. The Wamsutta beds are not traceable south of Providence. The Kingstown series of Dr. Foerste corresponds in part to the Mansfield, Cranston, Sockanasset and Pawtucket beds of Woodworth, and Foerste's Aquidneck shales correspond to the Westville, Seekonk and Tenmile River beds of Woodworth. The Dighton formation of the north is called Purgatory conglomerate in the southern field. It is unfortunate that the two parts of a field so small as the Narragansett Basin could not receive the same nomenclature. An insect fauna was found in the Pawtucket shales, which has already been described by Dr. S. H. Scudder. Special reference is made to this paper in the following article. The plants are also chiefly from the Pawtucket shales, and the list of species already reported in this Journal (3d series, vol. xxxvii, p. 229) were identified as equivalent to the flora of the Upper Carboniferous of Pennsylvania, by Lesquereux. w.

2. *View of the Carboniferous Fauna of the Narragansett Basin*; by ALPHEUS S. PACKARD, Proc. Amer. Acad. Arts and Sci., vol. xxxv, No. 20, fig. 1, pp. 399-405, April, 1900.—The author having recently studied new material, collected by Mr. J. H. Clark, of Providence, from this coal field, enumerates the following animal remains at present known from the Narragansett Coal Basin:

- Spirorbis carbonarius*. Pawtucket plant beds.
- Impression of an Annelid?* Pawtucket plant bed.
- Impression of a plant or worm?* South Attleboro, Mass.
- Sections of worm holes*. One mile south of East Attleboro.
- Anthracomya arenacea* (Dawson) Hind.

- Track of a gastropod mollusc?* Pawtucket plant beds.
Protichnites narragansettensis, n. sp. Pebble, dark arenaceous shales, North Providence.
Remains of a Crustacean? Black shales, Valley Falls.
Ostrakichnites carbonarius (*Protichnites carbonarius*) Dawson. Boulder, red shale, South Attleboro.
Anthracomartus woodruffi Scudd. Pawtucket plant beds.
Mylicris packardii Scudd. Bristol plant beds.
Etoblattina, 9 species, Scudd. Pawtucket plant beds, etc.
Gerablattina scapularis Scudd. Pawtucket plant beds.
Gerablattina fraterna Scudd. Silver Spring, East Providence.
Rhaphidiopsis diversipenna Scudd. Cranston plant beds.
Paralogus aeshnoides Scudd. Silver Spring, East Providence.

“The presence of the Spirorbis and of the tracks of two marine Arthropods suggest that the Rhode Island plant beds, even if in general of fresh-water origin, were deposited where the sea had access to them. The presence of these marine fossils, with the fresh-water naiad, *Anthracomyia arenacea*, strongly suggests that the horizon of the black shales of Providence and also of the red and greenish beds of Attleboro, Mass., belong to the same horizon as those of the South Joggins of Nova Scotia, which is Upper Carboniferous, the rocks there consisting of sandstones and dark carbonaceous shales, frequently becoming reddish. The South Joggins shales also contain the remains of Anthrapalæmon, which should be looked for in the Narragansett Coal Measures. Thus far, then, the animal remains confirm Lesquereux’s reference of the dark plant-beds to the Upper Coal Measures.

These beds also appear to be higher in the series than the Middle Carboniferous Mazon Creek beds of Illinois, which contain a larger number of marine animals, viz: Belinuridæ (Euproöps, Prestwichia and Belinurus), besides Anthrapalæmon and Acanthotelson, together with the impressions of marine annelid worms.”

3. *Geological Survey of Canada. Summary Report of the Geological Survey Department for the year 1899*; by GEORGE M. DAWSON, Deputy Head and Director, pp. 1–224, Ottawa, 1900. —The Canadian Survey was conducted with its characteristic energy during the year 1899. The finding of gold in the Yukon District, and the search for petroleum on the Saskatchewan has given these regions special importance, but activity has been shown in all parts of the country. Sixteen field parties were engaged during the summer, viz: in British Columbia, 3; Yukon District, 1; Great Slave Lake, 1; Alberta (boring operations), 1; Saskatchewan, 1; Ontario, 3; Ontario and Quebec, 1; New Brunswick, 2; Nova Scotia, 2; Ungava (East coast of Hudson Bay), 1; while several of the staff were engaged on special investigation upon collections, or materials, requiring laboratory study.

The investigations in the Yukon District, under direction of

Mr. McConnell, have been carried to considerable detail on account of the economic importance of the field. Gold is found in rich quantities in the stream-gravels, and "the product of a few of the 500-foot claims on Eldorado and Bonanza Creeks will exceed a million dollars each; while a considerable number on the same two creeks (in fact, the majority of the lower Eldorado claims and a few on Hunker Creek) will yield over half a million each, and claims running from a quarter to half a million are common on all these creeks and also on Dominion and Sulphur Creeks. Assuming a quarter of a million as the average, and that three-quarters of the claims in the district given above are rich enough to work, the total value approaches \$95,000,000, a figure which is well within the mark."

The borings in Northern Alberta, in attempts to reach the petroleum-bearing strata at the base of the Cretaceous, have still failed to reach the "tar-sands," which, it is estimated, lie at a depth of about 2000 feet in the Victoria region. Only 1840 feet have been penetrated, owing to the extreme difficulties of holding the bore-hole open at great depth. The soft and incoherent character of the great mass of the overlying rocks requires casing to be carried on *pari passu* with the drilling. It is estimated that the Victoria bore-hole is down to within 250 feet of the "Tar-sands." At Athabasca Landing, the bore-hole reached to very near the top of the Tar-sands. At the Pelican River locality the Tar-sands were reached, and penetrated 87 feet before the gas and tar closed the hole and the working. It is believed that the underlying Devonian limestones are the natural source of the petroleum or maltha accumulated in the "Tar-sands."

Gold was discovered in the alluviums of Serpentine River, New Brunswick, but in small quantities. In York County, New Brunswick, Silurian fossils have been discovered, indicating the probable inclusion of Silurian rocks in folds of older Cambro-Silurian rocks which are more or less altered by metamorphism. Gold-bearing rocks were exploited in several localities in Nova Scotia. The Plant beds of Harrington River, Nova Scotia, were shown to be of Devonian age. Upper Cambrian fossils were discovered in the older Paleozoic rocks of Cape Breton Island by G. F. Matthew. Trenton fossils from Akpatok Island, Ungava Bay, already reported by Dr. Bell in this Journal (June, 1899), are referred to. Marine fossils of the Windsor series from Cumberland County, Nova Scotia, were brought to light by Dr. Ami. Discussion of the age of the Riversdale, Horton and Plant beds of St. John of New Brunswick and Nova Scotia, tend to confirm their Carboniferous affinities. The fish of McArra's Brook, Antigonish County, N. S., are correlated with the Old Red hornstones of Hereford district of England, above the passage beds, by A. Smith Woodward. The appropriation used by the Survey was a little over one hundred and eighteen thousand dollars.

9. *Fossil Flora of the Lower Coal Measures of Missouri*; by
w.

DAVID WHITE, Monograph, vol. xxxvii, U. S. Geological Survey, pp. 1-467, plates i-lxxiii, 1899.—The material upon which this monograph is based was collected from the Coal Measures of Henry County, Missouri, by Dr. J. H. Britts of Clinton, Mo., and by Mr. Gilbert VanIngen and Dr. W. P. Jenney of the U. S. Geological Survey. The whole flora has been exhaustively studied, many new species described, and the old species subjected to careful criticism and revision.

Comparison of the flora with those of other regions in the United States and other countries has led to the following determinations of correlation of horizon for the Missouri flora, viz: "The Lower Coal Measures of Missouri, as represented by the coals of Henry County, were laid down soon after the Morris coal of Illinois, though probably earlier than the Upper Kittanning of western Pennsylvania, or very likely about the time of the formation of the D coal in the Northern Anthracite field."

In the European sections, close affinity is found between this flora and that of the zone of Bully-Grenay in the Valenciennes Basin, but as the author remarks, the fauna is perhaps in a measure transitional, "while it is probably contemporaneous with a portion at least of the upper zone of the Valenciennes Basin, as presented in the basins of Commentry or the Saar, . . . it may represent a slight paleontological transgression on the Stephanian ('Houiller supérieur')." It is represented in Silesia and Bohemia by the Schatzlar and Radnitzer Schichten, and corresponds with the "Transition Series" of Great Britain. w.

5. *La Face de la Terre (Das Antlitz der Erde)* par ED. SUESS, translated by EMMANUEL DE MARGERIE, vol. ii, plts. 1 and 2, figs. 1-128, pp. 1-878. Paris, 1900 (Armand Colin & Cie).—Those familiar with the original edition of Professor Suess' *Antlitz der Erde* will not need to be reminded of the extreme value of the French edition, the second volume of which has recently appeared. It furnishes a store-house of information regarding the structure of the land surfaces and the geological structure of the earth not met with elsewhere, except as distributed in a very large number of publications. The first volume has already been noticed in this Journal (see vol. v, p. 152). In the translation, M. de Margerie has had the assistance of MM. Bernard, Depéret, Kilian, Poirault, Six and Zimmerman. The translations are admirably made, and the editors have added 85 new cuts to the French edition. These are many of them charts, and add much to the value of the volume. The translators have also added, in foot-notes, a large number of new references to literature, bringing the bibliography up to 1899, which will increase the value of the book for investigators. The present volume discusses the oceans and the geological conditions of the land and ocean surfaces during the Paleozoic, the Mesozoic, the Tertiary; and in the latter part of the volume, a number of chapters are devoted to special subjects, such as the Temple of Sérapis, Baltic and North Seas, the Mediterranean during historic period, etc. w.

6. *The Gneisses, Gabbro-schists and associated Rocks of South-western Minnesota*; by C. W. HALL. (Bull. U. S. Geol. Surv. No. 157, 1899, pp. 160, pl. xxvii.)—The territory from which the material described in this paper was gathered embraces the valley of the Minnesota River and the southwest corner of the State of that name wherever gneisses are found. In the 120 miles of the river valley there project through its flood plain many exposures of gneisses, gabbro-schists, diorites and diabasic effusives. They often fill the entire valley one to two miles wide, rising in a profusion of knobs and hills 50 to 100 feet above the river. These exposures of crystalline rocks are described by districts, with maps of each district, on which they are located. Their geological relations are discussed and an account of their petrography, which appears to be that of well known types, is given with the addition of numerous colored plates. L. V. P.

7. *Brief Notices of some recently described Minerals*.—MÜLLERITE is a hydrosilicate of ferric iron described by Zambonini from Nontron, France, which is the original locality of the related nontronite (chloropal). It occurs in opaque incrusting masses of a yellowish green color; it is quite soft and has a specific gravity of 1.97. The mean of three analyses gave:

SiO ₂	Fe ₂ O ₃	Al ₂ O ₃	MnO	MgO	H ₂ O
48.82	35.88	4.30	0.63	0.35	9.66=99.64

This corresponds to the formula $\text{Fe}_2\text{O}_3 \cdot 3\text{SiO}_2 + 2\text{H}_2\text{O}$, which differs from the accepted composition of nontronite only in having three molecules less water.—*Zeitschr. Kryst.* xxxii, 157.

MELITE is a name given by the same author (l. c., p. 161) to partially investigated hydrosilicate of alumina and ferric iron known from a single specimen, labeled "Allophane, Saalfeld, Thuringia." It occurs in imperfect crystals (?) and stalactitic forms of bluish-brown color and opaque; the hardness is 3 and the specific gravity 2.18. An analysis gave: SiO₂ 14.97, Al₂O₃ 35.24, Fe₂O₃ 14.70, CaO 0.78, H₂O 33.75=99.64.

ROBELLAZITE. Briefly mentioned by M. Cumenge as occurring with the carnotite in Colorado. A partial chemical examination by M. Debiegne, it is stated, shows that it contains considerable vanadium and also niobium, tantalum and tungsten with alumina, iron and manganese as bases.—*Bull. Soc. Min.*, xxiii, 17, 1900.

CUPRO-GOSLARITE. A cupriferous variety of goslarite described by A. F. Rogers as forming a light greenish blue incrustation in an abandoned zinc mine at Galena, Kansas. An analysis gave: SO₂ [27.02], ZnO 23.83, CuO 6.68, FeO 0.13, H₂O 41.76, insol. 0.58=100.—*Kansas Univ. Q.*, viii, No. 2.

CUBOSILICITE is a name given by Bombicci to the bright blue form of silica in cubic crystals occurring at Tresztyan, Transylvania, and ordinarily regarded as pseudomorphous chalcedony, perhaps after fluorite. This he classifies as a definite form of silica (pseudo-isometric, mimetic) related to melanophlogite, sulfuricine and cristobalite.—*Mem. Accad. Bologna*, viii.

III. MISCELLANEOUS SCIENTIFIC INTELLIGENCE.

1. *American Association for the Advancement of Science—New York Meeting.*—The forty-ninth meeting of the American Association was held in New York City, during the week from June 25 to 30. The President of the meeting was Professor R. S. Woodward. The attendance was satisfactorily large, though perhaps not so numerous as the time and place of meeting might have seemed to promise; the total registration was about 450. In scientific spirit, however, the occasion was highly successful, the number of the papers presented being large and their character and interest above the average. The sessions of the Association were held in the buildings of Columbia University and the efforts of the officers of this institution and of those of the American Museum of Natural History contributed largely to the general success of the meeting.

The retiring President, Mr. G. K. Gilbert, delivered on Tuesday evening an admirable address upon the subject "Rhythms and Geologic Time." Other addresses were delivered by the Vice Presidents of several of the sections, as follows: Section A, on the teaching of Astronomy in the United States, by Asaph Hall, Jr.; Section B, on the Cathode Rays and some related phenomena, by Ernest Merritt; Section C, on the eighth group of the Periodic System and some of its problems, by Jas. Lewis Howe; Section D, on Some Twentieth Century problems, by William Trelease; Section E, on Precambrian sediments in the Adirondacks, by J. F. Kemp. These addresses are published in full in *Science* (issues of June 29 and following). In addition to the work before the different sections, fifteen affiliated societies had meetings in connection with the Association, and this fact added much to the scientific interest of the occasion.

The place selected for the meeting of 1901 is Denver, Colorado; the meeting will begin on Monday, August 26th. Pittsburg was recommended for the meeting of 1902. The officers elected for next year are as follows:

President: Charles S. Minot, Harvard Medical School.

Vice-Presidents: Section A, James McMahan, Cornell University; Section B, D. D. Brace, University of Nebraska; Section C, John H. Long, Northwestern University; Section D, H. S. Jacoby, Cornell University; Section E, C. R. Van Hise, University of Wisconsin; Section F, D. S. Jordan, Leland Stanford University; Section G, B. T. Galloway, U. S. Department of Agriculture, Washington; Section H, J. W. Fewkes, Bureau of Ethnology, Washington; Section I, John Hyde, Department of Agriculture, Washington.

Permanent Secretary, L. O. Howard of Washington; General Secretary, William Hallock, Columbia University; Treasurer, R. S. Woodward, Columbia University.

2. *British Association for the Advancement of Science.*—The coming meeting of the British Association will be held at Brad-

ford, England, commencing Sept. 5. The President is Professor Sir William Turner of Edinburgh.

3. *Catalogues of the Collections in the British Museum of Natural History*.—Recent issues of this series include the following :

The Cretaceous Bryozoa, volume i, pp. xiv, 457 with 17 plates ; by J. W. GREGORY, London, 1899. This follows an earlier work (1896) on the Jurassic Bryozoa, in which there was given an Introduction upon the structure and affinities of the group. The second volume upon the Cretaceous Bryozoa is promised for the present year.

Catalogue of the Arctiadae (Nolinae, Lithosiaceae) ; by Sir GEORGE F. HAMPSON. Pp. xx, 589 ; plates xviii-xxxv. London, 1900. This is the second of the volumes devoted to the Lepidoptera Phalænæ.

4. *The Norwegian North-Atlantic Expedition, 1876-1878*.—The following publications have recently been issued, containing the results of the further study of the zoological collections made by the Norwegian North Atlantic expedition.

XXV, Thalamophora by Hans Kiær, with 1 plate and 1 map.

XXVI, Hydroida by Kristine Bonnevie, with 3 figures, 8 plates and 1 map.

XXVII, Polyzoa by O. Nordgaard, with 1 plate and 1 map.

5. *The Grammar of Science* ; by KARL PEARSON, 2d ed. revised and enlarged, pp. 1-548, figs. 1-33 (Adam and Charles Black, London ; Macmillan Company, New York), 1900.—This book illustrates the undoubtedly strong trend of opinion among leaders of scientific thought toward some form of idealism. The inadequacy of the cruder forms of materialism to satisfy the questioning of the thinking mind has led men to look inward to the form of their conceptions of phenomena, in order to ascertain the relation these phenomena bear to each other.

The following passage expresses tersely the author's purpose : "The object of the present work is to insist . . . that science is in reality a classification and analysis of the contents of the mind ; and the scientific method consists in drawing just comparisons and inferences from the stored impresses of past sense-impressions, and from the conceptions based upon them. Not till the immediate sense-impression has reached the level of a conception, or at least a perception, does it become material for science. In truth, the field of science is much more consciousness than an external world. In thus vindicating for science its mission as interpreter of conceptions rather than as investigator of a 'natural law' ruling an 'external world of material,' I must remind the reader that science still considers the whole contents of the mind to be ultimately based on sense-impressions" (p. 52). The volume is full of keen observations and suggestions. The chief additions made in this edition are the chapters discussing the biological conceptions of science.

THE

AMERICAN JOURNAL OF SCIENCE

[FOURTH SERIES.]

ART. XVIII.—*On the Gas Thermometer at High Temperatures*; by LUDWIG HOLBORN and ARTHUR L. DAY.

Second Paper.

[Communication from the Physikalisch-Technische Reichsanstalt, Charlottenburg, Germany.]

IN a former paper* we described our measurements with the gas thermometer in some detail. The investigation was concerned principally with establishing the conditions under which this instrument may be used with certainty as a standard at high temperatures. We have shown that this is possible using pure nitrogen as the expanding gas in a bulb of platinum-iridium, this material being much superior to the porcelain so long in use for the purpose, both in the accuracy of the results obtainable and the convenience in handling.

Since then the work has been continued in the direction indicated at the close of the previous paper. In order to obtain the full advantages of the exactness which the use of the platinum-iridium bulb has rendered possible, the first step was to measure the undetermined coefficient of expansion of this metal at high temperatures. This was very necessary because the correction which the observations with the gas thermometer require on account of the expansion of the bulb increases more rapidly than the temperature to be measured. With the platinum-iridium bulb, for instance, it amounts to 10° at 500° , 30° at 1000° and 40° at 1150° , while the increase in the coefficient of expansion with the temperature, which of course is not taken into account in the earlier paper, affects the measurement 1° , 5° , and 7° at the above temperatures respectively.

Secondly, the influence of the pressure which the expanding gas exerts on the glowing walls of the bulb was investigated

* Ludwig Holborn and Arthur L. Day, this Journal (IV), viii, 165, 1899.

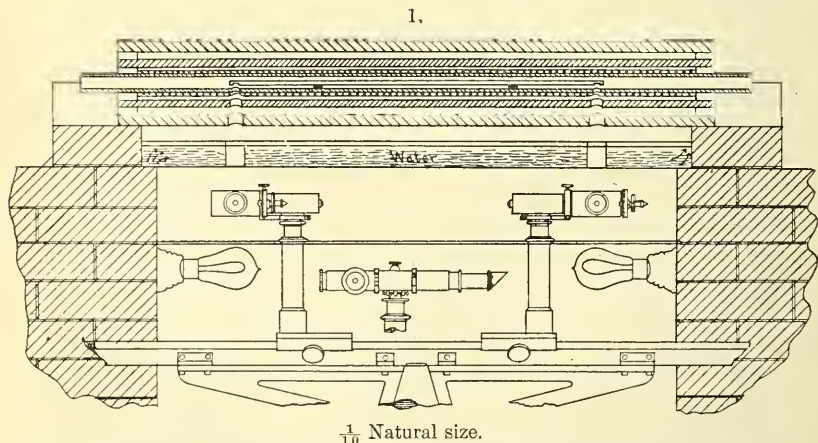
by a simple method, and finally a second platin-iridium bulb with walls of double the thickness of the first was used in a somewhat different oven in verification of the earlier results. The use of the same thermo-element throughout enabled the observations to be compared with those of the previous series. The thermo-electric force of this element (platinum—platin-rhodium), like the others involving the use of combinations of metals of the platinum group, may be represented by an equation of the second degree within the limits of temperature covered by the investigation (250–1150°). The form of the curve is therefore known when such elements are calibrated at three known temperatures.

To render such calibration independent of the gas-thermometer, a series of melting points of pure metals was determined in which we took advantage of the opportunity to investigate rather carefully the conditions which may affect the melting temperature in certain cases.

To this is further added a chapter on the measurement with thermo-elements, the accuracy obtainable and their lasting qualities.

1. *Determination of the Expansion.*

To determine the expansion of the platin-iridium bulb, a bar of the same alloy (80 Pt, 20 Ir) was obtained, 500^{mm} in length and 5^{mm} in diameter and its linear expansion studied up to 1000°. In connection with this we also undertook the measurement of the expansion of several other substances at high tem-



peratures and hope to communicate in detail the full results of the investigation at another time, restricting ourselves for the moment to such as are necessary to the discussion of the gas thermometer. Slight changes which the figures here commu-

nicated may undergo in the further investigation of the problem cannot sensibly affect the matter in hand.

The method may be described briefly as follows: The bar is laid for the measurement of its expansion symmetrically in an electrically heated fire-clay tube, 80^{cm} long, in which two lateral holes are bored 10^{mm} in diameter and 485^{mm} apart.

This tube is protected against loss of heat by radiation by two concentric fire-clay jackets with corresponding lateral openings, and the whole placed in a horizontal position upon two brick pillars in such a way that the openings are directed downward. Underneath these, two microscopes are mounted with the aid of which readings can be made of the position of certain marks cut upon a flat polished surface near the ends of the expanding bar.

The heating coil which is wound upon the innermost tube extends for a good distance beyond the ends of the bar so that the temperature opposite the opening differs in maximum only 10 per cent from that at the middle. On account of the increase in the coefficient of expansion with the temperature an average temperature for the bar is not sufficient if considerable local differences exist, hence it is eminently desirable that the differences of temperature along the bar be not too great. We are accordingly making an effort to improve upon our present coil in this direction in the hope that an even greater uniformity in the temperature distribution may be secured.

The microscopes were mounted upon sliding bases upon a long horizontal support and could be fixed in the proper positions with clamping screws. To protect them from the radiation of the oven a diaphragm through which water circulated constantly was introduced between.

With this apparatus the expansion was then measured with the eyepiece micrometers of the microscopes, the reading before heating being verified after the oven had cooled down again. No measurable displacement of the microscopes was observed throughout the series but displacements of the bar occurred now and then at the higher temperatures; symmetrical observations upon both ends, however, eliminate any error from this cause in the results.

The observations were usually made at 250°, 500°, 750° and 1000°, and for the same temperature upon different days agree to within $\pm 0.01^{\text{mm}}$, whereby it should be remarked that one whole turn of the micrometer screw corresponds to about 1^{mm}.

The divisions, which were situated in the axis of the bar to diminish the effect of a possible slight bending or warping, were cut with the dividing engine, rather deep but with sharp edges, and rubbed with Parisian red (iron oxide). They were

illuminated from below with incandescent lamps. After the temperature has reached a bright red heat where the bar itself gives out light, the divisions still show very sharply by contrast with the polished metal surface as bright marks upon a dark ground.

But with the metals which oxidise and even with platinum-iridium which tarnishes before reaching 1000° , it is better to illuminate the glowing rod from outside also. The divisions then appear as dark lines upon a bright ground exactly as at the lower temperatures.

The temperature of the bar thus heated was measured at nine different points with a thermo-element, both wires being insulated with thin porcelain tubes so that the junction could be brought to any desired point along its length.

Platin-iridium (80 Pt, 20 Ir).—The following values were obtained for the expansion λ_s of the platin-iridium bar, which together with a similar bar of pure platinum mentioned below, was kindly loaned by the firm of Heräus in Hanau:

t	λ_s (mm)	
	Observed.	Calculated.
250°	1·033	1·033
500	2·150	2·151
750	3·357	3·355
1000	4·645	4·645

The mean length included between the two groups of five divisions each at the two ends of the bar was $483\cdot07^{\text{mm}}$ at 0° . From these figures the expansion λ of a bar of unit length at 0° may be represented by the formula

$$\lambda = [8198t + 1\cdot418t^2] 10^{-9}.$$

Platinum.—We had no bar of the alloy 90 Pt, 10 Ir, of which the second bulb was made, but one of pure platinum was measured under the same conditions and yielded the following results, which we combined with the foregoing and used the mean for correcting the gas thermometer measurements made with the bulb of the lower alloy:

t	λ_s (mm)	
	Observed.	Calculated.
250°	1·114	1·113
500	2·309	2·304
750	3·570	3·571
1000	4·909	4·914

The length at 0° was $483\cdot52^{\text{mm}}$ and the corresponding formula for a unit bar takes the form

$$\lambda = [8889t + 1\cdot274t^2] 10^{-9}$$

The difference between the observed expansions and those calculated according to the formula amounts in both cases to less than 0.01^{mm} .

The largest measured difference between the observations upon the cold bars before and after heating to 1000° was 0.025^{mm} ; it was generally much smaller.

Porcelain.—The expansion of Berlin porcelain was also measured upon a bar of similar dimensions—unglazed in order that it might afterward be used at still higher temperatures. Plane surfaces were ground upon each end in the same positions as for the metallic bars, and upon these, divisions about 0.15^{mm} wide but very sharply defined were cut upon a lathe with a thin copper disc, five at each end. The readings at the highest temperature were at first somewhat difficult, for the divisions are but little brighter than the background. With illumination from outside, however, excellent readings could still be obtained after the candle power had been increased.

This bar had a mean length at 0° of 483.47^{mm} and yielded the following results for the expansion :

t	λ_s (mm)		λ (mm)
	Observed.	Calculated.	
250°	0.382	0.391	0.790
375	0.628	0.612	1.299
500	0.845	0.850	1.749
625	1.093	1.105	2.261
750	1.307	(1.377)	2.703
875	1.615	(1.666)	3.340
1000	1.977	(1.972)	4.089

Observations were made with the porcelain bar at temperatures between the original four, when it was found that the observed expansions could not be represented by a curve of the second degree. The following formula is approximately true between 250° and 625° ,

$$\lambda = [2954 t + 1.125 t^2] 10^{-9}$$

and yields the observed value again at 1000° , while for 750° and 875° the differences are considerably larger than the errors of observation. Below 250° also the values obtained from the formula are too large. The coefficient of expansion clearly increases more rapidly after red heat is reached than below 500° , and for the interval between remains nearly constant. The observed value at 1000° coincides with the calculated value again, which shows this temperature to be the point of intersection of two curves.

It may be further remarked that the various series of observations, after the preliminary heating to remove any tension

remaining from the grinding, agree quite as well for porcelain as for the metals. These observations were all made with a rising temperature, however, and it is not impossible that different values would be obtained for the expansion of porcelain if the temperatures observed were approached from the other direction, i. e. in the order 1000° , 875° , 750° , etc. It also remains to be investigated what differences exist between different bars of the same manufacture.

The expansion of Berlin porcelain formed the subject of an earlier investigation by Holborn and Wien,* who heated porcelain plates 9^{cm} in length in a gas furnace and measured them at the temperatures 550° and 1050° nearly. The mean value then arrived at, $44 \cdot 10^{-7}$ for the linear expansion coefficient, agrees with the present value at the higher temperature within the limits of accuracy then obtainable; at the lower temperature the earlier value is too large.

More recently Bedford† determined the expansion of French porcelain (Bayeux) between 0° and 800° . His formula is

$$\lambda = [3425t + 1 \cdot 07t^2] 10^{-9}$$

A comparison with our results is not instructive, however, for the expansions of porcelains from entirely different sources are not properly comparable.

Our measurements with the gas thermometer and porcelain bulbs, for the calculation of which the Holborn and Wien value was used, require only a slight correction for the new determination of the expansion, partly because the absolute expansion is very small and partly because the use of too large a value for the bulb expansion between 0° and 100° makes it also necessary to correct (diminish) the coefficient of expansion of the gas about $\frac{1}{100}$. The required correction, therefore, remains less than 1° for the whole range from 500° to 1000° .

Jena Glass 59^{III}.—For the measurement with the gas thermometer below 500° a bulb of Jena glass 59^{III} and a nitre bath were used. We have, therefore, undertaken to measure the expansion of this material up to 500° , but the results which have so far been obtained only serve to emphasize how much depends upon the way the glass is treated. We were able, however, to establish the fact that the mean error of the gas thermometer observations in this region, $\pm 0 \cdot 5^{\circ}$, is not exceeded in assuming for all temperatures up to 500° the constant value $18 \cdot 10^{-6}$. For more accurate measurements it is certainly advisable to employ the platin-iridium bulb for these temperatures also, the expansion of which is much better defined. In the nitre bath it will be necessary to protect it with a thin glass tube, however, as the bath becomes alkaline with use.

* L. Holborn and W. Wien, Wied. Ann., xlvii, 121, 1892.

† T. G. Bedford, Phil. Mag. V, xlix, 90, 1900.

2. *Influence of Pressure on the Gas Thermometer Bulb.*

At the highest temperatures measured the inside pressure in the platin-iridium bulb exceeded that outside by nearly an atmosphere. Although no permanent change in the volume of the bulb after heating could be detected, it seemed, nevertheless, not impossible that a temporary yielding of the glowing walls under the comparatively high pressure might take place. We have investigated this question by redetermining the higher temperatures with a much diminished pressure.

This was accomplished by allowing a portion of the expanding gas to pass over into the closed tube of the manometer, thus varying pressure and volume simultaneously. The expansion-coefficient of the gas suffers a small change thereby which we estimated to be about 0.1 per cent and therefore neglected it.

The closed manometer tube had already been provided with two points (see first paper—loc. cit.) at which readings could be made for volumetric purposes, so that by bringing the mercury tangent to the lower point a definite volume could be added to that of the bulb. To enable a more accurate measurement of the temperature of this supplementary volume to be obtained than was possible with the small thermometer in the cap, a second and more sensitive thermometer was fastened to the tube and the whole carefully packed in a thick layer of cotton batting containing only the two small openings necessary to make the readings above and below.

As a rule the observations were made in sets of three, the first and last with the mercury adjusted to the lower, the second to the upper point. Since it was necessary to wait some time between observations until the temperature conditions in the manometer became uniform, the temperature in the oven had opportunity to change a few degrees. These changes (δ) were measured with the thermo-elements. The supplementary volume between the two points was measured by weighing the mercury required to fill it, the results being as follows:

50.630 ^{cm³}
50.609
50.641
50.658

Mean 50.635^{cm³}

The differences between the determinations are largely due to an imperfectly formed meniscus at the lower point, the adjustment being necessarily made with a falling column as the mercury was drawn off through a fine capillary tube. The adjustment with a rising column as used for the temperature measurements gives much more exact results. Also the deter-

mination of the distance between the two points, 161.60^{mm} contained variations of only 0.02^{mm} in the individual readings. An error of 0.1 per cent in the determination of the supplementary volume would cause an error of 1° in the temperature at 1000° as determined at the lower point.

For the verification of the method several measurements of the volume of the bulb at 0° were made. For the first platinum bulb the determination by filling with water and weighing yielded 208.49^{cm³}, the volumetric determination 208.45 and 208.51^{cm³}; for the second bulb the results obtained by the two methods were 195.87 and 195.91^{cm³} respectively.

Table XII contains the temperature measurements at the two different pressures. *H* and *t* represent the pressure meas-

TABLE XII.

1899	<i>H</i> (mm.)	<i>t</i>	<i>H'</i> (mm.)	<i>t'</i> - δ	δ	<i>t</i> - <i>t'</i>
July 22	871.71	553.0°	524.13	550.9°	+2.0°	+0.1°
			526.38	556.6	-3.4	-0.2
	877.37	558.7	526.80	557.8	+1.5	-0.6
			528.35	562.0	-2.9	-0.4
	1339.50	1022.7	663.06	1015.4	+6.3	+1.0
			666.71	1030.1	-7.6	+0.2
Aug. 2	848.68	531.1	514.76	527.9	+2.5	+0.7
			517.07	533.8	-2.9	+0.2
	855.11	537.4	517.76	535.5	+1.6	+0.3
			519.05	539.1	-1.6	-0.1
1047.86	728.2	581.57	721.9	+5.3	+1.0	
		585.81	734.9	-7.3	+0.6	
1055.69	736.1	585.80	734.8	+1.1	+0.2	
		586.89	737.9	-1.9	+0.1	
Aug. 3	952.97	633.9	551.09	629.4	+3.5	+1.0
			558.78	651.4	-18.4	+0.9
	974.06	654.8	559.11	652.5	+1.6	+0.7
			560.58	656.7	-2.4	+0.5
1242.43	924.5	638.41	921.4	+1.5	+1.6	
		639.87	926.7	-3.1	+0.9	
1243.62	925.8	639.62	925.5	-0.6	+0.9	
		639.63	925.6	-0.6	+0.8	

ured upon barometer and manometer, and the temperature calculated from it for the readings made at the uppermost point; *H'* and *t'* - δ , the corresponding magnitudes for the readings

below. If to $t' - \delta$ the change in the oven temperature δ be added, the results contained in the column $t - t'$ or the difference between the results obtained at the two pressures is obtained. The mean value of this difference is $\pm 0.7^\circ$ and exceeds 1.0° only in a single instance.

It follows from this that within the given limits of accuracy, the measurements with the gas thermometer equipped with the platin-iridium bulb were not disturbed by the pressure upon the walls of the hot bulb. The initial pressure of the gas at 0° was 294.40^{mm} , nearly the same as in the earlier measurements, and its coefficient of expansion 0.003666 .

It has been suggested that this method be employed for obtaining the volume of the bulb at high known temperatures. It would then be possible to arrive at its coefficient of expansion in this way. In addition to the usual equation for the calculation of the temperature with constant volume,

$$\frac{HV}{1+at} = H_0V_0$$

a second relation is, to be sure, obtainable

$$H' \left(\frac{V}{1+at} + \frac{V_i}{1+at_i} \right) = H_0V_0$$

if t_i be used to represent the temperature of the supplementary volume V_i . The equations are nevertheless not independent of each other and are therefore not sufficient for the determination of the two unknown quantities V and t .

3. Comparison of the Thermo-element with the Gas Thermometer.

Second Platin-iridium Bulb.—The second platin-iridium bulb had nearly the same form and size as the first. It was, however, made from the alloy 90 Pt. 10 Ir., and its walls (1^{mm}) were twice as thick as those of the earlier bulb. Its volume at 0° was 195.87^{cm^3} .

After being boiled with concentrated nitric acid and several times rinsed with distilled water the bulb was carefully dried and attached to the manometer. Then it was evacuated with a mercury pump and maintained for several hours at a temperature of 1300° , being "rinsed" from time to time with fresh nitrogen and again evacuated. The gas for the final filling was also admitted at the high temperature and the exact pressure regulated after the bulb had cooled.

The new bulb was used for several series of measurements and proved to be quite as satisfactory as the old. The gas pressure at 0° remained constant to within 0.1^{mm} throughout.

The comparisons of the gas thermometer and the thermo-element T_2 , which are contained in Table XIII agree among themselves exceedingly well and with the earlier comparisons also. The differences, which amount to about 1.0° , are of the same order of magnitude as in all the other measurements made under the given conditions, the bulb with its extended volume and the minute thermo-electric junction being contained in an air bath. The highest temperatures are most favorable in this respect, as the radiation acts most strongly there to equalize the temperature distribution.

Below 500° where the air bath gave place to nitre, the uniformity is also greater of course.

The conditions under which the second bulb was heated were also considerably varied in order that the substantiation of the earlier results might be as complete as possible. The observations under "October 12" in the table, for example, were made with the bulb in the same oven which had served for the measurements with the first bulb, the others in a newer and larger one, but with two different heating coils on the different days—both wound logarithmically as described in the former paper. The tubes carrying these new coils were 42cm long and 6cm inside diameter while the older one was only 35cm long and 4.8cm in diameter. The fire-clay jackets were also correspondingly larger.

TABLE XIII.

Platin-iridium Bulb No. II (90Pt, 10Ir) Gas-Nitrogen.

 $V_0=195.87\text{cm}^3$ $V_1=0.904\text{cm}^3$ $H_0=276.35\text{mm}$ $a=0.003666$

1899	t	$e_2(\text{MV})$	Obs.— Calcul.	1899	t	$e_2(\text{MV})$	Obs.— Calcul.
Oct. 9	562.1°	4759	0.0°	Oct. 12	822.9°	7497	-1.5
	571.7	4861	-0.5		822.5	7496	-1.9
	573.2	4877	-0.6		1093.5	10558	-0.7
	829.5	7568	-1.5		1096.1	10598	-1.5
	1080.9	10423	-1.8		1097.4	10616	-1.7
" 12	552.0	4668	-1.0	" 21	669.1	5836	+1.1
	552.2	4670	-1.0		917.4	8526	-0.2
	552.6	4674	-1.0		1063.3	10189	+0.5
	821.2	7478	-1.5				

In Table XIII, t represents the temperature observed with the gas thermometer, e_2 the thermo-electric force of the element T_2 in microvolts. The last column (Obs.—Calcul.) contains the difference between t and the temperature in degrees calculated from the curve for e_2 . This curve will be referred to again further on.

The Observations with the first *Platin-iridium Bulb*.—The earlier observations (loc. cit. Tab. IX, p. 189) with the first bulb (80 Pt, 20 Ir), which were calculated with a constant coefficient of expansion, can now be corrected for the increase in the expansion with the temperature. The corrected values are contained in Table XIV, the numbers printed in italics in Table IX being here omitted. They were made with a uniformly wound oven coil, and a comparatively large (former paper, Table VIII, p. 189) fall in temperature along the bulb, roughly corrected with the help of the thermo-elements.

TABLE XIV.
Platin-iridium Bulb I (80Pt, 20Ir). Gas-Nitrogen.

After filling.	<i>t</i>	<i>e</i> ₂ (MV)	Obs.— Calcul.	After filling.	<i>t</i>	<i>e</i> ₂ (MV)	Obs.— Calcul.
2d day	529·1°	4429	+0·3°	22d day	1052·8°	10073	-0·1°
	614·8	5289	+0·1		1102·5	10653	+0·4
	1008·8	9555	+0·6	24th "	541·2	4552	+0·1
20th "	625·6	5401	-0·1		542·8	4569	-0·2
	812·1	7356	+0·6		615·5	5297	0·0
	907·3	8396	+1·2		717·3	6343	+0·7
	1014·8	9622	+0·9		814·3	7385	-0·2
	1113·9	10770	+1·7		919·7	8550	+0·2
22d "	511·8	4252	+1·1		1017·6	9661	+0·1
	512·7	4268	+0·2	1106·6	10700	+0·4	
	551·5	4655	-0·2	25th "	516·4	4304	+0·2
	659·3	5754	-0·8		616·3	5303	-0·1
	703·8	6217	-0·9		721·3	6382	+0·7
	757·0	6775	-0·5		821·9	7462	+1·6
	859·4	7881	-0·3		913·4	8458	+1·8
	949·2	8890	-0·6		1026·0	9756	+0·5
	1001·5	9496	-1·6	1134·5	11031	+0·7	
	1050·4	10044	-0·1				

The temperatures *t* are raised by the new coefficient of expansion 0·7° at 500°, 4·9° at 1000° and 7·5° at 1150°.

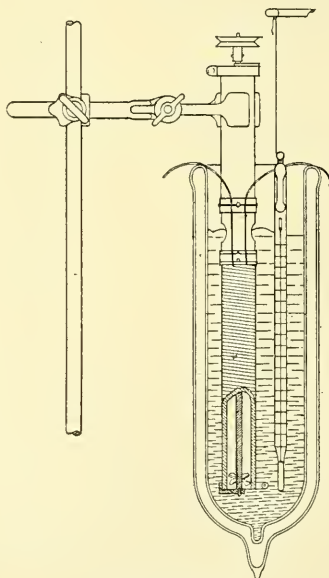
In the recalculation a small correction has also been added for the slight over-compensation of the heating coil (Table VIII)—13 microvolts at 500°, and 5 at 1150°. This corresponds to the measured temperature distribution according to which the junction of T₂ opposite the middle of the bulb reads smaller than the integral temperature—at 620° 1·2° smaller, at 820° 1·0°, at 1010° 0·5° and at 1150° 0·4°.

Observations below 500°.—After the observations in the electric oven between 500° and 1150° were completed, the thermo-element T₂ was compared with the gas thermometer

equipped with a bulb of Jena glass 59^{III} between 250° and 500° in a nitre bath. These observations were designed partly to verify the earlier observations below 500°, with which the element T_2 was connected only indirectly by comparison with T_1 ; and partly to enable us to compare the absolute values obtained from observations in the nitre bath and in the electrically heated air bath.

This nitre bath was considerably smaller than the one used for the earlier observations and was heated electrically, whereby we were enabled to secure constant temperatures more quickly and more exactly. A wrought iron cylinder 12^{cm} in diameter, such as is used for the transportation of mercury, was cut off at a height of 27^{cm} and fitted with a cover 2^{cm} thick, which carried a turbine driven by an electric motor for stirring. The heating coil was of bare constantan wire 1.5^{mm} in diameter wound in two parallel coils upon the cylinder, the insulation being provided by a layer of asbestos board, and held in place by smearing with clay. The whole was enclosed in a larger fire-clay jacket and the space between closed in with asbestos wool, giving a layer of air a centimeter or more thick about the coil. This bath can be used up to 700°.

2.



$\frac{1}{6}$ Natural size.

For temperatures below 300° where insulating liquids may be used for the baths, an arrangement which did us exceedingly good service certainly deserves mention in passing. A

cylindrical glass vessel, with double walls enclosing a vacuum, such as is commonly used for liquid air (Prof. Dewar's design), was filled with olive oil. Into the oil dipped a porcelain tube with two spiral furrows upon the outer surface, carrying coils of bare constantan wire. Inside the tube a turbine was inserted and held in place above and below with brass clamping rings. The stirrer and heating coil were thus in one piece and could be introduced into any insulating bath at pleasure. A vessel with a vacuum jacket is preferable however, on account of the very small quantity of electrical energy required to heat it, even when the walls are left unsilvered for reading a completely immersed thermometer. With 1.2 l. of oil only 100 watts were necessary to maintain a temperature of 250° . Although in using this apparatus we did not have an accident from explosion of the vessel, still, as this now and then occurs with such vessels at low temperatures without any apparent cause, it is well to make provision against possible personal injury in such an event. Table XV contains the comparison of thermo-element and gas thermometer in the nitre bath. The bulb was filled with nitrogen under a pressure of 471.84^{mm} at 0° .

TABLE XV.
Observations in the Nitre Bath.

	t	e_2 (MV)	Obs.— Calcul.		t	e_2 (MV)	Obs.— Calcul.
Feb. 6	340.7°	2639	-0.8°	Feb. 8	395.8°	3153	-0.8°
	340.5	2635	-0.5		395.8	3151	-0.6
	387.5	3076	-0.9		492.6	4082	-0.7
	386.9	3071	-1.0		492.8	4084	-0.7
	439.8	3573	-1.0		March 29	359.6	2803
	439.8	3572	-0.9	359.5		2802	$+0.4$
	501.3	4171	-1.3	400.5		3195	-0.6
	500.9	4168	-1.4	402.4		3217	-1.0
Feb. 8	267.4	1970	-0.6	498.5	4140	-1.0	
	291.6	2191	-0.8	498.3	4139	-1.1	
	291.7	2193	-0.9				

4. *Formulae for Thermo-elements.*

In a previous paper* we were able to show that the relation between the thermo-electric force and the temperature in metals of the platinum group together with gold and silver could be represented within wide limits with an accuracy of $\pm 1^{\circ}$ by a function of the second degree.

* Ludwig Holborn and Arthur L. Day, this Journal (IV), viii, 46, 1899.

The measurements made at that time were upon wires of pure rhodium, iridium, palladium, gold and silver together with alloys of platinum with palladium and ruthenium, whose thermo-electric force referred to pure platinum was measured up to 1300°. The elements with palladium or palladium alloys differed from the others in that the range within which the parabolic formula applies is smaller.

The formula for the normal element T_2 derived from the comparisons with the gas thermometer follows, together with certain of the others; those metals being chosen which are best suited to the measurement of high temperatures and which offer a basis for extrapolation.

These have been compared with the normal element from 50° to 50° approximately and the thermo-electric forces corresponding to the exact temperature intervals as contained in Table XVI, obtained from the observed data by graphical interpolation.

Since the publication of the earlier paper some observations above 1300° have been added with the aid of an oven equipped with a coil of platin-iridium wire.

Normal Element Pt-90 Pt, 10 Rh.—The thermo-electric force e_2 in microvolts of the platinum—platin-rhodium standard element T_2 is given by the following equation from 250° upwards

$$e_2 = -310 + 8.048t + 0.00172t^2$$

when the hot junction is maintained at the temperature t° and the cold junction at 0° . The formula is based upon the observations with the platin-iridium bulb above 500° and with the glass bulb and nitre bath below, as contained in Tables XIII, XIV and XV. In the columns "Obs.—Calcul." of these tables will be found the differences between the observed values and those obtained from the formula.

It will be recalled that the element T_1 was compared (first paper, Tables I to VII) above 500° with the gas thermometer equipped with porcelain bulbs. These observations differ among themselves somewhat, due in part to the behavior of the porcelain bulbs and in part to the less perfect conditions of measurement then obtaining. In the normal curve these observations are therefore not included, the curve being based entirely upon the temperatures measured with T_2 and the platin-iridium and glass bulbs. It is nevertheless not without interest to compare the temperatures obtained with T_1 and the porcelain bulbs, with the later ones.

This may be done by referring the normal values for T_1 (Table II former paper) to those of T_2 , by adding the differences $T_2 - T_1$ (Table XI, former paper) which were obtained by

direct comparison of the two elements, or in other words if T_2 be referred to the observations made with the gas thermometer as used with porcelain bulbs.

In Table XVI in the first column the values of the electromotive force for T_2 calculated from the formula are contained. The column Δ contains the differences between these values and those obtained from $T_1 + (T_2 - T_1)$ as suggested above. If now these values of Δ be compared with the earlier ones (Table XI, former paper) it will be seen that the differences have become very small now that the proper expansion coefficients of both porcelain and platin-iridium have been applied.

From this the conclusion follows that the porcelain bulbs lead to substantially the same results as the platin-iridium as long as the former are not used above 1100° . Under these conditions of course only bulbs glazed inside and out need be used, as the glazing does not melt up to that point.

Regarding the precautions necessary to be observed in the use of porcelain bulbs and the degree of accuracy obtained in temperature measurements with them, reference is made to the former communication.

Element Pt-90 Pt, 10 Ru.—The element platinum—platin-ruthenium follows a curve very similar to that of the normal element. Its equation is

$$e_3 = -359 + 9.260t + 0.00150t^2$$

Element Pt-Ir.—For the element platin-iridium a formula has also been calculated which like the others holds from about 300° on. Like the following curves it is based on the thermoelectric force of two pure metals and possesses a much greater curvature than those of the alloys. It is—

$$e_4 = -248 + 7.282t + 0.00554t^2$$

Element Pt-Rh.—For the combination platinum-rhodium two specimens of pure rhodium were at our disposal. The first was from Heräus, the second was especially prepared in the *Reichsanstalt*. The two formulae are

$$e_5 = -228 + 7.230t + 0.00660t^2$$

$$e_6 = -235 + 7.410t + 0.00660t^2$$

The second equation is derived under the assumption that the differences between the two curves e_5 and e_6 can be represented by a linear relation. The observed results between 300° and 1150° —the range of the gas thermometer—are made the basis for the derivation of this as well as of the other formulae.

Table XVI contains the observations upon these elements as well as the differences between them and the values calculated from the above formulae in microvolts and degrees.

TABLE XVI.

<i>t</i>	Pt—90Pt, 10Rh.			Pt—90Pt, 10Ru.			Pt—Ir			Pt—Rh ₁			Pt—Rh ₂			<i>t</i>
	<i>e</i> ₂ Calcul. MV	MV	Degrees	<i>e</i> ₃ Obs. MV	Obs.—Calcul. Degrees	MV	<i>e</i> ₁ Obs. MV	Obs.—Calcul. Degrees	MV	<i>e</i> ₅ Obs. MV	Obs.—Calcul. Degrees	MV	<i>e</i> ₆ Obs. MV	Obs.—Calcul. Degrees	MV	
250°	1810	-10	-1.1°	2062	+12	+1.2°	1966	+47	+4.7°	2004	+12	+1.1°	2042	+12	+1.1°	250°
300	2260	+3	+0.3	2550	-4	-0.4	2474	+39	+3.7	2536	+1	+0.1	2580	-2	-0.2	300
350	2718	+9	+1.0	3056	-10	-1.0	3002	+23	+2.1	3112	+1	+0.1	3160	-7	-0.6	350
400	3185	+9	+1.0	3580	-5	-0.5	3550	-1	-0.1	3714	-6	-0.5	3780	-5	-0.4	400
450	3661	+11	+1.1	4112	+1	+0.1	4156	+5	+0.4	4362	0	0.0	4432	-4	-0.3	450
500	4145	+5	+0.5	4646	0	0.0	4778	0	0.0	5032	-5	-0.4	5118	-2	-0.2	500
550	4638	+4	+0.4	5184	-4	-0.4	5424	-9	-0.7	5744	-1	-0.1	5834	-3	-0.2	550
600	5139	+3	+0.3	5738	+1	+0.1	6102	-14	-1.0	6484	-2	-0.2	6582	-5	-0.3	600
650	5649	-1	-0.1	6290	-4	-0.4	6810	-16	-1.1	7260	0	0.0	7372	+2	+0.1	650
700	6168	-6	-0.6	6860	+2	+0.2	7564	0	0.0	8076	+9	+0.6	8192	+6	+0.4	700
750	6695	-9	-0.8	7430	0	0.0	8320	-10	-0.7	8906	-1	-0.1	9038	+3	+0.2	750
800	7231	-10	-0.9	8012	+3	+0.3	9116	-7	-0.4	9790	+10	+0.6	9928	+11	+0.6	800
850	7775	-11	-1.0	8596	0	0.0	9948	+4	+0.2	10688	+2	+0.1	10840	+8	+0.4	850
900	8328	-11	-1.0	9196	+6	+0.5	10798	+5	+0.3	11644	+19	+1.0	11802	+22	+1.1	900
950	8890	-9	-0.8	9804	+8	+0.7	11670	0	0.0	12614	+17	+0.9	12782	+21	+1.1	950
1000	9460	-8	-0.7	10410	+9	+0.7	12588	+14	+0.8	13600	-2	-0.1	13774	-1	-0.1	1000
1050	10039	0	0.0	11018	0	0.0	13518	+12	+0.6	14616	-24	-1.1	14800	-22	-1.0	1050
1100	10626	+13	+1.1	11646	+4	+0.4	14482	+16	+0.8	15692	-19	-0.9	15884	-18	-0.8	1100
1150	11222			12274	0	0.0	15462	+9	+0.5	16800	-15	-0.7	16976	-39	-1.7	1150
1200	11827			12918	+5	+0.4	16482	+21	+1.0	17946	-6	-0.3	18130	-31	-1.3	1200
1250	12440			13556	-4	-0.3	17498	-5	-0.2	19106	-17	-0.7	19280	-60	-2.5	1250
1300	13062			14208	-6	-0.5	18580	+7	+0.3	20314	-11	-0.5	20492	-60	-2.5	1300
1350	13692			14864	-12	-0.9	19662	-9	-0.4	21556	-5	-0.2	21728	-69	-2.7	1350
1400	14331			15528	-17	-1.3	20766	-20	-0.9	22800	-30	-1.2	22988	-87	-3.4	1400
1450	14979			16220	-2	-0.2	21904	-35	-1.5	24136	+4	+0.2	24276	-110	-4.1	1450
1500	15635			16914	+8	+0.6	23056	-74	-3.0	25486	+19	+0.7	25684	-46	-1.7	1500

Attention should perhaps be called to the fact that in the comparisons between these elements and the normal element, different sections of the same curve are separated in particular cases by considerable intervals of time so that the distribution of temperature from junction to junction was not always the same even through the observations of a single series. Small consistent differences between "Obs." and "Calcul." may be fairly attributed to this cause, though for the temperature range 300° to 1150° when the gas thermometer was used, the differences rarely exceed 1° ; above 1150° where the figures are based upon an extrapolation of the formula for T , they are sometimes larger, though nowhere exceeding 5° .

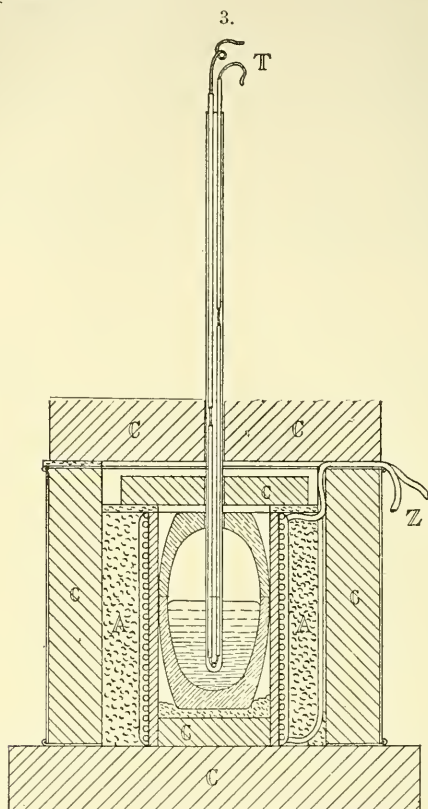
5. *Melting Points of Metals.*

To render the calibration of thermo elements independent of the gas thermometer we have determined the melting points of various metals lying between 300° and 1100° . Two methods were used: According to the first, a short wire (about 1cm long) of the metal to be melted was introduced into the hot junction of the thermo-element itself, and the thermo-electric force at the moment of the melting and consequent interruption of the circuit, observed. This method is simple but leads easily to doubtful results if the melting point is at all affected by the surrounding atmosphere, as is the case with most of the metals which oxidize in air. The element also becomes considerably shortened by the continual cutting off of the drop of melted metal after the observations and renewing the junction.

By the second method a larger quantity of metal is melted in a crucible and the thermo-element, protected by light porcelain tubes, inserted. When the heating is properly regulated the beginning of the melting or freezing point is readily recognizable as the temperature remains stationary for a considerable length of time. In the following, the words "wire method" and "crucible method" may serve to distinguish the two processes.

For the latter method the apparatus was arranged as indicated in fig. 3. A porcelain or graphite crucible is contained within a short fire-clay tube carrying a coil of bare nickel wire whose separate turns were insulated by smearing with clay. The whole is protected against loss of heat by a layer of loose asbestos A and a thick protecting tube also of fire-clay. The wires of the thermo element are insulated from each other by means of thin porcelain tubes and from the molten metal by a somewhat thicker tube of the same material, 5mm in inside diameter and 1.5mm in thickness, which dipped at least 4cm into

the melting metal (1^{cm} from the bottom of the crucible) and was clamped in position there.



$\frac{1}{4}$ Natural size.

The figure shows the oven with a graphite crucible in position and a second slightly smaller one with a hole bored in the bottom inverted over it as a cover.

The porcelain crucibles were similar in size and form, 1^{mm} thick and without cover.

Inasmuch as the electric heating furnished a much more uniform distribution of temperature than the gas oven and as no systematic differences between the melting and freezing temperatures were observed even with the most varied conditions of current and quantity of metal, we did not attempt to stir the metals in general. Only in one section of the observations with silver an effort was made to expose the melted metal more thoroughly to the air by stirring.

In the preliminary experiments the coil was wound directly upon a crucible of fire clay, but that was soon given up—first

because the clay was too porous to hold the melted metal in all cases (molten Cu for instance penetrated the wall and short-circuited a section of the coil), and second, the change from one crucible to another which happened nearly every day was much more convenient with the coil wound upon a separate tube. This oven consumed about 600 watts at 1000°

Some caution was necessary in heating up to the highest melting point (copper), otherwise the oven coil (1^{mm} in diameter) burned out—its renewal is a simple matter however.

For the temperature measurement thermo-elements were used which were cut from the same wire as T_2 and which have shown themselves to be perfectly identical with it in their results. Temperatures above 500° are derived from the normal curve (Table XVI) and below from the measurements in the nitre bath.

In each determination the thermo-electric force was observed from minute to minute and afterward plotted as a function of the time (mentioned hereafter as the "time curve"). A change of 10 microvolts could be compensated directly on the compensation apparatus by inserting 0.1 ohm. Smaller differences were interpolated from the galvanometer deflection, one scale division corresponding to from 0.5 to 1.4 microvolts according to the resistance in the branch circuit.

The accuracy of the crucible method is only limited by the accuracy with which it is possible to measure with the thermo-element, in regard to which more will be said further on. The wire method is less exact because the interruption of the circuit which indicates the melting point is not independent of any slight tension which may exist in the wires of the thermo-element, particularly the platin-rhodium wire—a factor which varies from one observation to another.

The highest temperatures observed for a particular metal by this method may therefore be expected to come nearest the truth. The inserted wire may even at times be seen to break just before the real melting takes place.

Gold.—The melting point of gold was determined by the wire method only. Two different specimens of gold had been placed at our disposal, one (I) from the *Gold und Silber Scheideanstalt* in Frankfurt-am-Main, the other (II) specially prepared and purified in the Reichsanstalt. No difference could be detected in their melting temperature.

The observations on the first two days (following table) were made in the gas thermometer oven, the others in the small oven used for the determinations by the crucible method (fig. 3) in which an empty crucible had been placed.

The arrangement of the thermo-element in the latter case was the same as above described for the crucible method with the omission of the porcelain protecting tube. The heating

was so regulated that the slowly rising temperature could be easily observed at any time. It was even possible to observe the stationary temperature at the melting point just before the interruption of the circuit. The observed melting temperatures in microvolts and in degrees follow :

	Microvolts.	Degrees.
Oct. 17 (I)	10209	1064·5°
	10209	1064·5
	10206	1064·2
	10208	1064·4
Oct. 18 (II)	10202	1063·9
	10201	1063·8
Nov. 15 (II)	10211	1064·6
	10211	1064·6
	10211	1064·6
Jan. 2 (II)	10190	1062·9
	10190	1062·9
	10200	1063·7
	10189	1062·8
Jan. 3 (II)	10203	1064·0
	10197	1063·5
	10213	1064·8

The mean is $1064\cdot0^{\circ} \pm 0\cdot6^{\circ}$.

Silver—Wire Method.—The melting point of silver was also first determined by the wire method and two specimens were used, both from the Frankfurt *Scheideanstalt*. One of these was in the form of wire $0\cdot5^{\text{mm}}$ and $0\cdot25^{\text{mm}}$ in diameter and the other a piece of sheet silver $0\cdot25^{\text{mm}}$ in thickness, from which thin strips were cut and inserted in the junction of the element. Here again no difference in the melting temperature of the two specimens could be established. The same ovens were used as for the gold determination. The results follow :

	Microvolts.	Degrees.
Sept. 25	8928	953·3°
	8936	954·0
Oct. 16	8937	954·1
	8922	952·8
	8922	952·8
Oct. 18	8914	952·1
	8927	953·3
	8930	953·5
	8935	953·9
Nov. 15	8939	954·3
	8958	956·0
	8925	953·1
	8941	954·5
	8924	953·0

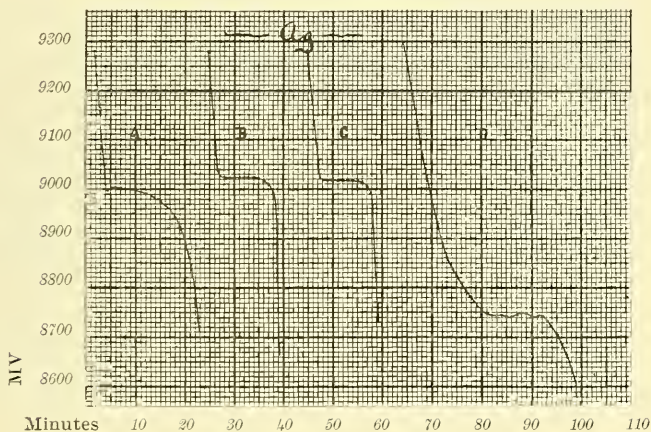
The mean is $953\cdot6^{\circ}$ with a mean error of $\pm 0\cdot9^{\circ}$.

Callendar as well as Heycock and Neville* have called attention to the fact that silver in an oxidizing atmosphere melts and solidifies at a lower temperature than in a reducing atmosphere, the difference being caused by the fact that silver absorbs oxygen in melting which is given off again below the melting temperature with considerable violence ("spitting").

Silver—Crucible Method.—In order to ascertain whether the melting point as determined by the wire method differs from the normal melting temperature we made determinations of it by the crucible method also.

(1) *In an Oxidizing Atmosphere.*—At first the silver was melted in an open porcelain crucible without stirring. Under these conditions the results differed from those with other metals in that the silver showed no definite melting temperature. In melting as well as in solidifying the "time curve" takes the form A (fig. 4), i. e. throughout its length it contains

4.



no horizontal section but falls or rises slowly over an interval of 6° to 8° , between which limits one point may be assumed as the melting point nearly as well as another. After substituting mica sheets for the clay oven-covers so as to be able to see the metal during the process, it was observed that it remained liquid through the greater part of this interval. The melting and solidifying occurred at about 954.5° (8940 microvolts). It may be added that the substitution of three mica covers separated by air layers, for the two clay covers, produced no essential change in the conditions within the oven.

If the thermo-element with its protecting tube were moved about or the liquid metal otherwise stirred, the form of the

* C. T. Heycock and F. H. Neville, Journ. Chem. Soc., 1895, 160 and 1024.

time curve remained in general the same except perhaps that the limits of the slow rise or fall of the temperature were moved closer together, and if the stirring was energetic, could be lowered some degrees.

Table XVII contains a number of results partly with stirring and partly without. *G* indicates the weight of the melted metal in grams, t_1 and t_2 the upper and lower limits of the slow change of temperature about the melting point, and t their arithmetical mean. The letters *F* and *M* distinguish freezing and melting points. The strength of the oven current i may serve to give an idea of the comparative expenditure of energy in the different determinations though only a rough one, as the resistance of the coil on different days was not always the same. The mean value of t for the observations where the metal was stirred is 954.9° and differs only 1.3° from the result

TABLE XVII.
Melting Point of Silver in Air.

1899. Novem'r.	<i>G</i> (gr.)	<i>i</i> (Amp.)		t_1		t_2		t			
				MV	Degrees.	MV	Degrees.				
Without stirring. 16	350	6.0	F	8940	954.4°	8983	958.2°	956.3°			
			M	8960	956.1	9016	961.0	958.6			
			F	8973	957.3	9000	959.6	958.5			
			M	8970	957.0	9014	960.9	959.0			
			F	8926	953.2	9006	960.2	956.7			
			M	8970	957.0	9014	960.9	959.0			
			17	500	5.0	F	8968	956.8	9001	959.7	958.3
						M	8963	956.4	9018	961.2	958.8
F	8942	954.6				8995	959.2	956.9			
M	8950	955.3				9011	960.6	958.0			
With stirring.	500	5.0	F	8931	953.6	8947	955.0	954.3			
			20	380	F	8950	955.3	8967	956.7	956.0	
					M	8931	953.6	9004	960.0	956.8	
					F	8890	950.0	8943	954.6	952.3	
					M	8931	953.6	9004	960.0	956.8	

obtained by the wire method. We assume therefore 955° as the mean melting point of silver in air, but at the same time we wish to emphasize the fact that the point is not well defined and, especially with the wire method, may suffer considerable variations with the nature of the surrounding atmosphere.

The "spitting" was also frequently observed through the mica covers. It took place between 933° and 940° , in the mean at about 936° , and was much more violent when the melted metal had been previously stirred.

The attempt to saturate the melted silver by allowing oxygen to bubble through it and thereby to reach a definite point of solidification for this condition led to no satisfactory result. The thermo-element showed considerable temperature oscillations throughout the solidification, but the mean, as was to be expected, lay much below the other determinations. Four observations gave the temperatures 939° , 938° , 942° and 938° — approximately the same temperature at which the absorbed oxygen is given off as noted above. Curve D (fig. 4) shows one of the time curves taken under these conditions, though it is by no means characteristic of all. In one case for instance with the temperature falling gradually, oxygen began to be given off at about 940° , whereupon the temperature began to rise, and reached 963° in the next five minutes, a violent spitting going on throughout the interval. At the latter temperature the metal became suddenly rigid and closed the porcelain tube through which the oxygen was entering.

(2) *In a Reducing Atmosphere.*—Higher results and better defined are obtained by preventing the access of oxygen. We sought to accomplish this first by melting the silver under a layer of common salt and afterward in a plumbago crucible over which a second crucible was inverted as a cover.

The time curves in both cases, for both melting and solidifying points, showed well defined horizontal segments, as may be seen from fig. 4, curves B (under NaCl) and C (in graphite).

Table XVIII contains a series of such observations from which 961.5° is obtained as the melting or solidifying tempera-

TABLE XVIII.
Melting Point of Silver in Absence of Oxygen.

Date.	G. (gr.)	i. (Amp.)		t.	
				MV.	Degrees.
In Porcelain Crucible under NaCl.					
Nov. 23	400	5.3	F	9023	961.7°
		10.7	M	9026	961.9
		6.7	F	9023	961.7
		11.3	M	9026	961.9
In Graphite Crucible.					
Dec. 8	352	6.7	F	9018	961.2
		9.5	M	9019	961.3
		6.0	F	9018	961.2
		9.3	M	9018	961.2

ture of pure silver. No spitting was observed under these conditions.

For the wire method neither of these courses is open for making determinations in the absence of oxygen. The bare junction cannot be placed in melted salt on account of its conductivity nor exposed to the reducing atmosphere of a plum-bago crucible without danger of affecting the thermo-element.

Copper.—The melting point of copper was not determined by the wire method on account of the oxidation of the metal. In the crucible, the material used was pure copper in 100 gr. blocks from the *Haddernheim Kupferwerk* and determinations were made in both oxidizing and reducing atmospheres.

TABLE XIX.
Melting Point of Copper.

Date.	G. (gr.)	<i>i.</i> (Amp.)		<i>t.</i>	
				MV.	Degrees.
In Oxidizing Atmosphere.					
Nov. 7	900	—	F	10216	1065.1°
29	330	7.5	F	10213	1064.8
		11.0	M	10215	1065.0
		7.5	F	10213	1064.8
		10.8	M	10214	1064.9
Dec. 5	350	10.3	M	10219	1065.3
		8.0	F	10210	1064.6
		10.5	M	10216	1065.1
6	350	7.5	F	10211	1064.7
		10.3	M	10220	1065.4
		7.0	F	10211	1064.7
		9.9	M	10219	1065.3
Mar. 19	370	6.8	F	10213	1064.8
		10.3	M	10213	1064.8
		6.8	F	10211	1064.7
		10.3	M	10211	1064.7
21	370	6.5	F	10212	1064.7
		8.5	M	10214	1064.9
		8.5	M	10216	1065.1
23	370	8.2	M	10216	1065.1
		8.2	M	10219	1065.3
In Reducing Atmosphere.					
Dec. 7	355	7.5	F	10442	1084.3
		10.2	M	10440	1084.2
		7.2	F	10440	1084.2
		10.4	M	10437	1083.9

(1) *In an Oxidizing Atmosphere.*—Copper differs distinctly from silver in that when melted in an open porcelain crucible it possesses perfectly definite and coincident melting and solid-

ifying points. This seems to indicate that the copper under these conditions becomes saturated with a definite quantity of oxygen or rather with a copper oxide.

To assure ourselves regarding the constancy of this point we have made a great many observations of it without being able to notice any essential variation.

It was also quite indifferent whether the copper was fresh or had been several times used for the same purpose before.

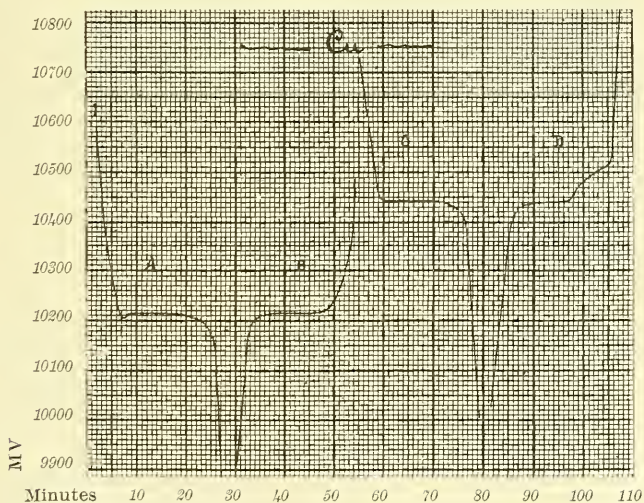
After many determinations with a particular mass of metal it seemed to melt with increasing difficulty, a phenomenon which points to a decrease in its conductivity for heat, but should receive rather more investigation. Table XIX contains the results of the observations, giving 1064.9° as the melting point of copper in air.

(2) *In a Reducing Atmosphere.*—If the same metal which has been repeatedly melted in air be heated in a plumbago crucible it gradually reduces and the melting point rises, ultimately reaching 1084° after the metal has remained for a long time in a molten condition and become thoroughly reduced.

Melting points obtained between 1065° and 1084° where the reduction is still incomplete give poorly defined curves.

At the close of Table XIX some observations with the pure metal are given. Fig. 5 shows time curves for observations in

5.



porcelain crucibles (curves A and B) as well as in graphite (curves C and D).

If the melted copper contained oxygen, the unglazed porcelain tube which served to protect the thermo-element was

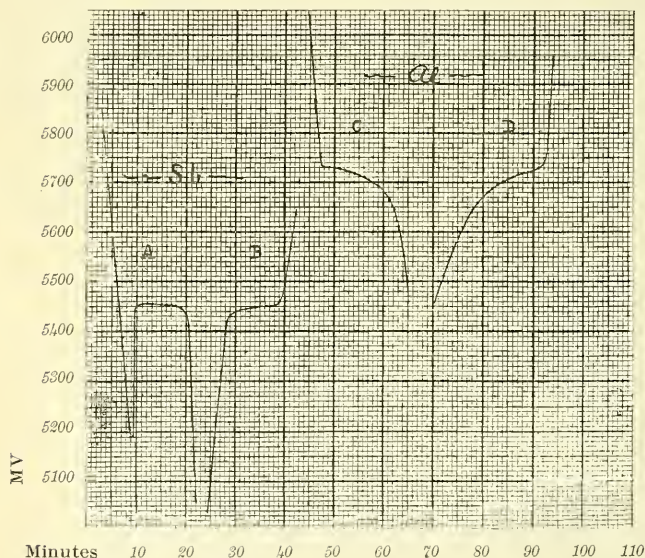
always colored black while pure copper had no noticeable effect upon it.

Antimony.—The antimony with which our determinations were made, as well as the following metals, were obtained from C. A. F. Kahlbaum, chemist, in Berlin; with the exception of aluminium they may be regarded as pure.

For melting the antimony only graphite crucibles were employed, and even then a slight superficial oxidation was observed after cooling. Reducing gases were not introduced on account of the thermo-elements.

In solidifying, the temperature first sank far below (as much as 20°) the melting point, to which it then rose suddenly (see fig. 6, curve A). The time curves are otherwise normal. The mean value for the melting temperature was 630.6° .

6.



Aluminium.—Aluminium was melted in both porcelain and graphite crucibles. As has been noticed already by former observers, this metal shows no sharply defined melting point. Its time curves are similar in form to those obtained from silver in air. We give the melting point as 657.3° —the mean of the results obtained in the porcelain crucible. The melting temperatures observed in graphite were slightly lower to be sure, but the difference is more probably due to the smaller quantity of metal used.

The melting point of aluminium is not a sufficiently well-defined point for use in calibrating thermometrical apparatus;

should a temperature in this region be desired, the melting point of antimony is preferable.

Zinc.—The melting temperature of zinc can also be determined in an open porcelain crucible provided the same zinc be not too often used—with too much oxide present, the curves were less well-defined. In graphite the form of the time curve and the temperature are the same as in porcelain. The mean melting temperature may be given as 419.0° .

Lead.—The only noticeable feature in the determination of the melting point of lead is its small melting or "latent" heat, which makes some attention to the regulation of the oven temperature necessary. The mean melting point is 326.9° .

Cadmium.—The melting heat of cadmium is much greater and sharp curves are obtained with little trouble. Its mean value is 321.7° .

In Table XX are contained the determinations of the melting and freezing temperature of antimony, aluminium, zinc, lead and cadmium.

Tables XXI and XXII contain examples of the time curves as they were observed, the numbers representing the thermo-electric force in microvolts as observed from minute to minute.

Alloys.—In conclusion certain alloys remain to be mentioned whose melting points we investigated by the wire method only.

The metals were in thin strips cut from sheets of perhaps 0.2 to 0.3^{mm} thickness, prepared especially for the purpose by the Frankfurt *Scheideanstalt*.

As the investigation could only be undertaken in an oxidizing atmosphere, alloys of silver and copper gave no definite results whatever. The silver melts and separates out of the partly oxidized copper, leaving it to interrupt the circuit by breaking at a considerably higher temperature.

The two alloys 95Ag5Au and 90Ag10Au were especially prepared to ascertain whether the addition of a small quantity of gold would prevent the action of oxygen on the melting point of silver. The observed results seem to negative this assumption.

The temperature scale for the highest melting points is based upon extrapolation from the curve of the normal element T_2 (Table XVI).

6. Measurement with Thermo-elements.

Some ten years ago when the first platinum—platin-rhodium thermo-elements were obtained from the firm of Heräus for the measurement of high temperatures, both the platinum and the rhodium were so imperfectly purified chemically that the duplication of an element with approximately the same thermo-

TABLE XX.

Date.	G (gr.)	<i>i</i> (Amp.)	<i>t</i>			
			MV	Degrees		
Antimony.						
Dec. 14	260	0.0	F	5453	630.8°	In graphite.
		8.5	M	5449	630.4	
		5.0	F	5452	630.7	
		7.5	M	5448	630.3	
Aluminium.						
Dec. 4	180	0.0	F	5729	657.7°	In porcelain.
		8.5	M	5723	657.1	
		4.5	F	5727	657.5	
		7.5	M	5723	657.1	
Dec. 9	110	4.3	F	5725	657.3	In graphite.
		0.0	F	5717	656.6	
		8.0	M	5707	655.6	
		5.0	F	5718	656.6	
		7.5	M	5705	655.4	
Zinc.						
Dec. 2	389	4.2	F	3377	419.2°	In porcelain.
		7.9	M	3375	419.0	
		0.0	F	3376	419.1	
		7.0	M	3374	418.9	
Mar. 30	293	0.0	F	3376	419.1	In graphite.
		7.3	M	3375	419.0	
		0.0	F	3376	419.1	
		7.3	M	3373	418.8	
Lead.						
Dec. 15	500	0.0	F	2507	326.9°	In porcelain.
		7.0	M	2508	327.0	
		0.0	F	2505	326.7	
		6.0	M	2505	326.7	
Mar. 26	493	0.0	F	2505	326.7	In graphite.
		5.8	M	2508	327.0	
		4.2	M	2510	327.2	
		0.0	F	2510	327.2	
		4.2	M	2510	327.2	
Cadmium.						
Dec. 2	430	0.0	F	2461	321.9°	In porcelain.
		6.0	M	2460	321.8	
		0.0	F	2458	321.6	
		6.7	M	2459	321.7	
		0.0	F	2457	321.4	
Mar. 27	430	6.5	M	2459	321.7	In graphite.
		0.0	F	2459	321.7	
		4.3	M	2458	321.6	
		0.0	F	2461	321.9	
		4.3	M	2460	321.8	

electric force was impossible. The platinum was the first to be successfully purified, then, sometime after, the rhodium. Thermo-elements from different meltings now show variations of only 1 per cent, when formerly 10 per cent, and more was observed. The wires are also much more homogeneous now than formerly.

TABLE XXI.
"Time Curves" (MV).

Min.	Ag under NaCl	Ag under NaCl	Ag in Graphite	Ag in Graphite	Cu in Porcelain	Cu in Porcelain	Cu in Graphite	Cu in Graphite
1	9423	8513	9439	8766	10590	10003	10680	10081
2	9330	8599	9354	8861	10475	10106	10609	10179
3	9247	8684	9255	8956	10370	10183	10530	10268
4	9130	8775	9156	9016	10310	10201	10455	10353
5	9039	8857	9056	9018	10240	10207	10444	10409
6	9024	8937	9018	9018	10210	10210	10442	10426
7	9024	9017	9017	9019	10210	10210	10442	10431
8	9023	9023	9018	9019	10212	10211	10442	10433
9	9022	9024	9018	9019	10214	10213	10442	10436
10	9022	9024	9018	9019	10213	10213	10442	10437
11	9022	9025	9017	9022	10213	10213	10443	10438
12	9022	9026	9017	9033	10214	10213	10443	10439
13	9021	9026	9016	9043	10213	10214	10443	10440
14	9021	9026	9013	9055	10213	10214	10442	10440
15	9020	9026	9006	9064	10212	10215	10442	10441
16	9018	9026	8976	9074	10212	10215	10441	10444
17	9014	9027	8763	9121	10211	10216	10439	10456
18	9007	9028		9290	10210	10219	10437	10472
19	8990	9027			10210	10227	10433	10484
20	8800	9029			10208	10245	10427	10491
21	8600	9029			10203	10268	10410	10497
22		9029			10200	10300	10342	10503
23		9031			10193	10340	10195	10509
24		9034			10177	10460	10070	10516
25		9037			10097		9968	10564
26		9050			9900			10720
27		9069			9740			
28		9232						
29		9350						
30								

Since the platinum wire at temperatures above 1000° is considerably less stable than the platin-rhodium, much would be contributed in the direction of the durability of the elements if the platinum could be replaced by some related metal or alloy of greater stability.

Up to the present, however, no satisfactory substitute has been found. The alloys of platinum with iridium and ruthenium, in part also with palladium, differ little in their thermo-electric properties from platinum-rhodium.

Pure palladium does possess a higher thermo-electric force when combined with platinum-rhodium than platinum, but can-

TABLE XXII.
"Time Curves" (MV).

Min.	Sb in Graphite	Sb in Graphite	Zn in Porcelain	Zn in Porcelain	Pb in Porcelain	Pb in Porcelain	Cd in Graphite	Cd in Graphite
1	6216	5223	3551	3265	2718	2309	2852	2392
2	6080	5294	3500	3316	2671	2353	2809	2425
3	5960	5360	3449	3344	2621	2401	2767	2442
4	5840	5408	3400	3351	2574	2450	2726	2449
5	5720	5430	3376	3355	2529	2496	2685	2454
6	5600	5437	3376	3359	2506	2503	2646	2456
7	5495	5441	3376	3362	2506	2504	2607	2457
8	5390	5443	3376	3365	2505	2505	2570	2458
9	5280	5445	3376	3368	2505	2505	2533	2459
10	5190	5445	3376	3369	2505	2505	2496	2459
11	5450	5446	3376	3369	2505	2506	2462	2460
12	5454	5447	3375	3370	2504	2507	2460	2460
13	5454	5448	3375	3372	2503	2510	2459	2460
14	5453	5448	3375	3372	2502	2514	2459	2461
15	5454	5448	3375	3373	2501	2521	2459	2461
16	5452	5449	3375	3373	2497	2529	2459	2463
17	5452	5449	3374	3374	2488	2572	2459	2469
18	5451	5450	3373	3374	2473	2676	2459	2485
19	5451	5456	3374	3374	2444		2459	2499
20	5448	5468	3373	3374	2365		2458	2525
21	5438	5488	3373	3375			2458	
22	5310	5512	3372	3375			2457	
23	5060	5555	3370	3376			2456	
24		5655	3366	3389			2454	
25			3357	3404			2452	
26			3342	3433			2449	
27			3300				2422	
28							2367	
29								
30								

not be considered in this connection, for it is even less stable at high temperatures than platinum.

Thermo-elements made from different alloys of platinum-rhodium such as were formerly used by Holborn and Wien* for temperatures above 1000° are also unsuited for general use on account of their low sensitiveness.

* L. Holborn and W. Wien, loc. cit.

TABLE XXIII.
Melting Points of Alloys.

Alloy.	<i>t</i>		Mean.	Alloy.	<i>t</i>		Mean.
	MV	Degrees.			MV	Degrees.	
95 Ag, 5 Au	8989	958·7	958·5	95 Au, 5 Pt	10805	1115·0	1116·0
	8990	958·8			10799	1114·5	
	8980	957·9			10826	1116·8	
		10827	1116·8				
90 Ag, 10 Au	9067	965·5	965·7	90 Au, 10 Pt	11551	1177·2	1176·3
	9070	965·8			11522	1174·9	
	9080	966·7			11533	1175·7	
	9060	964·9			11495	1173·6	
	9070	965·8			11578	1179·7	
40 Ag, 60 Au	9753	1025·3	1025·1	85 Au, 15 Pt	11482	1171·5	1237·7
	9756	1025·5			11599	1181·2	
	9730	1023·3			12285	1237·4	
	9760	1025·9			12282	1237·1	
	9757	1025·6			12273	1236·4	
95 Au, 5 Pt	10821	1116·4			12317	1240·0	

Thermo-elements were first intended only for the measurement of temperatures from red heat upwards; afterwards it became evident that lower temperatures could be measured by this means as well. There is no advantage gained by going below 250°, however, as the errors due to the decrease in the sensitiveness at the lower temperatures begin to be more noticeable.

The accuracy of thermo-electric measurements in the earlier investigation of Holborn and Wien (*loc. cit.*) was stated to be 5°. The present results carry it further. It could fairly be said that, for temperatures up to 1150°, the present error would be about 1°; that is to say, for any distribution of temperature between the junctions, the temperature of the hot junction is defined by the thermo-electric force to within $\pm 1^\circ$.

This does not in the least mean however, that with constant conditions considerably more accurate relative measurements cannot be made. The determinations of the melting points of the metals contain many examples of this.

The chief reason for this progress lies in the facilities for electric heating whereby the whole investigation of high temperatures gains in accuracy and simplicity.

For the thermo-elements the change is especially important since the electric oven allows a definite temperature distribution to be exactly reproduced at any time, whence it becomes

possible even with inhomogeneous wires to carry out accurate temperature measurements by always working under the same conditions. It should be borne in mind however that this accuracy is only relative and that the same thermo-elements with a different temperature distribution between the junctions may show considerable variations.

To obtain the limits of absolute accuracy the fall in temperature along the wire must be varied as much as possible.

In the investigation here presented this was accomplished in two ways, either by measuring the temperature at a given point with two elements having a common hot junction and so arranged that the fall in temperature along the wire from the hot to the cold junctions could be varied; or secondly, by using as many different points of an element as possible successively, as junctions, and comparing them in a very short oven under exactly similar conditions of temperature distribution.

For example, by the first method the temperature near the end of a long electrically heated tube would be measured with two elements joined at their hot junctions, and the one passing out at the end near by, while the other extended through the middle of the tube which is hotter than near the ends, and out at the far end. It is clear that the distribution of temperature from hot to cold junction along the wires of the two elements would be very different and could be varied at will by simply moving the common junction to and fro.

Any lack of homogeneity in the wires would then become apparent at once in differences between the readings of the two elements.

We rather preferred the second method, the small melting-point oven (fig. 3) being well adapted to the purpose.

The wires of the elements are best simply bound together with a short platinum wire when it is desired to use several points provisionally as hot junctions. It has however no influence upon the thermo-electric force if the wire be cut in different places and afterward joined in the oxy-hydrogen flame.

Thick lumps along the wires are in the way to be sure, but with a little practice a junction can be made which can hardly be detected.

Thermo-elements as they are now prepared in a diameter of 0.6^{mm} are very homogeneous and over a length of several meters show no differences which could affect the temperature measurement 1° . Thin wires reduced from the above by drawing are often less uniform. Two wires, for example, which had been drawn down to 0.25^{mm} showed about double this variation, nor could this difference between the thin and thick wires be removed by glowing electrically.

Repeated heating to high temperatures may also under certain circumstances render the wires less homogeneous.

Attention has been called to this before, but now that the accuracy of thermo-electric measurements has been so materially increased it becomes necessary to go into the discussion of this matter in considerable detail.

The changes are such as Holborn and Wien observed in platinum and palladium due to the action of combustion gases or hydrogen in the presence of silicium, and which affect the resistance even more than the electromotive force. Even with electric heating where combustion gases should no longer be feared, the walls of an oven which has not been previously heated to a higher temperature than that to be measured, and kept there until burned through and through, develop gases which pass through platinum.

We have called attention to this in the discussion of a platinum bulb for the gas thermometer in our earlier communication. The thermo-elements are also affected by exposure to these gases but may usually be protected by the use of porcelain, which does not itself develop such gases and is very impervious to them even in its unglazed form. For temperatures below 1100° still further protection is afforded by glazing of course. Tubes of more refractory material intended for temperatures where porcelain can no longer be used and necessarily imperfectly burned we have found to be porous and when first heated to give off gases themselves at high temperatures which in the presence of silicium affect platinum strongly.

If the wires of a thermo-element which have been exposed in this way be afterward used for temperature measurement, there is often little or no change in the electromotive force from the normal values *so long as the temperature distribution along the wire remains unchanged*, but considerable variations always appear as soon as it is varied.

We made observations for example with such an element C which had been exposed in an imperfectly burned oven, the hot junction being combined with that of a second element A which had not been so exposed. At 550° the following differences (C-A) in microvolts were observed in an electric oven 40^{cm} long in which the common junction of the two elements could be moved from the center toward either end as described elsewhere:

Hot Junction.		I.	II.
At the center (elements symmetrical)		7	7
6^{cm} to the right	} 2 wires heated 14^{cm} } } " " " 26^{cm} }	- 73	+ 3
6^{cm} " left		+126	+36

In column I the elements were crossed while in II the two wires of each element passed out of the oven at the same end.

Originally the two elements had agreed for all positions to within 5 microvolts.

Further examples could be cited, but we will confine ourselves to two where efforts were made to restore the damaged elements.

The element C' , cut from the same wire as T_2 , was in use for a long time for high temperatures in a long oven and there exposed to these gases, after which the following differences were obtained by comparison with the normal elements G and H under *unchanged* conditions of temperature distribution.

t	$H-C'$	$G-C'$
200°	+ 5 MV	+5 MV
300	+ 5	+6
400	+ 4	+5
500	-11	+6
600	-13	+7
700	-14	+7
800	-13	+5
900	-12	+7
1000	-13	+7
1100	-10	+8
1200	- 9	+8

When this same element C' was introduced into the short oven containing melted copper however, where the conditions of temperature distribution were very different, it gave 10,092 MV as the melting temperature while the two normal elements showed 10,212 microvolts.

C' was then taken out, glowed electrically at full white heat for several hours and again introduced into the copper crucible (with proper porcelain protection of course), when it gave 10,170 under the same conditions as before.

The same measurement was then a third time repeated after 5^{cm} had been cut from both wires at the hot junction—the electromotive force now proved to be 10,220 microvolts. In this condition further measurements were made of the melting points of lead, cadmium and zinc with results 2513, 2460 and 3371 microvolts respectively, all in good agreement with the normal values.

As has been said the glowing was done electrically, as much as 17 amperes being sent through the wires. Connected in series both wires glow with equal intensity at high temperatures, the smaller resistance of the platinum when cold being offset by its larger temperature coefficient.

An element B of the same group was used several times in an oven tube of a new material which was being tried for the first time, up to a temperature of 1600°.

The platinum wire became very brittle in the vicinity of the junction, broke in several places and was melted together in the oxy-hydrogen flame. Afterward when used for measurement at about 1000° in the long oven (unchanged temperature distribution) its reading was some 90 microvolts too low, but for the melting point of copper (different distribution) it gave only 9840 MV, i. e. some 400 MV too low. Sixteen centimeters were at once cut off at the hot junction, first from the platinum and afterward from the platinum-rhodium wire, when the melting temperature of copper was given at 10,290 and 10,150 MV respectively. Further shortening did not serve to change these figures and the pieces cut out were again inserted. In this condition the element was glowed as in the previous case for some hours at white heat, when the melting point of copper was read at 10,216 MV—almost exactly the normal value.

The glowing also restores the bright surface to wires (platinum particularly) which have become dull after exposure to these gases.

It is probable that the standard element which represents the earlier temperature scale of Holborn and Wien and which has been used many times for comparisons up to 1600° has undergone similar changes.

Its indications at 500° are now 10° , and at 1000° , 18° too high, while the new determinations of the melting point of gold differ only 8° from the old. This melting point in particular was well determined at that time, the greatest variations from the mean in twenty-five determinations under varied conditions being $\pm 4^{\circ}$ while the corresponding variations for silver and copper amounted to $\pm 10^{\circ}$.

At the lower temperatures no fixed points were then determined so that no further comparison is possible.

We draw the conclusion from this that the temperature scale once established can be maintained with certainty only with the help of fixed temperatures such as the melting points given above. It is further advisable to divide a set of thermo-elements into two groups, one to be used only in porcelain tubes up to say 1200° and the other under other conditions or in higher temperatures.

For the latter, up to the present at least, the same accuracy is neither to be obtained nor expected until observations with the gas thermometer have been made and the technical resources much extended.

Especial attention ought again to be called to the fact that in the measurement of the fall in temperature in an extended space such as is necessary for the calculation of its mean temperature, the utmost care should be taken that the elements be

in perfect order. The best arrangement is always to measure with two elements having a common hot junction and entering the space in question from opposite sides, as has already been described in our former paper.

Furthermore, inasmuch as thermo-elements have come to be employed very generally for technical purposes during the past few years, we take this opportunity to say that it is impossible to give the limits to which the electromotive force can be affected by the breaking of the protecting tube or other accident through which the wires become exposed to combustion gases in the manner above described—the higher the temperature the greater the danger. In case of such an accident the element should at once be tested and restored as elsewhere indicated.

7. Conclusion.

By way of conclusion we will bring together the temperatures which we have obtained for the melting points of the metals based upon observations with the gas thermometer and some recent results of other observers employing other methods.

These include the measurements with platinum resistances as carried out by Callendar and by Heycock and Neville* which are also based upon the gas thermometer up to 445° , the boiling point of sulphur, and the higher temperatures extrapolated from a formula of the second degree.

The melting points of gold and silver have also been measured with thermo-elements by Berthelot,† using the wire method and determining the temperature optically from the decrease in density of an air column in an electrically heated, open, porcelain tube.

	Gas Thermometer.	Platinum Resistance.	Optical Method.
Cadmium	321.7°	320.7°	
Lead	326.9	327.7	
Zinc	419.0	419.0	
Antimony	630.6	629.5	
Aluminium	657.	654.5	
Silver, in air	955.	955.	962
Silver, pure	961.5	960.7	—
Gold	1064.0	1061.7	1064
Copper, in air	1064.9	—	
Copper, pure	1084.1	1080.5	

Charlottenburg, May, 1900.

* Most of the figures are taken from Heycock and Neville's paper—loc. cit.; Callendar's values—Phil. Mag. V, xlviii, 519, 1899—differ in several cases 1° .

† D. Berthelot, C. R., cxxvi, p. 473, 1898.

ART. XIX.—*Notes on Certain Schists of the Gold and Diamond Regions of Eastern Minas Geraes, Brazil*; by ORVILLE A. DERBY.

IN a recent communication in this Journal (May, 1899) brief reference was made to certain schists of the region that were supposed to be in some way related to the curious schistose partings in the quartz veins that were more particularly discussed. The single type analyzed presented such an abnormal composition that it appeared desirable to make analyses, kindly executed by Dr. Hussak, of such material, unfortunately scanty, as was at hand as seemed suitable for that purpose. This material includes, in addition to the rock already analyzed by Hillebrand, two of the rocks mentioned in the above cited paper from the diamond mine of São João da Chapada. To these has been added a perfectly sound schist found loose in a topaz washing near Ouro Preto and presumed to represent the decomposed material in which the nests of topaz occur. In addition four analyses by Gorceix of schists of the Ouro Preto region are given from the Annaes da Escola de Minas de Ouro Preto (vols. 1 and 2, 1882, 1883).

	I.	II.	III.	IV.	V.	VI.	VII.	VIII.
SiO ₂	38.32	37.77	47.83	31.18	54.1	47.4	62.4	36.4
Al ₂ O ₃	28.16	21.37	26.75	20.39	27.3	31.2	22.7	26.1
Fe ₂ O ₃	2.24	28.36	8.51	36.41	7.7	6.7	3.8	23.1
FeO	4.02	---	---	---	---	---	---	---
CaO	0.32	tr.	tr.	tr.	---	---	---	1.3
MgO	12.04	1.05	2.43	0.79	0.8	2.0	1.3	2.1
K ₂ O	1.11	8.96	10.42	7.18	3.0	4.5	4.2	6.4
Na ₂ O	0.16	---	---	tr.	3.6	2.7	1.5	1.0
Loss on ignition..	8.01	3.89	5.33	3.83	3.8	5.6	3.4	3.9
Other elements ...	5.69	---	---	---	---	---	---	---
Total.....	100.07	101.40	101.27	99.78	100.3	100.1	99.3	100.3

- I.—Greenish schist with large crystals of cyanite; Serra do Gigante, north of Diamantina; loose block; analyst, Hillebrand (this Journal, May, 1889). "Other elements" include TiO₂ 4.93, ZrO₂ 0.09, MnO 0.16, (NiCo)O 0.04, P₂O₅ 0.47 and traces of SrO, Li₂O, S and F.
- II.—Bluish schist found loose in the diamond mine of São João da Chapada but presumed to come from a schistose layer in the lower quartzite (itacolumite); analyst, Hussak. TiO₂ occurs rather abundantly but was not separately determined.
- III.—Soft greenish schist found loose in the diamond mine of São João da Chapada but presumed to come from a schistose layer in the upper conglomeritic quartzite; analyst, Hussak. A strong trace of P₂O₅ was noted in the analysis.

- IV.—Black schist found loose in a small topaz washing near the Caxambú topaz mine, Ouro Preto, and which possibly represents the original matrix of the topaz occurring in nests in a clay similar to what would be given by the decomposition of this rock; analyst, Hussak.
- V.—Unctuous greenish schist from Boa Vista near Ouro Preto; analyst, Gorceix.
- VI.—Fibrous schist containing at times crystals of pyrophyllite and altered pyrite; Boa Vista near Ouro Preto; analyst, Gorceix.
- VII.—Whitish unctuous scaly rock from Boa Vista near Ouro Preto; analyst, Gorceix.
- VIII.—Compact violet schist in form of dike in limestone at Gandrela, north of Ouro Preto; analyst, Gorceix.

From macroscopic examination, No. I is evidently a chloritic rock with enormous crystals of cyanite up to 5^{cm} or more in length. The portion insoluble in sulphuric acid (43.41 per cent) calculated, according to Hillebrand's suggestion, as normal cyanite, muscovite, quartz and rutile, gives: cyanite 16.46 per cent, muscovite 8.46 per cent, quartz 13.66 per cent, rutile 4.73 per cent, with a doubtful excess of 0.09 per cent referred to zircon, which could not be found in the heavy residue, while a doubtful trace of zircon oxide was noted in the hydrochloric acid solution indicating that this element probably belongs to a soluble constituent. The soluble portion (56.66 per cent) reduced to 100 agrees very closely with Eaton's analysis of corundophilite from Chester, Mass., the most important differences being an excess of 1.88 per cent SiO₂, 1.13 per cent Al₂O₃, and a deficiency of 4.15 per cent of iron oxides. It is interesting to note that this Chester type of chlorite is also associated with an aluminous mineral, corundum, and that the somewhat similar type with the same association from the Calsagee mine, N. C., carries Mn, Ni and Co like the soluble portion of the Brazilian rock. Another element, monazite, revealed in minute quantities by washing, is probably indicated by the phosphoric acid of the analysis.

In the paper above cited an argument for the eruptive origin of this and other schists of the region was drawn from the occurrence of autigenetic monazite, which was considered as an original element that had passed unchanged through the process of metamorphism. This argument must now be withdrawn since, as will be shown in the following paper, evidence has recently come to hand that monazite may be formed by secondary processes; one of the proofs being furnished by the rock in question, which shows rutile included in monazite in the same manner as in the associated chlorite and cyanite. The probable origin of this rock will be discussed farther on.

No. II, as shown by microscopic examination, is essentially a micaceous rock heavily charged with a fine dust of hematite, which is easily removed by hydrochloric acid, leaving a considerable amount of a poorly characterized titanium mineral, rutile or anatase, which should give a small percentage of titanium not determined in the analysis. The alumina is nearly 8 per cent short of the amount required by the potash for normal muscovite and this deficiency is too great to be supplied by iron. Calculated on the basis of the Hallgarten sericite, however, the deficiency of alumina is reduced to a reasonable limit, and there can be no doubt that the rock is composed of 70 per cent more or less of sericitic mica with perhaps 2-3 per cent of chlorite and a small percentage of rutile, or anatase, and probably also of quartz. Washings reveal the merest trace of autigenetic monazite, which, in a measure, confirms the identification of this rock with the decomposition product (micaceous clay rich in iron oxide and with rutile) of a sheared dike in the vicinity.

No. III is in appearance a purely micaceous rock with no evidence, even in the heavy residue, of more than the merest trace of free quartz and hematite. It is so soft as to go easily to a slime in water and is apparently considerably decomposed, but this appearance is not confirmed by the high percentage of potash, which for normal muscovite requires 7.25 per cent more alumina than is given by the analysis. The rock probably contains over 80 per cent of an iron-bearing sericite with, perhaps, 7 per cent, more or less, of chlorite and a small percentage of quartz and earthy iron oxide. Washings reveal a small amount of microscopic tourmaline, of which some grains appear to be secondarily enlarged, and worn zircons of a size and abundance that seem extraordinary in a rock of such fine grain and of so purely argillaceous character. A trace of phosphoric acid was noted in the analysis and a very decided reaction was obtained from the heavy residue, so that it is certain that monazite is also present although in its rolled state it cannot be distinguished from the zircon.

No. IV is very similar to No. II in appearance and composition with a somewhat less deficiency of alumina for normal muscovite and a considerably greater percentage of free iron oxide. A trace of phosphoric acid indicates the presence of a phosphate, of which, however, no trace could be detected in the heavy residue, so that it probably is neither apatite nor monazite. The heavy residue shows a relatively small amount of autigenetic rutile and tourmaline and in this respect agrees with that of the decomposition clays of the immediate vicinity with nests of topaz, which, so far as can be made out, must have

resulted from the decay of a rock very similar to the one analyzed.

Nos. V, VI and VII were analyzed by Prof. Gorceix for the express purpose of determining whether the denomination "chloritic" or "talcose," ordinarily applied to the predominant schists of the region, was applicable or not. They may, therefore, be presumed to be the most typical unctuous schists that could be found in a sufficiently sound state to be analyzed. All three show a considerable amount of soda, which in No. V is in excess of the potash, so that the micaceous element cannot be a normal muscovite. Alumina is somewhat in excess of what is required to satisfy the alkalies as normal muscovite and it was noted that macroscopic pyrophyllite occurs in the bed from which No. VI was taken, so that it is probable that the excess of alumina may, in all three rocks, represent a small percentage of this mineral, or of cyanite, which is also very frequent throughout the region. Free iron oxide was noted in No. VI in the form of altered pyrite and it is probable that No. V also carries a small amount in the same state, or in that of hematite. All three show a considerable excess of silica, doubtless in the form of quartz, and the small percentage of magnesia probably represents a slight admixture of chlorite. Nothing is known of the heavy residues of these rocks or of the details of their mode of occurrence.

No. VIII has almost exactly the amount of silica and alumina required to satisfy the alkalies and magnesia as normal muscovite and chlorite, leaving an excess of about 8 per cent of silica for quartz. As in Nos. II and IV the iron is probably, for the most part, in the state of hematite.

All of these rocks show a low to extremely low percentage of silica and a high to extremely high one of alumina, while all without exception are poor in lime and, except No. I, in magnesia as well. On the other hand, all but No. 1 show tolerably high percentages of alkalies.

From these analyses no certain conclusion as to the original mineralogical composition and mode of origin of these rocks can be drawn. On the somewhat risky assumption that the present chemical composition of metamorphic rocks is, as regards non-volatile constituents, essentially identical with the original composition, no satisfactory comparison with known types can be made for the majority of these rocks. Nos. V and VII agree fairly well in composition with some of the trachytic and phonolitic tuffs cited in Zirkel's *Petrographie*, vol. iii, pp. 675 and 680. If, as is unusually done with such rocks, they be considered as metamorphosed arkose, the material must have been derived from a syenitic (or porphyritic) type rich in soda rather than from an ordinary granite or gneiss.

Neither of these hypotheses, however, will apply to Nos. III and VI, which, if original clastics, must have been almost purely micaceous sediments with, in the case of the latter, a slight admixture of kaolin. No. III is certainly clastic as proved by the heavy residue and it is presumed to represent a thin, shaly parting in a conglomeritic quartzite derived from an underlying micaceous quartzite and for this such an hypothesis is not improbable. Nos. II, VI and VIII also, if clastic, must likewise have been essentially micaceous but heavily charged with limonite. It is, however, improbable that in five representatives, taken by chance, of original argillaceous sediments, all should prove to be essentially micaceous and none essentially kaolinitic, and the hypothesis may be suggested that in the case of marine clays a gain in alkalis may possibly take place from soluble salts imprisoned from the sea water. No. I, regarded as a clastic, must have been a singular mixture of kaolinitic and magnesian clays.

On the other hand, the hypothesis of eruptive origin for any of these rocks, with the possible exception of Nos. V and VII, involves that of unknown and improbable types, or that of an important loss of lime and magnesia with a consequent concentration of alumina, iron and alkalis. In the paper above cited arguments for the possible eruptive origin of Nos. I and II were deduced from the presence of autigenetic monazite, the lack of recognizable allothigenetic elements and, in the case of No. II, from a presumed connection with a decomposed sheared dike in the immediate vicinity as well as of traces of original structure. The first argument has proved fallacious but the others still hold good and the second one is equally applicable to No. IV, while No. VIII (very similar in composition to Nos. II and IV) was taken to be a dike by Gorceix in his field examination, though it is probable that if he had known its composition a more detailed examination of this point would have been made.

In the same paper reference was made to a peculiar rock from near the fall of the river Dattas which presents many analogies with No. II but with more decided eruptive characteristics. Both of these rocks were supposed to be allied to No. I, and when that on analysis proved to be chloritic rather than sericitic, as was at first supposed, it was assumed that this might also be the case with the two rocks in question. A qualitative* analysis of the Dattas rock, however, shows that

* As the rock is heavily charged with tourmaline, which is evidently an introduced element, it was considered that its original composition had been so changed that the work of a quantitative analysis would not be compensated by the results.

like No. II, it is very poor in lime and magnesia and that it is composed essentially of a sericitic mica heavily charged with hematite dust. The two rocks are therefore closely similar, chemically as well as mineralogically, and it can safely be affirmed that the principal differences revealed by a complete analysis would be only such as correspond to the presence of the characteristic elements of tourmaline and to a greater abundance of titanium and probably also of iron oxide, both in the free state. These two rocks also agree in presenting on a polished section transverse to the plane of shearing, appearances of original structure that in the Dattas rock are perfectly well-defined but somewhat indistinct in the other. This appearance is that of the mottled aspect of a rock composed of an intimate mixture of white and colored elements, as, for example, a diabase or a basalt. In the Dattas rock the minute white areas are perfectly defined rectangles like those of the feldspar of the rocks mentioned, while in that from São João da Chapada (No. II) they appear to have lost their original sharpness of outline through shearing. Under the microscope, these rectangular areas of the Dattas rock are seen to be composed of a sericitic aggregate free from the iron and titanium dust with which the rest of the section is thickly sprinkled. On dissolving out the iron, the distinction between the white and colored areas disappears and the whole slide presents a sericitic aggregate sprinkled with a dirty white dust of titanium oxide, rutile or anatase, which before had been concealed by the iron. The appearance is that of a basaltic rock composed of a well crystallized white element (plagioclase or melilite) in a groundmass containing bisilicates (pyroxene?) or in a basic glass rich in iron and titanium oxides, which on alteration has settled in the place of the original minerals, or glass, but without invading the areas of the white element. In addition to this secondary iron (hematite) and titanium dust, the rock also contains these two elements as primary constituents in the form of crystals, often of considerable size, of magnetite and of minute octahedrons of an altered titanium mineral that was most probably perovskite.

The peculiar structure and characteristic primary accessories of this rock are strongly suggestive of an original basaltic character and, in view of the perovskite, most probably a melilite-basalt. On this hypothesis, however, the original rock must have been rich in lime and magnesia and its present chemical and mineralogical composition can best be explained by the hypothesis that before metamorphism it had been decomposed and leached *in situ*, the resulting residual clays and oxides retaining their original positions without mingling. A confirmation of this hypothesis is apparently to be found in

the peculiar manner in which the rock has been invaded by tourmaline. This mineral besides lining the shear and fracture planes with beautiful and symmetrical microscopic rosettes of radiating needles, also appears in the body of the rock, where, however, it is confined to the white areas, the needles being sharply cut off by the rectangular limits of the original crystals wholly or in part substituted by tourmaline. This invasion evidently took place after the decomposition and shearing but perhaps during the process of the recrystallization of the rock.

There is nothing intrinsically improbable in the hypothesis that some metamorphic schists may result from the alteration of rocks, both massive and clastic, that had been decomposed and leached *in situ*, and thus many of the anomalies of composition that this group presents may be satisfactorily explained. This hypothesis is involved in that presented by Vogt and Van Hise of the derivation of micaceous iron schists of Norway and of the Lake Superior region from original carbonates, and in another place evidence pointing in the same direction will be presented in relation to the extensive itabirite beds associated with the schists here discussed. If, as seems probable, the itabirites are derived from decomposed and leached carbonates, the process must have taken place on a gigantic scale in the region in question and the associated silicate rocks, whether massive or clastic, must have been affected to a greater or less extent.

The hypothesis of leaching cannot, however, be applied satisfactorily to the rock No. I, which, if an original basic eruptive of any of the ordinary types, must have lost heavily in lime and iron while magnesia was retained. This, however, would be contrary to what is supposed to have taken place in the other rocks here discussed and to what is generally observed in the modern cases of decomposition and leaching. Moreover the presence of monazite in this rock is, so far as present experience goes, incompatible with such types. It seems more probable that in this case the composition has remained comparatively unchanged and that the original rock must have been one of low silica, iron and lime, but of high alumina and magnesia associated with a considerable variety of rare elements. No type of rock with these characteristics has as yet been clearly defined, but that such may exist, either independently or as local phases or segregations in other types, is suggested by the occurrence of corundum and aluminous silicates in association with olivine rocks in North Carolina and Georgia, and that of alumina-magnesia silicates such as cordierite, prismatic and sapphirine as segregation (?) masses in the granulite of Saxony and the mica schist of Fiskernäs, Greenland. If, as may be presumed from the scanty information at hand,

these latter minerals, with their associates, form independent rock masses, a possible prototype for such schists as No. I can be imagined. The question still remains of the mode of origin of such a prototype if it exists, but the occurrence of corundum in olivine rocks shows that the association of alumina and magnesia in ultra basic eruptives is not an impossibility.

The probable occurrence in this rock of a soluble compound of zirconia is also suggestive of a possible prototype in the nepheline- or augite-syenite groups, in which such compounds have thus far proved to be most frequent and in which rocks with corundum indicating an excess of alumina have lately been recognized. Thus far, however, no member of this group at all approaching in composition the Serra de Gigante rock appears to have been described. That rocks of this group may have existed in the district and have contributed to the metamorphic schist series, is perhaps indicated by the perovskite- and, possibly, melilite-bearing schists from the neighborhood of Dattas.

As to how far the above analyses are typical of the phyllites of the series in question can only be conjectured. These, which as regards thickness are evidently more important than the quartz, iron-mica and calc-schists with which they are associated, present almost universally an aspect that has led to their being generally denominated as chloritic or talcose, that is to say they are predominantly of micaceous texture and of a character that suggests magnesian minerals. True chloritic and talcose as well as amphibolitic schists undoubtedly occur in the series, but it is almost certain that they are subordinate to sericitic schists, agreeing more or less perfectly in character with Nos. II to VIII of the above analyses. So far as can be judged from a superficial examination of material which for the most part is profoundly decomposed, the more quartzose types represented by Nos. V. to VII are the most abundant and characteristic, and these, from their composition and intimate association with the quartz schists, may be presumed to be, for the most part, of elastic origin. If so, however, the original clays must have been sufficiently rich in alkalies to form a sericitic mica as the predominant element of their metamorphosed state. Judging from the above analyses and from the almost universally micaceous character of the phyllites of the region, the proportion of alkalies must have been very uniformly above 5 per cent.

It is interesting to compare with the proportion of alkalies above deduced for these phyllites that given in the extensive list of analyses of ceramic materials in vols. 16 and 18 of the annual reports of the U. S. Geological Survey. Out of 550

analyses in which the alkalis were determined, only 67 (12 per cent) have over 5 per cent of potash and soda; 99 (18 per cent) over 4 per cent, and 177 (30 per cent) over 3 per cent. In the 87 that are given as slate, shale or shale-clay, the proportion having over 5 per cent is the same (11=12 per cent), but below this figure it rises rapidly, the numbers being 22 (25 per cent) with more than 4 per cent of alkalis and 62 (72 per cent) with more than 3 per cent.

In the low proportions of alkalis the difference above noted may reasonably be attributed to the greater amount of water in the earthy types, but this explanation will not apply in the comparison of the shales with the Minas schists, which, considered as original clastics (as a large, and probably the greater, part undoubtedly were), appear to have been uniformly richer in alkalis than the normal clay deposits of more recent times. If this conclusion is correct, these ancient argillaceous deposits must either have been composed quite uniformly of comparatively sound felspathic or micaceous material rather than of kaolin, or in some way the original proportion of alkalis must have been increased. The hypothesis above given that some of these schists are of eruptive origin will in part explain the predominance of micaceous rocks in the region, but, for the present at least, there are no good reasons for supposing that these constitute more than a small fraction of the whole. The case of No. III shows that a high proportion of alkalis is not necessarily indicative of eruptive origin and a number of other highly micaceous rocks have been washed with the same result of a residue with well-characterized allothigenetic elements.

An excess of alumina over that required by the alkalis to form mica, and which goes to form aluminous silicates, seems also to be a common feature of the schists of the region, as cyanite is a very common mineral both in the rock outcrops and in the loose material of the surface, as well as an element, often predominant, of the concentrates of the gold and diamond placers. In general, however, it appears to come from the more quartzose rather than from the more micaceous members of the series and in these cases it may be presumed to represent kaolin in original clastic sediments. Other aluminous silicates that resist decay so as to appear in the washings, as staurolite, are not at all prominent. From some of the above analyses it may be suspected that pyrophyllite will prove to be a common and widespread element, but from its liability to be confounded with mica it can only be detected in perfectly sound rocks, that are rare in the region. As above indicated, this excess of alumina may be attributed to the normal presence of kaolin in original clastics, or to the leaching of original eruptives.

A high proportion of iron oxide, more frequently in the state of hematite than of magnetite, is also a very common feature in the micaceous as well as in the quartzose and calcareous rocks of the region. In the latter case and in that of the actinolitic and a part at least of the quartzose iron schists (itabirites), it may be presumed to come from original carbonates. In other cases it may be attributed to iron sand in clastic deposits, to original elements either oxides or silicates, in eruptives, or to a leaching of limonite into rocks of any character. It is almost invariably accompanied by a certain proportion, often high, of titanium oxide as rutile (more rarely as anatase) which after quartz and the iron oxides is the most abundant and constant mineral of the gold and diamond concentrates. This, whether in elastics or eruptives, is probably indicative of original ilmenite. Chrome mica occasionally appears in the quartzites of the region and in one case (a pebble from the conglomerate of the Cavallo Morto diamond mine near Diamantina) this could be traced very satisfactorily to grains, evidently clastic, of chromic iron contained in the same rock and which probably indicate that a peridotitic rock has contributed to the original sandy deposit.

ART. XX.—Notes on Monazite; by ORVILLE A. DERBY.

Solubility in acids.—In order to test the relative solubility of the mineral in different acids a sample of monazite sand from Prado, Bahia, was freed from all admixture by careful picking and after grinding was submitted in parcels of about half a gram each to the action of 100 centigrams of the ordinary laboratory acids each diluted with one part of water. After standing for 68 hours without heating, phosphoric acid was determined in each solution as follows: in nitric acid solution, 2.51 per cent; in hydrochloric acid, 1.62 per cent; in sulphuric acid, 1.39 per cent.

Magnetism.—In cleaning up residues with the electro-magnet it has been found that monazite can be quite successfully separated from its usual associate, zircon, and that when monazite and xenotime occur together the latter can, by a proper graduation of the points, be almost entirely drawn away from the former. In one case in which the instrument had been used with an opening between the points of about a centimeter, for the separation of pyroxene of the acmite type which may be presumed to carry about 30 per cent of iron oxide, a mixed sample of monazite and zircon was very neatly and quickly separated without reducing the distance between the points, both minerals being equally free from iron staining or inclusions.

Microchemical reactions—A single granule of the mineral, no matter how minute, can be rapidly and securely identified by moistening it with sulphuric acid on a slip of glass and burning off the acid over a spirit lamp. The characteristic crystallization of cerium in double ball-shaped clusters of radiating needles or minute cucumber seed-shaped isolated crystals, can usually be detected after this operation in the ring of evaporated material about the granule, but better after adding a drop of water and allowing it to evaporate in a desiccator. Another drop of water with a slight admixture of ammonium molybdate solution added to the same preparation gives on evaporation a very satisfactory reaction for phosphoric acid. The reagents are best applied by means of a small loop on the end of a very fine platinum wire and an excess should be avoided, especially with the acid, as too large a drop is liable to run in a very annoying manner in the heating. The same reactions are given by the recently discovered cerium-aluminium phosphate, florencite of Hussak and Prior, but this can be distinguished by its form and cleavage when these are recognizable. The microcrystalline forms of cerium and yttrium

sulphates are so similar that a confusion with xenotime is also to be guarded against. It has recently been found also that soluble silicates containing zirconium give very similar forms in the sulphuric acid test, but these are readily distinguishable by the absence of phosphoric acid. Confirmatory tests with oxalic acid have not proved uniformly successful and can usually be dispensed with. The Florence test by crystallizations in a blowpipe bead requires from a half a dozen to a dozen grains of the usual size and is more successful with the salt of phosphorus than with the borax bead in which the presence of phosphoric acid appears to exercise a disturbing influence, although with patience the crystals characteristic of cerium can be obtained.

When, as is usually the case, the grains are transparent, the micro-spectroscope will also usually give a very satisfactory test on a single grain of the usual size, by means of the absorption band of didymium. This can also be obtained by either reflected or transmitted light with an ordinary hand, or rain-band, spectroscope, when a number of grains can be brought close together in balsam on a microscopic slide.

Natural etching.—The grains of monazite from decomposed pegmatites or muscovite-granites are frequently etched, though not to the point of completely obliterating their original form and faces, but the phenomenon has rarely been observed in the decomposition-products of other types of rocks. In a miner's residue from the Cavallo Morto (Dead Horse) diamond mine near Diamantina the extraordinarily abundant monazite grains (the mineral is lacking or rare in the residues of most of the mines of the district) are extremely fresh in appearance, presenting a strong contrast with the well-worn aspect of the associated zircons, and under the microscope this difference is seen to be due to the profound natural etching of the grains. The deposit is reported by a competent observer to be a decomposed metamorphosed conglomerate like others of the vicinity, and in this case the monazite grains should be equally worn with those of zircon. On the contrary, however, they have been completely rejuvenated in appearance and amongst thousands passed in review under the microscope, none showing what could be positively identified as original or worn faces were seen except when, as in the case of the latter, these had been protected by the secondary enlargements described below. In similar deposits in the neighborhood in which the cement of the ancient conglomerate was evidently highly argillaceous, a similar etching of the included quartz grains is almost universal and is often very beautiful.

Secondary enlargement.—In the above-mentioned residue from the Cavello Morto mine, many grains show a darker cen-

tral portion surrounded by perfectly clear material having the same optical orientation, but, like all the grains of the residue, without distinct crystalline form. The appearance is that of a rounded and worn grain with secondary enlargements principally in the form of clusters of spindle-shaped prolongations of the two poles of an original ovate grain. So far as can be made out from the color and from rather unsatisfactory optical tests, the clear outer shell is of monazite material, and this conclusion is apparently confirmed by microchemical tests, since grains in which the central nucleus appeared to be entirely protected from the action of the acid gave very satisfactorily the cerium and phosphoric acid reactions. As there was a possibility that the enlargement might be the newly discovered cerium-aluminium phosphate, florencite, which, when the form is not distinct, is readily confounded with monazite, a small quantity of the residue was tested by wet analysis for alumina, but with a negative result. It thus seems tolerably certain that in the clastic rock from which the residue comes there had been a new formation of monazite such as has been shown for tourmaline and several other minerals in similar rocks. The main objection to this conclusion is that monazite, as an autigenetic element, is thus far only positively known in rocks of eruptive or presumably eruptive origin, but this objection is weakened by evidence presented below indicative of a possible secondary origin.

Inclusions.—The monazite thus far extracted from granites, gneisses and porphyries, and that found in secondary deposits derived from these rocks, is free from inclusions, and in these cases the mineral, together with the associated accessories (zircon, magnetite or ilmenite, or both, and occasionally the phosphates xenotime and apatite), has evidently been one of the first to crystallize out of an eruptive magma. Quite recently, however, evidence has come to hand, that it may also be formed in other ways and probably by secondary processes. In addition to the secondary enlargements above noticed, which were evidently formed in a metamorphosed clastic rock, the following cases have been noted.

From the residues of diamond washings at São João da Chapada and Sopa near Diamantina, and from a gold washing at Bandeirinha in the same neighborhood, the monazite crystals of the peculiar prismatic habit described in a recent number of this Journal (vol. vii, p. 353) are frequently found heavily charged with scales of hematite and more rarely with minute needles and grains of rutile. This type of monazite has been traced home to some peculiar schists of the region that are presumed (in part on the evidence of the perfectly fresh condition of the monazite crystals), to be of eruptive

origin. In the schists from Sopa and São João da Chapada, which are almost purely sericitic, the monazite is free from inclusions, but in the chlorite-cyanite schist from the Serra do Gigante the monazite, like the accompanying chlorite and cyanite, is full of minute grains of rutile and no essential difference in the character and mode of occurrence of the inclusions in the three minerals can be detected. Whatever may have been the original nature of this rock, the silicate elements and the rutile are clearly secondary, and there is no escaping the conclusion that the monazite must be so also.

In a small gold-working called Ogó* a few hundred meters away from the great diamond mine of São João da Chapada, monazite of the ordinary type occurs with rutile in the peculiar partings, composed almost exclusively of bright green muscovite, of a quartz vein in diabase. The crystals are here of the type known as turnerite and both they and the associated muscovite and rutile closely resemble those of the microscopic residue obtained from a specimen of the Binnenthal gneiss faced on one side with the well known macroscopic crystals of that locality. In the case of the Brazilian crystals, however, rutile and muscovite are included in them in such a way as to show contemporaneous crystallization. The rock is a peculiar one and in its present condition may be suspected to be a secondary product, but this cannot be proven. A decomposed massive quartz-muscovite rock from the same mine, which in appearance is of the same nature as the partings except for being finer-grained and more quartzose, did not give monazite. The exposure is unsatisfactory and as in the field examination the greenish rock was taken to be a chloritic alteration phase of the diabase, no special attention was given to it beyond taking a specimen that quite unexpectedly proved of great interest. The quartz vein with its peculiar micaceous accompaniment suggests a comparison with the quartz-kaolin (pegmatitic?) veins of the same vicinity and may be suspected to be a special type of granitic apophyses.

If the possibility of a secondary formation of monazite be admitted, a more plausible explanation than that given in the paper above cited (that of an intimate interlamination of eruptive and elastic material, can be given for the micaceous partings and selvages of the Sopa quartz veins that afford perfectly fresh monazite mixed with well-worn zircons.

* Ogó, which is apparently a term of African origin, is the name given to the fine heavy concentrate of microscopic minerals obtained in panning for gold and diamonds, when the predominant color is yellow. It may consist of monazite, xenotime or rutile, or a mixture of all three, the first generally predominating. In panning the miner, seeing the formation of a heavy yellow concentrate, is often deceived and on cleaning up exclaims in disgust "the gold turned to ogó".

Occurrence in basic rocks.—Up to a very recent period the rule seemed to be very firmly established that it was useless to look for monazite as an autigenetic element in any but highly acid rocks. In the scores of rocks examined, both sound and decomposed, it had been found to be almost universal in the muscovite granites, or their porphyritic and gneissic equivalents; frequent in the biotite granites, but lacking in the amphibole-granites and all other more basic rocks that had been examined. The monazite-bearing schists above mentioned from Sopa, São João da Chapada and Serra do Gigante, however, indicate that this rule is not general, though as the type of mineral that they present is peculiar and perhaps of secondary origin, it may be that it only requires to be somewhat modified rather than set entirely aside. The Sopa schist (that one free from zircon) have not been analyzed, but it is evidently composed almost exclusively of sericite and there is no difficulty in considering it as a modified porphyry. That of São João da Chapada, though basic in the low silica (37.77 per cent) and high iron (28.36 per cent, for the complete analysis see No. II of the preceding paper) contents, may, as I attempted to show in the paper above cited, be plausibly considered as a basic phase of a mixed dike of essentially granitic character. No such hypothesis is, however, possible with the Serra do Gigante schist with its 38.32 per cent of silica, 28.16 per cent of alumina and 12.04 per cent of magnesia. As already stated under the head of inclusions, all the elements of this schist are apparently secondary and it might therefore be considered as a metamorphosed sedimentary clay. The high percentage of magnesia is, however, anomalous for such a clay and in the preceding paper in which this rock is more fully discussed the hypothesis of the decomposition and leaching, prior to metamorphism, of an original eruptive rock is suggested. Even so, however, it seems necessary to imagine a type of eruptives that has not yet been clearly recognized though there are indications of its possible existence.

ART. XXI.—*The Spectra of Hydrogen and the Spectrum of Aqueous Vapor*; by JOHN TROWBRIDGE.

It is customary to consider that there are two spectra characteristic of hydrogen—a four-line spectrum, so-called, and another consisting of many lines widely distributed through the spectrum, known as the white spectrum. The four-line spectrum appears when a condenser discharge is employed with what is called dry hydrogen: it is also readily produced in steam and water vapor. From the fact that a condenser discharge seems necessary to excite it in dry hydrogen, it is supposed to indicate a higher temperature than the white spectrum.

The four-line spectrum is found in the atmosphere of the sun; and is a characteristic spectrum of certain types of stars. There are also other lines attributed to hydrogen in the stars which are supposed to indicate conditions of pressure and temperature which perhaps can be imitated and studied in laboratories.

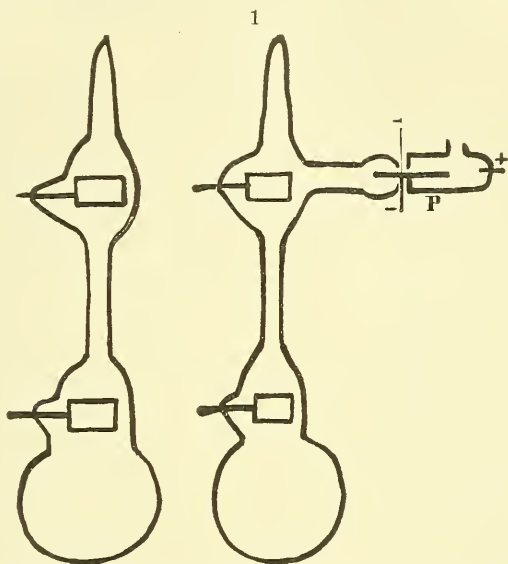
I propose to show, in this paper, that conclusions in regard to temperature and pressure of hydrogen in celestial bodies, deduced from observations on hydrogen enclosed in glass vessels, are untrustworthy; and that electrical dissociation must be considered as a most important element in determining the characteristics of a gaseous spectrum—more important indeed than the question of pressure and apparent temperature.

I shall give my reasons for believing that the four-line spectrum of hydrogen in the atmosphere of the sun is an evidence of aqueous vapor in that atmosphere and therefore is an evidence of the existence of oxygen in the sun. The conviction is forced upon me that the term dry hydrogen is a misnomer, when the gas is subjected in glass vessels to condenser discharges or to sufficiently powerful steady currents of electricity.

The bibliography of the subject of the spectra of hydrogen, is so extensive that I must, with due regard to the limits of this article, refer the reader to the reports of the committee of the British Association, and to the recognized bibliographical authorities on this subject, and I do this because I feel that in my experiments I have exceeded the experimental limits of previous investigators; for I have been enabled to employ more powerful electrical discharges than have been hitherto possible. My work, therefore, does not trench in this respect upon that of previous investigators.

The source of the electrical energy I employed was twenty thousand storage cells of the Planté type. The direct current from these cells through a liquid resistance was used to produce the white spectrum; and a glass condenser consisting of 300 (three hundred) plates of glass—each plate having a coated surface of 16x20 inches—with a total capacity of about 1.8 microfarads—was charged by the cells to produce the four-line spectrum.

I had great difficulty at first in obtaining tubes which would stand such powerful discharges. I began my work with end-on tubes which were closed by plates of quartz luted on by silicate of soda. The electrodes were hollow cylinders of aluminum connected to thick pieces of platinum wires. These



wires passed through the walls of the glass tubes, and were immersed in large vessels of distilled mercury. After considerable experience tubes were constructed which would resist the disruptive and heating effects of the discharges. This form of tube, however, was abandoned for another form, shown in fig. 1. The chief peculiarity of this form of tube is an X-ray bulb in place of the end covered with the quartz plate. The reason of the adoption of this form of tube is this: it was necessary to heat the entire tube to a high temperature for a long period during the process of exhaustion, to drive out the air and aqueous vapor before it was filled with hydrogen, and this heating was impossible with a luted-on end.

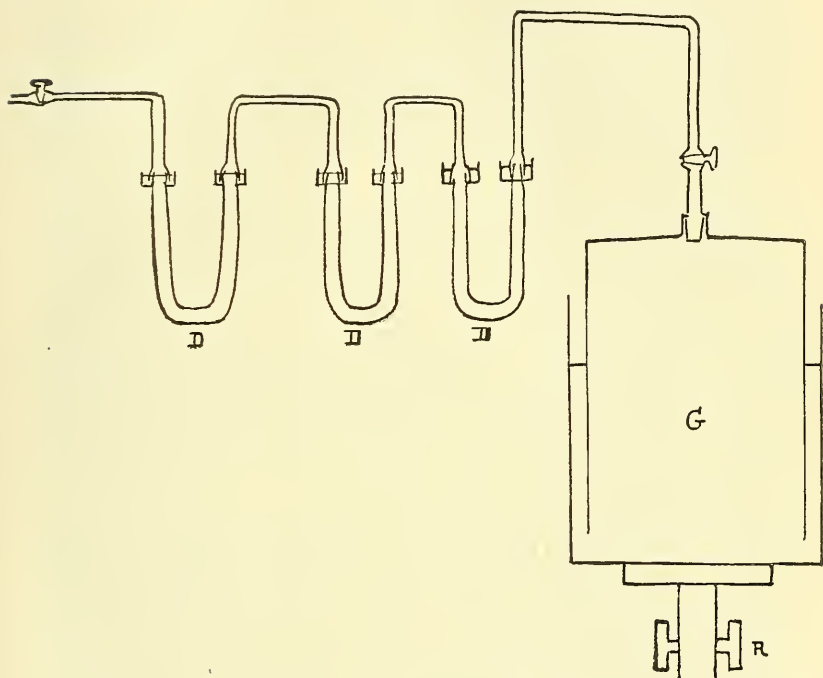
Moreover the thin glass of the bulb did not appreciably absorb the ultra-violet rays until the wave length 1800 was reached, and this absorption did not affect my conclusions since my work was confined to the portions of the spectrum studied by astrophysicists. The diameter of the narrowest portion of the tubes was about 1cm . I employed still another form of tube, shown also in fig. 1. At one side of this tube there was an adjunct consisting of a palladium tube P on the outside of the spectrum tube joined hermetically to a platinum wire inserted in the walls of the tube. A glass vessel filled with dilute sulphuric acid slipped over the palladium tube. At one end of this glass vessel was inserted a platinum wire. The palladium wire was made the cathode and the platinum wire the anode of a battery. In this way hydrogen is liberated on the surface of the palladium tube, is then occluded and can be made to pass into the spectrum tube.*

The tubes which were not provided with palladium tubes were filled with hydrogen or nitrogen, through a series of tubes filled with caustic potash, and through a number of tubes filled with phosphoric pentoxide provided with partitions of glass wool. After many attempts I adopted the following arrangement of these tubes, which proved satisfactory. In the first place I found it necessary to discard all gums or other adhesive material and to use ground-glass joints and mercury seals for these joints, having satisfied myself that all other forms of connections consume time and are worthless. The tubes were mounted on a board and after the joints had been adjusted the final adjustment with the gas holder could be readily and safely accomplished by means of a rack and pinion which raised or lowered by slow degrees the gas holder fig. 2. This drying arrangement could be exhausted to a high degree and would hold the rarified gas as long as desired. The method of filling the tubes was as follows: the spectrum tube was exhausted to about $\cdot 1\text{mm}$, having been repeatedly filled with hydrogen. This operation was continued until the hydrogen spectrum appeared. Finally the tube was exhausted to the X-ray stage by long heating and by employing a strong condenser discharge. Then the dried gas was admitted until a pressure of about $\cdot 1\text{mm}$ remained in the tube. The vessels with palladium tubes were carefully exhausted to the X-ray stage (and beyond, a spark preferring to jump six inches in air to passing through the tube). When these tubes with the palladium adjunct were connected to the terminals of the battery a slight heating was sufficient to start a discharge and to cause the occluded hydrogen to appear in the spectrum tube.

* A device due to Dr. William Rollins, Boston.

The operation of heating the spectrum tubes is highly important; moreover the discharge from a condenser should be employed in the process of exhaustion and the operation should be continued for several hours. In this subject of the exhaustion of tubes I am much indebted to suggestions of Dr. William Rollins and Mr. Heinze of Boston, who have had long experience in the preparation of X-ray tubes. This experience I believe is indispensable to one who essays to investigate the spectra of gases in glass vessels.

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The vacuum pump was a modification of Rapp's pump. It was automatic. There was but one valve and the number of ground-glass joints with mercury seals was reduced to a minimum. The hydrogen was produced by the electrolysis of phosphoric acid and doubly distilled water. The gasholder contained many liters of the gas. The nitrogen was prepared by passing air through ammonia water, then through heated copper gauze, then through tubes of caustic potash; and finally through tubes of phosphoric pentoxide.

The spectroscope consisted of a Rowland grating with a ruled surface five inches by two inches, and a radius of curva-

ture of sixty inches. The dispersion was small on account of the small number of lines to the inch; but it had the advantage of exhibiting the salient features of the spectra on one plate, with a dispersion of the order used by astrophysicists; and thus enabled comparisons to be readily made. The distance between the great H lines in the second spectra was approximately 2^{mm}. The camera was on an arm pivoted at half the distance between the slit and the grating. This arrangement was adopted for compactness. The grating and camera were enclosed in a light-tight box and a lever enabled me to expose the plate when this box was closed.

Various liquid resistances were employed to regulate the strength of the electric current. I finally used distilled water, which was kept running through glass tubes. Since my object was not to measure the current, but rather to keep it within safe limits, I did not find it necessary to employ cadmium terminals in a solution of iodide of cadmium. The direct battery current gave me the white type of hydrogen spectrum, while the second type was given by the powerful condenser discharges. A revolving drum provided with suitable metallic strips enabled me to safely charge and discharge the condensers through the spectrum tubes.

In working with powerful condenser discharges I found that I obtained essentially the same spectrum whether the tube was filled with hydrogen, rarified air or nitrogen, notwithstanding the great care taken in heating the tubes and drying the gases; this characteristic spectrum was that of water vapor—containing lines of hydrogen and of oxygen—with traces of nitrogen lines.

The employment, therefore, of a condenser discharge in glass tubes drives off aqueous vapor from the glass walls, even if extraordinary care is taken to expel such vapor in the process of exhausting such tubes. My experiments entirely confirm Professor Crookes' statement on this point. Glass seems to be of a permeable nature, and in the process of manufacture is doubtless filled with gases which reveal themselves when disruptive electrical discharges are employed. It does not seem possible that dry hydrogen can exist as such in a tube when it is subjected to a condenser discharge; and I therefore believe that the spectrum of hydrogen called the four-line spectrum is an evidence of the presence of aqueous vapor, for it can be produced with the greatest ease when we are certain that this vapor is present. This spectrum, to my mind, is an evidence of the presence of oxygen in the sun's atmosphere.

The true spectrum of hydrogen seems to be the white spectrum. This is developed by a steady current which should

not exceed a certain strength; for a powerful one drives off also rarified air from the walls of the tubes and there is a process of occlusion set up if large metallic terminals are employed. Such large terminals are necessary if very strong currents are used. With aluminum electrodes under the action of strong currents, spectra of the combinations of nitrogen with oxygen and hydrogen are produced in tubes presumably filled with pure dry hydrogen. When the tubes have been subjected to very strong currents it seems impossible to refill them with pure hydrogen. In one case I filled a tube with pure dry nitrogen and passed a current through it sufficiently powerful to melt down one of the aluminum hollow cylinders which formed an electrode. The aluminum surface was filled with cavities as if gases had bubbled from the interior. When this tube was filled with dry hydrogen it gave again spectra of compounds of nitrogen and no trace of the hydrogen spectrum. A powerful condenser discharge, however, produced the aqueous vapor spectrum. Hydrogen, thus, is extremely elusive when subjected to electrical dissociation in the presence of other gases in glass vessels with metallic electrodes, and indeed, a steady battery current can be employed at a certain stage of exhaustion to occlude hydrogen to such a degree that the vacuum is apparently raised and no discharge can be forced through the tube until it is subjected to external heat. This occlusion or destructive electrical dissociation is very much in evidence when the terminals of a battery of ten or twenty thousand storage cells are connected to an X-ray tube through a resistance of several megohms. If this tube is properly exhausted, that is by repeated heating and by the use of electrical discharges, no current can be forced through it unless the tube is repeatedly flashed with a powerful Bunsen burner. Suddenly, however, under the action of the heat a blue cloud rises in the tube. The cathode beam appears, the anticathode grows intensely hot and the exterior resistance must be quickly increased to save the tube. The pressure has apparently been greatly reduced; but the experiment of Dr. William Rollins with two connecting X-ray tubes, one of which is heated and the other not, the heated one carrying a current and the other resisting an eight-inch spark, shows that this is not the case and that we have to deal with electrical dissociation and not with change of pressure. Let us follow this experiment further. Presently the blue cloud in the tube grows smaller, the cathode beam disappears and the cloud slowly sinks into the positive electrode. At that moment the X-rays flash out with great brilliancy; but in a few seconds the light in the tube is totally extinguished, and a further heating from an external source is necessary before the phenomena can be repeated. It can be

repeated as often as desired, although the apparent resistance of the tube undoubtedly is raised by a more or less permanent occlusion of what I believe is water vapor. This experiment strongly recalls the familiar one of the glowing of platinum under a stream of non-ignited hydrogen. The electrical dissociation of the aqueous vapor evidently produces an intense heat at the anticathode and the X-rays are emitted during this dissociation.

If we suppose that there are material bodies circulating about the sun and charged negatively—the sun being charged positively—we might conceive of a similar action of a difference of potential on rarified aqueous vapor, which would be competent to produce a corona.

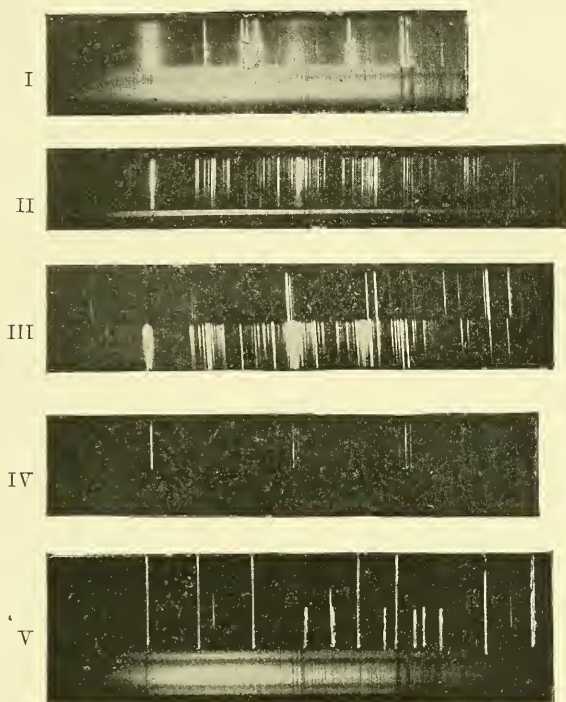
The white spectrum of hydrogen produced by hydrogen coming from palladium is in its main feature similar to that obtained from electrolytic hydrogen which is passed through the drying apparatus I adopted.

There are some bands, however, which need examination to determine whether they arise from impurities conveyed into the tube from the pump. The palladium tube apparently supplies hydrogen more or less continuously to the tube to repair the loss from the process of electrical dissociation, and is therefore more reliable than the form of tube which is filled with hydrogen through drying apparatus which necessarily cannot be subjected to heat, and through which rarified air is conveyed with hydrogen into the spectrum tube.

The study of the effect of powerful electrical discharges on hydrogen led me to endeavor to find the lines in the star ζ Puppis discovered by Professor E. C. Pickering: since their approximate wave lengths satisfy a modification of Balmer's formula, Professor Pickering attributes them to hydrogen. The wave lengths of these new lines are comprised in the spectral region extending from about 4200 to 3700. I have plotted them as short lines contiguous to the normal solar spectrum in fig. 3, No. V. The long lines correspond approximately to the most intense lines or bands in the spectrum of hydrogen produced in the tube provided with the palladium adjunct. The hydrogen spectrum when regarded as a whole on the scale of small dispersion I have employed, seems to be made up of lines spaced according to a certain order, very much as if two sets of lines spaced according to a certain arrangement should be superposed on each other; the fingers of one hand shifted over those of the other. The hydrogen lines are more or less intensified bands or dark accumulation of lines almost obscured by the spectra produced by the compound of nitrogen and aqueous vapor if a very strong battery current is employed in a tube provided with aluminum electrodes. It is possible that certain

hydrogen bands may be narrowed and reudered lines in a spectrum of small dispersion such as astrophysicists are compelled to employ, by an electrical dissociation of water vapor in the presence of an excess of oxygen, and that these new lines may thus be evidence of the presence of aqueous vapor in the stars of this new type. To test this theory I filled a tube with

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oxygen, and submitted it to a powerful condenser discharge. The resulting spectrum was of the general type obtained by these strong discharges in hydrogen, rarified air and nitrogen. It is shown in fig. 3, No. IV. If we compare this spectrum with I, in which the broad bands represent hydrogen lines in hydrogen at atmospheric pressure; with II, the spectrum of rarified air; also with the lower one in III, the spectrum obtained in pure nitrogen, we see that the main hydrogen lines are narrower and certain hydrogen lines are so faint as to be hardly visible even on the negative.

The general character of the spectral lines of hydrogen in respect to breadth and intensity seems not to depend so much

upon mere pressure and apparent temperature (temperature deduced from calculations of electrical energy in electrical discharges) as upon electrical dissociation of gases: of aqueous vapor, for instance in the presence of an excess or a small supply of oxygen. I believe, therefore, that any conclusion in regard to temperature and pressure of hydrogen in the stars deduced from laboratory experiments with electrical discharges in glass vessels are misleading, since we have mainly to do with phenomena of electrical dissociation and not of pressure and apparent temperature—that is temperature which can be measured by calorimetric methods or which can be calculated from the electrical dimensions of the circuit which is employed.

It would seem, therefore, that the study of electrical dissociation is necessary for the solution of problems relating to the conditions of gases in the atmosphere of the sun and in that of the stars. This seems to be a logical necessity, since the electromagnetic theory is far reaching and sensible heat is only one of the manifestations of electrical energy.

My conclusions are as follows:—

1. When a condenser discharge is sent through a rarified gas which is confined in a glass vessel, the gas can no longer be considered in the dry state: for aqueous vapor is liberated from the glass. When a sufficiently powerful condenser discharge is employed, dry hydrogen, dry nitrogen and rarified air give substantially the same characteristic spectrum. When a very powerful steady battery current is used to excite the tubes filled with these gases—various compounds of nitrogen and oxygen—nitrogen and hydrogen are formed if aluminum electrodes are employed.

2. The four-line spectrum of hydrogen in the solar atmosphere is an evidence of aqueous vapor, and therefore of oxygen in the sun.

3. Conclusions in regard to the temperature of the stars exhibiting hydrogen spectra are misleading, if based upon conditions of pressure and temperature in glass vessels; for conditions of electrical dissociation of aqueous vapor, for instance, the presence of an excess or lack of supply of oxygen, are the controlling ones rather than conditions of the mere pressure of the gases.

4. X-ray phenomena produced by a steady battery current strongly suggest an electrical theory of the origin of the sun's corona.

Jefferson Physical Laboratory, Harvard University.

ART. XXII.—*On a New Effect produced by Stationary Sound-Waves;* by BERGEN DAVIS.

IN the course of an investigation upon the action of sound-waves, under certain conditions, I have found that if a small cylinder, which is closed at one end, is placed in a stationary sound-wave, it will not only arrange itself perpendicularly to the motion of the wave, but will actually *move* across the wave in a direction perpendicular to the stream-lines. The force producing this motion is of considerable magnitude and it acts normally to the closed end of the cylinder, causing it to move in the direction of the closed end, i. e., the closed end is driven toward the wall of the pipe.

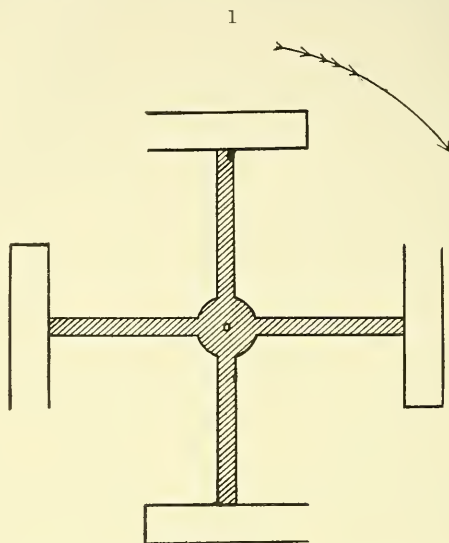
In investigating this effect I have used an organ pipe for the production of the stationary wave, its length being 70^{cm}, and the breadth and thickness 6.35^{cm} by 5^{cm}, respectively. The pipe when giving its first overtone made 358 vibrations per second. From the dimensions above given, it follows that a node will exist at 22^{cm} from the mouth-piece, and also at the stopped end. One wall of the pipe was made of glass. A thin rubber diaphragm was placed across the pipe near the node next to the mouth-piece, thus converting the remainder of the pipe into a closed chamber. The air in this chamber was in this way kept free from any currents that might arise from blowing the pipe. A diaphragm of this kind at the node, does not alter the intensity of the sound appreciably. The pipe, when blown very hard furnished a very strong overtone.

The small cylinders were made from "No. 00 gelatine capsules" such as are used for medical purposes. The hemispherical ends of these capsules were removed, and flat paper discs substituted for them at one end, the other end being left open. Four of these cylinders were arranged as shown in fig. 1. This system of cylinders was supported by a small glass pivot, which was mounted on the point of a fine needle. The needle was attached to a glass rod which was destined to carry the system into different regions of that part of the pipe located behind the protecting diaphragm.

The strength of the air blast was measured by a mercurial manometer, which was used to obtain the same force of blast throughout a series of experiments. The pipe was placed vertically with the stopped end downward, and the glass rod bearing the system of cylinders was introduced from below through an airtight opening. It could be moved up and down in the pipe, and the system of cylinders placed in any desired position in the stationary wave. The plane of the system of

cylinders was perpendicular to the direction of the stream-lines. A scale was placed along the pipe in such a position that the zero was at the stopped end, and the node was at division 48 and the middle of the loop at division 24.

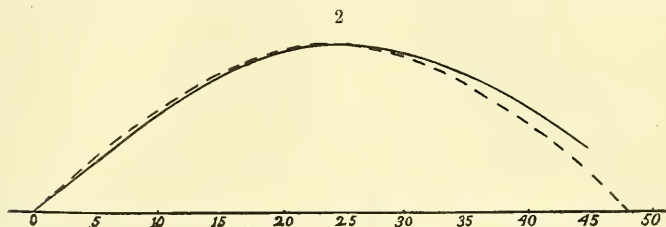
A stroboscopic disc was used to determine the rate of revolution.



Matters being thus arranged, it was found that the system of cylinders when placed in different parts of the closed portion of the pipe rotated with different velocities. A maximum was obtained at the middle of the loop, viz: at 24 and a minimum at either of the nodes, viz: at 0 and 48. The mean results of a set of experiments are given in the table below.

Position of cylinders in pipe.	Rates of revolution per second.
2	·77
7	3·53
12	6·43
17	7·70
21	8·74
24	9·00
26	8·80
29	8·47
32	8·12
37	7·00
42	4·69
47	·92

The above table is plotted in a curve shown in fig. 2. The abscissæ are distances along the pipe and the ordinates are rates of revolution per second. The full line is the observed curve; the broken line is a sine curve. The agreement of the two curves is sufficiently close to suggest that the rate of revolution at any point in the stationary wave is proportional to the velocity of the vibrating air particles at that point. The



curve rises slightly above the sine curve near the dividing membrane, viz: from 30 to 48.

So far as I am aware a rotation of this kind has not before been observed.

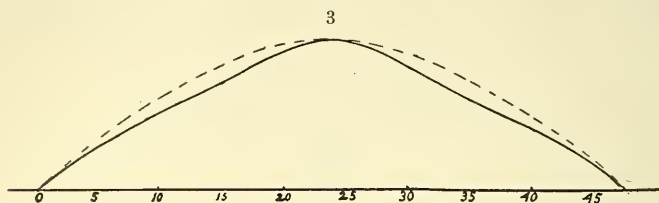
The force involved in the above reaction was measured with the aid of a double cylinder supported by a fine wire, which went up through the wall of the organ pipe inside of a glass tube, 20^{cm} long. This double cylinder was carried at the end of a torsion arm and each cylinder was .757^{cm} in diameter and 3.1^{cm} in length. The area of cross-section of each cylinder was .45 sq. cm. The pipe was placed in a horizontal position during the torsion experiments.

The amount of torsion necessary to bring the cylinders back to their original position was read on a graduated circle at the top of the glass tube.

The mean results of a number of experiments are given below.

Position of cylinders in pipe.	Torsion.
2	12°
7	72°
12	190°
17	355°
21	480°
23	555°
27	490°
32	315°
37	165°
41	72°
45	14°

The force exerted upon the cylinders increases very rapidly from the nodes to the middle of the loop. The table is plotted in a curve shown in fig. 3. The ordinates are the *square roots* of the plate readings. The curve corresponds not badly to a sine curve. Experiments with cylinders of the same length but of smaller diameters gave almost an identical curve.



Experiments made in various gases.—The body of the pipe between the diaphragm and the stopped end forms a closed chamber and into this chamber various gases were introduced. The pipe was still set into action by an ordinary *air* blast. The sound wave within the pipe was transmitted through two media, air and the gas, the diaphragm forming the division between them. The force of the air blast was kept constant throughout the experiments. The torsional measurements were made only at the middle of the loop. Although the number of vibrations executed by this composite mass of gas changed very much, the position of the node at the diaphragm was but slightly altered.

Several readings were taken in air, before and after each experiment with the gases. A fresh supply of gas was introduced into the pipe before each reading. The mean results of these experiments in different gases are given in the annexed table. The readings in air, as well as its density are taken as unity.

	Torsion deflections.	Density of gas.
Air	1.	1.
CO ₂	1.47	1.52
Illuminating gas....	.77	.75
Hydrogen064	.069

As will be seen, the force exerted is proportional to the density of the gas.

The rates of revolution of the system of cylinders mentioned above were also taken in the different gases and the mean result of several experiments are given below, the rate of revolution in air being taken as unity.

Air.	Illum. gas.	CO ₂ .	Hydrogen.
1.	1.	1.04	.59

The rate may be considered as approximately a constant for all the gases, although the force producing this rotation varies with the density of the gas. This can be explained by the circumstance that the resistance to motion is also proportional to the density. The deviation from this result in the case of hydrogen is explained by the fact that the friction on the supporting pivot was not an inconsiderable quantity compared to the total resistance experienced by the cylinders, and this factor is proportionately larger in the case of hydrogen.

Prof. Hallock first suggested to me the explanation of the effect described in this paper.* He pointed out that Daniel Bernoulli has shown that a gas in motion is virtually less dense than the same gas at rest. The air in the capsule cylinder is at rest, while that on the outside of it is in motion. The energy exerted is due to the difference in density on the two sides of the closed end of the cylinder.

An inspection of the curve in fig. 3 shows that the force is approximately proportional to the square of the velocity of the vibrating air.

I have applied this effect to the problem of determining the amplitude of the vibrating air particles in the middle of the loop of the stationary wave. It is evident from the formula given below, that if the force per unit of area on the closed end of the cylinder is obtained experimentally, the change of density may be found, and the amplitude of vibration calculated from it.

Prof. R. S. Woodward has kindly assisted me in applying the proper principles of hydrodynamics to the problem. The relation between the velocity, pressure and density of a moving fluid may be expressed by Bernoulli's equation, since, for this case the average state of the fluid only is under consideration.

$$\int \frac{dp}{\rho} = R - \frac{1}{2} u^2$$

R is the potential due to external forces, and is negligible in this case; hence

$$\int \frac{dp}{\rho} = - \frac{1}{2} u^2$$

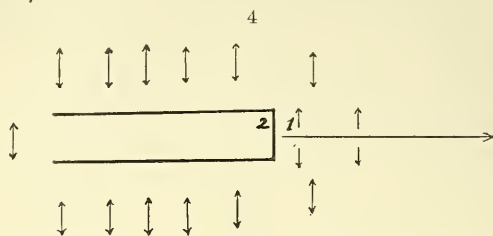
The adiabatic gas relation gives, $p = c\rho^r$, where c is the gas constant.

* This effect is somewhat similar in nature to the acoustic attractions and repulsions of Guthrie and the rotating mill of Dvořák that is, they all depend on the change in density arising from motions of a fluid. The effect herein described is of a different order of magnitude than that of the rotating resonators B D.

By substitution and integration I obtain,

$$\frac{1}{n} \frac{p_2}{\rho_2} \left[1 - \left(1 - \frac{p_2 - p_1}{p_2} \right)^n \right] = \frac{1}{2} (u_1^2 - u_2^2),$$

where $n = \frac{\gamma - 1}{\gamma}$.



The cylinder is in the moving air as shown in fig. 4; u_1 is the velocity of the vibrating air outside of the cylinder, u_2 is the velocity inside the cylinder near the closed end. p_2 and p_1 are the pressures on the two sides of the closed end of the cylinder. If the air in the cylinder next to the closed end is at rest, then u_2 becomes zero.

By expansion and reduction, the above equation becomes ;

$$u_1^2 = 2 \frac{p_2 - p_1}{\rho^2}$$

$p_2 - p_1$ is the quantity measured on the torsion balance. The absolute value of the pressure as determined from the experiments gave a force of 21 dynes on an area of one sq. cm. This gives a linear velocity to the vibrating air of 187^{cm} per second. The amplitude so obtained is 2.61^{mm}. This value for the amplitude is somewhat less than a corresponding value obtained by means of the sound-wave anemometer recently devised by myself, an account of which will be shortly published.

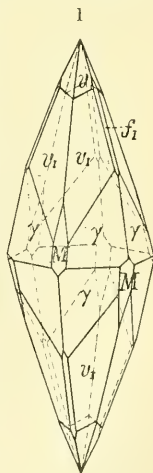
Physical Laboratory, Columbia University,
June 1st, 1900.

ART. XXIII.—On some Interesting Developments of Calcite Crystals; by S. L. PENFIELD and W. E. FORD.

1. CALCITE FROM UNION SPRINGS, CAYUGA COUNTY, N. Y.

THE crystals under consideration were found during the summer of 1899 by Mr. J. M. Clarke, Director of the Geological Survey of the State of New York, and were sent to New Haven for examination. Mr. Clarke had observed that the crystals presented certain features of unusual interest, and it was his wish that they should be described and that the specimens should be deposited in the Yale Collection. The crystals occur in the Onondaga limestone, in a region where slight tectonic disturbances have taken place, giving rise to fissures in which calcite has deposited as vein material. The most interesting feature presented by the crystals is their diversity of habit, shown often on a single hand specimen, and due to different methods of twinning together with peculiarities in the development of certain crystal faces.

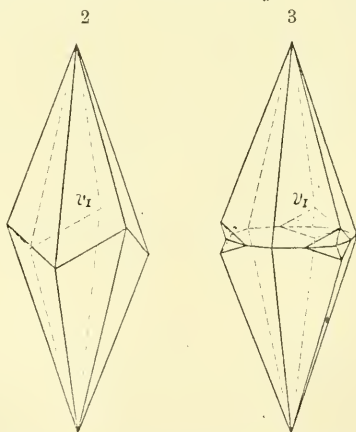
Most of the crystals were not well adapted for measurement with the reflecting goniometer, but, using one of the smaller ones, about 5^{mm} in length and 2^{mm} in diameter, it was possible to identify the prominent forms by means of their angles. The small crystals are quite highly modified and their development is represented by fig. 1. The terminal faces are the brightest and best developed, and are those of the common scalenohedron v , $21\bar{1}1$. There was measured for the identification of this form $r \wedge v$, $10\bar{1}1 \wedge 21\bar{1}1 = 28^\circ 56'$, calculated $29^\circ 1' 30''$. In the zone r , v , and making a very small angle with v is the scalenohedron v_1 , $7\cdot4\cdot\bar{1}\bar{1}\cdot3$, which is especially prominent on the crystals from this locality. This form was identified by von Bournon on crystals from Derbyshire and the Dauphiné Alps, and appears as form No. 37, plate 31, of his *Traité de Minéralogie*, published in 1808. The form was identified by its position in the zone r , v and the measurement $v \wedge v_1 = 3^\circ 23'$, calculated $3^\circ 55'$. On the crystals under consideration the faces of the scalenohedron v_1 have a vicinal development, and thus the contrast between them and the better developed faces of the scalenohedron v is generally quite marked. A negative rhombohedron, f_1 , truncates the edges of v_1 and appears always as a narrow face with vicinal development from which no reflection could be obtained. A rhombohedron in this



position would have the symbol $0 \cdot 12 \cdot \bar{1} \bar{2} \cdot 5$, and is a little steeper than the common rhombohedron f , $02\bar{2}1$, which truncates the pole edges of the scalenohedron v . The pyramid of the second order γ , $8 \cdot 8 \cdot \bar{1} \bar{6} \cdot 3$, was identified by the measurement $8 \cdot 8 \cdot \bar{1} \bar{6} \cdot 3 \wedge 8 \cdot 8 \cdot \bar{1} \bar{6} \cdot 3 = 25^\circ 40'$, calculated $24^\circ 46'$, and, further, by its being truncated by the positive rhombohedron M , $40\bar{1}1$. This rare pyramid was first identified by vom Rath* on crystals from Andreasberg in the Harz, and, as pointed out by the authors,† this same pyramid is the prevailing form of the siliceous calcites from the Bad Lands of South Dakota. On crystals from Union Springs there is a tendency for the upper and lower faces of the pyramid γ to round into one another, owing to vicinal development, and because of this rounding it was impossible to obtain an accurate measurement between the upper and lower γ faces.

On the majority of the specimens the crystals are not so highly modified as the one just described, but as already stated the variation in habit due to twinning and the unequal development of certain faces, gives to the specimens a peculiar interest. All the types to be described occur on a single specimen having a surface about half the size of one's hand covered with crystals. The crystals on this specimen were not suitable for measurement and therefore no angles will be given, but the forms were evidently like those identified on the small crystal previously described.

Scalenohedral type.—The scalenohedron v_1 , $7 \cdot 4 \cdot \bar{1} \bar{1} \cdot 3$, fig. 2, is apparently very common at the locality. It should be stated

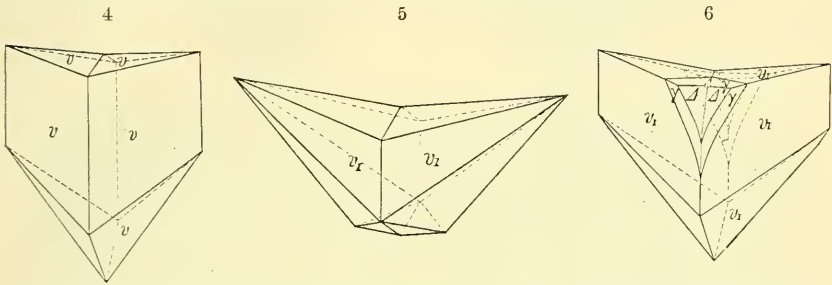


that this form has the same middle edges as the rhombohedron r , $10\bar{1}1$, and the common scalenohedron v , $21\bar{3}1$, but is somewhat steeper than the latter form. A twinning about the basal plane, fig. 3, is quite common.

* Pogg. Annalen, cxxxii, p. 521, 1867.

† This Journal, ix, p. 352, 1900.

Twins, with the rhombohedron e , $01\bar{1}2$, as twinning plane.—The habit resulting from this kind of twinning is like that of the well known Guanajuato calcites, fig. 4, described by Pirsson,* and it should also be stated that as early as 1837 Lévy† also described and figured calcite twins of this same type from Streifenberg, Nertschinsk, Siberia. Fig. 4 is analogous to the figures of Pirsson and Lévy, though drawn in a different position, and represents the common scalenohedron v , $21\bar{3}1$, drawn with the twinning plane vertical and having a position like that of the side face of a cube, or the pinacoid 010 of any of the three axial systems. This position has been adopted for representing the twin crystals as it gives the best idea of their peculiar development. Fig. 5 represents the scalenohedron v_1 ,



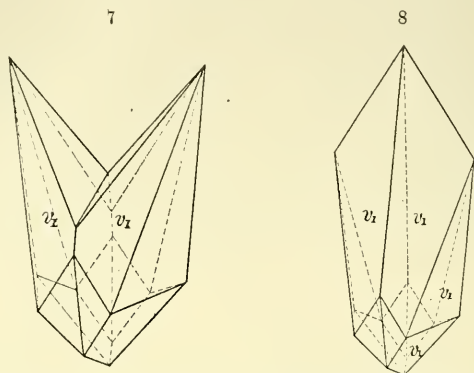
$7\cdot4\cdot\bar{1}\bar{1}\cdot3$ twinned without distortion, a type which has not been observed on any of the specimens, but the figure is introduced in order to show how, by the extension of the two lettered faces in front and the parallel faces behind, together with the suppression of the four small faces below, the Guanajuato type, fig. 4, results. Most of the Union Springs crystals of the Guanajuato type show in addition to the scalenohedron certain modifications at the reëntrant angle, fig. 6. The faces forming the reëntrant angle are the pyramid of the second order γ , $8\cdot8\cdot\bar{1}\bar{6}\cdot3$, and a rhombohedron designated as Δ , apparently in the zone with v_1 and γ , which would cause it to have the symbol $80\bar{5}3$. The surfaces forming the gash or reëntrant angle, however, are curved to such an extent that exact symbols cannot be assigned to portions of them.

Twins with the rhombohedron f , $02\bar{2}1$, as twinning plane.—The rhombohedron f is one of the rare twinning planes of calcite, and the habit presented by the crystals from Union Springs is very striking. The scalenohedron v_1 , $7\cdot4\cdot\bar{1}\bar{1}\cdot3$, twinned about f , and drawn with the twinning plane vertical, as previously

* This Journal III, xli, p 61, 1891.

† Description d'une Collection de Minéraux formée par H. Heuland, vol. i, p. 10, fig. 5, Plate 1.

described, is represented by fig. 7. In the Union Springs crystals representing this twinning law the reëntrant angle at the top wholly fails, and a peculiar, pointed, spear-head devel-



opment, fig. 8, results from the extension of the two front lettered faces of fig. 7 and the corresponding faces at the back. The crystals observed have always been attached at the lower end. Several crystals of this peculiar type were observed on the specimens sent by Mr. Clarke, and they are said to be quite common at the locality. On a crystal with a broken point the reëntrant angle measured from the rhombohedral cleavages was found to be $35^\circ 38'$, calculated $35^\circ 27'$.

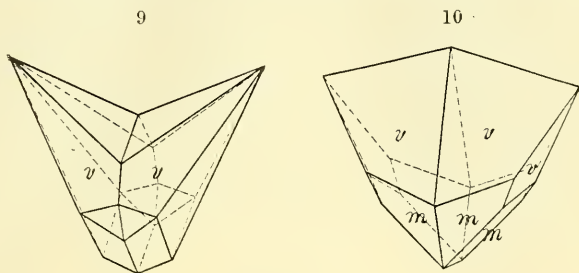
Lévy, in fig. 6 of the atlas to his work already cited, gives a representation of a crystal from Kongsberg in Norway, of identically the same habit as fig. 8 of this article; however, the habit is apparently a very unusual one, and it is interesting to record it at a new locality.

On a single specimen or even at one locality, as a rule, all crystals of a certain mineral have the same or nearly the same habit, resulting undoubtedly from crystallization under uniform conditions, and therefore it seems a matter of more than usual interest to note on a single hand specimen from the Union Springs locality, the occurrence of simple scalenohedrons, fig. 2, and of three distinct types of twinning, figs. 3, 6 and 8. The calcite crystals seem to be all of one generation. Associated with them are a few crystals of dolomite, apparently of later growth.

2. BUTTERFLY TWINS FROM EGREMONT, CUMBERLAND, ENGLAND.

The so-called butterfly twins from Egremont are well-known and are figured in many mineralogies. Lévy in his work, already cited, gives three figures of them, No. 17, 68, and 69 of his atlas. A few words concerning them and new figures

are introduced in the present article for the sake of comparison with the two types of rhombohedral twinning previously described. The twinning plane in these crystals is the rhombohedron r , $10\bar{1}1$, and the common scalenohedron v , $21\bar{3}1$, thus twinned and drawn, as in previous cases, with the twinning plane vertical, is represented by fig. 9. Fig. 10 represents a



crystal of the butterfly twin type in the Brush collection, and, by comparison with fig. 9, it may be seen that the upper faces of the crystal result from the development of the two front, lettered faces of fig. 9 and corresponding faces behind, to the complete obliteration of the reëntrant angle. The faces at the lower extremity of fig. 10 are those of the prism m , $10\bar{1}0$.

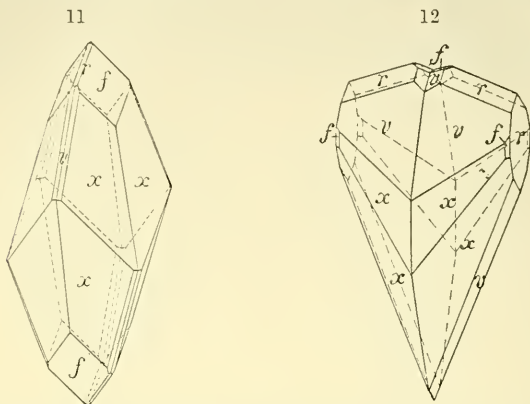
It is a matter of interest to observe how the scalenohedron when twinned as described according to the three rhombohedral laws gives apparently simpler shapes by distortion, or unequal development of some of its faces, than if the distortion had not taken place.

3. CRYSTALS FROM PALLAFLAT, CUMBERLAND, ENGLAND.

A feature of the crystals from this locality, as represented by specimens in the Brush collection, is the prominent development of the negative scalenohedron x , $13\bar{4}1$. This form, as shown by fig. 11, has its shorter pole edges bevelled by the common scalenohedron v , $21\bar{3}1$, and has the same middle edges as the negative rhombohedron, f , $02\bar{2}1$. Fig. 11 was drawn by Mr. W. Valentine of the Sheffield Laboratory. It presents nothing new, and is practically identical with fig. 674 of von Bournon's *Traité de Minéralogie* published in 1808. The figure is introduced in the present article, because by understanding its simple zonal relations, the same forms can be easily identified as they occur on a twin crystal to be described.

Fig. 12 represents the development of two beautiful twin crystals in the Brush collection, both occurring on the same hand specimen. The twinning plane is the unit rhombohedron, and the development is analogous to that of the butterfly

twins from Egremont, fig. 10. A prominent feature of the twins is the vertical zone v , f and x of the individual to the



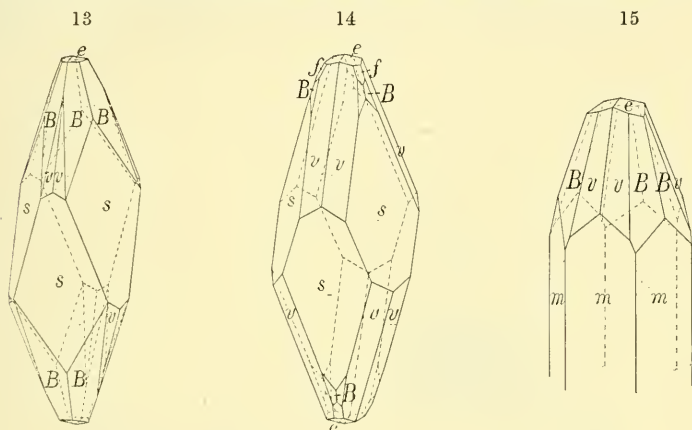
right, extending over the twinning plane to x , f , and v of the left hand individual, and so on around the crystal. Thus with this method of twinning four x faces, two in front and two behind, form as it were a vertical prism, analogous to the prism formed by four faces of the scalenohedron v , fig. 4, when the flat rhombohedron e , $01\bar{1}2$, is the twinning plane.

In figs. 4 and 12 the rhombohedral symmetry is not apparent, and the habit is like that of twin crystals of the monoclinic system, having the vertical faces v and x , respectively, as prisms and a pinacoid as twinning plane. The twin crystals represented by fig. 12 are so attached that only a portion of the lower x and v faces are visible.

4. CRYSTALS FROM THE STANK MINE, LANCASHIRE, ENGLAND.

This locality is represented in the Brush collection by a suite of eleven groups of crystals. A form which appears on all of the crystals, and which seems to be especially characteristic for the locality, is the negative scalenohedron B , $2\cdot8\bar{1}0\cdot3$, having the same *middle edges* as the rhombohedron f , $02\bar{2}1$. On all of the specimens the faces of the scalenohedron have a decided vicinal character. Some of the specimens are simply groups of scalenohedral points without modifications. Generally, however, as shown by fig. 13, the flat negative rhombohedron e , $01\bar{1}2$, is slightly developed, and long narrow faces of the positive scalenohedron v , $21\bar{3}1$, modify the pole edges, while some portions of the negative rhombohedron s , $05\bar{5}1$, may be seen near the attachment of the crystals. Another prominent type which seems to be common is represented by fig. 14. In

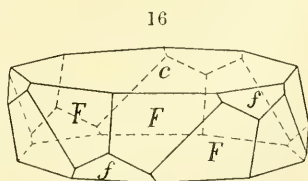
this type of crystal the rhombohedron s is prominent and is bevelled by the scalenohedron v ; the pole edges of v are truncated by the rhombohedron f , and B and e appear as slight modifications only. The type of crystal, however, which it is desired especially to call attention to is represented by fig. 15, there being several almost ideally developed crystals of this type in the collection. The peculiar feature presented by the crystals consists in the termination of the prism m , $10\bar{1}0$, by twelve faces of almost equal size and shape, thus giving the appearance of a prism terminated by a dihexagonal pyramid. Six of these terminal faces belong to the positive scalenohedron v , and the remaining six to the negative scalenohedron B , while the negative rhombohedron e appears at the end of the crystal. The distinction between the v and B faces is at



first glance not at all apparent, but the forms may be distinguished by their relative positions with reference to the rhombohedron e and by slight differences in their surface markings, the v faces being finely striated parallel to their intersection with the unit rhombohedron, while slightly undulating, smooth, vicinal surfaces characterize the form B .

5. CRYSTALS FROM ELEOLITE-SYENITE, MONTREAL, CANADA.

The crystals under consideration are small tabular individuals, fig. 16, measuring about 8^{mm} in diameter, and were found by Mr. W. F. Ferrier of Rossland, B. C., at the Corporation Quarry, Outremont near Montreal. Only a few crystals were found, and one of these was presented to the Brush collection by Mr. Ferrier. A tabular habit is not at all uncommon for calcite, but the crystals deserve brief men-

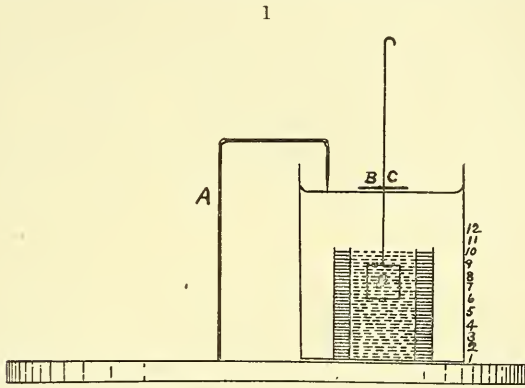


tion because of the peculiar combination of forms. Besides the basal planes only two forms are present, the negative rhombohedron f , $02\bar{2}1$, a very common form, and the positive scalenohedron F , $42\bar{6}1$, a rare form for the species. The scalenohedron F has the same parameter relation on the horizontal axis as the common scalenohedron v , $21\bar{3}1$, but is twice as steep, F intersecting the vertical axis at 6 and v at 3. Fig. 16 was drawn by Dr. H. E. Gregory while a student in the laboratory.

Sheffield Laboratory of Mineralogy and Petrography,
Yale University, New Haven, May, 1900.

ART. XXIV.—*A Method of Measuring Surface Tension*; by
 JAMES S. STEVENS.

IN the methods of measuring surface tension which have formerly been employed there seems to be two difficulties: the weights applied to break the tension cannot be added in continuous increments, and they afford a certain accelerating force beside their own mass when they are dropped on to a scale pan; and secondly, it is difficult to apply the pull exactly at the center of inertia of the body used to break through the surface skin. The second difficulty was overcome in a method described by T. P. Hall, where filaments of glass were detached from a liquid surface by weights applied to a scale pan which formed the opposite end of a balance.



A, hook gauge; BC, disk; D, sinker.

The method described here is free from difficulties of the above nature. An iron wire is bent into the form of a T (fig. 1) and the cross allowed to float on the surface of water. The body of the T projects a little distance into the center of a coil of copper wire through which a current may pass. The part of the T forming the plunger is adjusted so that its weight is nearly sufficient to break through the liquid surface. A slight pull towards the center of the coil due to the passage of a current will then accomplish this. The system is raised and lowered by a rack and pinion so that it may be placed on the liquid surface as gently as possible.

By means of a continuous resistance the current may be applied in small increments, and the amount required to break through read from an ammeter. When the current is applied the plunger immediately moves to the center of the coil and

the pull is then steadily applied to the center of inertia of the system.

The magnetic force may be calibrated at leisure by use of a balance or it may be calculated from the constants of the coil. This in proper units must be added to the weight of the system.

In some experiments already made by this method a circular disk of thin ferrotype plate was used and a rather heavy sinker of soft iron fastened to its center. The variation of surface tension with temperature in case of distilled water, the effect of adding a drop of oil, the surface tension of a soap solution and the results with disks of different areas have been tested.

In order to test the law of inverse squares a hook gauge was employed so that the height of the water surface, and hence the floating disk, could be adjusted very closely. By means of a scale fastened to the side of the vessel containing the liquid, the weight of the surface could be changed by successive centimeters. If we assume that the force required to break the tension is constant, and if we call the distance between the center of attraction of the coil and the magnetic center of the system attracted, when they are nearest together, x ; then $x+1$, $x+2$, etc., will represent other distances used. Let c' , c'' , etc., stand for currents corresponding to these distances and we have

$$\frac{c'}{x} = \frac{c''}{x+1} = \frac{c'''}{x+2}, \text{ etc.}$$

Substituting values and solving we have for a given experiment $x = 6^{\text{cm}}$.

Applying the values to other measurements we have 1.88, 1.81, and 1.95 as ratio numbers expressing the value of the same force, namely that required to break through the surface tension.

Very little work has been done with this apparatus and this paper is designed to describe the method rather than announce results.

Orono, Maine.

SCIENTIFIC INTELLIGENCE.

I. GEOLOGY AND MINERALOGY.

1. *The Glacial Gravels of Maine and their associated deposits*; by GEORGE H. STONE. U. S. Geol. Survey, Monograph XXXIV, pp. 1-499.—Ever since the early writings of Agassiz on the glacial geology of New England, the "horsebacks" and terraces of Maine have been generally known, but the great extent and remarkable perfection of the glacial deposits is a recent discovery. Prof. Stone's work, extending over a period of thirteen years, has resulted in a series of maps showing the distribution of the water-laid sediments of glacial time over the whole state. The deposits from each glacial stream have been mapped and grouped into some sixty gravel "systems." Each "system" is composed of trunk stream (usually an esker) with tributaries and distributaries (delta branches). Under separate heads a detailed description is given of composition, shape, etc., of each group. The gravel "systems" are usually 10-25 miles in length; the Houlton "system" is 115 miles long and the Katahdin stretches out for 125 miles. The longer "systems" extend for many miles like an artificial roadway carefully constructed, and with the exception of Sweden constitute the finest examples of gravel ridges in the world.

This reconstruction of the river systems of glacial time has been done with great labor, for there are no adequate maps of Maine and much of the region is a succession of swamps and forests. The fifteen maps accompanying the monogram show at once the difficult nature of the work and the remarkable development of the forms of drift in the state. The chapters on Classification and Genesis are valuable not so much because of the theoretical conclusions drawn as for the great mass of detailed description and discussion of the abundant material at hand. Much has been done in glacial geology since the manuscript for this monograph was completed (1889), but we know of no more important contribution to the general discussion of eskers. Besides the treatment of eskers space is given to descriptive and theoretical matter in regard to kames, coast gravels, valley drift and tills. A study of the till and the marine clays was made to determine whether there was an interglacial period in Maine as in the Mississippi Valley, but the evidence was found to point to a single glacial period.

H. E. G.

2. *The Illinois Glacial Lobe*; by FRANK LEVERETT. U. S. Geol. Survey, Monograph XXXVIII, pp. 1-797, with 18 maps and 15 illustrations.—This is the most elaborate volume on Glacial Geology yet published and renders accessible the mass of facts gathered during many years of research by a force of geologists under the direction of Prof. Chamberlain. The problem before Mr. Leverett was the investigation of the complicated phenomena

arising from the movements and deposits of the successive advances of the ice front in the Lake Michigan region. Not only detailed study but more refined methods of investigation were demanded, and the development of these methods will be useful in all future glacial work. The various drift sheets are described in detail under the following heads: 1. The Illinoian Drift Sheet—the most extensive in the region, which was preceded by an interval of deglaciation (Yarmouth) and followed by the Sangamon interval of deglaciation. 2. The Iowan Drift Sheet, in the discussion of which the whole loess problem is gone over and conflicting views compared, without however arriving at a definite general conclusion as to the origin of the loess. The Iowan deposits are overlaid in places by the “Peorian soil” of another interval of deglaciation. 3. Two Wisconsin Drift Sheets complete the series.

Perhaps the most generally interesting part of the monograph are chapters XI and XII, dealing with beaches of Lake Chicago and the Chicago outlet and with the many cases of reconstructed drainage throughout the drift-covered area. Each stream is treated separately and the past and present conditions of its drainage basin is discussed. In Chapter XIV all the reliable well records of Illinois are given and conclusions drawn as to character of soil, relation of drift to water supply, etc. From the well data and from comparison of the drift areas to adjoining driftless areas the average thickness of the Illinois drift sheet is given as 100 to 130 feet.

H. E. G.

3. *The Geography of the Region about Devils Lake and the Dalles of the Wisconsin with some notes on its Surface Geology*; by ROLLIN D. SALISBURY and WALLACE W. ATTWOOD. Wisconsin Geol. and Nat. Hist. Survey, Bulletin No. 5, pp. 1-146. 85 plates and illustrations.—This bulletin of the Wisconsin Survey begins an educational series designed primarily for use in the schools. It marks the beginning of a plan to make the work of the geological surveys directly useful to all intelligent people of a state instead of to the few specialists. The volume was evidently prepared by skillful teachers, for the region chosen for description is clearly marked and typical and interesting and the plan of presentation is that of the most successful instructors. Part I is a description of the topography of the area—the quartzite ridges and the sandstone plains. Part II deals with the history of the topography. The origin of the rocks which show themselves at the surface is explained and their scenic features which result from erosion are described. The glacial period and its effect on topographic form is treated of easily. In connection with each division of the subject there is given just enough theoretical matter to put the abundance of local facts in proper perspective. The volume contains no geological facts regarding this interesting region which have not been published before, but it is timely and valuable as a plan for presenting a geographical subject to schools and as a suggestion to state surveys.

H. E. G.

4. *The Moraines of Southeastern South Dakota and their Attendant Deposits*; by JAMES E. TODD. U. S. Geol. Survey, Bulletin 158, pp. 1-168.—This report is a description of the Pleistocene deposits of the Dakota-Nebraska region. The deposits are described as a first or outer moraine which extends down to the Missouri River, a second moraine bordering a region farther north, and various forms of drift outside the moraines. The drift region beyond the outer moraine is largely loess and Prof. Todd describes at some length the distribution character of these interesting deposits. The Missouri River is believed by the author to have been largely excavated since the ice age and, furthermore, that the river is still deepening its trough.

H. E. G.

5. *Preliminary Report on the Copper-bearing Rocks of Douglas County, Wis.*; by U. S. GRANT. Wisconsin Geol. and Nat. Hist. Surv. Econ. Series, No. 3, Bulletin VI. Pp. 55, maps and plates. Madison, 1900.—This work has for its purpose a study of the copper-bearing rocks of the Keweenaw series in the region southeast of Duluth and Superior, the so-called Douglas Copper Range, to ascertain the character and value of the copper deposits. The rocks and the occurrence of the ores are quite similar to what is found in the well-known region of Keweenaw Point. The promise of valuable deposits is not, however, very great, the copper being scattered so much as to make in general very lean ores. While the value of the work is chiefly local, it will also prove of interest to all who are interested in the origin and occurrence of ore deposits.

L. V. P.

6. *Bulletins of the Geological Survey of Finland, No. 9. Esquisse hypsométrique de la Finlande*; by J. J. SEDERHOLM. Pp. 17, 8vo. Helsingfors, 1899.—This is a short description of a relief map of Finland on a scale of 1 : 2,000,000, in which the main features of the topographic relief are brought out and described.

No. 10. *Les depots Quaternaires en Finlande*, by J. J. SEDERHOLM (pp. 28, 1899). This little work contains a brief but very good description of the glacial deposits of Finland and is accompanied by an excellent map on the same scale as the preceding one. The differences of relative level of land and sea at different times, the movement of the ice and the remarkable and celebrated osar, colossal glacial embankments crossing the country in long lines and well shown on the map, are described and discussed.

No. 11. *Neue Mitteilungen über das Ijolitmassiv in Kuusamo*, by V. HACKMANN (pp. 45, 8vo, 1900). Ijolite is a name given by Ramsay and Berghell to the rock forming a mountain mass composed of pyroxene and nephelite and of granular structure. Recently Hackmann has studied this occurrence in the field and gives a full description, showing it to be a complex of alkaline rocks ranging from ijolite through essexite into pyroxene-syenite. The ijolite is the main type. Several new analyses are given and the whole is an interesting contribution to the petrography of the alkaline rocks.

L. V. P.

7. *Analyses of Rocks, Laboratory of the U. S. Geological Survey, 1880 to 1899*; F. W. CLARKE. Bull. 168, U. S. Geol. Surv., 1900, pp. 308. This is a new edition of Bulletin 148 of the Survey publications, containing a large number of new analyses, chiefly executed by W. F. HILLEBRAND and H. N. STOKES. All told, some 1,400 analyses of rocks, clay, soils, etc., are given. The order of the analyses has been changed, the more important constituents being given first, as is usual with many petrographers. This makes the analyses much more convenient for use. As before, the arrangement is geographical. The former chapter on rock analysis by Hillebrand is omitted and it is stated that this, in an enlarged form, will soon appear as a separate bulletin. One cannot but be impressed by the enormous amount of painstaking, careful and skillful work represented by these analyses, so complete and so detailed. From no other source could such a mass of similar analyses be gathered, and they doubtless represent the most important work done in American chemical geology.

L. V. P.

8. *The Bjurböle, Finland, Meteorite*.—Portions of this meteorite are now on exhibition in the Finland pavilion at the Paris Exposition. The fall took place March 12, 1899. The mass, which had a total weight of about 340 kilos, fell into the Gulf of Finland, making a hole about nine yards in diameter in the ice (which had a thickness of about 28 inches) and penetrating the submarine clay to a depth of six yards. The removal of the mass has been a matter of considerable difficulty and was accomplished only in sections. The portions on exhibition at Paris are the largest secured and weigh 83 and 22 kilos respectively. The interior of the stone is gray with yellowish spots. The exterior is covered with a black crust. Prof. Stanislas Meunier gives in a recent issue of "Le Naturaliste" the results of a detailed study of the meteorite. He describes it as friable in texture and in structure of the variety to which he applies the term *öolitic*. The mineralogical composition is largely chrysolite and pyroxene, with which are associated grains of nickel-iron, iron sulphide and chromite. Sp. gr. = 3.54. Chondrules varying in diameter from .5 to 2.5^{mm} are easily separated. They are of several types, which are fully figured. The most remarkable are those of a fibrous, probably feldspathic, mineral often enclosed in a metallic shell, and those containing porphyritic crystals of augite. The meteorite is classed by Meunier as belonging to his Type 38, Montrejtite. The stones of Hesse, Aussen, Prairie Dog Creek and Searsmont will be remembered to be other examples of this type.

O. C. F.

OBITUARY.

Professor G. H. F. ULRICH, F.G.S., Director of the School of Mines of Otago University, Dunedin, New Zealand, died on May 26, at the age of seventy years. By his active investigations, first in Victoria and later in New Zealand, he contributed much to the knowledge of the Mineralogy and Geology of those countries.

THE

AMERICAN JOURNAL OF SCIENCE

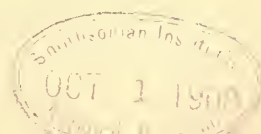
[FOURTH SERIES.]

ART. XXV.—*Notes on the Colorado Canyon District*; by
W. M. DAVIS.

A VISIT to the district of the Colorado Canyon leaves the traveler with a deepened impression of his indebtedness to the three explorers from whom an understanding of that marvelous region has come. Newberry, crossing the plateaus south of the canyon in 1858, recognized the deep erosion of the fundamental crystallines before the deposition of the Paleozoic series, and the extensive denudation of the plateau uplands where the retreating escarpments succeed each other in order of age; he made explicit statement that not only the canyon but the escarpments of the plateaus "belong to a vast system of erosion, and are wholly due to the action of water."* Powell's adventurous expedition down the river in 1869, justly classed with the most daring explorations of the continent, added an account of the division of the plateaus into huge blocks by faults and monoclines trending about north and south, in contrast to the cliffs of erosion which in a general way trend east and west; he clearly stated that "the cliffs of erosion are very irregular in direction, but somewhat constant in vertical outline; and the cliffs of displacement are somewhat regular in direction, but very inconstant in vertical outline."† Dutton's more elaborate surveys of the canyon and of the plateaus to the north of it in 1879 and 1880, aided by Holmes' wonderful drawings, have made the canyon district classic ground, the type of all that is gigantic in displacements and denudation, the region to be cited, at home and abroad, as

* Ives Expedition: Report upon the Colorado River of the West, Washington, 1861. Part III, Geological Report, 45.

† Exploration of the Colorado River of the West, Washington, 1875, 191.



the standard example, best studied, best described, and best illustrated, of river work on a grand scale.* It should be added that the topographical maps made by Bodfish and Renshawe in 1879, although of small scale and large contour interval, are of great value in giving the location and altitude of the more important features.

The wish that I have had for many years to visit the canyon was fortunately brought to fulfilment in June. Our party consisted of Prof. R. E. Dodge of Columbia University, Dr. H. E. Gregory of Yale University, Mr. R. L. Barrett of Chicago, Mr. Richard Wetherill of Pueblo Bonito, New Mex., Dr. Tempest Anderson of York, England, and the writer. Great assistance in plans and outfit was given by Mr. F. E. Hyde, Jr., of New York, who was unhappily unable to go with us. We left the Sante Fe Pacific Railroad at Flagstaff, Arizona, on June 4, and followed a very irregular route northward, partly with wagons, partly with horses and pack train, and on June 26 reached Milford, Utah, whence the Oregon Short Line carried us to Salt Lake City. Postponing a fuller account of our observations to a later date, I desire to make brief statement here of certain points on which other conclusions than those announced in the Survey reports seem to be admissible.

The unconformities in the Canyon.—The Kaibab section of the canyon discloses the nearly even floor on which the horizontal Paleozoic strata rest. The floor is of complex structure. The fundamental schists with granitic dikes, which are exposed in the western Kaibab section, are overlaid in the eastern section by the heavy Unkar and Chuar series, dipping eastward and measuring according to Walcott about two miles in thickness.† The wedge in which the tilted formations terminate westward is a most remarkable geological structure, alike for its distinctness and for its significance. It is easily recognized near the bottom of the northern cliffs of the canyon when viewed from the southern rim at Hance's or at Cameron and Berry's; these primitive stopping places being at the ends of the road from Flagstaff, which forks a few miles before reaching the rim. The wedge is still better seen by descending the Grand View trail from Cameron and Berry's to a promontory of the Red Wall limestone north of the "copper mine." The crystalline floor of the Unkar series is of extraordinary evenness; no inequalities in it were detected in the view from the Red Wall promontory, except a small pre-Paleozoic fault (see

* Tertiary History of the Grand Canyon District, with Atlas, U. S. Geol. Survey, Monogr. II, Washington, 1882.

† 14th Ann. Rept., U. S. G. S., 1894, 508-512.

diagram by Gilbert, fig. 52, in Walcott's article, above referred to). Close to the base of the Unkar is a sheet of basalt, described as a contemporaneous flow by Walcott (l. c., 516-518), but to our view seeming rather to be an intrusive sill, inasmuch as it broke from one layer to another, instead of lying conformably between the under and overlying beds.

The floor of the Paleozoic series is somewhat less regular than that of the Unkar. A few miles down the canyon there is a mound of schists that rises rather higher than the basal Tonto sandstone, and the harder layers of the lower Unkar also interrupt the Tonto for a mile or more up stream from the apex of the wedge on the northern side of the canyon. It is notable that the deformed schists beneath the Unkar series descend by a continuous and very steep slope to the river; hence their whole mass must be about equally resistant; but west of the apex of the wedge and for a number of miles down

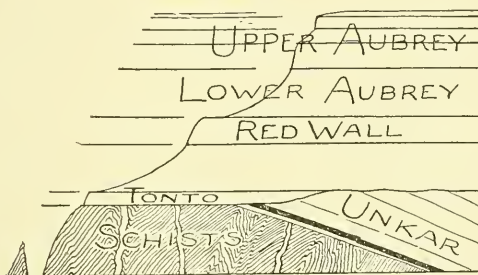


FIG. 1. Diagram to illustrate the section and profile of the canyon in the Kaibab

stream, the schists show a distinct bench just beneath the Tonto before steepening to their precipitous lower slope; hence here the upper part of the schists must be weaker than the lower. It occurred to me that this weakness may be ascribed to a pre-Tonto subærial weathering of the schists. If so, the gently undulating sub-Tonto floor may be, with renewed confidence, regarded as an ancient lowland of subærial denudation, as described by Powell, only slightly modified by marine erosion previous to Tonto deposition; while the sub-Unkar floor may have been planed down by the sea to smooth and firm rock from some unknown but probably moderate measure of relief, before the Unkar series was laid down on it.

The erosion of the plateaus.—The retreating cliffs of Permian, Triassic and other formations, on the north and east of the Grand Canyon district, have been taken to prove extensive erosion over the plateaus in which the canyon has been cut. Dutton concludes that this erosion was accomplished during a

lower stand of the region; in other words, that the plateaus are parts of a vast peneplain produced in a cycle of erosion anterior to that in which the canyon has been eroded, the two cycles being separated by an uplift of several thousand feet. It has been pointed out by other observers (although I believe that no statement of their opinion has been published) that the supposed peneplain surface coincides roughly with the resistant Upper Carboniferous strata over large areas, and hence that the plateaus might be explained as a surface of halting or at least of hesitating erosion in a single cycle, after the stripping of the weaker overlying formations. It is not easy to make definite choice between these two explanations, but I believe Dutton's to be the more probable of the two, for such reasons as the following. There is a mesa in the valley of the Little Colorado, southeast of the road from Flagstaff to Tuba, capped by a horizontal sheet of lava which rests unconformably on a bevelled surface of tilted Permian shales; but the shales are elsewhere denuded to a lower level. Again, there are numerous and extensive landslides at the base of the Echo and Vermilion (Triassic) cliffs south and west of Lee's Ferry; the slides seem to indicate a revival of sapping after a long pause. Further, the considerable length of many streams that flow against the dip of the strata suggests the need of the most favorable conditions for their development, such as would be associated with two cycles of erosion rather than with only one. Finally, the strong contrast between the gigantic young ravines of the steep canyon walls and the shallow mature valleys of the plateaus seems beyond explanation as dependent on the resistance of the Upper Carboniferous strata during a single cycle of erosion. But much more work should be done before this question can be definitely settled.

The Esplanade.—The canyon, where traversing the Uinkaret plateau and the greater part of the adjacent Kanab, consists of a broad and flat-floored upper valley, beneath which a deep and narrow chasm has been cut. Dutton terms these two parts the upper (or outer) and the inner canyons, and gives the name, esplanade, to the well developed floor of the upper canyon. He explains the esplanade as the result of valley-broadening during a pause in the general uplift in consequence of which the plateaus have been dissected: in other words, after a first cycle in which the plateaus were broadly denuded, a second cycle was introduced by general uplift of the region; but this uplift was interrupted by a pause long enough for the erosion of the upper canyon (l. c., 121). Other observers have suggested that the esplanade is a structural bench determined by the resistance of the Red Wall limestone, from which the overlying weak lower Aubrey sandstones and shales have been stripped. Cer-

tainly the surface of the esplanade, as seen in the magnificent prospect from Vulcan's Throne, where the canyon cuts the Toroweap valley, coincides closely with the top of the Red Wall series. Further up the canyon, the topographical maps show the esplanade in Kanab canyon on the north and in Cataract canyon on the south, but its distinctness lessens eastward and it disappears before entering the Kaibab. Where we saw the canyon in the Kaibab, its cross section was altogether different from that where the esplanade is well developed in the western Kanab. In the Kaibab there are two structural benches, as shown in fig. 1: one is determined by the Red Wall limestone, the other by the Tonto sandstone; and of the two the latter is rather more pronounced than the former. The change in form from the Uinkaret to the Kaibab seems to result from a weakening of the strata that are included between the firm limestones of the Red Wall and the firm sandstones of the Tonto. This change in structure is fully recognized by Dutton, and he explains the difference of form between the Uinkaret and the Kaibab sections by the greater altitude of the Kaibab during the erosion of the canyon (l. c., 257, 258). He does not, however, explain the absence in the Kaibab section of all signs of the pause during the general uplift by which he explained the esplanade further west. To my view, the esplanade, like the Kaibab benches, seemed to be entirely of structural origin; it does not seem to be necessary to postulate a pause during uplift in order to account for it.

The Toroweap is a broad valley that has been eroded along the fault that divides the Kanab and the Uinkaret plateaus, north of the canyon. A similar valley, which I have designated in my notes as the South Toroweap, is seen on the southern side of the canyon. Both of these valleys are flat-floored and open on the canyon close to the level of the esplanade. Dutton concluded that the Toroweap is the broadened valley of an ancient river which disappeared by reason of change from humid to arid climate before the inner canyon was cut. He recognizes that the valley has been somewhat shallowed by lava flows from the volcanoes of the Uinkaret on the west, but does not attach much importance to this fact (l. c., 92, 99). It does not seem necessary to adopt Dutton's conclusion as to climatic change, for the lavas in the floor of the Toroweap appeared to be in sufficient volume to account for its shallowness in contrast to the depth of Kanab and Cataract canyons a little further east.

Date of faulting.—The displacements by faults and monoclines, resulting in the division of the plateau region into a number of huge blocks, Shiwits, Uinkaret, Kanab, Kaibab, etc., have been variously dated by different observers. Powell

states that when the great denudation began, there were no faults and no benches: the first uplift was broad and even, without differential displacements; the displacements began later, and so slowly as not to affect the course of the original streams (l. c., 200, 201). Dutton dates the Hurricane fault as later than the earlier lava flows and earlier than the later ones (hence later than the great denudation and before the later canyon cutting); the Toroweap fault was not begun until the upper canyon had attained very nearly its present form, so that the floor of the esplanade as a topographic feature was dislocated by the fault; the Kaibab began to have a distinct existence only after the great Miocene denudation, when the Colorado was beginning the erosion of the upper canyon; the Echo cliffs monocline is rated as very nearly coeval with the East Kaibab monocline, and the latter began with the Pliocene (l. c., 117, 94, 192, 205). Walcott places the uplift of the Kaibab before the completion of the great (Miocene) denudation,* while there were still some heavy overlying strata upon it, because the layers of the monoclines are flexed without manifest fractures. The dating of the displacements therefore still offers an interesting subject for discussion.

While subordinate movements on the fault lines and monoclines may have occurred during the erosion of the inner canyon, the chief movements seem to be much older. This conclusion is enforced by several observations. First, there is the arrangement of the Permian and Triassic escarpments with respect to the northern extension of the fault lines: had the faults been of recent date, later than the great denudation by which the escarpments were pushed back for scores of miles from the original extension of their strata, the borders of the several formations as shown on a map ought not to be significantly out of line on the two sides of a displacement; for in the post-displacement period there would not have been time for much retreat of the cliffs on one side of a fault in excess of the retreat on the other side. But as a matter of fact, the Permian and Triassic cliff-makers have been greatly denuded with respect to and subsequent to the displacement of the blocked plateaus. The Permian cliffs retreat northward as they approach the Hurricane fault south of Toquerville; the Permian and the Triassic cliffs are out of line by several miles on the two sides of the fault that passes near Pipe Spring, west of Kanab. This style of arrangement was recognized by Powell, who stated that the higher blocks have been more eroded than the lower ones, and hence "the cliffs of the higher blocks stand further back from the axis of upheaval than those of the

* Bull. Geol. Soc. Amer., i, 1890, 60, 64.

lower blocks" (l. c., 191). Secondly, the Vermilion cliffs of the Trias make a great detour around the northern end of the Kaibab arch, such as could not be accounted for if the uplift of the Kaibab were as recent as the erosion of the canyon. Thirdly, the Trias margin along the Echo cliffs is in close accordance with the monocline of the same name; this proves that time enough must have elapsed since the bending of the monocline for the Triassic sandstones to retreat from whatever irregular front they may have had while they still lay horizontal, to their present well-defined structural alignment. The time that sufficed for the erosion of the canyon by the active Colorado does not seem long enough for the recession of the Triassic cliffs from an irregular to a regular front under the attack of dry-climate weather alone. Fourthly, the relatively small retreat of the Aubrey cliffs from the fault lines, as along the Hurricane and the Toroweap faults, may be explained by the relatively recent uncovering of these structures, long ago faulted.

The origin of the drainage system.—Both Powell and Dutton attribute an antecedent origin to the larger members of the drainage system in the plateau region. It is possible to conceive of the conditions necessary for the maintenance of such a drainage system through all the movements of upheaval and displacement that the plateau region has suffered; but it seems to me doubtful if these necessary conditions have been actually provided; and it is very difficult to find decisive proofs of an antecedent origin. During the great denudation by which thousands of feet of Mesozoic strata were stripped from the plateaus, there must have been opportunity for many spontaneous rearrangements of river courses. This opportunity would have been increased if there had been several successive cycles or partial cycles of denudation, separated by movements of elevation, as is entirely possible: the opportunity would have been still further increased if, as is eminently probable, displacements with slight tilting occurred during the denudation, for, as Hayes has shown, the greater rearrangements of drainage are dependent rather on gentle tilting movements of elevation than on inequality of rock resistance. The denudation of the region is usually dated as beginning in early Eocene time, because it is postulated that late Eocene strata once stretched all over the canyon district; but if a true-scale section be drawn, it does not appear improbable that the cañon district may have been slightly uplifted so as to undergo some denudation while the Eocene deposits were accumulating in basins on the north and east. The whole stretch of Tertiary time, with repeated uplifts and displacements, would

certainly offer abundant opportunity for the development of new stream courses, even if the drainage of the region had been chiefly antecedent at the beginning. Moreover, some of the valleys actually follow the strike of weak monoclinical structures, and are therefore to be explained as of subsequent rather than as of antecedent or consequent origin; other valleys slope against the dip of the strata, and may come to be classed as obsequent instead of antecedent. Altogether, this problem is so complicated that its solution seems out of reach at present. The main trunk of the Colorado may be antecedent, but such an origin seems improbable for the side streams.

Climatic changes.—The change from a moist Miocene climate to an arid Pliocene climate, already referred to in connection with the Toroweap valley, does not seem to be essential to the development of existing forms. House Rock valley, for example, is explained by Dutton as having been eroded by a member of the original antecedent drainage system, whose waters dried up at the time of the change from the moist Miocene to the dry Pliocene period; but may this valley not be equally well explained by headward erosion under a persistently dry climate along the weak Permian monoclinical strata that lie between the resistant Carboniferous strata of the Kaibab arch on the west and the heavy Triassic sandstones of the Paria plateau on the east? The valley seems to belong to the class of subsequent valleys above referred to; the depth of its floor having been determined by the sill of resistant Carboniferous strata in the platform of the Marble canyon on the southeast. The valley is not yet significantly deepened in response to the uplift which has permitted the erosion of the main canyon by the full-bodied Colorado.

The ravines which dissect the flanks of the Kaibab and the uplands of the Paria plateaus are ascribed by Dutton to the increased rainfall of the Glacial period, by which the aridity of Pliocene time was interrupted (l. c., 196, 202, 228); but in view of the active although intermittent erosion and transportation now going on in at least some of the flanking ravines of the Kaibab, it seems inadvisable to limit the origin of ravines of this class to any special period of time. Effective work is done on them to-day when thunderstorm torrents flush their channels, and similar work has probably been done ever since the uplands were stripped of the weak overlying Permian strata and exposed to erosion.

It may be added that the longitudinal summit valley of the Kaibab, classed by Dutton with House Rock and Toroweap valleys as the work of Miocene streams that became extinct with the coming of a dry Pliocene climate (l. c., 193-195),

seems explicable as a result of underground drainage. Sink holes of various sizes are found on the Kaibab uplands, and Dutton reports that large springs issue from the walls of the Kaibab canyon (l. c., 138, 196).

It was impressed upon me throughout our excursion over the plateaus that a multitude of problems connected with the history of the canyon district await solution. The earlier explorers gathered a great harvest there, and presented their results so vividly that newcomers may enter the region as if already acquainted with it. But the gleanings that the earlier explorers left are abundant; local structural studies as well as such considerations as are involved in dates of faulting, origin of stream courses, cycles of erosion and past climatic conditions, offer plentiful material to investigators for years to come.

Harvard University, July, 1900.

ART. XXVI.—*On the Determination of Minerals in thin Rock-sections by their maximum Birefringence*; by L. V. PIRSSON and H. H. ROBINSON.

IN the determination of minerals in thin sections of rocks, the common species possess certain diagnostic characters by which they are as a rule readily recognized. In doubly refracting minerals the interference colors or the birefringence between crossed nicols is probably the most generally useful character for diagnosis and this combined with the refractive index is sufficient in many cases to determine the mineral at a glance, when taken in connection with its color, cleavage, etc.

In the case of the less common or rare minerals, the approximate determination of the maximum birefringence becomes in many cases of the highest importance and in combination with other characters may be sufficient to definitely determine the species.

It is thus a matter of the greatest value to the beginner in microscopical petrography that he should learn the values of the maximum birefringence exhibited by the common minerals and associate these in his mind with the colors exhibited between crossed nicols. To do this he should have a simple and relatively accurate method. The more experienced petrographer has of course learned these relations for the common minerals by practice, though generally without any definite method.

In the case of birefringent minerals the strength of the interference colors shown by sections between crossed nicols depends on three conditions: (1) on the orientation of the section with regard to the ellipsoid of elasticity, (2) the thickness of the plate and (3) the difference in values between the axes of greatest and of least elasticity.

If the axial ellipsoid be cut perpendicular to an optic axis, and this in uniaxial crystals is of course a basal section, the section of the ellipsoid has the form of a circle. Hence for parallel light the elasticities are alike in all directions, the plate is isotropic and shows no interference-colors between crossed nicols. As soon, however, as the ellipsoidal cross-section is inclined to this position it takes an elliptical form, the light elasticities are not alike in all directions and hence interference colors are shown between crossed nicols. As the cross-section is inclined more and more the eccentricity of its elliptical form continually increases and with it the strength of the interference colors until the maximum of both is reached; in uniaxial crystals this is in the direction of the prismatic zone parallel to

the vertical axis and in biaxial crystals in a direction parallel to the plane of the axes of greatest and least elasticities, that is to say the plane of the optic axes. In these cases the difference between the indices of the axes of greatest and of least elasticity is the maximum birefringence, since it indicates the maximum eccentricity that can be obtained in any elliptical section of the ellipsoid of elasticity; for this, therefore, the strongest interference colors are shown. We have prepared a table of the values of maximum birefringence which is given on page 264. These values, which have been carefully collated from the best and most recent determinations, are given in descending order and the list embraces the great majority of rock-forming minerals.

The recognition of sections of a mineral in a rock slide which exhibit the maximum birefringence is somewhat inadequately treated in most handbooks, it being generally assumed that the student will form his own method from the optical data given. It involves the determination of the "order" of the interference color displayed. It rarely happens that the bounding walls of a mineral section in the slide are exactly perpendicular to the plane of the object and cover glasses; on the contrary they are apt to be inclined or in many cases have the form of a thin wedge. From the thin edge of such a wedge, especially by the aid of a high power and where the colors are not interfered with by the overlapping of some other mineral, as at the edge of the section, or near a crack or hole, or against some isotropic mineral or one cut nearly normal to an optic axis, it is often possible to count the successive colors in ascending order, finally arriving at the top and noting the color given by the full thickness of the section. This is essentially the same process as moving up the thickening quartz wedge and noting the colors it yields. If the ascending slope to the top of the section is extremely steep the stronger colors of the first order may be so crowded as to appear almost like a dark line. If by any chance this process cannot be used, then the section must be brought to extinction by the quartz wedge, the section then removed and the color in the wedge noted and its order ascertained by slowly withdrawing the wedge and noting the successive orders of colors down to zero.

The colors also vary with the thickness of the section. Good sections as made to-day should not vary sensibly in thickness in different portions, that is the two surfaces should be parallel. They should also average about 0.03^{mm} or 0.001 inch in thickness, or the maximum color exhibited by quartz should not be above clear white or at most pale straw-yellow.

With this thickness for the slide the majority of rock-making minerals show maximum colors in the first and second orders, colors easily remembered and determined.

In a rock slide containing a number of sections of the same mineral, the ellipsoid of elasticity will be cut in various directions and some of them must approach the maximum of birefringence, the probability increasing with the number of the sections. While it will rarely happen that one is thus oriented exactly parallel to the axial plane or optic axis, sensible divergences may exist without perceptibly altering the birefringence or vitiating the method to be presently described. In a number of haphazard sections of a mineral then, the one showing the maximum of birefringence may be assumed to be nearly parallel to the optic plane or axis thus giving approximately the highest birefringence of the mineral. This determination may be confirmed by the fact that such sections should not yield any loci of optic axes or axial bars or bisectrices in convergent light. The color in plain light, the pleochroism with one nicol, the cleavage and crystal form often help, in addition, to confirm the orientation.

If now we consider only those sections of minerals which give the maximum of birefringence, we may eliminate the factor of the orientation of the section in considering the strength of the interference colors and this will then depend on the other two factors, the thickness of the plate and the value of birefringence. There is thus a constant relation between them and if we know the thickness of the plate and the value of the maximum birefringence, we may readily calculate the interference color; or, given the thickness and the color, the maximum birefringence may be found; or, finally, with this and the color, the thickness of the plate.

These relations have been put into graphic form by Michel-Lévy in the beautiful colored plate in "*Les Minéraux des roches*" by Michel-Lévy and Lacroix. We have not found them expressed in graphic form in any other handbook with which we are acquainted.* The advantage of having these relations shown in a diagram where they can be readily seen, followed and used is very great. It is probable that the uncertainty, difficulty and cost of the colored plate may also have deterred some authors from using it. It is, however, by no means necessary that the plate should be colored, the diagram is the essential part and this we have thought would be of service to petrographers and students in this country, many of whom are unacquainted with the colored plate of Michel-Lévy, or are deterred from using it by its cost and inconvenient form. We have put it into compact shape and have simplified the color-names, by indicating the place of the distinct and well-

* Such as Rosenbusch, Rosenbusch-Iddings, Zirkel, Luquer or Harker. The last named, however, gives some very useful practical hints in this direction.

known colors rather than the transition tones. If anything is sacrificed to accuracy by this method, it is more than compensated for by simplicity and convenience in use.

In using this diagram (p. 265) for the determination of an unknown mineral, the thickness of the rock-section is first determined. This is done by taking some well-known mineral, such as quartz, for instance, which furnishes a number of sections in the slide and observing the highest color given by any of them. The maximum birefringence for quartz is 0.009 as seen in the table and by following this line down toward the left-hand lower corner until it intersects the vertical line of the highest interference-color observed, and by then following out to the edge on a horizontal line the thickness of the section can be told. The numbers represent hundredths of millimeters.

For the determination of an unknown mineral the highest color given in numerous sections is observed and the thickness having been determined as above, by means of the diagonals the numerical value is noted which corresponds to the given color in a section of the determined thickness. The maximum birefringence of the unknown mineral having been determined the table of birefringences is referred to and the mineral usually found to be one of a group of several having approximately the same values. Which particular one it may be, is in the vast majority of cases readily told by its comparative refractive index in plain light, cleavage, color and other optical properties.

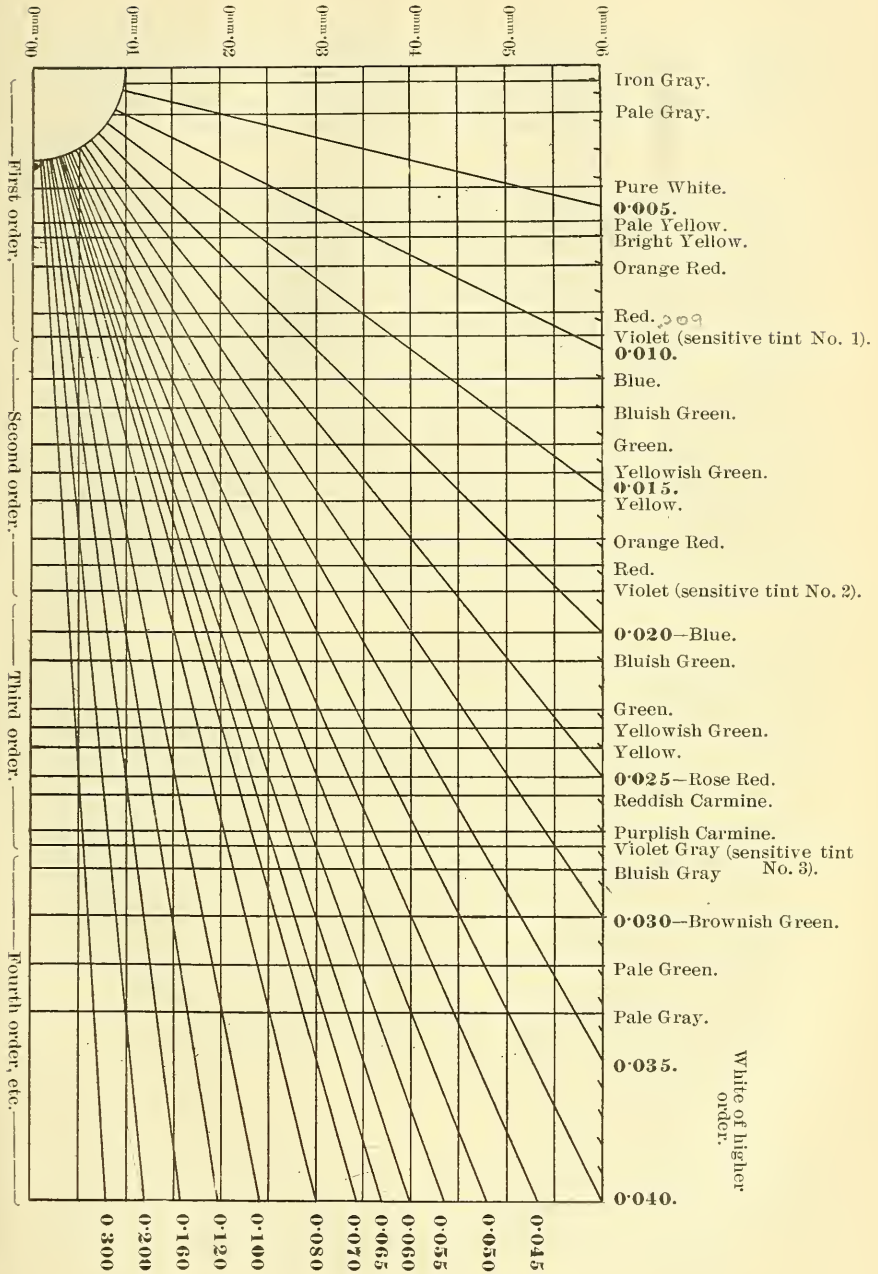
Thus, for example, in a slide having numerous quartz sections the highest color shown is a pale yellow. The birefringence of quartz is 0.009. The diagonal 0.009 crosses the line of pale yellow at about 0.035 which gives the thickness of the section. In the slide are numerous grains of an unknown mineral, colorless, of strong refraction and whose highest color, given by several of them, is an orange-red of the second order. But there passes through the intersection of the thickness 0.035 and the line of orange-red of the second order the diagonal 0.029 which is the approximate value of the birefringence of the mineral. Referring to the table this is found to approximate to that of allanite, chondrodite, diopside, cancrinite and tremolite. The refraction, color, cleavage, etc., show at once that of these, diopside is the only one which corresponds and thus the mineral is determined.

BIREFRINGENCES (MAXIMUM).

0·287	Rutile.	0·028	Tremolite.	0·011	Clinochlore.
0·179	Dolomite.	0·025	Actinolite.	0·010	Dumortierite.
0·172	Calcite.	0·024	Anthrophyllite.	0·010	Staurolite.
0·126	Aragonite.	0·024	Diallage.	0·010	Topaz.
0·121	Titanite.	0·024	Hornblende,	0·009	Bronzite.
0·107	Cassiterite.		common.	0·009	Labradorite.
0·072	Hornblende,	0·023	Gibbsite.	0·009	Quartz.
	basaltic.	0·023	Tourmaline.	0·008	Albite.
0·062	Zircon.	0·022	Glaucophane.	0·008	Andesine.
0·061	Anatase.	0·021	Augite.	0·008	Cordierite.
0·058	? Biotite, black.	0·021	Barkevikite.	0·008	Corundum.
0·055	Epidote, dark.	0·021	Brucite.	0·008	Kaolin.
0·055	Astrophyllite.	0·021	Sillimanite.	0·008	Oligoclase.
0·050	Talc.	0·019	Pargasite.	0·007	Soda-orthoclase.
0·049	Aegirite.	0·018	? Tridymite.	0·007	Soda-microcline.
0·045	Monazite.	0·016	Cyanite.	0·006	Orthoclase.
0·043	Anhydrite.	0·015	Dipyre.	0·006	Microcline.
0·038	Muscovite.	0·015	Ottrelite.	0·005	Eudialyte.
0·037	Epidote, light.	0·014	Wollastonite.	0·005	Melilite.
0·036	Meionite.	0·013	Hypersthene.	0·005	Nephelite.
0·036	Olivine.	0·013	Serpentine.	0·005	Zoisite.
0·035	Biotite, light green.	0·012	Anorthite.	0·004	Apatite.
0·032	Allanite.	0·012	Natrolite.	0·002	Leucite.
0·032	Chondrodite.	0·012	Seybertite.	0·002	Vesuvianite.
0·029	Diopside.	0·011	Andalusite.	0·001	Chlorite.
0·028	Cancrinite.	0·011	Antigorite.		

INDICES OF REFRACTION (MEAN).

2·711	Rutile.	1·669	Bronzite.	1·566	Brucite.
2·534	Anatase.	1·667	Sillimanite.	1·56	Biotite,
2·53	Brookite.	1·654	Seybertite.		light green.
2·38	Perovskite.	1·65	Dumortierite.	1·559	Labradorite.
2·029	Cassiterite.	1·645	Apatite.	1·553	Andesine.
1·952	Zircon.	1·644	Glaucophane.	1·551	Quartz.
1·930	Titanite.	1·644	Anthrophyllite.	1·551	Talc.
1·812	Pyrope.	1·641	Hornblende,	1·55	Dipyre.
1·811	Monazite.		common.	1·545	Nephelite.
1·791	Aegirite.	1·638	Andalusite.	1·542	Oligoclase
1·78+	Allanite.	1·635	Tourmaline.	1·542	Gibbsite.
1·770	Almandine.	1·633	Aragonite.	1·541	Cordierite.
1·765	Corundum.	1·632	Melilite.	1·54	Serpentine.
1·751	Epidote.	1·627	Actinolite.	1·54	Kaolin.
1·747	Grossular.	1·626	Wollastonite.	1·54+	Canada balsam.
1·741	Staurolite.	1·624	Pargasite.	1·535	Albite.
1·721	Vesuvianite.	1·622	Tremolite.	1·526	Microcline.
1·721	Augite.	1·622	Dolomite.	1·523	Orthoclase.
1·720	Cyanite.	1·622	Chondrodite.	1·507	Cancrinite.
1·719	Hornblende,	1·62	Topaz.	1·507	Leucite.
	basaltic.	1·609	Eudialyte.	1·496	Hauyne.
1·718	Ottrelite.	1·601	Calcite.	1·487	Analcite.
1·715	Spinel.	1·591	Biotite, black.	1·486	Natrolite.
1·705	Astrophyllite.	1·59	Clinochlore.	1·483	Sodalite.
1·701	Barkevikite.	1·587	Anhydrite.	1·476	Tridymite.
1·70	Zoisite.	1·585	Muscovite.	1·46	Nosean.
1·699	Hypersthene.	1·583	Meionite.	1·44	Opal.
1·688	Diallage.	1·580	Anorthite.	1·434	Fluorite.
1·683	Diopside.	1·577	Chlorite.		
1·679	Olivine.	1·567	Antigorite.		



ART. XXVII.—*Studies in the Cyperaceæ*; by THEO. HOLM.
 XIV. On a collection of *Carices* from Alaska with remarks upon the affinities of *Carex circinata* C. A. Mey. and *C. lejocarpa* C. A. Mey. (With figures in the text.)

AMONG the very extensive botanical collections, which Dr. Walter H. Evans brought from Alaska during the summers of 1897 and 1898, were several *Carices*, which were kindly donated to the writer, and upon which we have prepared the following paper. The Alaskan coast with adjacent islands has for many years been known to possess an unusually rich vegetation of *Cyperaceæ*, especially *Carices*, and a considerable number of highly interesting species are recorded in the works of Boott, Hooker, Ledebour, Macoun and Meyer.* Several of these species are absolutely local and their geographical range is thus confined to some of the small islands or to a few localities on the main land; some others are found, also, on the Asiatic side of Bering Sea; a few are distributed throughout the northern part of Asia, Europe and this country, while some of these species are even arctic and circumpolar. The following species were collected by Dr. Evans:

Carex macrocephala Willd.

Many gigantic specimens with spikes 7^{cm} in length and over 4^{cm} in thickness; the species is considered as diœcious by Franchet, who examined a large number of specimens from eastern Asia, but among those collected by Dr. Evans there were several individuals with purely staminate and pistillate spikes on the same plant.

var. *bracteata* Holm nov. var.

The bracts of the inflorescence developed into green leaves, about 10^{cm} in length.

Common in sand along shore, Kussiloff.

C. Macloviana d'Urv. var. *Haydeniana* (Olney).

On flat at Feeny's ranch, five miles from Kadiak.

C. canescens L.

Common about Sitka.

var. *polystachya* Bail.

In low places, Unalaska.

* For references consult the bibliography appended to this article.

C. cryptantha Holm nov. sp. (fig. E).

Roots very slender; rhizome short, cespitose; culms numerous until 45^{cm} in length, triangular, scabrous above; leaves shorter than the culm, about 25^{cm} in length, narrow, flat, scabrous along the margins and on the lower surface, their sheaths short, pale brown, thin in texture, not fibrillose; inflorescence very short, from 1½ to 2^{cm} in length, consisting of 2 or 3 contiguous, sessile, reddish-brown spikes, mostly androgynous with a few female flowers at base; bracts scale-like, obovate, brown, membranaceous, the midrib sometimes excurrent as a bristle-like awn; scales of staminate and pistillate flowers (fig. F) broadly elliptic, acute, shining reddish brown, pointed, the margins hyaline; utriculus (fig. G) rather broadly elliptic, attenuated at both ends, pale brown, glabrous, coriaceous, with several, distinct nerves, the beak short, entire; caryopsis dull and brown, roundish in outline, almost flat; stigmas 2.

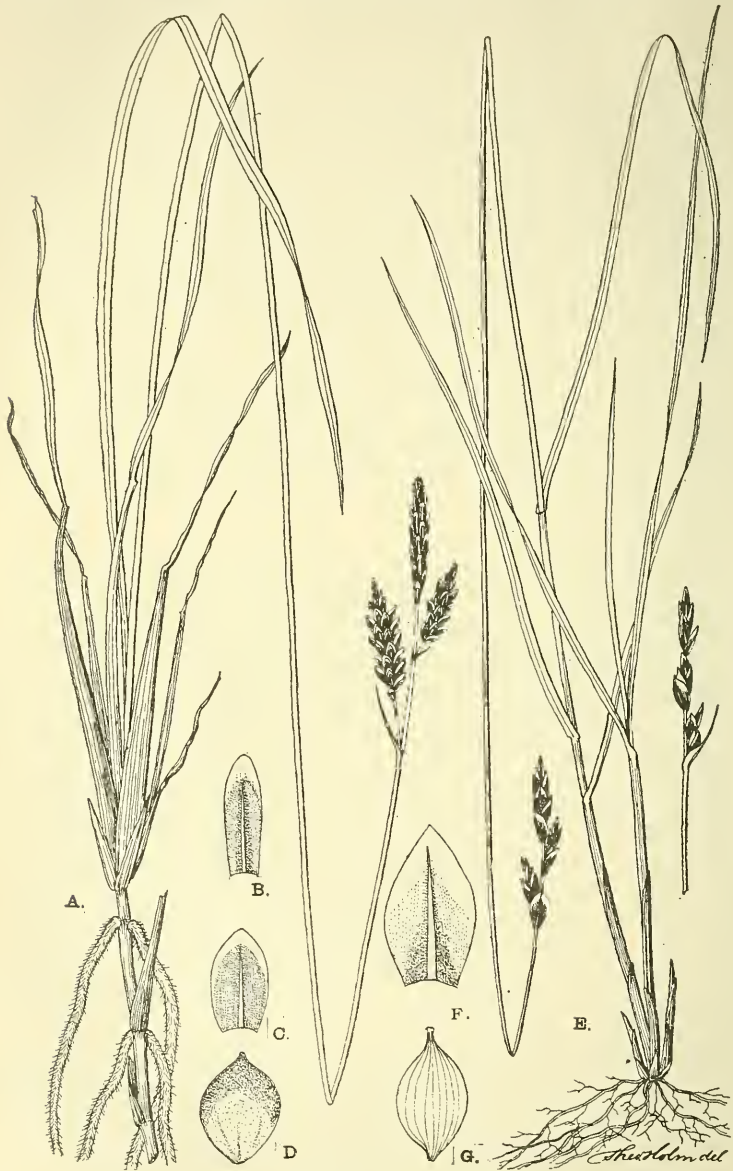
In dense mats along small streams between Kussiloff and Kenai. Its nearest ally is *C. heleonastes* Ehrh., from which it differs in its narrow inflorescence and shining reddish brown scales, which are larger and much broader, almost entirely concealing the utricule.

C. læviculmis Meinsh.

This species has been known for many years, but has not been well understood; it was referred to *C. elongata* by Bongard; to *C. remota* by Treviranus, and as “*potius C. glareosa*” by Ruprecht, while L. H. Bailey described it as a variety *sparsiflora* of *C. Deweyana* Schw. It was not until 1893 that Meinshausen segregated it as a species, and his diagnosis reads as follows: “*Radix cæspitosa, fibrosa; culmis tenuissimis acute triquetris, 1-1½ ped., flaccidis lævissimis inferne foliatis; foliis linearibus planis, infimis brevissimis, summis dimidio culmi subæquantibus; spica elongata, spiculis 4-8 subglobosis remotius interruptis, infima sæpe valde remota, ferrugineo—viridibus, basi masculis; bracteis squamæformibus, ima setacea longe aristata rarius foliacea spiculæ 2-3—plo. excedente; perigyniis ovatis rostro acuminato, ore minimo integro, plano convexis marginibus acutis scabris, utrinque rubro nervatis patulis, apice paulo incurvatis demum ferrugineis, squama ovata acuta v. acuminata ferruginea acute carinata paulo breviora.*”

The specimens examined by Meinshausen were from Sitka and Kamtschatka (legit Mertens). Dr. Evans collected it on edges of bogs on cleared lands, Wrangell; along the beach above high tide, Howkan; in woods, common, Kussiloff.

The species has, furthermore, been collected by Thos. Howell in marshes by Yes Bay, Alaska, and by John Macoun in the vicinity of Comox, Vancouver Island. Rev. G. Kükun-



Carex lugens (A-D) and *C. cryptantha* (E-G) natural size.

thal informs us (in litteris) that he has, also, seen specimens from Idaho, Washington and Oregon, which he considers identical with this species.

C. polytrichoides Muhl.

Edge of marsh, Kussiloff.

C. tenella Schk.

In marsh, not abundant, Kussiloff.

The specimens are somewhat shorter and more robust than the typical plant from Scandinavia and Rocky Mountains, but the structure of the spikes and utriculus is the same.

C. circinata C. A. Mey.

On rocky hilltops, 900 to 1200 feet alt., Kadiak.

This species is one of the very rarest, and has only been found a few times before, perhaps not very far from where Dr. Evans collected it. Meyer, who is the author of the species, gives the locality only as "in rupibus Unalaskæ," while Prescott, who described the Cyperaceæ for Bongard, states that it was found on "the island of Sitcha"; according to Ledebour, Redowsky is credited with having collected the plant somewhere in Kamtschatka, though with a query. In later years Professor John Macoun has reported the species as having been found by Barclay near Sitka, but these specimens, which we have seen, do not belong to *C. circinata*, but to *C. pyrenaica* Wahlbg.; the species occurs, however, on the Shumagin Islands, where it was collected by Mr. Harrington in 1871-72, and of which a few specimens are preserved in the National Herbarium. But we find no record of it in Trautvetter's works, in Kjellman's or Franchet's, nor is it included in James Macoun's list of plants from the Pribilof Islands. The species must be very local, and seems to be confined to the Alaskan coast with adjacent islands.

C. Gmelini Hook.

Among rocks on beach, Sitka; on flats, Kussiloff; abundant at Homer, Cook inlet.

C. nigella Boott.

A very few specimens were collected on Kadiak, where it grew together with *C. circinata*.

C. vulgaris Fr.

Growing in small tufts in wet places, Wrangell.

C. interrupta Bœckl.

Abundant at Wrangell, grows in dense tufts.

C. lugens Holm, nov. sp. (fig. A).

Roots thick, very hairy; rhizome ascending, stoloniferous; culm from 30 to 45^{cm} in height, slender, but stiff, sharply

triangular, scabrous; leaves shorter than the culm, narrow and flat, scabrous along the margins and the lower surface; spikes 3 to 4, contigous: one terminal staminate, 2 or 3 lateral pistillate; staminate spike peduncled, about 2^{cm} in length, the scales (fig. B) obovate-oblong, obtuse, dark brown with hyaline margins, the midrib pale, not excurrent; pistillate spikes erect, very dark, short, only 1 to 2^{cm} in length, densely flowered, the upper sessile or nearly so, the others peduncled; bracts not sheathing, very short and filiform; scale of pistillate spike (fig. C) obovate and acutish, black with very narrow hyaline margins, midrib not excurrent, barely visible; utriculus (fig. D) glabrous, a little longer than the scale, thin in texture, round obovoid with a minute entire beak, purplish above, pale green below, nerves barely visible; caryopsis obovate, flat; stigmas, 2.

Common in drier places on bay, Kussiloff.

Related to *C. cespitosa* Good., but is at once distinguished from this by its stoloniferous habit and darker colored spikes. *C. cespitosa* grows always in dense tufts and the culms are at the base surrounded by purplish, bladeless sheaths, besides that the utricule is pale green.

C. cryptocarpa C. A. Mey.

Abundant near villages, Sitka; on sands with *Elymus*, Kussiloff; in wet places near beach, Wrangell and Howkan.

C. macrochaeta C. A. Mey.

In bogs at Sitka, 1000 feet above sea-level; on mountain top about 1200 feet, Kadiak.

C. salina Wahlbg. var. *cuspidata* Wahlbg.

On flat in rich soil, Kussiloff. The specimens were low, from only 10 to 17^{cm} in height, but agree very well with the Norwegian plant; it is very interesting to notice that the species exhibits exactly the same variation in Alaska as in Norway, thus the var. *reducta* (Drej.) has been collected by Mr. James M. Macoun on St. Paul Island, while the var. *haematolepis* (Drej.) was collected on the same island by Mr. Thos. H. Kearney Jr.

C. stygia Fries.

Common on flats near edge of bog, Kussiloff; in bogs and wood meadows, Howkan, Sitka.

C. livida Willd.

In rather dry places, Sitka. Specimens very tall, reaching until 30^{cm} in height.

C. pauciflora Lightf.

Abundant in bogs, Sitka.

C. utriculata Boott.

On bog, Sitka.

C. physocarpa Presl.

Common at Kadiak on rather dry mountain side at an elevation of 500 feet. The specimens seem better referable to this species than to *C. compacta* R. Br., but neither this or *C. physocarpa* is well understood, thus the identification is very uncertain.

In considering the geographical distribution of some of these species, the following are circumpolar: *C. canescens*, *C. vulgaris*, *C. salina* and partly, also, *C. livida*; *C. cryptocarpa* is very common on the western and eastern coast of Bering Sea, and extends as far south as Japan and China, but it is otherwise known only in Greenland, Iceland, Farø Islands and a few places on the Norwegian coast; *C. stygia*, which seems to be very abundant in several parts of Alaska has not, so far, been recorded with certainty from other countries than Finmark in Norway. *C. tenella*, which has a very wide distribution in North America, from the Atlantic to the Pacific Ocean, does not occur in Greenland, is rare in Scandinavia and Arctic Russia, has not yet been reported from Arctic Siberia, and is first met with again in Kamtschatka. *C. macrocephala* and *C. Gmelini* are only known from the eastern and western coast of Bering Sea but both extend as far south as Japan, the former even to China; *C. macrochaeta* has been collected in many places on the Alaskan coast, but not, so far, on the Asiatic side, unless Kjellman's specimens from Konyam Bay may be referred to this rather than to *C. podocarpa* R. Br. *C. Macloviana* shows an exceedingly wide distribution ranging from Alaska southwards to Magellan, besides that it occurs, also, on Sandwich Islands, in Greenland and Arctic Europe. *C. nigella* is only known from Alaska and the Rocky Mountains in middle Colorado; *C. interrupta* is only known on this side of Bering Sea: Alaska and Oregon.

The very considerable material of the rare *C. circinata*, which was placed at our disposal, induced us to study its affinities, inasmuch as we received, also, a number of excellent specimens of its nearest ally, the equally rare *C. lejocarpa* C. A. Mey., for which the writer is indebted to Mr. James M. Macoun, who, some years ago, detected this species on the Pribilof Islands in Bering Sea.

These two species resemble each other very much, but are, nevertheless, readily distinguished by the following characters:

The rhizome is horizontal and creeping in *C. lejocarpa*, but cespitose in the other; the leaves are flat and straight in the former, but almost filiform and "circinate" in the other; the utricle has a hyaline bidentate beak in *C. circinata*, but a merely truncate in the other. As regards the distribution of the sexes, the spike was found to be androgynous in all the specimens, several hundreds, which we have examined of *C. circinata*; in *C. lejocarpa*, on the other hand, we have found specimens that were truly *diœcious*, and others in which the spike was androgynous, the latter being quite scarce. The number of stigmas varies in both, and we observed two and three in the same spike of *C. circinata*, three being, however, the commonest; in *C. lejocarpa* we noticed two stigmas in only two specimens.

In making our first disposition of these species we did not hesitate to place them among the "*Carices genuinæ*," not only on account of the number of stigmas being mostly three, but especially because the external structure and coloration of the spike suggested affinities to certain species of *Carex* proper rather than of *Vignea*. However it is often very difficult to decide where to place some of these little, monostachyous species, and one is often most inclined to consider the number of stigmas as much more important than it really is. The distigmatic monostachyæ are thus generally placed among the "*Vignæ*" as representing small sections of their own, as, for instance, the *Capituligeræ* of Rev. G. Kükenthal, to which are referred *C. nardina*, *C. capitata*, *C. pulicaris*, and others, yet the *C. nardina* of the Rocky Mountains in Colorado has most often three instead of two stigmas, and as we have stated above a similar variation exists also in both *C. circinata* and *C. lejocarpa*. We should prefer to compare the structure of utriculus besides the shape and relative size of the scales and bear in mind that the formerly so-called "*Psyllophoræ*" in most cases may be referable to the higher developed "greges" not only among "*Carices genuinæ*," but also among "*Vignæ*," as illustrated by Drejer and Boott in various instances. And in regard to our two Alaskan species, no Caricologist can possibly deny that their spikes, considered by themselves, show a striking resemblance to those of *C. sempervirens* Vill. and its allies. We have, also, but so far in vain, tried to establish some connection between them and *C. nigricans*, *C. pyrenaica*, *C. macrostyla* and *C. pulicaris*; the affinities of these species with their deciduous squamæ and reflexed utricles may, perhaps, be sought among some of the higher forms with squarrose spikes.

In our further disposition of *C. circinata* and *C. lejocarpa* we have considered them as "*formæ hebetatæ*" of that section

to which *C. sempervirens* is supposed to belong, the affinities of which may naturally be sought among such boreal or alpine forms as *C. misandra* R. Br., *C. frigida* All., *C. ferruginea* Scop., etc. The systematic position of these species is not defined in Drejer's work, but he makes some allusions to a possibly existing affinity between *C. circinata*, *C. Davalliana* on the one side and *C. sempervirens*, *C. misandra* and *C. frigida* on the other, while he enumerates these as "*formæ desciscentes*" under "*Melananthæ*." However, on his table of affinities III Drejer places these between "*Sphaeridiophoræ* and *Lamprochlænæ*," and he evidently considered them as representing a section or grex of their own; the work, however, which this author began in such an admirable manner, unfortunately, never finished. In Tuckerman's "Enumeratio methodica" we find some of these species classified as "*Fuliginosæ*" (*C. fuliginosa* and *C. frigida*) and as "*Ferrugineæ*" (*C. ferruginea*, *C. sempervirens*, *C. brachystachys* and *C. firma*), while he does not seem to have known *C. lejocarpa* or *C. circinata*. A similar disposition is, furthermore, to be noticed in Dr. Christ's Catalogue, where these same species are arranged in one section "*Frigidæ* of Fries," together with *C. mucronata* All. and *C. brachystachys* Schrank. The *Frigidæ* are taken up again by Rev. G. Kükenthal, placed between *Sphaeridiophoræ* and *Hymenochlænæ*. Still another, but most unsuccessful classification is proposed by Professor Bailey, who includes all the *Frigidæ* in *Microrhynchæ*, which he, furthermore, combines with *Ærostachyæ*,* *Melananthæ* and *Trachychlænæ* (*C. glauca* Scop.) of Drejer, while *C. circinata* and *C. lejocarpa* are placed under a new section *Leptocephalæ* Bailey including *C. polytrichoides*! It does not seem necessary to discuss further such combination, which is in no respect in accord with the system proposed by Drejer, to whom Professor Bailey refers as author of these various sections.

It appears as if the *Frigidæ* may really constitute a special section, as proposed by Dr. Christ and Rev. G. Kükenthal, but we do not know how far Mr. Kükenthal extends this section, since this author has only treated the South American *Carices*, of which only one, *C. Vallis pulchræ*, is enumerated. However, by examining the European and boreal species of this section, we have noticed that certain and, indeed, very striking analogies exist between them; moreover, the utriculus may be well referable to one single type in all of them. This organ is in these species more or less trigonous in transverse section; it is membranaceous, usually glabrous or sometimes a little

* The sectional name *Ærostachyæ* (*αἰώρα*) as proposed by Drejer is by Professor Bailey constantly cited as "*Ærostachyæ*."

scabrous above; it is generally attenuated at both ends, and the orifice of the beak is mostly bidentate with erect teeth, or entire, oblique. The nerves are seldom distinct, and the coloration may vary from green to deep purple or brown; thus the spikes in some species attain the same dark aspect as is characteristic of the *Melananthæ*. Another character common to these European and boreal species is the position of the pistillate spikes, being more or less remote and nearly always borne on very conspicuous peduncles, in some species of quite considerable length and nearly filiform. The sheaths of the bracts are well developed, and in regard to the general habit, these species exhibit but a few divergences, which, however, may be considered as analogous deviations, such as we know exist in some of the other sections. The terminal spike is sometimes gynæcandrous or androgynous in *C. brachystachys* and, as it seems, constantly gynæcandrous in *C. misandra*; in all the other plioistachyous, which we have examined, the terminal spike is purely staminate. Such varied distribution of the sexes is also met with among the *Melananthæ*, the *Hymenochlænæ* and the *Spirostachyæ*. And if we, moreover, consider *C. lejocarpa* and *C. circinata* as members of the *Frigidæ*, we then perceive a still more gradual variation among the species themselves: from the diœcious *C. lejocarpa* to *C. circinata* with androgynous spike, and to *C. sempervirens* and its very nearest allies with unisexual spikes, borne on the same culm, to *C. brachystachys* with occasionally androgynous or even gynæcandrous terminal spike and *C. misandra* with its invariably gynæcandrous. Abnormally decomposed pistillate spikes, as these are frequently met with in several of the other sections, have also been noticed in some of the *Frigidæ*; Boott records such cases from *C. brachystachys*, and we have, furthermore, found these in *C. ablata* Bailey and in one single inflorescence of *C. circinata*.

In looking over the various sections, as these are described by Drejer, we have not been able to find any under which the *Frigidæ* might be properly arranged. They seem to possess some characters common to the *Hymenochlænæ*, yet lacking the essential character of these: "perigynia plerisque nervata, rostro apice hyalino bilobo—bifidove, vel distincte bifido," besides "squamæ membranaceæ albidæ" and "color plerumque læte viridis." And in comparing them with the *Spirostachyæ* certain analogies seem to exist; for instance, the darker color of the spikes, the mostly erect and remote, pistillate spikes, while "perigyniis distincte bifidis" and "spicis densifloris quasi squarrosis" are not possessed by the *Frigidæ*. We have not, on the other hand, been able to detect any intergradating forms between these three sections, by which the *Frigidæ*

might be connected as "lesser" or "higher" developed types; on the contrary, it seems as if we in the *Frigidæ* themselves might trace some of their earlier types represented by, for instance, *C. lejocarpa*, that we might count *C. sempervirens* and *C. frigida* among their central forms, while *C. misandra* may be looked upon as a "*forma desciscens*," yet possessing the most essential characters of the section.

While thus fully recognizing the maintenance of these species as representing a section of their own, we should prefer not to adopt "*Frigidæ*" as the sectional name but one in Greek which, moreover, would be in stricter conformity with those suggested by Drejer for the other sections. And when we suggest "*Stenocarpæ*" as a more proper name, this refers not only to the characteristic shape of their utriculus, but also to one of the central forms, "*Carex stenocarpa*" of Turczaninow, the Asiatic homologue of *C. sempervirens*. The principal characteristics of the section, drawn from the central forms, are as follows:

Carices stenocarpæ.

Terminal spike staminate; lateral ones pistillate, distant, peduncled, erect or finally drooping. Bracts distinctly sheathing, mostly filiform, short. Utricle membranaceous, glabrous or a little scabrous above, narrowly attenuated at both ends; orifice of beak bidentate with erect teeth or obliquely cut, often hyaline; stigmas 3, achenium trigonous. Boreal or alpine species with mostly dark-colored spikes. The lesser developed, monostachyous types (*C. lejocarpa* and *C. circinata*) remind of certain *Vignæ*, the central of *Spirostachyæ*; at the limits of the section are types (*C. misandra*, *C. brachystachys*) with androgynous or gynæcandrous terminal spikes, which remind of certain *Melananthæ* (*C. atrata*, *C. Mertensii*).

Among the species which represent this section may be enumerated:

<i>Hebetatæ</i>	{	<i>C. lejocarpa</i> C. A. Mey.
		<i>C. circinata</i> C. A. Mey.
		<i>C. sempervirens</i> Vill.
		<i>C. frigida</i> All.
<i>Centrales</i>	{	<i>C. ferruginea</i> Scop.
		<i>C. ablata</i> Bail.
		<i>C. firma</i> Host.
		<i>C. hispidula</i> Gaud.
<i>Desciscentes</i>	{	<i>C. brachystachys</i> Schrank.
		<i>C. misandra</i> R. Br.

To these may possibly be added some others, of which we have seen no specimens, but judging from the diagnoses and

illustrations, we suppose that they may be regarded as members of this section: the monostachyous *C. hakkodensis* Franch., *C. rhizopoda* Maxim., *C. heteroclita* Franch. and *C. grallatoria* Maxim. all from Japan; *C. Vallis pulchræ* Phil. from South America, and *C. acicularis* Boott from New Zealand; finally the plioistachyous *C. hæmatostoma* Nees, *C. psychrophila* Nees and *C. macrogyna* (Turcz.) from northern India. Boott himself considered *C. macrogyna* as being an ally of *C. sempervirens*, but states that in the former the uppermost spikes are sometimes androgynous. In regard to *C. hirtella* Drej., *C. hæmatostoma* and *C. psychrophila* Boott compares, also, these with the *Ferrugineæ* Tuckin., as representing on the mountains of northern India those analogons species found on the Alps of Europe. *C. grallatoria* is, according to Maximowicz, a close ally of *C. lejocarpa*.

The *Stenocarpæ* constitute a small, very limited section, and the types are either alpine, arctic or confined to the coasts of the Pacific from Bering Sea to Japan. *C. misandra* exhibits the widest range in geographical distribution; besides being circumpolar, this species occurs on the mountains of both northern and middle Europe and on the Rocky Mountains as far south as middle Colorado. The central forms, excepting *C. ablata* from Vancouver Island, are confined to the European Alps or, if we include those cited by Boott, to the mountains of India. Finally, as mentioned above, there is a species in South America and one in New Zealand, both with a single, androgynous spike, which Mr. Kükenthal considers as belonging to this same section, with him the *Frigidæ* of Fries.

There is, thus, considered from a geographical viewpoint, a wide gap between the "*hebetatæ*" and "*centrales*;" yet, as we have stated above, there is one of the latter which inhabits Vancouver Island; moreover by including the Japanese representatives of "*hebetatæ*" these are not so very remote from their higher developed allies on the mountains of northern India. The Pyrenees and middle European Alps possess, on the other hand, types which are very isolated and are only connected with their northwest American allies through *C. misandra*. The most gradual development of the section may, thus, be traced in Asia and in the northwestern corner of America, where all the types occur: *hebetatæ*, *centrales* and *desciscentes*. But we have no species of this section in Europe, which as *C. lejocarpa* may point towards some fundamental types of the recent time, and the *formæ centrales* of the European mountains, *C. sempervirens*, etc., may be considered as representing the remnants of an old center of which the earlier types have disappeared long ago, but leaving a group of species, the highest developed in the section.

While thus the *Stenocarpæ* represent a very natural little group of *Carices* in respect to their external, morphological, characters, we have, furthermore, examined the anatomy of some of these species, which seems to be very uniform. Whether uniformity in anatomy, however, may be considered as being an absolute necessity for the establishment of sections of *Carices* to which morphological peculiarities are in common, is by no means certain. We will, no doubt, meet with many exceptions by extending our studies to several of the other and larger sections, as for instance the *Melananthæ* and *Microrhynchæ*, the members of which exhibit so many diverse types from the extreme north and south, and living under very different conditions as to climate and soil. It is, also, very possible that several of the smaller sections may be found to possess a like structure in anatomical respects, yet being apparently distinct when considered from a morphological viewpoint. This is readily noticed from the literature on this subject, the works of Mazel and Lemcke. The former of these authors has examined the structure of the root, the stem and the leaf of 43 species of *Carex*, nearly all from Europe, and his final conclusion is thus expressed: "On ne peut en aucun cas se baser sur des caractères anatomiques pour grouper systématiquement les espèces dans le genre *Carex*"; yet it does not seem as if this author has made any attempt to deduct a classification from his anatomical results. The other author, Alfred Lemcke, who has examined the rhizome and the above-ground stem of about 160 species of *Carex* from Europe and this country, believes, on the contrary, in the possibility of classifying *Carices* from the structure of their rhizomes, and considers Dr. Christ's system as the most practical for this purpose. The *Frigidæ* Fr. including *C. mucronata* All. are, thus, considered as being closely related to each other, while we find *C. circinata* enumerated under *Psyllophoræ* among such species as: *C. gynocrates*, *C. capitata*, *C. polytrichoides* and even *C. Fraseri*. However we doubt very much whether any conclusion may be drawn from the structure of the rhizome alone or from the stem so as to establish sections or smaller groups of *Carices*. Nevertheless the works of these two authors constitute a valuable contribution to our knowledge of the anatomy of *Carex*, even if their results do not bear directly upon the classification.

In regard to our *Stenocarpæ* we are, thus, well aware of the possibility that the anatomical characterization which we present as supplemental to the morphological, may also contain several points that are common to some of the other sections. These anatomical notes may, nevertheless, become useful to further studies of the genus, at least as a contribution

to similar investigations. As regards the material which we have examined, the specimens were collected in the following localities:

C. circinata C. A. Mey., rocky hilltops, 360–375^m alt., Kadiak; *C. lejocarpa* C. A. Mey., grassy banks, St. George Island, Bering Sea, and marshes at Yes Bay, Alaska; *C. sempervirens* Vill., Alps of Switzerland, Tyrol; *C. ferruginea* Scop., along streams, 2000^m alt., *C. frigida* All., 2300^m alt., *C. firma* Host., 2000^m alt., *C. hispidula* Gaud. 2100^m alt. and *C. brachystachys* Schrank, 1260^m alt., all from the Alps of Switzerland; *C. ablata* Bailey, 600^m alt., Vancouver Island; *C. misandra* R.Br., St. Mathew Island, Bering Sea; along mountain streams on Gray's Peak, 4000^m alt., Colorado; on dry, grassy mountain slopes or swamps, West Greenland; Rendalen, Spitzbergen; in swamps on rocks, Nova Zembla; Alps of Tyrol at 3000^m alt.

The root.

In the species examined we found the roots very strongly built with a more or less thick-walled hypoderm inside the epidermis. The outermost strata of the cortex show a similar thickening and constitute thus a firm sheath around the inner bark, which is thin-walled and usually collapsed throughout to the innermost stratum that rests on endodermis. In *C. sempervirens* the outer six strata of the cortex were especially heavily thickened, also in *C. hispidula*, in which even the innermost four or five strata exhibited a like thickening, while in *C. firma* the entire cortical parenchyma was developed as a solid mass of stereids. Endodermis is thick-walled in all the species but not to the same extent; thus an U-endodermis may be found in *C. lejocarpa*, *C. circinata*, *C. misandra*, *C. hispidula* and *C. sempervirens*, while it appears as a V-endodermis in the remaining species; the thickening is especially heavy in *C. firma* and *C. sempervirens*. The pericambium is most often thin-walled, or in some roots of *C. firma*, *C. frigida* and *C. sempervirens* the cell-walls may be observed as being somewhat thickened. It consists of only one stratum, which forms a closed ring in *C. sempervirens* and sometimes also in *C. firma*, but is, in the other species, interrupted by the protohadrome vessels in a more or less regular manner. We have noticed the following cases: In *C. circinata*, *C. lejocarpa*, *C. misandra*, *C. frigida* and *C. ablata* the pericambium was constantly interrupted by all the protohadrome vessels, there being mostly four or five pericambium cells between each two of these vessels. In *C. hispidula* we found in some roots the pericambium interrupted by all the protohadrome vessels or by only five out of twenty-five, and these roots were from the

same individual; the number of pericambium-cells between each two vessels varied in this species from four to seven, but four and five was the commonest. This peculiar case was, moreover, noticed in other specimens from a still higher elevation (2800^m Mt. Ryssel), in which sometimes most of the proto-hadrome vessels were situated inside the pericambium, and only a very few bordered on endodermis. In *C. firma* a similar and very irregular position was noticed, since the pericambium was either not interrupted at all or it was broken by the majority of the vessels, by 14 out of 15. In *C. brachystachys* we found no case where the pericambium was not interrupted, but the position of the proto-hadrome was very irregular even in the same root, cut at different places. The number of interruptions varied thus between 8 and 13, while the distance between the places where the sections were made was about 5^{cm}; in most of the roots of this species, however, we observed that the majority of the proto-hadrome vessels were located inside the pericambium, and only in one root did we notice that fourteen out of nineteen bordered on endodermis. In *C. ferruginea* Scop. we observed that all the proto-hadrome vessels bordered on endodermis in roots of very different thickness, but in a single root the following position was noticed: 25 out of 26 vessels had broken through the pericambium, but at the other end of the same root there were only 18 of these 25 vessels that bordered on endodermis.—*C. hispidula*, *C. ferruginea* and *C. brachystachys* illustrate thus the singular fact that the pericambium is very irregularly interrupted by the proto-hadrome, but, on the other hand, we found no cases in these species where it was not interrupted at all. *C. firma*, on the contrary, possesses roots in which the pericambium may sometimes form a closed ring or it may be broken by nearly all the proto-hadrome vessels. This singular instance is, moreover, to be observed in roots of *C. supina* Wahlbg. in some specimens which we collected in West Greenland; the species belongs, however, to another section, evidently to the *Lamprochlanæ* Drej., which we intend to discuss in a subsequent paper.—But we have not so far noticed any such variation in any of the other *Cyperaceæ* which we have heretofore examined. In regard to the other vessels in the roots of these species we have not found any occupying the center of the central-cylinder, with the exception of small, lateral roots, which usually possessed only one central vessel. The conjunctive tissue which thus occupies the innermost part of the central-cylinder is distinctly thick-walled in all the species, with the exception of *C. ablata*, where it is thin-walled even in such roots of which the endodermis shows the cell-walls to be very considerably thickened. The leptome, includ-

ing the proto-leptome, was very well developed in all the species and was located between each two proto-hadrome vessels.

In examining the above-ground stem we find that it does not exhibit such striking divergences as was noticed in the roots. Sections taken from about the middle part of the culm showed the outline to vary from cylindrical to sharply triangular as follows: it was found to be terete and nearly glabrous in *C. circinata*, *C. sempervirens* and *C. brachystachys*; pentagonal and almost glabrous in *C. ferruginea*; obtusely triangular and glabrous in *C. ablata*; triangular and nearly glabrous in *C. lejocarpa*, *C. misandra* and *C. frigida*, while triangular and very scabrous in *C. firma* and *C. hispidula*. The cuticle is quite thick and perfectly smooth, not wrinkled in these species. The epidermis shows generally a distinct thickening of the outer wall, especially in *C. sempervirens* and *C. ablata*; the radial cell-walls are a little thickened in *C. ferruginea*, *C. frigida* and *C. ablata*. Cone-cells were found outside the hypodermal stereome in all species; stomata occurred frequently in that part of epidermis which covered the bark and showed the same structure as in the leaves, which will be described later. A somewhat peculiar epidermal structure is represented by *C. hispidula*, where the outer wall in nearly all the cells was noticed to be extended into roundish papillæ, especially near the stomata, although not covering these. The cortical parenchyma is mostly developed as short palisades, radiating towards the center of the stem, and contains lacunes, sometimes of quite considerable width in *C. circinata*, *C. lejocarpa* and *C. hispidula*. The stereome occurs as hypodermal groups, covering the larger mestome-bundles, but is also to be found on the hadrome side of these, bordering on the pith. The smaller mestome-bundles have usually only a minute support of stereome, this being merely developed as a few cells on either face of these. This tissue, the stereome, is very thick-walled, at least on the leptome-side, in *C. circinata*, *C. sempervirens*, *C. hispidula*, *C. firma*, *C. frigida* and *C. brachystachys*, much less so in the remaining species, and it is generally rather open on the hadrome-side of all the mestome-bundles.

The mestome-bundles are developed as larger, in transverse sections oval, or as smaller, nearly orbicular, and are arranged in almost regular alternation with each other. They constitute in most of the species only one peripheral band, but in *C. hispidula*, *C. ablata* and *C. brachystachys* we noticed an inner band of a few, 3 to 5, large bundles, almost entirely imbedded in the pith. The parenchyma-sheath is invariably thin-walled and seems often to contain chlorophyll; a mestome-sheath with the inner walls distinctly thickened is also noticeable in all the

species; but we found no trace of the inner green sheath, which we remember is known to occur in many genera of the *Cyperaceæ*, but not so far known from any species of *Carex*.

There is a central, thin-walled pith in all the species, and this seems to break down, leaving a wide central cavity in all the species examined, with the exception of *C. circinata*.

The leaf.

The leaf exhibits a somewhat greater variation than is noticeable in the stem. The narrowest leaf-blade is possessed by *C. circinata* and *C. brachystachys*, but if we consider the sections taken from the middle of the blade, this is in none of these so narrow that it might be described as semi-cylindrical. The surface of the leaf is very smooth in most of the species, but in *C. hispidula* the blade is somewhat constricted between the mestome-bundles, besides that the epidermis in this species is very scabrous on either face. In *C. brachystachys* the lateral parts of the blade are much thicker than the mediane, and very large prickle-like projections are developed on the upper surface on each side of the bulliform-cells. The leaf-surface is thus mostly smooth with no deep furrows, but it is scabrous from prickle-like projections in *C. sempervirens*, *C. ferruginea* and *C. brachystachys*; in *C. hispidula* both surfaces are quite scabrous from the numerous papillæ, which also characterized the stem of this species. Bulliform-cells occur on the whole upper face of *C. circinata* and *C. lejocarpa*, but these cells are in the other species restricted to a single group just above the midrib; the outer cellwall of epidermis is very thick in some species, for instance, *C. hispidula*, *C. frigida* and *C. firma*, but not so in *C. ablata*. Some certain variation exists also in the relative length and breadth of the epidermis-cells outside the stereome, and this is especially distinct when we examine the leaf in superficial sections; we notice, for instance, in *C. ferruginea*, *C. brachystachys* and *C. misandra* that these cells are narrower, but not shorter, than the surrounding epidermis, while in *C. sempervirens*, *C. frigida* and *C. hispidula* they appeared to be much shorter but not narrower. In *C. firma* and *C. ablata* these same cells, covering the stereome, were distinctly both narrower and shorter than the stomatiferous strata.

In regard to the stomata, these are in the *Stenocarpæ* confined to the lower surface of the leaf, and show a very uniform structure; they are free in all the species, slightly projecting in *C. misandra*, *C. firma* and *C. ablata*, but level with epidermis in the others; in *C. hispidula*, as mentioned under the

stem, the stomata are surrounded by mostly four papillæ, which, however, are quite short and not bent over the stoma, thus this is almost free as in the other species. This structure of the subsidiary cells with their papillæ in *C. fimbriata* has already been described by Schwendener in his paper on the stomata of the *Gramineæ* and *Cyperaceæ*.*

The mesophyll is either represented as a homogeneous tissue throughout the leaf, as in *C. circinata* and *C. lejocarpa*, or it is differentiated into a more or less distinct palisade-tissue on the upper face and a more open, pneumatic on the lower, as in *C. ferruginea*; in the other species the palisades are short and not so regular, but appear nevertheless to be placed vertically on the blade, with the exception of *C. ablata*, in which the palisade-cells show a tendency to radiate towards the center of the mestome-bundles. Lacunes were observed in all the species, one between each two mestome-bundles, and they were especially wide in *C. circinata*, *C. lejocarpa*, *C. misandra* and *C. hispidula*. The stereome occurs as in the stem on both faces of the mestome-bundles and as hypodermal in the larger of these; it is, moreover, to be observed as a small, isolated group in each of the two leaf-margins. It is very thick-walled in *C. sempervirens* and *C. ferruginea*, moderately so in the other species; it is rather weakly developed in *C. ablata*.

The mestome-bundles lie in one plane, and the midrib is generally somewhat larger than the others and more prominent. We find in the leaf as in the stem two forms of these bundles, larger and smaller, which show the same alternation in regard to their arrangement and the same structure as we have described under the stem.

* The position of the stomata in *Carex*, whether these are free and exposed or sunk below the surrounding epidermis and sometimes partly covered by papillæ, does not seem to be of any importance in regard to the classification of these species. On the contrary, Mazel has shown that species from the various sections of *Carex* may exhibit the same structure and position of stomata, and there does not appear to be any relation between stomata, being free or not, and the local environment, climate and soil. This author has shown, for instance, that *C. Davalliana*, *C. Pseudocyperus* and *C. hirta*, all from swamps, possess superficial stomata, while these are sunk in *C. paniculata*, *C. vulgaris*, *C. paludosa*, etc., from similar wet localities. In species which inhabit drier stations, woods and hills, a similar variation exists; thus the stomata are superficial and perfectly free in, for instance, *C. divulsa*, *C. brizoides*, *C. præcox*, *C. sylvatica*, etc., but not so in *C. maxima*, *C. glauca*, etc. A still more peculiar case may be illustrated by *C. misandra*, which we have examined from very different and remote localities: arctic and alpine regions of America and Europe, still maintaining the same structure and free position of the stomata, while in *C. capitata*, which we found growing almost side by side with this species, the stomata are protected by long and almost ramified papillæ.

It does seem as if the position of the stomata is not always an indication of the conditions of the surroundings under which the plants live, and it is, no doubt, one of the many inherited characters which are explainable only in a few species to which the natural surroundings are unchanged.

Utriculus.

The anatomical structure of this little organ appears to be very uniform in these species of the *Stenocarpæ*; it is very thin on account of the very little mesophyll, which forms here only one stratum between the two mestome-bundles, while it may be represented as two or three in the immediate vicinity of these. The outer wall of epidermis, on the dorsal face, is, however, thick in these species, while that on the ventral face is invariably very thin, like the radial walls. Stereome is also developed in the utricle of these species, but it is not very thick-walled and does not occur in groups of any considerable size either; it accompanies the two mestome-bundles, and it occurs besides as isolated groups between these, but not in all the species. In *C. misandra* and *C. frigida* the stereome was restricted to the mestome bundles, but in the other species we found from 5 to 14 isolated groups between these; the utricle of *C. brachystachys* and *C. ferruginea* appear to possess the strongest mechanical support of these species.

In comparing these anatomical details of the *Stenocarpæ*, the species examined seem to possess a very uniform structure in spite of the fact that they were collected from very remote stations, Alaska, Colorado, Middle-Europe and the arctic region, and in the case of *C. misandra*, the specimens from Alaska exhibited the same anatomical characteristics as those from the alpine slopes of Rocky Mountains (Colorado) and Switzerland besides from the arctic region. The internal structure considered by itself is very much in conformity with that of a number of arctic plants of various orders which we have had opportunity to examine, and which seems to indicate that these species lived under conditions that were influenced by a moist rather than by a dry atmosphere. The development of the *Stenocarpæ* seems thus to have taken place, at least, in recent time, in alpine or boreal regions, where they are yet in existence, and if we combine the geographical distribution, including the nature of the surroundings, with the external and internal peculiarities of these *Carices*, we see no objection to consider *C. circinata* and *C. lejocarpa* as "*formæ hebetatæ*" of this section.

Brookland, D. C., April, 1900.

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ART. XXVIII.—*Experiments on High Electrical Resistance:*
Part I; by OGDEN N. ROOD, Professor of Physics in
Columbia University.

IN making electrical measurements of high resistance, if the "direct deflection method" is used, under ordinary circumstances a resistance of 5000 megohms can be measured with some degree of certainty, and if the conditions are very favorable, 50,000 megohms may be reached. The "loss of charge method" is limited by the insulation of the condenser employed, this insulation itself being equal only to about 10,000 megohms.

In the method described in this paper 5000 megohms is small resistance, and it is not till 15,000 or 20,000 megohms has been reached that the manipulation becomes particularly simple. After that, the only limit is found when the resistance is so great that the electrometer contrived by me refuses to give any indications, even with electricity of very high potential. Just what this limiting resistance is, I cannot at the present moment state, being cut off by the dampness of the air from such determinations; meanwhile, it is certain that the figure is quite high.

In the older methods of measuring resistance, a flowing current of electricity is made to act on a system of astatic needles, but in the method here proposed the electricity is allowed to slowly accumulate, till it reaches a certain potential: in the older methods the electricity is treated as though its velocity were infinite, but the present plan is based on the fact that when the resistance is very great the electricity may require ten minutes or even an hour before raising the potential of the electrometer to the required degree. Thus, while in the older methods of work the electricity in its flow resembles a quick-moving incompressible fluid, in the one here described it may be compared to that of a thick pitchy fluid, moving slowly in a long open tank, and attaining its level at the farther end of the tank after the lapse of a longer or shorter time. After its level has been attained, if a leak is made anywhere in the floor of the tank, over this leak there will be a depression of the fluid; if the hole is made in the middle of the tank, some of the fluid drawn off and the hole then stopped, the pitchy fluid will flow to the depression from both sides, but it may take minutes before the old level is once more established.

Again, if we imagine the walls of the tank to be cold and the pitchy fluid to be at a higher temperature, some of it will

harden on the walls, and to a certain extent obstruct the flow of the fluid by contracting their cross-section. This illustration corresponds to a polarizing action which often sets in, and gradually diminishes the flow of the electricity, through or over a substance of very high resistance, and the greater the resistance the larger will be this counter action. What has been said so far applies to the cases of glass, mica, paraffin paper, silk and most artificial insulating substances.

If we now imagine our tank to be quite narrow, offering thus great resistance to the passage of the pitchy fluid, then at the end where the fluid is introduced, the height or potential will equal that of the supplying source, but as the distance from this end increases, the potential falls and finally becomes zero. This case corresponds to the surface conduction of electricity by black sealing wax, jade and gutta percha. They all furnish to the electrometer only small quantities of electricity, which can only be drawn from the neighborhood of the reservoir, and here the polarization is very marked.

Finally, imagine the walls of the tank to be still farther contracted and that the pitch-like fluid refuses to enter them in perceptible quantities, and we have the cases of ebonite, ordinary rosin and exfoliated mica (in very dry weather).

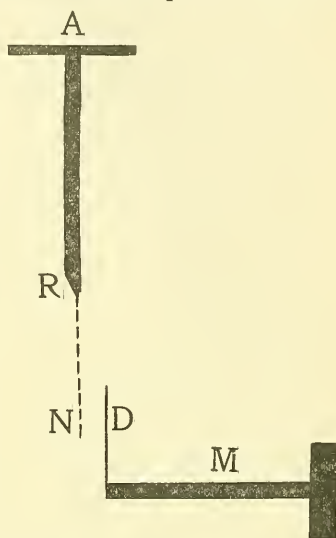
Evidently, if two tanks of the same length and cross-section deliver equal amounts of the pitchy fluid in equal times, then their conducting powers should be set as equal, and in general, if there be no loss in transit, the amounts delivered per minute, by tanks differing from each other, will be a measure of their relative conducting powers. Finally, if the tanks almost refuse to deliver up any of the pitchy fluid, still, if their potentials or the heights of the fluid at various points can be measured, in a manner analogous to that employed in the study of the conduction of heat by solid bodies that have reached the stationary condition, then it becomes possible to determine their relative conducting powers, and by comparison with somewhat better conductors, an estimate of their resistance in standard units can be formed.

Apparatus.—I modified a single leaf electroscope, so that it became a serviceable electrometer for this kind of work. Its plan may be seen in the diagram. A brass rod AR was supported on a bar of ebonite, which was placed on two rods of the same substance fastened on a base of ebonite. The rods and the bar were painted with melted rosin. The rod AR carried an aluminum leaf, bluntly pointed below, and attached above to the point R by a gold leaf hinge. Aluminum leaf was used, it having been found that under the action of the minute but very numerous electrical discharges, the gold leaf slowly wasted away at its corners, which it always presented to

the plate D with an undesirable twist. D is a plate of sheet brass faced with platinum foil, and receives the charge from the aluminum leaf when its potential has been sufficiently raised. The distance of this plate from the leaf can be varied by a micrometer screw, M, and thus the sensitiveness of the apparatus regulated. When a Leyden jar was used as the source of the electricity, the plate D was connected with the earth; when the street current

1

was employed, D was connected with one of the electric wires, a resistance of two or three thousand ohms being interposed, and as only thousands of megohms were dealt with, this amounted to nothing. An arrangement by which a slight vibration could be communicated to D was added, and by a single motion D was vibrated and the electrometer discharged by contact with A. When the electrometer gives a stroke the leaf usually remains attached to the plate D, which is the reason of the above mentioned contrivance. At A, various adjuncts can be placed, binding screws, etc. The whole arrangement amounts to a kind of unit jar which discharges itself when charged to a certain potential. Over it was placed a glass case to avoid air currents, but the case was in contact only with the base of the apparatus. Two such electrometers were made by me, one rather large for Leyden jar experiments; the other was smaller, especially in capacity, and was used only with currents of 110 to 112 volts. Some small electroscopes on the above mentioned plan were also made; they could be set to indicate certain potentials in bodies either in contact with them or at fixed distances from them. Their capacities were quite small. I also made a Bohnenberger electroscope with its dry pile, the delicacy of which could be varied at will by moving the brass plates connected with the terminals of the pile. It was used only in the study of good insulators. The source of electricity was sometimes a Leyden jar charged to a potential of 12,000 volts, but in the measurements given in megohms a current of only 110 to 112 volts was used. This same current was employed in the preparation of standard resistances. It is to be understood that the experiments were performed in dry winter



weather, the hygrometer having a range from 12 to 30 per cent of moisture. At the end of this paper certain determinations that were made in damp weather are added.

Preliminary experiments.—The case of a sealed glass tube 1^m 30^c in length, will illustrate partially what has been said about the behavior of electricity moving over a bad conductor. When the tube was made to establish connection between a Leyden jar charged to a potential of 12,000 volts and the electrometer, the electricity advanced slowly over its entire length, the electrometer not being at all affected for some time. In about five minutes it gave regular strokes; the tube had assumed the stationary condition and had nearly the same potential over its whole length, except near the electrometer, where for 3 or 4^{cm} it was lower. For half an hour it was allowed to deliver charges to the electrometer with much regularity, the rate being observed with a stop-watch. The electrometer was then connected with the middle of the tube; at first the rate of discharge was greatly increased, as the potential was higher here than at the drained end, and the electricity was flowing into the electrometer from both sides, but in about five minutes the flow had fallen nearly to the old rate and become quite regular. An air space of a centimeter was now made between the tube and the electrometer; the latter continued to give discharges, but more slowly; the potential of the middle part that had been somewhat drained was rising and acted inductively on the electrometer, and this process continued for some minutes, till the original level had been established. The electrometer was then again connected with the farther end of the tube and matters allowed to get into a stationary condition. The strokes having become regular, the middle of the tube was connected with the ground by wet paper and wire. The electrometer strokes however continued for a considerable time, but with diminishing frequency, till the supply of electricity at the farther end of the tube had exhausted itself. In this and in all similar experiments it is to be understood that the two ends of the conducting body were provided with metallic armatures, tinfoil fastened with thread. A great many experiments were made on conductors of this class without any armatures at all; the tinfoil however facilitates matters, and makes the flow more regular.

All this illustrates the case of a glass tube with rather good insulating properties, and I pass on now to more ordinary samples. A sealed tube of hard Bohemian glass, 1^m 30^c in length, was washed in water, dried in sunshine and connected with jar and electrometer. In a second or two the electrometer began its discharges, the rate being about three seconds per stroke. The entire tube was then heated quite hot with a

Bunsen burner; for two minutes the flow of the electricity was stopped, but it then began again, the intervals between the strokes being 8, 7, 4, 3, 3, 3 seconds. Two days afterwards the experiment was repeated with a result of 10 strokes in 32 seconds. Ten days later, its rate was 10 strokes in 42 seconds. A sealed tube of soft German glass of the same length that had been carefully cleaned and heated gave 10 strokes in 20 seconds; heated again for three minutes and cooled, its rate was 10 strokes in 36 seconds and on the following day this had fallen to 10 in 15 seconds. Ten days afterwards it gave 10 strokes in 20 seconds. These slight irregularities are however thrown into the shade by an observation made on a tube of the same kind of glass, where the resistance of one-half of it was twelve times as great as that of the other, and it remained in this condition for two weeks, or until it was cleaned; then the whole of the tube had the lower resistance. A tube made from the glass that is now used in Germany for Leyden jars, was found to have a high resistance, so that the electricity did not seem to reach its farther end at all. After being in contact for three-quarters of an hour with a Leyden jar having a potential of 12,000 volts, at a distance of 16^{cm} from the jar, the potential of the tube was only 600 volts, and an additional exposure for the same length of time extended this distance only to 22^{cm}. At the time these experiments were made I had not constructed a set of units of high resistance, or it would have been easy to have measured the resistance of these tubes in megohms when traversed by a current of 12,000 volts. Later on I measured the resistance of a tube of soft German glass, 27^{cm} in length, which had been for some days in a box dried with sulphuric acid. This at first proved to be 250,000 megohms, then under the influence of the current it increased to 380,000 megohms. On this day the hygrometer indicated 43 per cent of moisture. Some time afterwards, the hygrometer indicating 53 per cent, the resistance of the tube had fallen to 180,000 megohms. The current here employed, as in all cases where actual figures are given, had a potential of only 110–112 volts.

Silk.—A double thickness of rich brown silk, 45^{cm} long, 2^{cm} broad, conducted electricity at a potential of 12,000 volts very well; the electrometer strokes began immediately, and continued hour after hour. A band of white silk was then taken, a meter in length and 3^{cc} in breadth; the discharges were at the rate of about one per second. The silk was then boiled twice in water, and dried, without affecting its rate. It was then boiled in alcohol and dried; the discharges were a little slower at first, one in four seconds, but the rate soon increased to one in two and a half seconds. Two threads of white sewing silk

failed to affect the electrometer, but with the electroscope it was found that the charge had extended over 15^{cm}. When twenty-four threads of this silk, half a meter in length, were used, the electrometer gave slow strokes needing three minutes each. A single thread of this silk was heated in a bath of one part gutta perch, three parts rosin; a large quantity of vapor bubbles escaped, and it became a good non-conductor. When it was connected with the Leyden jar, charged *positively* to 12,000 volts, not only did it not conduct, but became *negatively* electrified, with a potential of about 200 volts. Negative electrification under such circumstances will be referred to later on.

Paraffin paper.—Ordinary commercial paraffin paper, viz: writing paper which had been passed through melted paraffin wax, conducted quite well. A strip a meter in length, in fifteen minutes had acquired over its whole length a potential of about 12,000 volts. Some measurements were made of short strips of this paper provided with metallic armatures, a potential of only 110 volts being employed.

Length.	Breadth.	Resistance in megohms.
25 ^{mm}	18 ^{mm}	80,000
63 ^{mm}	“	125,000
132 ^{mm}	“	300,000

It is evident that the resistance of the connecting armatures has influenced these results.

Mica.—Using a potential of 12,000 volts, a strip of mica (muscovite) 30^{cm} long and 2^{cm} broad, was compared with a strip of paraffin paper of the same dimensions; the resistance of the mica was sixty times that of the paraffin paper. The hygrometer at the time indicated 30 per cent of moisture. Two strips of mica from the same specimens were prepared; the length of each was 27^{cm} the breadth 2^{cm}. Both were heated till their edges had become somewhat white, and on the following day their relative resistances were determined; these turned out to be nearly as 1 to 2. Knowing that mica when heated gives off traces of water and fluorhydric acid, a piece of this mica was heated in a Bunsen flame till its structure was destroyed, and on the following day, having been provided with a tin-foil armature, it remained for twenty minutes in contact with a Leyden jar charged to a potential of 12,000 volts. It was then removed and presented to an electroscope capable of indicating 100 volts, but failed to show any charge.

In the case of all the substances that thus far have been mentioned except two, it was found that when they were connected with the Leyden jar, that within at least fifteen minutes the

part farthest from the jar had assumed its potential, provided that no electricity was conducted off from it. This is not true of the substances that follow; their case is similar to that of rods conducting heat and slowly attaining the stationary condition.

Jade, gutta percha, black sealing wax.—There was placed at my disposal a fine bar of jade, 30^{cm} long with a square cross section of 12^{mm}. This was simply washed in water and sun-dried. Afterwards it was exposed to 12,000 volts for fifteen minutes. The electricity had traversed its whole length with a diminishing potential, that of the farther end being only 600 volts. This would have been much more than enough to have furnished electrometer strokes, but the rate of travel and the rate of delivery was so slow, that the electrometer was not affected. The following experiment was then made with the jade, and also later, with black sealing-wax and gutta percha. The end of the jade in contact with the Leyden jar had of course a tin-foil armature to facilitate conduction from the jar to the surface of the jade; at a distance of two centimeters from this armature the jade was wrapt with a strip of tin-foil, and this was brought into metallic connection with the electrometer placed at some distance. This arrangement assumes that the electricity will leave the Leyden jar, traverse two centimeters of naked jade, and then reach the electrometer by a connecting wire, and that no electricity will pass through the air by convection, and reach the movable conducting strip of foil and hence the electrometer. This supposition appeared to be practically justified by repeated experiments, where the knob of the Leyden jar retained its short distance from the movable armature, but was disconnected from the fixed terminal armature of the jade. I repeatedly watched for ten minutes at a time without finding that the disconnected jar produced any effect at all on the electrometer. This fact makes it possible to study cases which otherwise would offer difficulty. Using two centimeters of naked jade in this way, the electrometer furnished strokes needing ten seconds. Using four centimeters, the time of the first stroke was not 20 seconds but 47 seconds, the next following stroke needing 54 seconds, indicating the polarization usually observed in cases of surface conduction in connection with high resistance. Gutta percha and black sealing wax were now treated in exactly the same way; the movable armature being at 4^{cm}; they both refused to affect the electrometer, but with a distance of 2^{cm} the times needed for electrometer strokes were as follows:

	Gutta percha.	Sealing wax.
1st stroke	1' 20"	2' 2"
2d " "	2' 9"	5'
3d " "	2' 25"	no stroke.

On this day there was 30 per cent of moisture in the air.

In all the cases thus far examined the substances in question conducted sufficiently well to affect the electrometer arranged with a medium sensitiveness, but there were other insulators having still higher resistances, at least in dry weather.

Ebonite.—One of the most interesting of these is ebonite or hard rubber. It being very doubtful according to my previous experiments whether ebonite conducted electricity at all in dry weather, a strip of this material, 40^{cm} long and 1.5^{cm} broad, was provided with tin-foil armatures and then tested with the Bohnenberger electroscope for accidental charges, and afterwards connected for half an hour with a jar charged positively to 12,000 volts. It was then removed from the jar, the tin-foil discharged, and passed over the Bohnenberger electroscope. It was found to be *negatively* charged for about 17^{cm} from the jar end, the potential being a little under 110 volts, and gradually rising near the jar end, probably not much above 200 volts. This negative charge, I take it, is probably due to a separation of the ions, owing to the prolonged inductive action of the strong field. The final result is, that if the ebonite conducted at all in the regular way, such conduction was masked by the inductive effects.

Gutta percha.—Exactly the same experiment was made with a similar strip of gutta percha: it showed a charge of positive electricity that extended 16^{cm} from the jar-end of the strip; beyond this the electrification was negative, with a potential of from 100 to 200 volts. There seemed to be a neutral zone dividing the two opposite electrifications.

Amber.—The amber was 3^{cm} in length; one-half of it was wrapt in tin-foil fastened down by thread. Being supported by *insulating* silk thread, it was left for half an hour in contact with the knob of a Leyden jar charged positively to 12,000 volts. On removal and discharge of the foil armature, it was found charged positively, with a potential of 300 or 400 volts.

Rosin.—A piece of freshly cut rosin of the same dimensions and treated in the same way, became negatively electrified. Other experiments given below tend to show, that taking dry and damp weather together, plain rosin is the best insulator that was examined.

Insulating silk thread.—Ordinary thick white sewing silk was passed through a bath of one part of gutta percha melted

with three of rosin. Four centimeters of this was treated as above and became feebly *negatively* electrified, with a potential of about 200 volts. The temperature of this bath is rather high, and one composed of wax and rosin probably would answer all practical purposes.

Is the conduction of electricity by glass entirely on the surface, or does the electricity penetrate a little below the surface?—One experiment bearing on this question is here given: a sealed tube of soft German glass conducted quite well; a portion of it was heated till it melted black sealing wax, and in this way a few centimeters of it were coated with the wax, but when it was cold it conducted about the same as before. An application of boiled linseed oil produced no effect on its conducting power; finally, a few centimeters of the surface were ground away on a stone, but after washing, drying and the application of heat, it conducted just as before.

Experiments in damp weather.—The results of a few experiments on the effects produced by dampness of the air on the resisting power of certain insulators may be of interest, and some are here given.

Ebonite.—Hygrometer at 78° ; a piece of the same sample previously mentioned was used; the distance from armature to armature was 17^{mm} ; the breadth 13^{mm} , thickness 2^{mm} ; it had been exposed to the air for several days. Employing 110 volts at first, it conducted quite well, its resistance being about 5000 megohms. It was then warmed, and while cooling showed a resistance as high as 1,200,000 megohms. In three minutes this fell to 15,000 megohms and then remained constant. Two hours later, the same figure was obtained.

Gutta percha of the same dimensions with the ebonite, without being warmed or dried in any way, had a resistance of 12,000 megohms.

Rosin, with armatures of tin-foil, the distance between them being only 9^{mm} , had a resistance of 1,500,000 megohms. It had been exposed to the damp air for days, and was not warmed or dried. These experiments were made on the same day, but on the following day when the air was somewhat drier, the hygrometer standing at 65° , the resistance of the ebonite was 340,000 and that of the gutta percha 440,000 megohms. The resistance of the rosin was too great to be measured without altering the distance of the plate P, which at the moment was undesirable. Freshly cut natural beeswax also appears to be a good insulator in damp weather. A few days later, the hygrometer standing at 60° , the resistance of the ebonite was again measured and turned out to be 1,100,000 megohms. It had in the meanwhile been exposed to the air

which had been damp. Wishing to improve the resisting power of ebonite against moisture, I painted a similar piece with boiled linseed oil, and allowed it to dry for two weeks. It was not improved as shown by my electrometer, and to avoid all uncertainty, I requested Mr. H. C. Parker to measure its resistance by the "direct deflection method" with a galvanometer; it turned out to be only 5300 megohms. It was then gently heated by a spirit lamp and while warm measured; the resistance had been raised to 90,000 megohms; it fell at first quite rapidly, and afterwards slowly, finally reaching, in four minutes, 6000 megohms. It was then heated to a temperature of 100° C. for four hours, and after free exposure to the air for three days, it was found that its resistance had been increased fifteen fold, the hygrometer standing at 54° when it was measured.

One more experiment may be given; on a somewhat damp day the hygrometer stood at 56°, the thermometer at 80° F.; pieces of mica and thin ground glass were prepared with metallic armatures; the pieces were 13^{mm} broad and the distance from armature to armature was 55^{mm}. The resistance of the glass was found to be nearly one thousand megohms, that of the mica 50,000 megohms.

In the preliminary measurements that have been detailed in this article, it has been assumed that during the transit of the electricity there is practically no loss by convection nor by leakage in the electrometer, so that the time consumed by a given charge in traversing a conductor and charging the electrometer to a given potential, would be directly proportional to the resistance of the conductor, and the few experiments that I have been able to make seem to justify this conclusion, but my set of high resistances was not completed in time to give it a rigorous examination. In dry weather the Leyden jar lost only two-thirds of its charge in twelve hours, although its necessary construction was far less favorable for insulation than was the case with all of the other arrangements.

The measurements given in this article were made with the help of certain standard units that I have recently constructed on a new plan. A considerable number were made, the resistance in each case not exceeding 5000 megohms. These were measured separately by Mr. H. C. Parker with a galvanometer and the "direct deflection method." Afterwards, they were joined together, and furnished suitable standards.

It may finally be remarked that almost all of the experiments described in this article are quite easy of execution, and many of them are well adapted for class illustrations.

June 18th, 1900.

ART. XXIX. — *On two new Occurrences of Corundum in North Carolina*; by JOSEPH HYDE PRATT.

WHERE formerly corundum was supposed to be rare in its occurrence and to be found in quantity only in the basic magnesian rocks, it is now known to occur in various types of rocks and in quantity, in syenites, gneisses and schists. Many new occurrences have been discovered during the past few years, some of which give indications of being of considerable economic importance, while others are only of scientific value. In the present paper two new occurrences are to be described, that have been observed in North Carolina, one in an amphibole-schist and the other in a quartz-schist.

Corundum in Amphibole-Schist.

At the Sheffield mine in Cowee township, Macon County, North Carolina, corundum has been mined in a saprolitic rock at various times for a number of years. While sinking a shaft eight feet square to penetrate the depth of the corundum-bearing saprolite, the solid unaltered rock was encountered. The shaft was 87 feet deep and showed the following sequence downward. The first 12 feet was through the saprolitic rock, in which there were seams, a few inches wide, of kaolin; the next two feet were corundum-bearing; from 14 to 28 feet the same saprolite was encountered and then another two feet that was corundum-bearing, followed by another ten feet of the saprolite and two more feet of the corundum-bearing rock; from 42 to 65 feet the rock began to be less decomposed and from 63 to 66 feet another seam that was corundum-bearing was encountered. From this point the rock became more and more solid, until at 77 feet the fresh rock was encountered. These various seams in the rock are very pronounced and are dipping 30° toward the west near the top, but become nearly horizontal nearer the bottom of the shaft. The seams of decomposed feldspar observed near the top of the shaft become less and less kaolinized downward, until in the solid rock the seams are of a pure plagioclase feldspar. In the hard rock exposed there are two seams of corundum similar to those above, although in the fresh rock the corundum seams are not as pronounced as in the saprolitic rock. There is often considerable of the feldspar bordering the seams of corundum. The general trend of the rock is about N. $5-10^{\circ}$ E.

From what could be seen of the solid and the saprolitic rocks the corundum occurs in seams a few feet in width at intervals in the rock, and while the corundum may be ten or more per cent in these veins, its percentage in the rock, that it would be necessary to mine, would not probably be over three or four. The actual width of the dike is not known, but the saprolitic

rock has been cut across for nearly 100 feet in a direction about at right angles to the strike.

The fresh rock at the bottom of the shaft is somewhat varied in appearance and while it all does not show any definite gneissoid structure, the more finely divided portions are distinctly so. There are streaks, a few inches thick in the rock, that are composed almost wholly of a plagioclase feldspar. Some portions of the rock are decidedly porphoritic and contain phenocrysts of a light gray amphibole, a centimeter in diameter, in a groundmass of feldspar. A large part of the rock is made up, however, of small roughly-outlined prismatic crystals of an amphibole, probably hornblende, and irregular fragments of plagioclase feldspar. The hornblende is almost black in color but in thin splinters it has a bronze luster and a deep resinous color. Biotite of a deep brown color occurs sparingly, and a pink garnet is rather abundant. It is this part of the rock that is of a gneissoid structure and in which the corundum occurs. The corundum is of a light to a purplish pink color and in nodules up to two or three centimeters in diameter. There are some streaks in the rock that are very highly garnetiferous, composed essentially of the garnet and plagioclase feldspar, or of the garnet and biotite. Chalcopyrite occurs very sparingly in these portions of the rock. Small particles of graphite have been observed in the coarsely crystallized portions.

Prof. L. V. Pirsson has kindly made a microscopical examination of this rock, the results of which are embodied in the following paragraphs.

In thin section the microscope disclosed the minerals, hornblende, plagioclase feldspar, garnet, biotite, muscovite, staurolite and rutile. Hornblende is the most common, forming about two-fifths of the section, while of the remainder, plagioclase and garnet occur in about equal quantities and the others in comparatively insignificant amount.

“The hornblende is formless but tends to irregular columns almost invariably extended in the plane of schistosity; it has very rarely a somewhat stringy tendency in its cleavage but is usually homogeneous in broad plates. Its color is a clear olive-brown and it is somewhat pleochroic but not strongly so. It is everywhere dotted by the small grains of garnet, which rarely show good crystal form. The garnet occurs associated also with the plagioclase.”

“The plagioclase occurs twinned according to the albite law only. In sections perpendicular to 010, the lamellæ show extinctions as great as 30° and the plagioclase is therefore rich in lime and as basic as labradorite, which it probably is. It shows strong evidence of shearing movement in the rock; it is

often broken, exhibits rolling extinctions and the albite lamellæ are curved and bent. It runs along the planes of schistosity between the feldspars and forms a mosaic of angular broken grains."

"Staurolite was found in rather broad irregular grains, and the rutile in small irregular grains and well crystallized prisms."

Prof. Pirsson has indicated that the character and structure of this rock, composed chiefly of amphibole, labradorite and garnet, suggests most strongly that it is a metamorphosed igneous rock of the gabbroid family. During metamorphism, the augite of the gabbro would be converted into the brown hornblende and any iron ore that was present would be taken up by the hornblende and garnet. The rutile would have resulted from the titanitic acid that is a regular component of the iron ores in these gabbro or diabase rocks. Staurolite is rather naturally expected, as it is usually a mineral of metamorphism, and its natural home is in the schistose rocks. The feldspar has suffered the least (except the corundum) chemically and shows only the shearing of dynamic processes.

The corundum does not occur in crystals but in small fragments and in elongated nodules, which are cracked and seamed, and appear to have been drawn out by the shearing processes. The general character and shape of the corundums would indicate that they were original constituents of the igneous rock and were not formed during its metamorphism.

The exact classification of this rock is not easy, but it will probably be nearer correct to bring it under the head of an amphibole-schist.

Corundum in Quartz-Schist.

In the crystalline rocks of the southwestern part of North Carolina and the northeastern part of Georgia an interesting occurrence has recently been observed, namely that portions or bands of these are corundum-bearing. These corundum-bearing bands are first encountered on the head waters of Tallulah river, in the northern part of Rabun County, Georgia, and can be followed in a northeasterly direction to the Yellow Mountain in Clay County, North Carolina. They are near the top of the Blue Ridge, at an elevation of from 3000 to 4000 feet.

The composition of these rocks vary from those that are a normal gneiss to those that contain no feldspar and can best be described as quartz-schist, composed of biotite mica and quartz. Some portions of the rock are rich in garnet, while others are almost entirely free from this mineral, and occasionally there are small bands of white quartz. They are distinctly laminated and are frequently intersected by granitic dikes, some of which are coarsely crystallized and of a pegmatitic character, that are often parallel with the beddings of the schists, although many

of them are cutting irregularly through them. Where these dikes are parallel to the bedding of the schists, the laminated structure of the latter is more apparent. The general strike of these crystalline rocks is N.E.-S.W. and with a dip of about 30° to the N.W.

Portions, or bands, of these schists are corundum-bearing, but they are irregularly defined and gradually merge into the normal rock. They have a similar relation to the normal schists that the garnet-bearing bands of a gneiss have to the normal gneiss in which they occur. They are not veins in any sense of the word, but are simply portions of the same mass of crystalline rocks in which corundum occurs as a constituent of the rock. These bands vary in width from a foot or two to 12 or 15 feet, but in these wider ones the corundum-bearing portion is not continuous but is intercepted by streaks of barren rock and granitic dikes.

These bands can be traced for a distance of five or six miles in a N.E.-S.W. direction, sometimes outcropping continuously for nearly a mile. There are at least two of these corundum-bearing bands which are parallel to each other and about two miles apart. The only variation that has been observed in them is the percentage of corundum and garnet, otherwise they are identical. The percentage of corundum is never high, and from determinations made on samples from various parts of the deposits, it varies from two to five per cent.

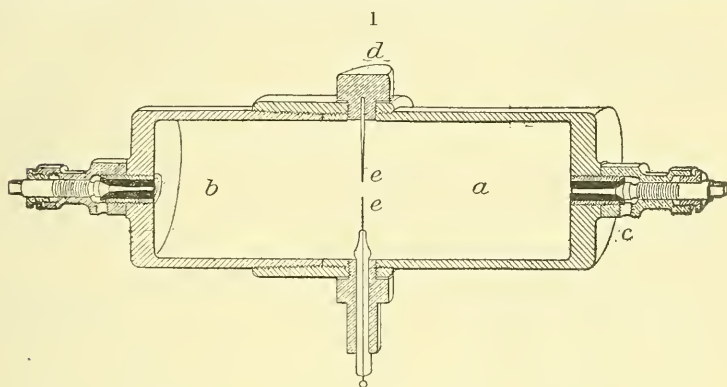
The corundum occurs for the most part in small particles and fragments that have no definite shape and are of a gray, white, and bluish-white color to almost colorless. It is also in crystals from minute ones to some that were observed two and a half inches long and about one-half an inch in diameter, which are usually fairly well developed in the prismatic zone.

It is probable that these schists are the result of the metamorphism of sandstones and shales formed from alluvial deposits of many thousand feet in thickness, that were formerly the bed of the ocean. By lateral compression these have been folded and raised into the mountain ranges of this section. That these were much higher than at the present time is very evident from the granitic dikes that are of deep-seated origin. By decomposition and erosion the mountains have been worn down to their present condition, thus exposing the schists in contact with granitic dikes which have aided in their thorough metamorphism. The shales were rich in alumina which not improbably was in the form of bauxite, and during their metamorphism the excess of the alumina crystallized out as corundum. This mineral has crystallized out along the planes of lamination so that during the subsequent weathering of the rock the corundum has been left in knotty nodules, studding the surface of the rock, giving it the appearance of containing a high percentage.

ART. XXX.—On the Products of the Explosion of Acetylene, and of Mixtures of Acetylene and Nitrogen; by W. G. MIXTER. Second Paper.

[Contributions from the Sheffield Laboratory of Yale University.]

THE bomb described in the first paper* was again used in the work unless otherwise stated. The figure of it is here reproduced (fig. 1) for convenience. The acetylene gas for the experiments was made by thrusting lumps of calcium carbide through the tubulure of a glass gas holder previously filled with water. The gas was dried as before by passing it through



a cylinder containing a kilo of small sticks of caustic potash. The determinations of acetylene were made as follows: The gas to be tested was passed into a eudiometer through a stop cock at the top of it. The acetylene was absorbed by an ammoniacal solution of cuprous chloride and the ammonia was finally washed out with water. The temperature and pressure of the gas remaining after the absorption of acetylene was made the same as before. This method is applicable only in absence of oxygen and carbonic oxide. The latter will be present when the acetylene exploded was mixed with water vapor or oxygen. Repeated tests for carbonic oxide failed to show the presence of that gas in the products of the acetylene exploded.

Experiment 36.—Gas, 99 per cent of acetylene; pressure, 2 atmospheres. The explosion was prompt. It extended through a brass tube $1\frac{1}{2}$ mm in diameter and 1 meter in length to the manometer. The determinations of acetylene in the residual gas gave 4, 4.3 and 3.7 per cent.

* This Journal, vol. ix, 5.

Experiment 37.—Gas, 98 per cent of acetylene; pressure, 1289^{mm}; temperature, 12°; electrodes, 2^{mm} apart. The gas was sparked repeatedly for several seconds. Carbon separated about the electrodes, but there was no explosion and the gas after the sparking was found to contain 98 per cent of acetylene.

Experiment 38.—Gas, 98 per cent of acetylene; pressure, 1312^{mm}; temperature, 15°. The electrodes were 3^{mm} apart. There was no explosion and no change in pressure during the sparking.

Experiment 39.—Gas, 96 per cent of acetylene; pressure, 1472^{mm}; temperature, 14°. The pressure after the explosion at 14° was 1425^{mm}, showing a condensation of 3.2 per cent. Two determinations of acetylene in the residual gas gave 3.4 and 3.3 per cent.

Experiment 40.—Gas, 98 per cent of acetylene; pressure, 5.1 atmospheres, which was the same after the explosion when the bomb had attained the temperature of the water in the tank in which it was immersed. The gas tested was taken before the pressure in the bomb had fallen below 4 atmospheres. Three tests yielded 2.3, 2.4 and 2.4 per cent of acetylene. The carbon nearly filled the bomb.

The nitrogen used in subsequent experiments was prepared from sodium nitrite and ammonium chloride. The mixtures of acetylene and nitrogen were dried as described by caustic potash.

Experiment 41.—Gas, 89.3 per cent of acetylene and 10.7 per cent of nitrogen; pressure 3.2 atmospheres before and after the explosion. Two estimations of acetylene in the gas after the explosion gave 1.4 and 1.4 per cent. In order to determine the amount of hydrocyanic acid, the gases (933^{cc} reduced to 0° and 760^{mm}) from the bomb were passed slowly through a dilute solution of pure potassium hydroxide. This solution was then titrated with a decinormal solution of silver nitrate, requiring 5.2^{cc} to produce a permanent precipitate of silver cyanide. This corresponds to 0.028 gram of hydrocyanic acid and 23^{cc} of the compound in the gaseous state. The bomb had a capacity of $1\frac{1}{10}$ liters, but for convenience the result may be given as the amount of hydrocyanic acid in a volume of 1 liter in the bomb. It was 0.083 gram and 63.3^{cc}. The potash solution after the titration was acidified with hydrochloric acid, boiled, cooled, then an excess of potash was added and the solution was again boiled. The escaping vapor was found to be free from ammonia. Nor was any found in another test similar in all respects except that the titration with silver nitrate was omitted.

Experiment 42.—Gas, 85 per cent of acetylene and 15 per cent of nitrogen; pressure, 3061^{mm}; temperature, 17·8°. The pressure after the explosion when the gas had cooled to 17·8 was the same as before the explosion. Each of two tests for acetylene in the residual gas gave 0·25 per cent. The residual gas was passed slowly through an absorbing tube containing a dilute solution of potassium hydroxide until the pressure in the bomb had fallen to that of the atmosphere; 20^{cc} of a decinormal solution of silver nitrate were used in titrating the solution of potassium hydroxide. Next 10^{cc} of a concentrated solution of potassium hydroxide were poured into the bomb to absorb the hydrocyanic acid remaining in it. The mixture of carbon and potassium hydroxide was thoroughly shaken and then washed on a filter. The filtrate required 14·4^{cc} of a decinormal solution of silver nitrate, making the total amount of the solution of silver used 34·4^{cc}, corresponding to 0·1858 gram of hydrocyanic acid, or 0·1688 gram and 139·6^{cc} (at 0° and 760^{mm}) per liter in the bomb.

Experiment 43.—Gas, 76·6 per cent of acetylene and 23·4 per cent of nitrogen; pressure 3818^{mm} at 14·8°. 1·7 per cent of acetylene was found in the first portions of gas drawn from the bomb and 2·5 per cent when the pressure was little more than that of the atmosphere. The attempt to estimate the hydrocyanic acid remaining in the bomb was unsuccessful. An alcoholic solution of potash was used, as this wets the finely divided carbon better than water. The filtered solution became turbid on adding silver nitrate and the titration was so unsatisfactory that the result was rejected. The solution of potassium hydroxide used to absorb the gas drawn from the bomb required in titrating 32·3^{cc} of a decinormal solution of silver nitrate corresponding to 0·1747 gram of hydrocyanic acid. As $\frac{4}{5}$ ths of the gas in the bomb were taken for this determination we may assume that $0·1747 \times \frac{5}{4} = 0·2187$ gram was the total amount of hydrocyanic acid in the gas in the bomb and not condensed on the carbon. This result calculated for a volume of 1 liter gives 0·200 gram and 165^{cc} (at 0° and 760^{mm}) of hydrocyanic acid. The gas after the explosion was alkaline to litmus and the solution of potassium hydroxide gave off ammonia on warming.

Experiment 44.—Gas, 87 per cent of acetylene and 13 per cent of nitrogen. The nitrogen was collected over an alkaline solution of pyrogalllic acid to absorb traces of oxygen, and then was mixed with the acetylene over a solution of ferrous sulphate in order to free the gases from possible traces of ammonia. As the bomb leaked after the explosion, it was only possible to determine approximately the ratio between the

quantities of hydrocyanic acid and ammonia formed. The potassium hydroxide solution through which the gaseous products of the explosion were passed required 5^{cc} of a decinormal solution of silver nitrate to form a permanent precipitate of silver cyanide, 5^{cc} more of the silver solution were added in order to remove the cyanides from the solution, and the precipitated silver cyanide was filtered off. The filtrate was then cautiously distilled into hydrochloric acid. The acid solution was finally evaporated to dryness. The residue was 0.0035 gram of ammonium chloride, equivalent to 0.001 gram of ammonia. As 5^{cc} of the silver solution correspond to 0.027 gram of hydrocyanic acid, the ratio between the ammonia and hydrocyanic acid formed is apparent.

In order to determine whether or not the gas after the explosion contained cyanogen, the alkaline solution from which ammonia was distilled was acidified with hydrochloric acid, boiled, then made alkaline with potassium hydroxide. This last solution gave when distilled no ammonia. If cyanogen had been present in the gas, there would have been formed in the solution of potassium hydroxide through which the gaseous products of the explosion were passed, cyanide and cyanate. The latter salt, as is well known, yields an ammonium salt when treated with an acid.

Experiment 45.—The mixture of acetylene and nitrogen containing 28 per cent of the latter gas was allowed to stand 24 hours over a solution of ferrous sulphate mixed with ferrous hydroxide in order to free the gas from oxygen and possible traces of ammonia. The pressure of the gas before the explosion was 3781^{mm}, a little less than 5 atmospheres, at 6.4°. The gas drawn from the bomb after the explosion when the pressure was nearly 5 atmospheres contained 3.2 per cent of acetylene, and that portion drawn at 1.3 atmospheres contained 5.6. The hydrocyanic acid estimated as in previous experiments was found to be .170 gram and 140^{cc} per liter in the bomb. The ammonia found in the same volume of gaseous products was .0047 gram. Here we have a ratio of 1 of ammonia to 36 of hydrocyanic acid in the products of the explosion.

Experiment 46.—The mixture of acetylene and nitrogen containing 10 per cent of the latter was purified as in the previous experiment. It was sparked when at a pressure of 3856^{mm}. The first portion of the gas drawn from the bomb was found to contain 1.7 per cent of acetylene and that taken at a pressure of 2 atmospheres 2.2 per cent. The hydrocyanic acid found corresponded to 0.145 gram and 120^{cc} per liter in the bomb. Likewise the ammonia was 0.0046 gram. Here the ratio of the ammonia to the hydrocyanic acid is 1 to 30.

Experiment 47.—Two liters of a dried mixture of acetylene and nitrogen containing 10 per cent of the latter gas were passed through a combustion tube which was heated to a temperature sufficient to decompose the acetylene. The gaseous products were passed through a solution of potassium hydroxide to absorb hydrocyanic acid. But not a trace was found by the prussian blue reaction. The same result was obtained a second time.

Experiment 48.—A mixture of acetylene and air was exploded in an open liter jar. Much soot separated and the flame passed slowly into the jar. No hydrocyanic acid was detected in the products, and none in several repetitions of the experiment.

Experiment 49.—The sides of a jar were moistened with strong ammonia water and the air in the jar was mostly displaced by acetylene. On applying a flame to the mouth of the jar there was a slight puff and an abundant separation of soot in the jar. Hydrocyanic acid was found in the products of four experiments, and it was also formed rather abundantly when illuminating gas was substituted for acetylene.

Experiment 50.—Acetylene was passed through concentrated ammonia water, then through a tube heated to dull redness by one burner. Hydrocyanic acid was formed.

The experiments 19 to 28 described in the first paper (*loc. cit.*) were made in U-tubes having a diameter of 17 to 20^{mm}. The residual gas contained much acetylene and the amount of condensation products was considerable. In order to learn if the character of the products is due to the expansion of the gas when exploded, the next two tests were made in an iron U-tube 7.5^{cm} in diameter. To the ends of this tube were attached the parts *a* and *b* of the iron bomb (fig. 1, p. 299). The apparatus was designed for observing the effect of explosions under the same conditions as in the bomb except that the gas was not subjected to constant volume. About 50 kilos of mercury were required in the apparatus and 1800^{cc} of dry acetylene measured at atmospheric pressure were used. The U-tube when charged for an explosion was hung in a tank of water in order to cool it quickly after an explosion. When acetylene is fired in such a U-tube the products cool more rapidly than in the bomb.

Experiment 51.—Gas, 99 per cent of acetylene; pressure, 2060^{mm} at 18°. The electrodes were 2^{mm} apart. The explosion was prompt and the pressure after it was 2097^{mm} at 18°. Two determinations of acetylene in the gas after explosion gave 3 per cent.

Experiment 52.—Gas, 99 per cent of acetylene; pressure, 2300^{mm}. The electrodes were 4^{mm} apart. Two estimations of acetylene in the gas taken at 3 atmospheres gave each 1.5 per cent, and when the pressure in the apparatus had fallen nearly to that of the atmosphere the portions tested were found to contain 1.9 and 1.8 per cent of acetylene.

The next two experiments were made in the iron U-tube in order to find the effect that the expansion accompanying the explosion has on the formation of hydrocyanic acid. The mixtures of acetylene and nitrogen were kept for a day over a solution of ferrous sulphate in which was suspended ferrous hydroxide.

Experiment 53.—About 1.5 liters, measured at atmospheric pressure, of gas containing 15 per cent of nitrogen were used. The mixture was sparked when under a pressure of 2.5 atmospheres. The residual gas contained 1.25 per cent of acetylene. 800^{°c} of it were passed through a solution of potassium hydroxide. This was then titrated with a decinormal solution of silver nitrate, requiring 0.4^{cc} to produce a permanent precipitate, a result not indicating satisfactorily the presence of cyanides.

Experiment 54.—The mixture contained 13.5 per cent of nitrogen. Two liters of gas were taken and subjected to a pressure of 2.5 atmospheres. The explosion was prompt. The estimation of acetylene in the residual gas was not satisfactory but it showed there was less than 3 per cent present. 1000^{°c} of the gas left by the explosion were passed slowly through a solution of potassium hydroxide. One drop of a decinormal solution of silver nitrate gave to the solution a permanent brown precipitate. Next, 200^{°c} of the gas were passed through a solution of potassium hydroxide, made from the metal. Finally a few drops of a solution of ferrous ferric sulphate were added and the mixture was warmed, cooled, and made acid with hydrochloric acid. No prussian blue was formed, a proof that no hydrocyanic acid had been formed.

These results in the iron U-tube prove conclusively that the large amounts of acetylene and condensation products in the earlier experiments in glass tubes are not attributable to expansion at the time of the explosion.

In order to learn whether the peculiar explosive wave of experiments 27 and 28 of the first paper can be transmitted to gas in a larger vessel and there produce the same effect, the next two experiments were made in the apparatus shown in fig. 2. The tube *a* is glass and has a diameter of 18^{mm}.

Experiment 55.—Gas, 98 per cent of acetylene; pressure, 2.6 atmospheres. The electrodes were 3^{mm} apart and less than 1^{mm} from the glass. The sparking at first produced a filament of carbon between the electrodes. A much stronger secondary current was then applied and after some seconds there was a sudden glow and carbon was deposited abundantly in a length of 10^{cm} of the tube. The explosion did not extend to the gas in the bomb. The residual gas contained 97.5 per cent of acetylene.

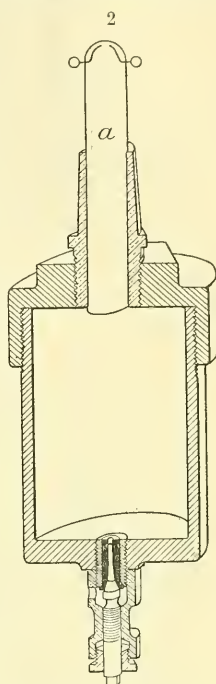
Experiment 56.—Gas, 98 per cent of acetylene; pressure, 2316^{mm} (3.3 atmospheres) at 15°. The electrodes were 4^{mm} apart and 3^{mm} from the glass. A momentary closing of the primary circuit produced an explosion. The pressure when the gas had cooled to 15° was the same as before the explosion. There were in the glass tube brilliant black rings of carbon separated by spaces less densely coated. Two estimations of acetylene in the residual gas taken at 2.7 to 3.3 atmospheres gave each 1.3 per cent, and two portions taken when the gas in the bomb was at 1.5 atmospheres and less gave 1.7 and 1.6 per cent.

If the peculiar explosive wave started in the tube extended to the gas in bomb it caused almost complete decomposition of the acetylene.

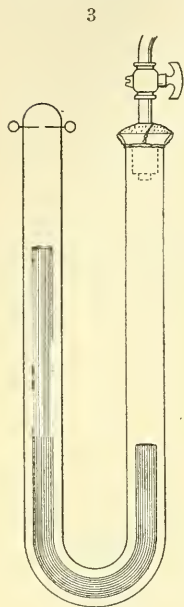
As the acetylene used thus far in the work was not free from traces of water, the following experiments were carried out in U-tubes such as shown in fig. 3 filled with gas containing 98.8 per cent of acetylene. Some phosphorus pentoxide was placed in the tube and also a clean piece of phosphorus; the latter was to remove traces of oxygen.

Experiment 57.—The volume of the gas diminished slowly on standing indicating an absorption, but there was no appearance of a tarry product. Five days after the tube was filled, the gas, measuring 60^{cc}, was condensed to 19^{cc} and then sparked. The explosion was prompt and violent. The condensation was 4 per cent and the acetylene in the residual gas was 4 per cent of the gas taken.

The tube was filled again with acetylene and the phosphorus pentoxide was made to coat the upper part of it. A piece of phosphorus was placed in the gas. The volume



diminished slowly and a reddish liquid appeared. A little water was introduced in order to find if phosphoric acid will condense acetylene. The change if any was slight in two days. Owing to an accident the experiment was not completed, but the absorption observed suggests doubt as to the value of phosphorus pentoxide as a drying agent for acetylene. In the remaining experiments no such absorption was observed.



Experiment 58.—After 8 days the phosphorus was removed and the pressure increased to 3.3 atmospheres and the gas was then sparked. The condensation was 8 per cent and the acetylene remaining was 8 per cent.

Experiment 59.—A smaller quantity of phosphorus pentoxide was used than before. However, some remained unaltered and after 8 days the pressure was made 3 atmospheres and the gas was sparked. The explosion was prompt and apparently as violent as any in the glass U-tubes. 6 per cent of the gas had formed condensation products and the acetylene in the residual gas was 24 per cent of that taken.

Experiment 60.—The acetylene stood over a solution of calcium hydroxide 3 days before it was put into the U-tube. The phosphorus pentoxide used filled a tube 2^{cm} long and 1^{cm} wide used to hold it. The phosphorus was also placed in a short tube. When the apparatus was filled the volume of the gas was noted and it did not change appreciably in the two weeks the gas was left to dry and to give up traces of oxygen. The gas, 57^{cc}, was then condensed to 19^{cc} and sparked by closing the primary for an instant. The explosion was prompt and violent. The condensation was 5 per cent and the acetylene in the residual gas was 21 per cent of the gas taken for the test.

The results obtained with dry acetylene indicate that water is not essential to the explosion of the gas, and that the wide variations noted in the first paper in the products of explosions in glass tubes are not to be ascribed to the presence of water.

The following table gives the number of cubic centimeters of acetylene and hydrocyanic acid per liter of gas in the bomb after explosion.

No. of Exp.	Pressure in atmospheres.	Per cent of C ₂ H ₂ in gas after explosion.	C. C. at 0° and 760 ^{mm} of C ₂ H ₂ per l. in bomb.	C. C. at 0° and 760 ^{mm} of HCN per l. in bomb.
29	3	4	120	
30	3	4.1	123	
36	2	4	80	
39	1.95	3.3	66	
40	5.1	2.4	122	
41	3.2 N 10.7 %	1.4	45	63
42	4. N 15 "	0.25	10	140
43	5. N 23 "	2.1	105	165
45	5. N 28 "	4.4	220	140
46	5. 10	2.	100	120
51 in U-tube	2.7	3.	81	
52 "	3.	1.7	51	
53 "	2.5 N 15 "	1.3	32	0
54 "	2.5 13 "	Not estimated.		0

The above are all of the results obtained with the bomb, fig. 1, and with the iron U-tube having the same diameter as the bomb, excepting experiments 31 to 35 described in the first paper. The amounts of acetylene and hydrocyanic acid found in the residual gas drawn from the bomb do not include quite all of these compounds, as there was some condensation on the carbon. This is evident from the fact that gas taken at 4 or 5 atmospheres contained less acetylene than that taken at lower pressures. The above data derived from the experiments are not sufficient to yield the relation between the pressure and the quantity of acetylene in the residual products. It is obvious, however, that at the lowest pressure at which explosion will occur the total quantity of acetylene left in the bomb is much less than when the initial pressure is 3 or more atmospheres. The small amounts found in experiments 51, 52 and 53 show that a more rapid cooling in the large iron U-tube than occurs in the bomb does not increase the acetylene in the residual gas. This fact supports the view expressed in the first paper that the acetylene in the gas after explosion is not an undecomposed portion of the original gas.

There is, however, no question about the synthesis of hydrocyanic acid in the explosions of mixtures of acetylene and nitrogen in the bomb. But in the U-tube the nitrogen did not combine. In the former instances the molecules of nitrogen either acquired energy adequate for combination or were dissociated, while in the latter they remained in the inert condition characteristic of the element. The velocity of explosion

of acetylene is 1000 meters and higher per second,* and in the iron U-tube the expanding gas must move a column of mercury about 80^{cm} in length, that is, the pressure on 1^{sq} cm must move one kilo of mercury. It is highly improbable that this mass is displaced materially before the explosion is completed. If this assumption be correct, the temperature at the instant of explosion is the same in the U-tube as in the bomb. In the former the expanding gas cools more rapidly than in the latter. In the bomb there is a sufficient number of collisions to impart to the molecules of nitrogen the energy adequate for combination or to dissociate them into atoms. In the U-tube the volume of the gas increases after the explosion and the collisions are fewer than in the bomb and the nitrogen does not combine. Thus we may conclude that, even at the temperature of exploding acetylene, a sufficient frequency of collisions is requisite to cause nitrogen to combine.

The amount of hydrocyanic acid remaining in the bomb does not appear to depend upon the proportion of nitrogen in the mixture exploded, for the result was the same with 15 per cent of nitrogen in the mixed gases at a pressure of 4 atmospheres as with 28 per cent at 5 atmospheres. It may be that a fixed quantity of hydrocyanic acid per unit of space remains when the products of the explosion cool and assume a state of equilibrium.

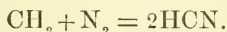
Acetylene and ammonia, as shown by experiments 49 and 50, yield hydrocyanic acid at a much lower temperature than is required to cause nitrogen to combine. It may be that ammonia is the first compound of nitrogen formed in the bomb, but the fact that a little ammonia is found in the products of an explosion is not conclusive, as this may have resulted from the decomposition of hydrocyanic acid. Dewar† passed a mixture of hydrogen and nitrogen through a carbon tube heated externally by an electric arc and obtained hydrocyanic acid. Here the conditions were similar to those in bomb, namely, carbon, hydrogen and nitrogen at a very high temperature. Dewar made no statement regarding the presence or absence of ammonia in the products of his experiment, and Berthelot‡ considered that ammonia played no part in the synthesis of hydrocyanic acid effected by sparking a mixture of acetylene and nitrogen. The results in the bomb indicate that less acetylene remains when nitrogen is present than when pure acetylene is exploded, but they are not conclusive. Berthelot (*loc. cit.*)

* Berthelot and Chatelier, *Comptes Rendus*, cxxix, 427; *Ann. Chim. et Phys.*, xx, 15.

† *Proc. Roy. Soc.*, xxix, 188, and xxx, 85.

‡ *Comptes Rendus*, lxvii, 1141.

observed that a mixture of acetylene and nitrogen diluted with hydrogen did not deposit carbon when sparked, and that all of the acetylene or nitrogen, an excess of one or the other being present, was removed by prolonged sparking of the mixed gases over a solution of caustic potash to remove the hydrocyanic acid. He considered that nitrogen united directly with acetylene and represented the reaction by the equation



As acetylene was not decomposed by the sparks, it must have combined directly with nitrogen. The reaction is of great interest, as it shows that the nitrogen molecules are rendered chemically active by conditions which do not dissociate endothermic acetylene.

ART. XXXI.—*Scapolite Rocks from Alaska*,* by J. E. SPURR.

THE following rocks containing scapolite as an essential constituent were collected by the writer in the summer of 1898, while making a geological reconnaissance in southwestern Alaska.

Andesine-oligoclase-scapolite-biotite rock.

This rock was found on the Yentna River (a branch of the Sushitna) some 12 or 15 miles above the junction with the Sushitna, where it forms large masses of uniform appearance, which are probably great dikes cutting an older rock. Specimens of this older rock, taken in different places, proved to be hornblende syenite and hornblende diorite. The scapolite rock is coarse-grained and granitoid in appearance, the bluish-gray of the scapolite suggesting quartz at first glance. Under the microscope the structure is hypidiomorphic granular, grading to panidiomorphic. The essential minerals are feldspar, scapolite, and biotite, abundant in the order of naming. The feldspars are mostly twinned according to the albite law, and one crystal examined by the Fouqué method showed on a section perpendicular to the negative bisectrix an angle of 70° between the plane of the optic axes and the albite twinning, proving it to be an andesine-oligoclase. The scapolite is both idiomorphic and hypidiomorphic, like the feldspar. Some sections are always dark, and give the dark cross of uniaxial minerals in convergent light. The cleavages, which are often broken, intersect at right angles and belong to the faces 100 and 010; another less perfect cleavage evidently belongs to the prism 110. The relief of the mineral is small, like quartz, while the double refraction is higher, giving a light-yellow color; it is optically negative. The biotite is fresh, and like the other minerals is idiomorphic or hypidiomorphic in habit; it contains grains of apatite. Small crystals of zircon, showing the prism and pyramid faces, are occasionally found.

Analysis of andesine-oligoclase-scapolite-biotite rock.

[Analyst, Dr. H. N. Stokes.]			Per cent.
	Per cent.		
SiO ₂	62.78	BaO04
TiO ₂56	SrO	trace
CO ₂	---	CaO	4.84
P ₂ O ₅15	MgO	2.32
S02	K ₂ O	2.15
Cl	trace	Na ₂ O	4.11
Al ₂ O ₃	17.16	Li ₂ O	trace
Fe ₂ O ₃	1.96	H ₂ O at 100°24
FeO	2.31	H ₂ O at 110°88
MnO06		
		Total	99.58

* Published by permission of the Director of the U. S. Geological Survey.

The above analysis bears out the results of the optical investigations and throws some additional light on the probable variety of scapolite. The high percentage of silica in the rock and the low percentage of lime indicate that the scapolite is probably of the variety dipyre. The rock therefore presents some analogy to the so-called dipyre-diorite of the Norwegian geologists.

Microcline-scapolite rock.

This is from the same locality as the rock just described and is doubtless a variation produced by differentiation. In the outcrop it has the appearance of a vein, forming a distinct tongue of light-colored material between walls of mottled rock, which appear to be essentially the same as the scapolite rock previously described, but of finer grain. This finer-grained biotite-scapolite rock grades off distinctly into the coarser-grained lighter-colored vein, which is about two inches thick. Under the microscope the structure is very coarse and the constituent grains are intergrown. Among the minerals microcline prevails, with considerable scapolite, which is largely altered to calcite. The scapolite gives in the basal section uniform darkness and the interference cross; it has abundant irregular cleavages and cracking, and shows yellow of the first order as an interference color.

Quartz-scapolite porphyry.

A rock called quartz-scapolite porphyry, for want of a better name, comes from a point far distant from the rocks just described, being found on the Kuskokwim River at the mouth of the Holiknuk, where it occurs among the light-colored dikes which cut the Cretaceous shales and shaly limestones. Under the microscope the rock shows phenocrysts of quartz, which are of small size and are bounded chiefly by the dihexahedral planes, but sometimes have the prism sparingly developed and sometimes show the faces of pyramids of two orders; the edges are often frayed by magmatic resorption. There are frequent large idiomorphic phenocrysts of scapolite, more or less broken down by decomposition. These phenocrysts give a uniaxial figure in convergent light. The outlines of the crystals show that they are of the tetragonal system, and they are characterized by low single and double refraction; they are optically negative. There is good cleavage parallel to the two lateral pinacoids, while that parallel to the prism is present but not so well developed. Irregular fractures are common. The mineral has decomposed along the margins and cleavages, partly to calcite, but mainly to an opaque kaolinic substance

which is colored brown with iron oxide. There are abundant microscopic inclusions arranged in zones within the crystal. Judging from the paucity of calcite among the decomposition products and also from the apparent small single and double refraction of these scapolites, the mineral is probably a soda scapolite (marialite end of the series). By exception, some almost entirely altered phenocrysts now consist chiefly of calcite, but what remains seems to be scapolite, and the outlines of the phenocrysts corroborate this inference. These decomposed scapolites are probably lime scapolites, or meionites. A few small phenocrysts, mostly altered to calcite, muscovite, etc., seem to have been originally soda-lime feldspar. The ground-mass is a fine-grained aggregate, consisting chiefly of quartz, orthoclase, and muscovite. In the fresher scapolite phenocrysts the cross sections are sometimes rectangles, but the favorite occurrence is in the form of a penetration twin of two such crystals along the composition plane 110; and the corners of the pinacoids whose faces constitute the rectangles are truncated by the prism. There are no true longitudinal sections in the slide, all being of small double refraction and having approximately the same orientation.

Occurrence of Scapolite Rocks elsewhere.

The scapolite rock from Norway in the vicinity of Oedegården has been described by Brögger and Reusch. This rock is essentially a mixture of scapolite and amphibole with accessory titanite, and was regarded by the describers as a facies of the gabbro (hyperite) in the neighborhood of veins of apatite which traverse the gabbro. The observers concluded, moreover, that the scapolite rock was formed by the alteration of the normal hyperite or gabbro, since they found remnants of diallage in the hornblende, as if the former were the original mineral, and of plagioclase in the aggregates of scapolite, as if the scapolite had formed from the plagioclase. The alteration, according to Brögger's view, was due to gaseous or pneumatolitic action.*

In Canada, also, granular rocks containing scapolite have been described from a number of localities.† In one of these rocks the principal constituents are pyroxene, hornblende, and scapolite, with accessory epidote, enstatite, pyrrhotite, and rutile, and the rock is classified as a scapolite diorite, it having a granular structure. Another rock is made up of an aggre-

* See Rosenbusch's *Mikroskopische Physiographie der Massigen Gesteine*, vol. i, p. 331.

† On some Canadian rocks containing scapolite, etc., by Frank D. Adams and Andrew C. Lawson: *Canadian Rec. Sci.*

gate of plagioclase, scapolite, and green hornblende, with accessory pyroxene, quartz, epidote, and pyrite, and this rock is termed a plagioclase-scapolite diorite. It is stated in the pamphlet referred to that "although the derivation of at least a part of the hornblende of these rocks from pyroxene is well-nigh certain, the derivation of the scapolite from plagioclase, which, as already stated, has been pretty clearly proved in the case of the Norwegian rocks, is not so evident in these similar rocks from Canada." Also, Dr. A. P. Coleman reports from Ontario a scapolite rock consisting of hornblende, plagioclase, and scapolite, in which the scapolite was evidently primary.*

Finally, scapolite rock occurs in the white limestone of New Jersey at many localities and in large amounts, the rock consisting in one locality of scapolite, hornblende, pyroxene, and sphene, and in another locality of the same minerals with the addition of a little plagioclase. These dikes are evidently igneous, and fumarolic or gaseous action has been suggested as taking part in their formation.†

In addition to these granular scapolite rocks, scapolite gneisses and amphibolites have been described from various places.‡

Origin and Classification of the Alaskan Scapolite Rocks.

In the locality on the Yentna River which has been described the main intrusive rock of the country was noted in the field to be a biotite granite, but of this, unfortunately, no specimen was retained. Dikes of quartz-feldspar rock or alaskite,§ which were found cutting the granite, have been carefully examined microscopically, and connected with this alaskite were rocks transitional into the granite, and quartz veins sometimes containing tourmaline. The specimen of scapolite rock was from a large mass forming bluffs 40 feet high, and seemed typical; it is certain that this rock occurs in notable quantity, but in just what proportion to the other dikes is not certain. All these dikes cut an ancient igneous rock, which was shown in one case to be a hornblende syenite and again a hornblende diorite.

In the second case which has been described, namely, the quartz-scapolite rock from the Kuskokwim River, the scapolite dikes were associated with a number of other siliceous dikes. Those which were examined microscopically, besides the scapolite porphyry, are an alaskite porphyry and a granite porphyry.

* See Zirkel's *Lehrbuch der Petrographie*, vol. ii, p. 783.

† F. L. Nason, *Annual Report of the Geological Survey of New Jersey*, 1890, p. 33.

‡ Zirkel, *op. cit.*, vol. iii, p. 339.

§ See author's paper on *Classification of Igneous Rocks according to Composition*; *Am. Geol.*, xxv, April, 1900, p. 210.

In both these cases the scapolite is regarded as original. There is no evidence of its derivation from feldspar, but rather in every case of contemporaneous formation. Certainly the mineral is not the product of weathering, for it is nearly always in a process of decomposition, showing itself as unstable under atmospheric conditions as are the feldspars. It is exceedingly probable, however, that in the formation of this, as in other occurrences of scapolite rock, gases have played an important part. The scapolites contain chlorine, otherwise they have essentially the composition of the soda-lime feldspars. The marialite scapolite corresponds to oligoclase, and the meionite scapolite to anorthite, with a series between like that of the feldspars, so that when scapolite is found instead of feldspar we may suppose chlorine gas to have been present at the time of the formation of the rock. This is, however, no good reason for considering the rock as of secondary origin, for in most igneous rocks the gases play some part in the formation and in many a very weighty part, the rocks nevertheless belonging to the class of original igneous rocks. The writer is inclined, therefore, to give the scapolite rocks a place among the primary igneous rocks wherever the scapolite shows evidence of having formed contemporaneously with the other rock minerals. In cases where it evidently is an alteration product from some other rock, the classification will of course be different.

In case classification is attempted, the scapolites might be considered, for the purposes of classification, as equivalent to the feldspars, when they occur in rocks where feldspar is present in an equal or greater amount. Thus the biotite-scapolite rock from the Yentna River might be called a biotite-scapolite-belugite,* since the feldspar belongs to the andesine-oligoclase series, and the equally important scapolite seems to be of a variety containing much lime and corresponding to anorthite in the feldspar series. The scapolite-belugites might then be considered a group of the belugite family.

The porphyritic scapolite rock from the mouth of the Holik-nuk on the Kuskokwim might be called a scapolite-adamellite porphyry, adamellite being the name adopted by Brögger for the quartz-bearing monzonites. In this case the scapolite appears to belong mainly to the soda end of the soda-lime series, and thus is the equivalent of oligoclase-andesine. The groundmass appearing to be mainly orthoclase, the feldspathic constituents (considering scapolite as such) would be on the whole intermediate between the granite and diorite families, or in the monzonite family. In this case the rock might be con-

* See writer in *Am. Geol.*, xxv, April, 1900, p. 233.

sidered to belong to what might be called the scapolite-adamellites, which would be a group under the monzonite family where the scapolite partly takes the place of the feldspar.

It would probably be best, however, to separate the scapolite-feldspar rocks into a distinct class, and to give the groups characteristic names, the analogy with the feldspar rocks being expressed by writing in the tabulation each group laterally opposite the corresponding group of feldspar rocks. Thus the rock first described may be called yentnite, from the Yentna River, instead of scapolite-belugite, but it may be written opposite the belugite group; the type would then be a biotite yentnite. Similarly the quartz-scapolite porphyry from the Kuskokwim* River might be called kuskite, instead of scapolite-adamellite; the kuskite group of the scapolite-feldspar class of rocks could be written opposite the adamellite (or quartz-monzonite) group of feldspar rocks, and the type might be called a kuskite porphyry.

* Eskimo *Kuska*, derivation uncertain, *kwi*k, genitive *kwi*m, river.

ART. XXXII.—*On the Qualitative Separation of Nickel from Cobalt by the action of Ammonium Hydroxide on the Ferricyanides*; by PHILIP E. BROWNING and JOHN B. HARTWELL.

[Contributions from the Kent Chemical Laboratory of Yale University—XCVII.]

SOME years ago F. W. Clarke* suggested a method for the separation of nickel from cobalt depending upon the solvent action of ammonium hydroxide upon the precipitated ferricyanides. The method may best be described by quoting from the original article: "To the slightly acid solution containing the two metals, I first add an excess of ammonium chloride. This causes the cobalt precipitate, which otherwise would run through the filter, to fall in a denser state, and also of a much darker color, often nearly black. I then add the potassium ferricyanide until the precipitation is complete, and afterwards agitate strongly with a considerable excess of ammonia. Upon filtering, all the cobalt remains upon the filter, being recognized by the characteristic color of the precipitate, and the nickel is readily detected in the filtrate, by means of ammonium sulphide. If, upon filtering, the portion at first running through is turbid, it may be disregarded, or returned to the filter, that which filters through subsequently being almost invariably clear."

In making a study of this method we found two serious objections; first, the practical impossibility of obtaining a good filtration from the cobalt ferricyanide, even in the presence of the ammonium chloride, and, second, the large amount of sulphur thrown down when ammonium sulphide was added to the filtrate containing the nickel with the excess of ferricyanide.

Our first attempt was to secure, if possible, a complete separation of the precipitated cobalt ferricyanide and the dissolved nickel by filtration. This we were able to accomplish by the addition of a small amount of a solution of an aluminum salt to the original solution which held back the cobalt, and, as experiment showed, allowed the complete solvent action of the ammonium hydroxide upon the nickel salt. Amounts of nickel as small as 0.0001 gram. were detected, when mixed with the aluminum salt, by precipitating as ferricyanide, extracting with ammonium hydroxide, and testing in the manner to be described.

On turning our attention to a possible improvement in the method for the detection of the nickel, it was found that when the ammoniacal solution of the nickel ferricyanide was treated

* This Journal, *xlvi*, 67.

with strong sodium or potassium hydroxide solution, in the presence of an excess of potassium ferricyanide, a black flocky precipitate formed which gave no test for ferro or ferricyanide, and gave every indication of being nickelic hydroxide. This reaction we found to afford us a most delicate test for nickel.

The method as modified by us may be described as follows: Dissolve not more than 0.1 gm. of the salts of the two elements in about 5^{cm}³ of water, add a few drops of a saturated solution of alum, destroy any free mineral acid by neutralizing with ammonium hydroxide, and make faintly acid with acetic acid. To this solution add about 0.5 gm. of potassium ferricyanide and agitate to effect the solution of the ferricyanide and the complete precipitation of the nickel and cobalt salts. Then add about 5^{cm}³ of strong ammonium hydroxide and filter. To the filtrate, which should have no reddish color, add a piece of sodium or potassium hydroxide about the size of a pea and boil. The appearance of a black precipitate, in case of very small amounts of nickel showing first as a dark coloration, indicates nickel.

The tables following give a record of the experimental results. With the precautions indicated, this method may be applied very satisfactorily.

TABLE I.

	CoSO ₄ ·7H ₂ O gram.	NiSO ₄ · 7H ₂ O gram.	KAl(SO ₄) ₂ saturated sol. cm ³	K ₃ FeC ₆ N ₆ gram.	NH ₄ OH (conc.) cm ³	NaOH, solid, about size of a pea.	Result.
1.	----	0.0100	2	0.5	5.	"	Heavy ppt.
2.	----	0.0050	"	"	"	"	"
3.	----	0.0010	"	"	"	"	"
4.	----	0.0003	"	"	"	"	Distinct
5.	----	0.0001	"	"	"	"	Plain

TABLE II.

1.	0.10	----	"	"	"	"	None
2.	0.10	0.0100	"	"	"	"	Heavy
3.	0.10	0.0050	"	"	"	"	Distinct
4.	0.10	0.0030	"	"	"	"	Very faint
5.	0.05	----	"	"	"	"	None
6.	0.05	0.0100	"	"	"	"	Heavy
7.	0.05	0.0050	"	"	"	"	Distinct
8.	0.05	0.0030	"	"	"	"	Plain
9.	0.05	0.0010*	"	"	"	"	Faint

* Equivalent to 0.0002 of the metal.

SCIENTIFIC INTELLIGENCE.

I. PHYSICS.

1. *Radio-activity of Uranium.*—Sir WILLIAM CROOKES has recently made an investigation in regard to Becquerel's rays, the most important result of which is the fact that he has succeeded in purifying uranium salts to such an extent that they are photographically inactive. There can be no doubt, therefore, that the radio-active rays are not a specific property of pure uranium, as has been previously supposed.

In the course of his experiments Crookes has tested a large number of specimens of barite and witherite from various localities without obtaining the slightest photographic action from them. He then tested every mineral in his extensive collection with the result that the following were found to be active: Pitchblende, uranite, autunite, orangite, thorite, euxenite, samarskite, alvite, bröggerite, monazite, xenotime, arrhenite, sipylite, fergusonite, chalcophile, hiemite. These minerals, which are arranged in the order of their intensity of action, all contain either uranium or thorium. Pitchblende, the most active mineral showed great variations in the activity of specimens from different localities, but roughly speaking the variation was proportional to the percentage of uranium present. A layer of powdered pitchblende a quarter of an inch in thickness gave as strong an action as one of the same material two inches thick, showing that the action does not pass through much thickness of the active material. The action was shown to be proportional to the time of exposure. Several salts and oxides of uranium were tested for their comparative activity upon the photographic plate with the result that they did not show much difference. A sample of metallic uranium from M. Moissan, however, showed a slighter action. With the intention of using it as a standard for comparison the author prepared a very pure specimen of uranium nitrate by treatment with ether and subsequent repeated crystallization. To his surprise this gave no action upon the sensitive plate. It was then shown by a series of experiments that no modification of physical or chemical conditions materially affects the radio-activity of a uranium compound when, to begin with, the salt experimented on possesses it; other similar experiments showed that, starting with an inactive uranium salt, nothing that can be done to it will cause it to acquire activity. It was found that when ordinary crystallized uranium nitrate is dissolved in ether a heavy aqueous liquid separates containing uranium nitrate that is much more active than that which remains dissolved in the ether. Systematic recrystallization from water showed a concentration of active material at the soluble end. Fractionation carried out by heating uranium nitrate until a certain amount of basic salt was produced, proved that the active body gradually

accumulated towards the basic end. By treating a solution of very active uranium nitrate with ammonium carbonate in excess, a small flocculent precipitate was obtained which showed great radio-activity; it was found, however, that the active body was not wholly insoluble in ammonium carbonate. The precipitate just mentioned gave in five minutes as strong an action upon the sensitive plate as is given in 24 hours by ordinary uranium nitrate. The author calls the active substance provisionally "UrX." Thus far he has not observed any special spectrum for the substance. He decides that UrX is certainly distinct from "polonium," because the rays emanating from it penetrate glass while polonium rays do not. It is not so easy to settle whether UrX is distinct from radium, although many arguments point to its not being radium.

Experiments have been commenced by the author, which tend to show that thorium may be separated into an active and inactive body.

It is the author's opinion that in the present state of our knowledge of radio-active bodies it is safest to retain an open, or even a slightly sceptical mind. He calls attention to the fact that we recognize them mainly by photographic and electrical tests—reactions which give strong results even when the active body is present in too small a quantity to be detected by the spectrum—one of the most delicate of tests. He remarks that the radiographic test is cumulative. If no action is apparent at the end of an hour, one may be shown after twenty-four hours. If a day's exposure will show nothing, try a week's. The article concludes with the following sentence: "Considering my most active UrX does not contain sufficient of the real material to show in the spectrograph, yet is powerful enough to give a good impression on a photographic plate in five minutes, what must be its dilution in compounds which require an hour, a day, or a week to give an action?"—*Chem. News*, lxxxix, 253, 265. H. L. W.

2. *Electrical conductivity of Gases traversed by Cathode Rays*.—Starting from the results reached by J. J. Thomson and Rutherford in regard to the conductivity of gases traversed by Röntgen or uranium rays, owing to the production of positive and negative ions within them, J. C. McLENNAN shows that cathode rays impress a similar condition upon a gas. The conductivity produced by the cathode rays is explained as due to the motion of the ions, positive and negative, produced in the gas by the radiation. The ionization by cathode rays was found to be about 300 times that due to intense Röntgen radiation. On experimenting with different gases, the result was established that to determine the relative ionization produced in two gases by cathode rays of the same intensity it is sufficient to determine the absorbing power of the two gases for the same rays. It is also probably true that the ionizations produced by rays of constant intensity traversing different gases at the same pressure are proportional to the densities of the gases. The calculated

values expressing the relation between the ionization by cathode rays in different gases (oxygen, nitrogen, etc.) compared with that of air (= 1) was found to be approximately the same as those observed for Röntgen rays, except in the case of hydrogen. —*Proc. Roy. Soc.*, No. 431, p. 375.

3. *Kleiner Leitfaden der Praktischen Physik* von FRIEDRICH KOHLRAUSCH. Pp. xix, 260; 8vo. Leipzig, 1900 (B. G. Teubner). —The many excellent features of the work on Practical Physics by Dr. Kohlrausch, now before the public for some twenty-five years and in its eighth edition, are well known. The same author has now prepared an elementary volume on the same lines, which is designed especially for beginners in Physics and which has a unity character impossible in the larger work. His long experience has enabled him to write a book eminently fitted to fill the place for which it is designed; it is indeed a model both in arrangement and in clearness of statement and conciseness of presentation. It should have a wide sphere of usefulness not only at home, but wherever laboratory work in physics is carried on.

4. *Photometrical Measurements and Manual for the General Practice of Photometry with special reference to the Photometry of Arc and Incandescent Lamps*; by WILBUR M. STINE, Swarthmore College. Pp. vii, 270. New York, 1900 (The Macmillan Company). —The introduction of new forms of illumination in recent years has much increased the importance of photometrical measurements and extended the range of their application. Hence the value of the present volume, which gives in compact form the physical basis of photometry, the various forms of photometers, the light standards in use and finally the particular application of photometrical measurements to electric lamps of the two types. The author has not only brought together a remarkable amount of material within small compass, but in the frequent and well selected references has practically placed before the reader the important part of the original literature of the subject.

II. GEOLOGY AND MINERALOGY.

1. *Status of the Mesozoic Floras of the United States*. First Paper: *The Older Mesozoic*; by LESTER F. WARD, with the collaboration of Wm. M. Fontaine, Atreus Wanner, and F. H. Knowlton. From the Twentieth Annual Report of the U. S. Geological Survey, 1898-99, Part II, General Geology and Palaeontology, pp. 211-748, with 158 plates. Washington, 1900. —This extended and timely collaboration brings into convenient form for reference the principal known data concerning the Triassic and Jurassic floras of the United States, including many new descriptions. The salient facts dealt with are here enumerated.

The structural determination of several new genera and species

of conifers from Triassic and Jurassic horizons is by Dr. Knowlton. The new genus *Pinoxylon* may be specially mentioned as a Jurassic pine from South Dakota.

The Triassic flora of North Carolina, as based upon the original collection of Emmons and most of his types, is the work of Professor Fontaine. The Emmons collection of fossil plants made some fifty years since, and long supposed to have been lost, was recently found in the geological collection of Williams College by Professor T. Nelson Dale. Their redescription accompanied by new figures, and made in the light of a knowledge of the several other Triassic plant-bearing horizons since discovered in the United States, is especially satisfactory. The Triassic horizons of North Carolina being among the earliest investigated in this country, represent classic ground. Moreover, from the wooded character of the plant-bearing region, and from the fact that there are few natural products to induce excavation, specimens of this flora are now seldom to be obtained, save from a few localities principally in Chatham County, where, as at Egypt and Deep River, intermittent coal mining is still being carried on in the same horizons which yielded the important vertebrate fossils *Dromatherium*, and *Microconodon*, and also *Rhytidodon rostratus* Marsh. The Deep River locality was visited by the reviewer a few years since, but unfortunately the mines and culm heaps were on fire and could not be examined at that time, though being again put in order for further work.

A rich new Triassic flora from York County, Pennsylvania is of much interest. This was discovered by Mr. A. Wanner, and is in large part figured and described by him, Professor Fontaine also having determined the several forms. While some new plants occur, the essential agreement of this flora with that of North Carolina gives to its description a fundamental value.

A brief notice of the Triassic flora of Virginia, made by Fontaine the subject of Monograph VI of the U. S. Geological Survey publications, is given by Professor Ward, who notes the agreement between this and the Triassic flora of Lunz in Austria later studied by Stur and referred by him to the Keuper.

A characteristic Jurassic flora from Oroville, California, is described and illustrated by Professor Fontaine.

The Jurassic cycadean trunks from Carbon County, Wyoming, forming the new genus *Cycadella* with twenty species, are described by Professor Ward. A review of his preliminary description of these cycads was given by the present reviewer in this Journal for April, 1900, pp. 383-387, and need not be repeated here. The 104 plates now devoted to the illustration of these fine fossils forms the principal feature of the present volume, as well as a most interesting addition to the series of plates showing various forms of isolated cycadaceous leaves so prominent in the several Triassic and Jurassic floras mentioned above.

In addition to these Wyoming cycads Professor Ward also

describes as *Cycadeoidea nigra* a remarkably well preserved and handsome new species of cycadean trunk of supposed Jurassic age from the vicinity of Boulder, Colorado. This trunk is of more particular interest as extending the geographical range of the genus *Cycadeoidea* one step further west. While the distinctness of the present species is quite certain, it might be noted that externally it bears a strong resemblance to *Cycadeoidea Uhleri* from the Potomac formation of Maryland, the latter being in all probability a close specific, or even varietal relative. There is likewise a close resemblance to *Cycadeoidea* (*Raumeria*) *Masseiana* Capellini, from the scaly clays of Italy.

If in the present invaluable contribution there is anything open to criticism, it is perhaps the absence, in a work otherwise so profusely illustrated, of maps and sections of the more important localities and areas dealt with, such for instance as appear in the present author's contributions on the lower Cretaceous of the Black Hills in the XIXth Annual Report of the U. S. Geological Survey.

G. R. W.

2. *La Flore Wealdienne de Bernissart*; par A. C. SEWARD. Memoires du Musée Royal d'Histoire Naturelle de Belgique, Vol. i, 1900.—The plants described in this memoir are the unfortunately somewhat fragmentary remains of the Flora occurring in the Wealden beds of Bernissart near the French frontier between Mons and Tournay, famous for the discovery in 1877 of numerous remarkably preserved *Iguanodon* skeletons.

The writer's conclusions are translated in part as follows: The plants of Bernissart are sufficient to demonstrate a closer resemblance to the Wealden facies than to any other Mesozoic type of vegetation.

I have elsewhere remarked that from the botanical point of view there exists an intimate resemblance between the Wealden type of Flora wherever well developed, as in England and Northern Germany, and the Flora of the inferior Oölite. There is not in fact a well marked break in the palaeobotanical continuity between the Jurassic and Wealden Floras of itself justifying the term Wealden as applied to plant beds such as those of Hastings on the coast of Sussex; comprised in the lower Wealden by the English geologists.

The composition of the flora of Bernissart is interesting for the marked preponderance of ferns, the apparently total absence of cycads, and the rarity of conifers. In other regions, for example in England, Portugal, and Germany the cycads play a preëminent role in Wealden vegetation. A collection containing many conifers and some ferns has recently been made in the Wealden strata at Bracquagnies, 30 kilomètres east of Bernissart, whilst specimens from Baume, 8 kilomètres further east, consist principally of conifers.

It is also remarkable that the specimens of *Weichselia Mantelli* and *Laccopteris Dunkeri* are more numerous in the Bernissart sediments than all others. Perhaps the deduction may be made

that the land drained by the rivers which laid down the plant debris in the Bernissart sediments was to a great extent covered by the two species first named and that the remaining vegetation also consisted principally of ferns. The aspect of the specimens and this abundance of ferns likewise suggests the idea that these plant-bearing sediments were derived from a low region chiefly occupied by ferns to the relative exclusion of larger plants or trees. These probably grew in the higher background or in the localities beyond the reach of the water which deposited the Bernissart strata. It is interesting, again, to note that the plant debris contained in these strata furnishes no facts favoring the presence of Angiospermous species. As we know from richer floras of the same age, it appears most probable that the higher classes of Phanerogams were not represented, or at least occupied a very secondary position in the vegetation of the Jurassic and Wealden periods.

G. R. W.

3. *Notes on Some Jurassic Plants in the Manchester Museum*; by A. C. SEWARD, M.A., F.R.S., from Volume xlv, Part III of "Memoirs and Proceedings of the Manchester Literary and Philosophical Society, Session 1899-1900.—This communication deals with Inferior Oölite species from the Gristhorpe plant beds, in part figured by Lindley and Hutton. It adds certain important data to our knowledge of Jurassic vegetation.

G. R. W.

4. "*The Maidenhair Tree*" (*Gingko biloba* L.); by A. C. SEWARD, F.R.S., and Miss J. Gowan, Newnham College, Cambridge, *Annals of Botany*, Vol. xiv, No. LIII, March 1900 (pp. 100-124, with plates viii-x.—This beautiful tree is here described with a completeness and precision which will be most pleasing and timely to every student of Botany or Palaeobotany.

Engler's subdivision *Gingkoaceæ* is adopted for the monotypic genus *Gingko*,—distinguished by the possession of motile male cells as well as by other characters of more or less importance. Following an historical sketch and diagnosis, the vegetative structures and fructification are described, reference always being had to the literature on the several subjects. A sketch of fossil *Gingkoaceæ* is appended, which adds greatly to the value of the contribution.

G. R. W.

5. *New minerals from Greenland*.—A recently published account* of minerals collected in 1897 by G. FLINK in the Julianehaab district of southern Greenland, contains descriptions of a considerable number of new species and also of other rare known species associated with them. This important paper consists of two parts. Part I (pp. 7-180), by G. Flink, is on the minerals of Narsarsuk on the Tunugdliarfik fiord in southern Greenland. These are found only in a very limited area of the prevailing syenite, where the rock has in spots a pegmatitic character. The prominent minerals are microcline and ægirite, which often occur as large crystal individuals; in drusy cavities between these minerals the rarer species occur in successive generations. Of

* Meddelelser om Grönland, xxiv, 1899.

the more interesting known species of which full descriptions are given, may be mentioned the following: Parisite, eudidymite, epididymite, ægirite, arfvedsonite, catapleite, neptunite, elpidite.

The new species described are nine in number. The descriptions of these leave nothing to be desired in the way of fullness and accuracy, but they can only be briefly characterized here.

CORDYLITE is a *barium-parisite*. It occurs in minute hexagonal crystals, club-like in form; they are related in angle to parisite; hardness 4.5; fracture conchoidal; sp. gravity 4.31; color wax-yellow. The composition is expressed by the formula $Ce_2F_2BaC_3O_9$.

ANCYLITE occurs in minute orthorhombic crystals with strongly curved faces. Color light yellow to orange or resin-brown; hardness 4.5; no cleavage; sp. gravity 3.95. In composition it is a hydrated carbonate of cerium and strontium for which the formula deduced is: $4Ce(OH)CO_3 + 3SrCO_3 + 3H_2O$.

SPODIOPHYLLITE resembles a chlorite; it occurs in crystals belonging to the rhombohedral system, combinations of the hexagonal prism and base; cleavage basal, micaceous; hardness about 3; sp. gravity 2.633; color ash- to pearl-gray. In composition it is a metasilicate related to ægirite; formula $(Fe,Al)_2(Mg,Fe,Mn)_3(Na,K)_2(SiO_3)_8$.

TAINIOLITE (Fæniolite) is a kind of mica occurring in elongated colorless crystals with sp. gravity = 2.86. An analysis by Mauzelius on 0.1 gram gave: SiO_2 52.2, Al_2O_3 2.7, FeO 0.6, MgO 19.1, K_2O 11.5, Na_2O 1.8, Li_2O 3.8, loss 8.7 = 100. The loss is referred to water (hydroxyl) and fluorine.

LORENZENITE occurs in needle-like orthorhombic crystals, colorless to violet or brown; luster adamantine; hardness 6; sp. gravity 3.42. An analysis (by Mauzelius) gave SiO_2 34.26, TiO_2 35.15, ZrO_2 11.92, Na_2O 17.12, K_2O 0.37, H_2O 0.77 = 99.59. The formula deduced is $Na_2(Ti,Zr)_2Si_2O_7$.

LEUCOSPHEENITE occurs in wedge-shaped monoclinic crystals; color white inclining to grayish blue; luster vitreous, on some faces pearly; hardness 6.5; sp. gravity 3.05. An analysis (Mauzelius) gave: SiO_2 56.94, TiO_2 13.20, ZrO_2 3.50, BaO 13.75, Na_2O 11.14, K_2O 0.56, H_2O 0.31 = 99.40. The formula deduced is $Na_4Ba(TiO)_2(Si_2O_6)_5$. The mineral seems to be related to eudidymite.

NARSARSUKITE occurs in tabular tetragonal crystals; color honey-yellow to brownish gray; hardness a little above 7; sp. gravity 2.751. An analysis (Christensen) gave: SiO_2 61.63, TiO_2 14.00, Fe_2O_3 6.30, Al_2O_3 0.28, MnO 0.47, MgO 0.24, Na_2O 16.12, F 0.71, H_2O 0.29 = 100.04.

CHALCOLAMPRITE occurs in small regular octahedrons, resembling pyrochlore; color dark grayish brown inclining to red; hardness 5.5; sp. gravity 3.77. An analysis (Mauzelius) gave: Nb_2O_5 59.65, SiO_2 10.86, TiO_2 0.52, ZrO_2 5.71, Ce_2O_3 (etc) 3.41, Fe_2O_3 1.87, MnO 0.44, CaO 9.08, K_2O 0.38, Na_2O 3.99, H_2O 1.79, F 5.06 = 102.76 (deduct O) = 100.63. It is regarded as consisting of equal amounts of $RNb_2O_6F_2$ and $RSiO_3$.

ENDEIOLITE is related to the preceding species and like it occurs in regular octahedrons, at first taken for pyrochlore. Color dark chocolate-brown; hardness 5, sp. gravity 3.44. An analysis (Mauzelius) gave (assuming the loss to be SiO_2): Nb_2O_5 59.93, SiO_2 [11.48], TiO_2 0.76, ZrO_2 3.78, Ce_2O_3 4.43, Fe_2O_3 2.81, MnO 0.37, CaO 7.89, K_2O 0.43, Na_2O 3.58, H_2O 4.14, F 0.69 = 100.29 (deduct O) = 100. The formula deduced is $\text{RNb}_2\text{O}_6(\text{HO})_2 + \text{RSiO}_3$.

Part II (pp. 181-213) contains descriptions of minerals from the nephelite-syenite of Julianehaab, by O. B. Boeggild and Chr. Winther. Three new species are included here and also an account of steenstrupite.

EPISTOLITE is a silver-white mineral described by Boeggild. It resembles brucite in its tabular crystals and pearly luster on the basal cleavage; it is, however, monoclinic in crystallization. Color white; hardness 1 to 1.5; sp. gravity 2.885. An analysis (Christensen) gave: SiO_2 27.59, Nb_2O_5 33.56, TiO_2 7.22, FeO 0.20, MnO 0.30, CaO 0.77, MgO 0.13, Na_2O 17.59, H_2O 11.01, F 1.98 = 100.35 (deduct O) = 99.52. No definite formula can be obtained, as the material used was somewhat altered.

BRITHOLITE, described by Winther, occurs in pseudo-hexagonal crystals, related in composite form to aragonite. Color brown, opaque; luster greasy to vitreous; hardness 5.5; sp. gravity 4.446. An analysis (Christensen) gave: SiO_2 16.77, P_2O_5 6.48, $(\text{Ce}, \text{La}, \text{Di})_2\text{O}_3$ 60.54, Fe_2O_3 0.43, CaO 11.28, MgO 0.13, Na_2O 1.85, H_2O 1.27, F 1.33 = 100.08.

SCHIZOLITE, also described by Winther, is characterized as a manganese pectolite. It occurs in columnar masses and prismatic crystals, elongated parallel the *ab*-axis; cleavage parallel to two faces in this zone. Color pink to brown; hardness 5 to 5.5; sp. gravity 3.089. An analysis (Christensen) gave: SiO_2 51.06, TiO_2 0.68, Ce_2O_3 1.47, FeO 2.79, MnO 12.90, CaO 19.48, Na_2O 10.71, H_2O 1.36 = 100.45. The formula deduced is $4\text{R}_2\text{O} \cdot 10\text{RO} \cdot 15\text{SiO}_2$, which does not agree very closely with that of pectolite.

OBITUARY.

JAMES EDWARD KEELER. The Director of the Lick Observatory, Professor James E. Keeler, died at San Francisco after a brief illness on August 12th. Few of his friends had a suspicion that his health was otherwise than in the excellent condition which was believed characteristic, until the widely published announcement of his death met their eyes.

James Edward Keeler, born at La Salle, Illinois, on September 8th, 1857, where he also received his early schooling, doubtless owed much of his intellectual alertness to his father whose career had been singularly varied and interesting. A paymaster in the navy during the civil war, his intimate connection with the monitor class of ironclads, including a participation in the famous fight between the Monitor and the Merrimac and subsequent

experience in the East, strengthened in him the innate taste for mechanical arts which he shared and enjoyed with his distinguished son. The removal of the family to Mayport, Florida, while the future astronomer was still quite young, profoundly modified his course of development in replacing an ordinary career of schooling by an employment of all his energies in adapting himself to a new life and a new environment. But even in this active life his natural inclination towards astronomy asserted itself. By the exercise of much self-denial he was able to purchase an achromatic telescope objective of about two and a half inches diameter; this he mounted and supplied with a thoroughly serviceable stand, and with the product of his ingenuity he commenced the acquiring of his remarkable store of knowledge of physical astronomy. This was followed by a singular consequence. Mr. Charles H. Rockwell, who has ever remained a kind and sympathetic friend, learned by accident of the young astronomer in Florida and, eager to advance scientific learning in every way, secured for him an opportunity to gain the advantages afforded by one of our greater universities. As Mr. Keeler's schooling was not of a character which fitted him immediately for entrance to university standing, he entered as a special student in the Johns Hopkins University, then in the initial stages of its evolution. Here it was that the writer became intimate with him and learned his extraordinary capacity for fruitful work. Shortly after securing his bachelor's degree here in 1881, he became an assistant to Professor Langley, aiding him most efficiently in his delicate researches with the bolometer, at the same time establishing a reputation for scientific ability which practically secured his future. With this experience, broadened by two years of study at Berlin and Heidelberg and further increased by four or five years passed at Mt. Hamilton as an astronomer of the Lick Observatory, his subsequent career is wholly natural. In 1891 he accepted the position as Director of the Allegheny Observatory, a position exceptionally congenial to him because of the many friends gained during the time in which he acted as assistant to his predecessor Professor Langley. The opportunity for more important work, however, which came to him in 1898, with an invitation to become Director of the Lick Observatory, obliged him to sacrifice this agreeable environment and to enter upon a life of great fruitfulness and activity which has been, so unfortunately for science, interrupted by his death.

This brief note does not offer the place to review the scientific achievements of Mr. Keeler; but it is gratifying to find that not only those engaged in allied branches of science recognized his merits, but more popular appreciation was evinced by his election as an Associate of the Royal Astronomical Society in 1898, and as a member of the National Academy at its last meeting. Quite fitting is it, however, that the writer should take this occasion to record the grief of a great number of his personal friends through whom his life was enriched in a degree never attained by men less selfish and less loyal than he.

C. S. H.

T H E

AMERICAN JOURNAL OF SCIENCE

[F O U R T H S E R I E S .]

ART. XXXIII.—*Elaboration of the Fossil Cycads in the Yale Museum*; by LESTER F. WARD. (With Plates II-IV.)

THE visitor to the Peabody Museum of Yale University, if he penetrates to the basement of that building, finds himself literally "in the woods,"—a petrified forest of Mesozoic cycads. A few of the finest specimens may be seen in the exhibition halls on the second floor, and many more will be ultimately placed there, but at present the bulk of the collection is undergoing elaboration below, in close association with the gigantic dinosaur bones that have made the names of Marsh and of Yale so justly celebrated in the history of science. In a very literal sense, cycads are to the vegetable kingdom what dinosaurs are to the animal, each representing the culmination in Mesozoic time of the ruling dynasties in the life of that age. Professor Marsh saw this, and, as the last act of his life, had the sagacity to make the Yale Museum for all time the Mecca for all who shall wish to gain a realizing sense of the fauna and flora of America in a period now forever closed.

It has chanced to be my fortune or misfortune to be situated, as it were, in the storm track of cycadean investigation on this continent, and upon me has devolved the duty of roughly blocking out the general line of study of that wonderful extinct vegetation that has been coming to light in this country in such rich profusion. Science demands a terminology and a nomenclature, and whatever may be said of the superficial character of all systematic work, it is now true and always has been true, nay, it must always be a need, that the systematist precede the structuralist and provide him with a language and a framework for his finer researches. This is all I profess to have attempted, and not

only am I keenly aware of the superficiality and defectiveness of my work, but I have urged on all occasions the importance of the exhaustive study of the wonderful structure of these cycadean trunks that reveals itself to the lens and even to the naked eye. I presented this aspect of the case in the strongest form I could command to Professor Marsh, and pointed out to him the "unlimited possibilities" of such a study of the great Yale collection. I am happy to record his warm appreciation of the fact, which led him actually to inaugurate it some months before his death, by inducing Mr. George R. Wieland to undertake it and by placing at his disposal every possible facility for the prosecution of this work. Professor Marsh's successor, Dr. C. E. Beecher, with the approval of the Trustees of Yale University Museum, has allowed no interruption in these valuable researches, so creditable to the institution, and has placed them on a permanent basis.

I am only to speak here of my own work, much of which I have already recorded and need only allude to,* confining myself to the additional results that have been reached by recent study. As stated in the Nineteenth Annual Report of the U. S. Geological Survey, Part II, pp. 546 and 547, I worked up all the cycad material from the Black Hills in the Yale Museum, in the months of March and June, 1898, and the 22 species described in that report were based on the material in the U. S. National Museum, and on 126 specimens in the Yale Museum, contained in two invoices, the first of 87 specimens and the second of 39. These were illustrated in 97 plates, and that paper constitutes the basis for subsequent and future investigations. It was known at that time that other

* For the benefit of any who may be specially interested, I herewith refer to the following papers of mine relating in whole or in part to fossil cycadean trunks: *Fossil Cycadean Trunks of North America, with a Revision of the Genus Cycadeoidea* Buckland. *Proc. Biol. Soc. Washington*, vol. ix, Washington, April 9, 1894, pp. 75-88; *The Cretaceous Rim of the Black Hills*, *Journal of Geology*, vol. ii, No. 3, Chicago, April-May, 1894, pp. 250-266; *Recent Discoveries of Cycadean Trunks in the Potomac Formation of Maryland*, *Bull. Torrey Bot. Club*, vol. xxi, No. 7, July 20, 1894, pp. 291-299; *Some Analogies in the Lower Cretaceous of Europe and America*, *Sixteenth Annual Report, U. S. Geological Survey, 1894-95, Pt. I*, Washington, 1896, pp. 463-542, pls. xcvi-cvii; *Descriptions of the Species of Cycadeoidea, or Fossil Cycadean Trunks, thus far discovered in the Iron Ore Belt, Potomac Formation, of Maryland*, *Proc. Biol. Soc. Washington*, vol. ix, March 13, 1897, pp. 1-17; *Descriptions of the Species of Cycadeoidea, or Fossil Cycadean Trunks, thus far determined from the Lower Cretaceous Rim of the Black Hills*, *Proc. U. S. Nat. Mus.*, vol. xxi (No. 1141), Washington, 1898, pp. 195-229; *The Cretaceous Formation of the Black Hills as indicated by the Fossil Plants*, *Nineteenth Annual Report, U. S. Geological Survey, 1897-98, Pt. II*, Washington, 1899, pp. 521-946, pls. lvii-clxxii; *Description of a New Genus and Twenty New Species of Fossil Cycadean Trunks from the Jurassic of Wyoming*, *Proc. Washington Acad. Sci.*, vol. i, pp. 253-300, pls. xiv-xxi; *Status of the Mesozoic Floras of the United States, First Paper, The Older Mesozoic*, *Twentieth Annual Report, U. S. Geological Survey, 1898-99, Pt. II*, pp. 211-748, pls. xxi-clxxix.

invoices were on their way, but it was impossible to wait for them to arrive and be described before going to press with the Nineteenth Annual Report. They came, however, during the summer, and added 44 specimens to the Yale collection, which I studied in November of that same year, but the results were not then published. In fact, the investigation was purposely left incomplete, because in the meantime, viz., in October of that year, I had been over the entire cycad-bearing area of the Black Hills in company with Mr. H. F. Wells, who had collected all the cycads for Professor Marsh, and had seen on the ground such an immense number of trunks and fragments that I had determined to make every effort to have these added to the collections already made. I had understood from Mr. Wells that Professor Marsh had declined to purchase any more, and I made a strong effort to induce the authorities of the U. S. National Museum to secure them on the extremely reasonable terms for which Mr. Wells offered to do the work. Failing in this, I appealed to Professor Marsh to secure them for Yale, which he did promptly, and before spring they had all arrived.

To give an idea of the extent and wealth of these latest accessions, it is only necessary to say that while, even after the arrival of the third invoice in the summer of 1898, the entire collection numbered only 170 specimens, it now numbers 731 specimens! It is true that many of these are fragments broken from larger trunks, but this was precisely what was needed to complete and perfect the collection. From a scientific point of view, fragments are often more valuable than perfect trunks, since they reveal the internal structure and throw light on the entire nature of the plants. Then again many of these fragments and *dissecta membra* are found to belong to specimens previously received, and add directly to their value. For example, the largest specimen in the collection, which, when the parts were gotten together and weighed, proved to be the largest trunk known in the world, came in different invoices and in four pieces. One branch, No. 145, was in the third invoice, and I had left it unassigned, well knowing that it was incomplete, and hoping that the remainder might ultimately be found. It did in fact arrive, and the parts have been brought into position and mounted in the exhibition hall on the second floor, where it may be seen of all men. Happily it proves to belong to the great branching species which I dedicated to Professor Marsh, and will stand forever as a tribute to his labors in this field—the *Cycadeoidea Marshiana*.

As was naturally to be supposed, and as I fully expected, the greater part of these numerous accessions have been found to belong with more or less certainty to one or other of the 22

species already described, but not only do they complete and supplement the previous collections in the manner above pointed out, but they greatly enrich it by adding many and often much finer specimens than any that existed before. First of all, they have added two species to the Yale collection which were formerly only represented in that of the U. S. National Museum. These are *Cycadeoidea excelsa* and *C. occidentalis*. In the second place, the beautiful *C. pubcherrima*, of which the type is at the National Museum, is now represented at Yale by at least one almost equally fine specimen, while the great *C. Jenneyana*, of which there were only a few broken pieces, now numbers its representatives by scores, some of them very fine trunks. The species which the new accessions have perhaps most richly endowed is the rare *C. Wellsii*, to which three or four huge and superb trunks are now added. Every species known from the Black Hills is now represented in the Yale collection, which was not previously the case, and it is safe to say that that collection constitutes the largest and finest assemblage of fossil cycadean trunks in the world.

Besides the specimens from the Black Hills, the Yale Museum possesses two specimens from the Jurassic cycad bed of the Freezeout Hills of Wyoming, treated by me in the paper above cited in the Proceedings of the Washington Academy of Sciences, and still further illustrated in the Twentieth Annual Report of the U. S. Geological Survey. These are Nos. 127 (*Cycadellu Reedii*) and 128 (*C. Beecheriana*), for the completion of which Prof. Wilbur C. Knight of the University of Wyoming, at my suggestion, has given to Yale the complementary fragment, No. 500.54, of his much larger collection from the same bed. There have also been discovered among the forgotten collections in the Yale Museum two fragments of cycads from the Iron Ore Clays of Maryland. When these trunks were refound, Professor Marsh at once remembered that he had secured them in 1867, from Mr. Philip Tyson, the original discoverer of cycads in Maryland. I have examined these specimens, and find them both to belong to the predominant species of that region, *Cycadeoidea marylandica*. They bear the numbers 729 and 730.

There have also now been found in the Yale Museum collections, 7 specimens of cycadean trunks from the original Purbeck forest beds of the Isle of Portland, that overlie the Portland stone. These are in two lots. The first consists of a single specimen, the gift of Dr. Gideon Mantell to Prof. Benjamin Silliman shortly after the attention of paleobotanists had been called to these objects, so long known to quarrymen as "crow's nests." It bears Mantell's label, *Mantellia nidi-*

formis Brongniart, with a reference to the figure in his Wonders of Geology, 1839, p. 365, fig. 70 (or is it to his Medals of Creation, in the later editions of which, but not in the earlier, the same figure occurs?). As the *Mantellianidi-formis* of Brongniart is the *Cycadeoidea megalophylla* of Buckland, it is evident that Dr. Mantell supposed the specimen to belong to this species, but it does not agree with his own figure, and seems rather to represent the small-scarred species, *Cycadeoidea microphylla* Buckland. This specimen has been numbered 732 of the Yale collection.

The other lot consists of 6 specimens which were received on Sept. 24, 1898, from Mr. A. N. Leeds, the collector and donor. They have now been given the numbers 733-738. Three of them (Nos. 733-735) belong without doubt to *Cycadeoidea microphylla* Buckland. No. 736 probably belongs to the same species. Nos. 737 and 738 have the large scars of *C. megalophylla*, and undoubtedly represent that species.

I made my fourth visit to the Yale Museum in May, 1900, and commenced work on the new material on the third of that month. My first work was to identify as many of the specimens as possible with previously described species. In many cases this was easy of accomplishment, but in dealing with the immense number of fragments and the imperfect, immature, dwarf, and depauperate specimens that necessarily accompany a complete collection, and which, in spite of the trouble they give, ought always to be collected, doubts constantly arise and cannot be removed. In the list which follows I have expressed such doubts by interrogation marks, and they are taken to indicate the tentative and incomplete character of the work. The future will doubtless greatly change matters in these respects, and the reader is asked to make due allowance for the unsettled state of our knowledge of these objects. A mere list of species with numbers attached is admitted to be unattractive reading, but it nevertheless affords an idea of the relative abundance of the species, and will have value as a guide or catalogue of the collection for any who may wish to study or examine it. An alphabetical arrangement of the species will probably prove more convenient than any attempt at systematic classification. The following are the 29 species thus far known, with the specimens in the Yale collection that have been assigned to each:

Cycadeoidea aspera, No. 104.

Cycadeoidea cicatricula, No. 118.

Cycadeoidea Colei, Nos. 12, 20, 25 ?, 28 ?, 48, 52 ?, 57, 68, 80 ?, 224 ?, 240 ?, 246 ?, 291, 321, 433, 444, 476 ?, 539.

Cycadeoidea colossalis, Nos. 2, 7, 10, 13, 17, 37, 40, 55, 133, 238, 333?, 351, 354?, 438? ?

Cycadeoidea dacotensis, Nos. 1, 3, 5, 6, 30, 39?, 43, 54, 61?, 62, 63, 95?, 106?, 134, 212, 213, 214, 215, 232, 307?, 324, 328, 353?, 355, 356, 358, 369, 372, 373?, 384?, 398, 425?, 435, 503?, 504?, 505?, 508?, 509, 510, 511?, 512, 513, 516?, 518, 523??, 525, 526, 528, 530, 538, 542, 543, 546, 547, 548, 549, 550, 716?, 719?, 720?

Cycadeoidea excelsa, Nos. 236, 239, 248, 370?, 453, 455, 461?, 479, 481, 498.

Cycadeoidea formosa, Nos. 89, 611, 615?, 639?, 641?

Cycadeoidea furcata, Nos. 18, 60, 718.

Cycadeoidea heliochorea, Nos. 722, 723, 724, 725, 726.

Cycadeoidea ingens, Nos. 92, 94, 99, 100, 103, 110, 117, 122, 123, 208, 554?, 557?, 560?, 561, 562, 563, 565, 566, 568, 572, 577, 578, 579, 580, 581, 582?, 584?, 592, 607, 609, 610, 612, 614?, 624, 625, 626, 628?, 633, 642?, 644, 646, 650, 653?, 656, 657, 661, 670, 671, 675, 676?, 678, 684, 690, 697, 705, 706?, 710.

Cycadeoidea insolita, Nos. 50, 64?

Cycadeoidea Jenneyana, Nos. 81, 87, 88, 90, 91?, 93, 96, 97, 98, 101, 102, 108, 109, 111, 112, 113?, 114, 115, 116, 120?, 121?, 124?, 125, 126, 171?, 173?, 174?, 176?, 192?, 193, 194, 195, 196, 197, 198, 199, 200, 201?, 202, 203, 204, 205, 206, 207, 551, 553, 555?, 556?, 558, 559, 564?, 567?, 569?, 570?, 571?, 573, 574?, 575, 576, 583, 585, 586, 587, 588, 589, 590, 591, 593, 594?, 595?, 596?, 597, 598, 599?, 600, 601?, 602?, 603?, 605?, 606, 608?, 613?, 616, 617, 618, 619, 621, 622?, 632?, 634?, 636?, 638?, 643, 645, 647, 648?, 649, 652, 654?, 655?, 658, 659, 660?, 663?, 665, 666?, 667?, 668?, 669, 672, 673, 674, 677, 680?, 681, 682??, 683?, 685?, 686, 688, 689?, 691?, 692?, 693, 694, 695, 696, 698, 699, 700, 701, 702, 703?, 708, 710, 721.

Cycadeoidea McBridei, Nos. 8, 19, 23, 26?, 27, 29??, 38, 42?, 46, 73, 76, 179, 180, 189?, 225?, 227, 228??, 231, 233?, 235, 245, 250, 252, 257?, 258??, 262, 268?, 271?, 288, 289, 298, 299, 301?, 305, 314?, 320?, 325, 326?, 329?, 338?, 339?, 340, 343?, 346?, 361?, 375, 379, 380, 381, 383?, 387?, 388, 407, 415, 422, 429, 431, 436?, 437, 439?, 449, 456, 462?, 467, 469?, 471, 472, 489?, 490, 492, 494, 495?, 496, 497?, 522?, 544?

Cycadeoidea Marshiana, Nos. 4, 9, 11, 33?, 44, 79?, 129, 143, 145, 161, 164, 167, 169?, 216, 219?, 221?, 256, 267, 276, 281, 283?, 285, 300, 392, 405, 410, 411, 418?, 430, 502, 541, 711.

Cycadeoidea minima, Nos. 53, 149, 150, 152, 153, 154, 155, 156, 157, 168, 426, 468, 474, 478, 714.

Cycadeoidea minnekahtensis, Nos. 14, 22, 24, 34?, 41, 47, 70, 71, 72, 83, 86, 140?, 142, 144, 148, 158, 160, 162?, 181, 182?, 184, 186?, 210, 211, 217, 220, 229?, 230, 242, 244?, 249?, 259?, 264, 265, 266?, 273??, 277, 279?, 286?, 290?, 292, 309, 310, 318?, 327?, 331, 332, 344?, 348?, 349, 350, 357, 362?, 365?, 366, 367?, 396?, 397, 401, 403?, 404, 412, 420??, 432?, 440, 445, 473?, 480, 491, 506?, 507, 520, 532, 533, 534, 535, 536?, 540, 728.

Cycadeoidea nana, No. 84.

Cycadeoidea occidentalis, No. 234.

Cycadeoidea Paynei, Nos. 52?, 58?, 69?, 132, 135, 163?, 165, 188?, 247?, 263?, 272, 280?, 284?, 293??, 316, 319??, 334, 336?, 337?, 364, 376, 386, 395?, 399, 413?, 423?, 428?, 434?, 447, 448, 451?, 452, 493?, 712, 713.

Cycadeoidea protea, Nos. 32, 185, 187, 241, 253, 296, 297, 303, 315, 359, 382, 414, 457, 458, 463, 466, 487, 499, 521, 529.

Cycadeoidea pulcherrima, Nos. 78, 159, 545.

Cycadeoidea reticulata, Nos. 254, 282, 287, 335, 342, 377, 378.

Cycadeoidea rhombica, Nos. 620, 623, 627, 629, 630, 631, 640.

Cycadeoidea Stillwelli, Nos. 16, 36, 56?, 105, 107?, 119, 175?, 226?, 515??, 517?, 552?, 651??, 662??

Cycadeoidea superba, Nos. 137, 146, 147, 218, 717.

Cycadeoidea turrita, Nos. 15, 35?, 45, 49, 51, 65, 66, 67, 74, 75, 82, 85, 139, 141, 151?, 166, 183, 190?, 191?, 223?, 255??, 261?, 269, 270?, 278, 295??, 304?, 313, 323, 330, 352?, 360, 368, 374, 394, 402, 406, 408?, 409, 446?, 450?, 482, 483, 484, 486, 519??, 731.

Cycadeoidea utopiensis, No. 727.

Cycadeoidea Wellsii, Nos. 21, 59, 130, 136, 138, 222, 243, 322, 391, 400, 500, 501, 537.

Cycadeoidea Wielandi, Nos. 77, 131, 393, 424.

Indeterminable, Nos. 31, 170, 172, 177, 178, 209, 237, 251, 260, 274, 275, 294, 302, 306, 308, 311, 312, 317, 341, 345, 347, 363, 371, 385, 389, 390, 416, 417, 419, 421, 427, 441, 442, 443, 454, 459, 460, 464, 465, 470, 475, 477, 485, 488, 514, 524, 527, 531, 604, 635, 637, 664, 679, 687, 704, 707, 709, 715.

The specimens marked as "indeterminable" are largely fragments, i. e., small pieces broken away from the trunks to which they belonged, and it is to be hoped that an exhaustive comparison with all the specimens in the collection may result in the restoration of a considerable number to the trunks now in the collection. As an aid in this work I have indicated in many cases the species to which they are most likely to belong, and given other hints as to where to look with the greatest prospect of success.

Considerable search has already been made to bring such fragments together, and it has been successful in about twenty cases. As a natural result this has changed a few of the assignments previously made and published; thus, No. 9, doubtfully referred to *C. colossalis*, proves to be a part of the trunk, No. 33, of *C. Marshiana*; No. 13, referred to *C. dacotensis*, belongs to Nos. 17 and 40, constituting a fine trunk of *C. colossalis*; No. 47, thought perhaps to be a branch of *C. Marshiana*, fits No. 142, and both together make up a good specimen of *C. minnekahtensis*; and No. 70, which resembles

C. turrita and was so referred, when joined to No. 480, to which it belongs, proves to represent *C. minnekahtensis*.

Other alterations made are as follows:

In one case, viz., that of No. 110, I have changed the assignment previously made and published in the Nineteenth Annual Report of the U. S. Geological Survey (pp. 614, 615, pls. xcix and c), referring it to *C. ingens* instead of *C. McBridei*. This change is important because this specimen was the only one from the Blackhawk region that I had referred to the latter species. After an examination of the much larger collection now in hand from that region, I am satisfied that this specimen belongs to *C. ingens*, and that *C. McBridei* is not represented in that section of the Black Hills.

The small specimen, No. 53, which I regarded (loc. cit., p. 615) as perhaps a "miniature" or "undeveloped" form of *C. McBridei*, I now make the type of a new species, *Cycadeoidea minima*, represented by 17 specimens (see infra, p. 341), and I do the same with another small specimen, No. 32, which I doubtfully referred (loc. cit., p. 608) to *C. minnekahtensis* as a dwarf representative similar to the specimen No. 19 of the U. S. National Museum. This new species is now represented by 20 specimens, all but one of which are in the Yale Museum, and may bear the name *Cycadeoidea protea* (see infra, p. 343).

After all possible assignments had been made and the indeterminate material separated out, there remained 60 specimens which, while exhibiting good specific characters, were not referable to any of the species described. These 60 specimens, however, proved to belong to very different types, and a careful classification of them shows that they constitute 7 distinct specific groups, or, in other words, 7 new species. The systematic description of these new species will therefore be our next and final task.

Cycadeoidea superba n. sp.

Trunks large (30–40^{cm} high, 30–50^{cm} in diameter, with a girth of over one meter), short-conical or somewhat globular, little compressed, unbranched; rock soft but not fragile, reddish brown, of low specific gravity; organs of the armor horizontal except near the summit where they are ascending and pass into a large terminal bud, which, however, is sometimes wanting, and a cavity, or crow's nest, occupies the summit; scars definitely arranged in spiral rows around the trunk, the angle made by the rows with the vertical axis diminishing above, those from left to right making an angle of 20–35°, those from right to left from 60–75°; leaf scars subrhombic, high in proportion to their width, the vertical angles usually

rounded, 30–40^{mm} wide, 20–25^{mm} high; leaf bases present, usually 1–3^{cm} below the surface, sometimes filling the scars or even projecting, soft, porous or spongy in structure; vascular bundles occasionally visible in a row some distance from the margin, appearing either as depressions or elevations; walls 5–8^{mm} thick, very distinct, lighter colored than the leaf bases, smooth but more or less grooved on the surface, sometimes divided by a commissure in two plates; reproductive organs numerous and well developed, sometimes tending to arrange themselves in vertical rows, raised above the surface, usually large, 4×7^{cm}, or even 6×9^{cm} in diameter, surrounded by numerous subtriangular bract scars covering much of the surface, the central portion relatively small, usually solid and amorphous, of a spongy consistency; armor 7–8^{cm} thick, obscurely attached to the axis; woody zone 5–8^{cm} thick, showing 2 or 3 rings; cortical parenchyma 3^{cm} thick; fibrous zone 3–5^{cm} thick, in one specimen consisting of two rings, the outer 12^{mm} thick, showing both longitudinal and radiate structure, the inner 8^{cm} thick, apparently subdivided into three subordinate rings; medulla 10^{cm} or more in diameter, of a homogenous structure, usually decayed at the base.

This fine species is represented in the Yale collection by 5 specimens, the numbers of which, with their weights and state of preservation, are as follows:

No. 137,	48·35 kilograms,	nearly complete.
No. 146,	45·36	“ “ “
No. 147,	39·90	“ a hemisphere.
No. 218,	56·12	“ nearly complete.
No. 717,	9·75	“ incomplete.

Nos. 137 and 146 must have grown close together, as each has a flattened area near the base, and these two surfaces exactly fit together. They are, however, covered by the appressed leaf scars and had no organic attachment. Moreover, each has its own independent base and axis, and there is no proof that they were connected otherwise than mechanically. No. 147 also has some similar flattened areas, and it seems to have been the habit of this species to grow in clumps or clusters.

Its nearest affinities are with *C. McBridei*, but it differs in its globular, symmetrical form, in the more open scars, and in their smooth well-defined walls, also in the finer structure. Its next nearest relationship is with *C. Wellsii*, but it lacks the thick walls and most of the other essential features of that species.

All the specimens are from the Minnekahta region.

On Plate III may be seen side views of Nos. 137 and 146, showing also the terminal bud.

Cycadeoidea rhombica n. sp.

Trunks of medium size (40–50^{cm} high, 20–30^{cm} in diameter), subcylindrical, tapering upward, more or less laterally compressed, unbranched; rock soft on the immediate surface, hard and flinty within, reddish brown on the weathered surfaces, whitish and somewhat chalcedonized in the interior, with rather high specific gravity; organs of the armor horizontal or somewhat descending; phyllotaxy clear and well marked, the spiral rows from left to right forming an angle with the axis of 25°, those from right to left of 60°; left scars strictly rhombic with parallel sides and all angles, vertical as well as lateral, sharp, 15^{mm} wide, 9^{mm} high, and very uniform; leaf bases present usually filling the scars, square across the top but lying at different levels, rough-granular and somewhat porous, without apparent bundle scars; walls very thin, rarely 1^{mm}, consisting of a white flinty substance, the surface smooth and longitudinally grooved or pitted, often not rising to the summit of the leaf bases, thus forming cracks between them, often split into two, more or less equal, thin plates; reproductive organs present but not specially numerous or conspicuous, fairly well developed, usually projecting or forming elevations, but occasionally decayed, leaving cavities, small, 15×25^{mm} in diameter or smaller, the involueral bract scars inconspicuous, the central portion solid and amorphous externally; armor very thin, 15–20^{mm}, joined to the axis by a more or less irregular line sometimes appearing definite; woody zone about 4^{cm} thick, usually showing three distinct layers; cortical parenchyma 12–15^{mm} thick, often conspicuously marked by the thick vascular bundles passing across it and curving upward and outward to the leaf bases, its outer wall marked by shallow, longitudinal grooves, 6–10^{cm} long, pointed at their extremities, lying side by side but overlapping one another; fibrous zone in two distinct rings, the outer 10–15^{mm} thick, the inner 12^{mm} thick, separated from the outer by a definite line appearing on the fractures as a seam or crack, its inner wall also definite, both rings appearing longitudinally striate on radial fractures, but both, and especially the outer, showing on transverse fractures a radiate structure with woody wedges; medulla 6–8^{cm} thick, porous in its outer, and cherty or flinty in its inner, portion.

This is one of the best-defined species in the collection, although none of the specimens are complete. It consists of Nos. 620, 623, 627, 629, 630, 631, and 640, all from the Blackhawk region. These specimens all came in the same box with a number of others of different species (*C. Jenneyana* and *C. ingens*), and it seems probable that they were found close together. Still a comparison of them shows that they

probably represent at least three different trunks. Nos. 620 and 623 are the bases of two different trunks. Nos. 627 and 630 fit together, and Nos. 629 and 631 probably belong to that trunk. No. 640 is a small fragment clearly showing the same characters, but not known to belong to any of the other specimens.

The only affinities shown by this species with any other are with *C. Stillwelli*, and here the resemblance is confined chiefly to one specimen, No. 105, yet the scars of that specimen are not only larger but have more or less curving sides, and the leaf bases are not so squarely truncated as in these forms. It is also more strictly cylindrical. Further material may tend to assimilate these specimens; if so the effect will be to remove that specimen from *C. Stillwelli* and enlarge the scope of this species.

The specific name relates to the exactly rhombic scars. On Plate II a view of the fragment No. 629 is represented, which shows the rhombic scars to good advantage.

Cycadeoidea heliochorea n. sp.

Trunks very large, the largest probably 50^{cm}, ellipsoidal or nearly globular, flattened at the summit where a small terminal bud is set in the center of a broad surface occupied by the small, spirally arranged scars of the upper leaves, laterally compressed to an unknown extent, all the specimens probably representing the broader sides, unbranched; rock soft, reddish brown, rather fragile, of low specific gravity; organs of the armor ascending in all the specimens, but these all belong to the upper part; phyllotaxy not traceable; leaf scars obscurely shown, subrhombic or rhombic, the upper vertical angle as sharp as the lower, 25^{mm} wide and 15^{mm} high in the larger examples, diminishing to the size of bract scars; leaf bases rough and structureless, often sunk 2–3^{cm} below the surface, showing in a few cases faint traces of bundles in the form of projections from the bottom of the scars or ribs in the sides above the bottom; walls very thick and irregular, of the same texture as the leaf bases, presenting a rough, jagged, and uneven surface with scarcely any longitudinal arrangement or subdivision into plates; reproductive organs very numerous and prominent, nearly covering the surface to the exclusion of the leaves, usually projecting 3 or 4^{cm} as decorticated cones, sometimes absent, having fallen out, leaving deep, bowl-shaped cavities, usually large, attaining a diameter of 4×7^{cm}, but smaller ones are also common, surrounded by numerous and conspicuous, concentrically arranged, involueral bract scars, the inner bracts themselves still present forming involucre to

the projecting spadices which they invest, the central portions consisting of obovate, longitudinally striate fruits having definite and peculiar pits and markings; armor 8^{cm} thick, the leaf bases appearing to blend with the large curving vascular strands that are seen passing through the cortical parenchyma; fibrous zone and medulla not represented.

This well-marked species consists, so far as known, of the 5 fragments Nos. 722-726, collected by Mr. Wells ten miles west of Sundance in Wyoming, 90 miles northwest of the Minnekahta localities. These specimens therefore possess an especial interest from the point of view of distribution, the locality lying between the old one and the Hay Creek region from which Professor Jenney made his fine collection of fossil plants. Nos. 722-724 consist of broad portions of the armor at the upper part of three different trunks, the first two including the terminal bud and broad, flat surrounding area. No. 722 is the largest specimen, measuring 28^{cm} the longest way. They all show a curvature on both sides, which if carried all the way round would make an immense trunk. It is altogether probable that they may represent the flattened sides of much compressed trunks. It is greatly to be hoped that more material may be obtained from this locality.

The specific name, from the Greek words for sun and dance, alludes to Sundance, the only name with which the locality has been associated.

Plate IV shows the outer surface of No. 722, with the small scars spirally arranged around the small terminal bud.

Cycadeoidea utopiensis n. sp.

Trunk small (21^{cm} high, 17×22^{cm} in diameter, with a girth of 63^{cm}), irregularly short-conical, elliptical in cross section but apparently not forcibly compressed, showing one small branch and a cavity from which a second has been removed, the surface well preserved, but having an area near the summit covered by what appears to be an outer coating of ramentum as in the genus *Cycadella*, more or less obscuring the organs and definitely broken away on three sides including the apex; terminal bud well developed, 5^{cm} high, 7×9^{cm} in diameter, rounded towards the summit where there is a depressed area 3^{cm} in diameter occupied by small rhombic scars surrounding the somewhat projecting axis 12^{mm} in diameter; base preserved in great part, somewhat hollowed out; rock soft except the central portion which is hard and fine grained, light drab color on the surface with a somewhat calcareous appearance, but containing no lime, jet black at the center, of rather low specific gravity, weighing 8.17 kilograms; organs of the armor nearly

horizontal in the middle part of the trunk, declined near the base and progressively ascending above to the erect terminal leaves; phyllotaxy faintly traceable over some small areas near the base, indicating that the spiral rows from left to right make an angle of about 50° , and those from right to left one of about 60° , with the axis; leaf scars subrhombic, the vertical angles rounded, the upper often reduced to a curve, only those near the base normal or distinct, the rest much reduced and distorted, crowded and massed together so as to present a peculiar wrinkled and gnarly appearance, small and very narrow in proportion to their width, normal ones 2^{cm} wide, 1^{cm} high; leaf bases always present, of an open structure, with deep, variously-shaped pores separated by thin partitions, some projecting, others depressed, often showing the leaf bundles, which are arranged in a row all the way round some distance from the margin, thus appearing near the center always as large distinct pits; walls 2–5^{mm} thick, white, flinty and fine grained, much grooved and divided longitudinally, those of the central and upper parts of the trunk consisting of two smooth plates separated by a deep furrow, rising above and surrounding the small leaf bases so as to enclose them and form numerous shallow cups or short tubes, the whole giving to the surface a peculiar wrinkled appearance; reproductive organs common but some distance apart and not the cause of the general distortion, normal and fairly well developed, rising a little above the surface, variable in size, some having a diameter of $25 \times 40^{\text{mm}}$, others only of $15 \times 25^{\text{mm}}$, surrounded by very numerous subrhombic involneral bract scars resembling and gradually passing into the leaf scars, the central portion showing markings that represent the essential organs; armor about 25^{mm} thick, somewhat definitely joined to the axis; woody zone 3–4^{cm} thick, the parts incapable of being distinguished or described, but consisting in large part of loose open structure; medulla $6 \times 10^{\text{cm}}$ in diameter, black within, hard and fine grained.

This species is represented by the single specimen, No. 727, purchased by Mr. Wells from a dealer in Hot Springs, who stated that it was obtained from some unknown person who claimed to have found it "50 miles west of Hot Springs in Wyoming." If the direction were due west, this point would fall in longitude $104^\circ 30'$ west from Greenwich and 12 miles west of the line of the Cretaceous border on the maps of the Black Hills. As the foot hills end about 20 miles west of Hot Springs and are succeeded here as in all parts of the Hills by the Upper Cretaceous, and these again by higher deposits, one would naturally suppose that this locality would fall far out on the Tertiary terrane. It therefore seems more probable that

the direction was northwest from Hot Springs, and this might locate it in the Lakota formation some distance north of Cambria and in the general region of the Newcastle coal field. This, could it be proved, would be interesting as supplying another link in the cycad chain which is fast encircling the Hills, and would partially close the wide gap between the Minnekahta and the Sundance regions. But, as it now stands, everything is in doubt, and it is greatly to be hoped not only that we shall ultimately learn the true locality, but that other specimens may be found. The patch of ramentum, if such it be, near the summit of the specimen raises the suspicion that it may belong to the genus *Cycadella*, and as all the specimens of that genus thus far known have come from the Jurassic, it is possible that the horizon of the bed holding this specimen may be lower than that of the other Black Hills cycads.

The specific name alludes to the alleged locality, which, as we have seen, would be a sort of geological nowhere.

Plate III shows a view of the best-preserved side of the specimen.

Cycadeoidea reticulata n. sp.

Trunks small (9–15^{cm} high, 10–20^{cm} in diameter, with a girth of 40–50^{cm}), globular or oblate-spheroidal, vertically compressed and also elliptical in cross section, unbranched, usually hollowed at the summit, but in one specimen showing the worn bases of the terminal leaves concentrically arranged, depressed at the base with a concave center surrounded by a groove between the armor and medulla; rock generally soft, but sometimes hard in the interior, reddish brown or lighter on the weathered surfaces, drab or dark within, of rather low specific gravity, the specimens all weighing less than 4 kilograms; organs of the armor normally horizontal; phyllotaxy more or less distinct, the rows of scars from left to right making an angle of 45°, those from right to left of 60°; leaf scars subrhombic, 15–20^{mm} wide, 6–9^{mm} high; leaf bases soft, porous or spongy, sunk 1–5^{mm} below the surface, occasionally with indistinct bundle scars near the center; walls very thin, often less than 1^{mm}, rarely exceeding 2^{mm}, hard and fine grained, presenting a smooth white surface contrasting strongly with the leaf bases, rising above them in plates or dikes or separating them with a delicate network of fine lines, thus giving the whole surface a reticulate appearance, sometimes divided by a commissure into two plates, or by two such into three; reproductive organs not abundant or conspicuous, rarely wanting, their places occupied by cavities, usually small, but variable in size, vaguely defined, averaging 15x20^{mm} in diameter, having scarcely any visible involucre bracts, the central portions porous or open, the hol-

low interior in one instance containing a collection of small round bodies of doubtful nature; armor 2–3^{cm} thick, somewhat definitely joined to the axis; woody zone 2–4^{cm} thick; cortical parenchyma 1–2^{cm} thick, generally of a loose porous structure; fibrous zone 1–2^{cm} thick in two rings, the outer thicker, dark, hard, and fine grained, the inner thin (5–7^{mm}), partitioned off by medullary septa and bounded on both sides by scalloped lines; medulla 5x8^{cm} in diameter, of a coarse and soft structure.

This species embraces 7 specimens, the numbers of which, with their weights and state of preservation, are as follows:

No. 254,	1.02 kilograms,	incomplete.
No. 282,	3.42	“ nearly complete.
No. 287,	3.86	“ “
No. 335,	3.18	“ “
No. 342,	1.25	“ incomplete.
No. 377,	3.42	“ “
No. 378,	2.61	“ “

These all came from the original Minnekalita locality. Their nearest affinities are with *C. turrita*, and they resemble certain of the branches or turrets of that species, but are all entire trunks and show no signs of branching.

The specific name refers to the reticulate appearance produced by thin white walls that form a network over the surface.

A side and top view of No. 342 are given on Plate IV. This specimen shows the leading characters as clearly as any example.

Cycadeoidea minima n. sp.

Trunks diminutive, the smallest known (6–12^{cm} high, 8–14^{cm} in diameter, with a girth of 20–39^{cm}), ovoid or obovoid, eccentric or oblique, sometimes vertically compressed or flattened at the top, simple, with a terminal bud or corresponding depression, the base usually hollowed out, but sometimes downwardly projecting; rock generally soft and porous, reddish brown or drab colored, of low specific gravity; organs of the armor horizontal except near the summit; phyllotaxy traceable in several specimens, the rows of scars from left to right making an angle with the axis that varies from 30–50°, those from right to left making an angle in all cases of 75°; leaf scars subrhombic to nearly rhombic, 15–20^{mm} wide, 7–9^{mm} high; leaf bases nearly filling the scars, sometimes rising above the walls, presenting a rough, spongy appearance, often having a pair of large pits near the center, and occasionally showing indistinct marginal rows of bundle scars; walls thin, often less than 1^{mm}, hard and flinty, light colored or white, wrinkled, striate, or grooved on

the surface, sometimes with a visible commissure; reproductive organs few and obscure, small and inconspicuous, never exceeding 3^{cm} in diameter, the bract scars very obscure and the central portion mostly solid; armor poorly exposed, 2–4^{cm} thick, separated from the axis by a somewhat definite line; woody zone 2–3^{cm} thick; cortical parenchyma 1^{cm} thick, traversed by the distinctly visible strands; fibrous zone 1–2^{cm} thick, not differentiated, somewhat definitely bounded on both walls, the outer often a scalloped surface, the inner marked by longitudinal ridges forming narrowly rhombic meshes; medulla 2–5^{cm} in diameter, porous or spongy, homogeneous and structureless.

This species embraces 15 specimens of the Yale collection and one from that of the U. S. National Museum, which are the smallest complete cycadean trunks known. The numbers, weights, and state of preservation, are as follows;

No. 53,	1.57 kilograms,	complete.
No. 149,	1.22	“ “
No. 150,	1.39	“ “
No. 152,	0.79	“ “
No. 153,	0.45	“ “
No. 154,	0.45	“ “
No. 155,	0.34	“ “
No. 156,	0.30	“ “
No. 157,	0.23	“ nearly complete.
No. 168,	2.15	“ complete.
No. 426,	0.30	“ fragment.
No. 468,	0.45	“ nearly complete.
No. 474,	0.45	“ incomplete.
No. 478,	0.11	“ fragment.
No. 714,	1.36	“ complete.
U. S. Nat. Mus. specimen,	0.62 kilograms,	nearly complete.

When I made my first study of the 87 specimens then in the Yale collection, I found No. 53 of this group, and rather than create a new species for it I doubtfully referred it to *C. McBridei*, as previously stated. Mr. Wells's next shipment, which I studied in November, 1898, contained 9 specimens so much like it that I was then certain that they constituted a specific group. The later accessions have added 5 others that belong to the same group.

In October of that year, when in the field with Mr. Wells, I picked up a small perfect specimen in the region southeast of Minnekahta station, and on account of its diminutive size I brought it back with me to Washington in my valise. It also belongs in this group. It now bears the number 2248 of the locality catalogue of fossil plants of the U. S. Geological Sur-

vey, and will form a part of the cycad collection of the U. S. National Museum. Its specific gravity is considerably higher than that of the other specimens, but beyond this there are no essential differences.

On Plate II is represented the original specimen, No. 53, which is typical of the species, but was not figured in the Nineteenth Annual Report of the U. S. Geological Survey.

Cycadeoidea protea n. sp.

Trunks small (8–12^{cm} high, 10–20^{cm} in diameter), low, sprangling, and of all shapes, often contracted at the base, the distortion probably only slightly due to compression, usually much branched, or consisting entirely of several somewhat equal systems or branches, often with a well-preserved terminal bud at the summit of each branch, usually nearly complete and little broken or worn; rock hard, firm, and fine grained, reddish brown on the surface with lighter stripes, darker within, of high specific gravity; organs of the armor wholly under the influence of the subordinate systems or branches, usually erect with reference to the terminal buds; phyllotaxy also relating wholly to the branches, around which and the often flattened summits the leaves are spirally arranged, sometimes with great regularity; leaf scars, where visible, of nearly normal subrhombic shape, but very variable in this respect due to the exuberant branching, often reduced to mere slits; very small, 6–13^{mm} wide, averaging 9^{mm}, 1–6^{mm} high, averaging 4^{mm}; leaf bases porous or showing a columnar structure but of a hard substance, usually sunk some distance below the surface, occasionally exhibiting a few large pits; walls very thin, often less than 1^{mm}, hard and flinty, of a lighter color than the leaf bases and divided into several plates the edges of which present a striate or grooved appearance, the middle plate sometimes divided by a commissure; reproductive organs small, few, and obscure, but certainly present and apparently functional, somewhat raised, 12×15^{mm} in diameter, surrounded by scars that scarcely differ from the leaf scars, the central portion heterogeneous, apparently showing the ends of the essential organs; armor where exposed about 2^{cm} thick, but certain leaf bases have a length of 4–6^{cm}, joined to the axis by a somewhat definite line; wood rarely exposed, appearing about 3^{cm} thick and divided into several rings, the outermost ring (cortical parenchyma) thin, the others showing no structure, the innermost wall (exposed in one specimen) marked with broad shallow grooves terminating in pits or scars and alternating with one another; medulla of a coarse sandy consistency, contrasting

with the surrounding hard tissues, often projecting downward below the armor, varying in size with the specimen.

I have already mentioned the specimen No. 32, and the similar one No. 19 of the U. S. National Museum Collection, both of which I took for immature forms of *C. minnekahtensis*. The latest invoices contained 19 additional specimens so like these as to render that theory no longer tenable, and I feel entire confidence in erecting this large and very distinctive group into a new species. In fact, notwithstanding the anomalous character of these forms and the consequent difficulty in finding terms by which to describe them, there is scarcely a species known to me that has less confusing relationships with other species than has this one. Though truly protean, as I have attempted to imply by the specific name chosen, none of its many forms is at all the same as those of any other species.

It now consists, as we have seen, of 21 specimens, 20 of which are in the Yale collection. The numbers, with their respective weights and state of preservation, are as follows:

No. 32,	2.61 kilograms,	complete.
No. 185,	0.94	" nearly complete.
No. 187,	0.14	" "
No. 241,	1.59	" "
No. 253,	1.02	" fragment.
No. 296,	1.25	" nearly complete.
No. 297,	0.34	" fragment.
No. 303,	1.59	" nearly complete.
No. 315,	0.91	" "
No. 359,	1.81	" "
No. 382,	1.25	" "
No. 414,	3.52	" "
No. 457,	1.47	" complete.
No. 458,	1.47	" incomplete.
No. 463,	0.56	" fragment.
No. 466,	0.68	" incomplete.
No. 487,	3.18	" nearly complete.
No. 499,	0.68	" " "
No. 521,	0.79	" " "
No. 529,	0.22	" " "
No. 19,	U. S. Nat. Mus.,	1.81 kilograms, nearly complete.

A careful study of all these specimens, and a reëxamination of No. 19 of the U. S. National Museum, have tended to strengthen the impression which the last-named specimen, considered alone, made upon my mind, and which I noted at the bottom of page 607 of my memoir on the Cretaceous Formation of the Black Hills, viz., that these small, gnarled, and branching forms have a decidedly rootlike appearance, and

suggest that the trunks, like those of the "coontie," may have developed entirely below the surface.

I have selected for illustration of this species the somewhat exceptionally regular, but still fairly characteristic specimen, No. 457, represented on Plate IV. This is a chiefly top view, showing the three nearly equal systems or branches, each with its apical scars and central bud perfectly preserved.

We now have, therefore, 29 species of fossil cycadean trunks from the Black Hills, all of which are represented in the Yale collection. To any one interested in the extinct floras of America this collection can scarcely fail to appeal in an especial manner. But we are probably only at the threshold of the subject. Not only will the raw material continue to accumulate, but as we penetrate deeper into the inner structure it is safe to predict that the results will be such as even the botanists proper cannot afford to ignore.

EXPLANATION OF PLATES.

PLATE II.

- No. 629.—*Cycadeoidea rhombica* Ward; type; side view. From near Blackhawk, South Dakota.
No. 53.—*Cycadeoidea minima* Ward; type; side view. Minnekahta, South Dakota.
(The numbers refer to the catalogue of the Yale Collection.)

PLATE III.

- No. 727.—*Cycadeoidea utopiensis* Ward; type; side view. From west of the Black Hills, Wyoming.
Nos. 137, 146.—*Cycadeoidea superba* Ward; type; side view. Minnekahta, South Dakota.

PLATE IV.

- No. 722.—*Cycadeoidea heliochorea* Ward; type; side view. From ten miles west of Sundance, Wyoming.
No. 342.—*Cycadeoidea reticulata* Ward; type; top view. Minnekahta, South Dakota.
No. 457.—*Cycadeoidea protea* Ward; type; top view. Minnekahta, South Dakota.

ART. XXXIV.—*On the Chemical Composition of Turquois* ;
by S. L. PENFIELD.

THROUGH the kindness of Mr. Ernest Schernikow of New York City, the writer has recently received a suite of turquois specimens from deposits in Los Cerillos Mts., New Mexico, and the Crescent Mining District, Lincoln Co., Nevada, and one fragment of exceptionally fine quality from the last named locality was presented with the special request that it should be used for chemical analysis. The material was very fine-grained, of a beautiful robin's-egg blue color, and broke with a smooth fracture. A thin section of the material appeared translucent and almost colorless, and when examined under the microscope, the turquois seemed to be perfectly uniform, showing no evidence of being made up of two substances, such, for example, as an aluminium phosphate, mixed with a copper salt as coloring material. The material was so fine-grained that no clue as to its crystallization could be made out, other than that it acted somewhat on polarized light. The specific gravity, taken by suspension in the heavy solution, was found to be 2.791.

In considering the chemical composition of turquois, it should be borne in mind that analyses have been made of only massive, cryptocrystalline fragments, and although they may be selected ever so carefully, no such guarantee of the purity of the material can be given as when, for example, a well crystallized mineral is analyzed. In order to show, however, that turquois is a material of nearly uniform composition, the new analysis is given below in connection with analyses made by other investigators. Analyses have not been included which show a large proportion of foreign constituents other than silica. The analyses are as follows:

	I. Lincoln Co., Nevada. Penfield.	II. Nichabour, Persia. Church.*	III. Karkaralinsk, Russia. Nicolajew.†	IV. Fresno Co., California. Moore ‡	V. Los Cerillos, New Mexico. Three analyses by Clarke.§	VI.	VII.
P ₂ O ₅	34.18	32.86	34.42	33.21	31.96	32.86	28.63
Al ₂ O ₃	35.03	40.19	[35.79]	35.98	39.53 ^b	36.88	37.88
Fe ₂ O ₃	1.44	2.45 ^a	3.52	2.99	---	2.40	4.07
CuO	8.57	5.27	7.67	7.80	6.30	7.51	6.56
H ₂ O	19.38	19.34	18.60	19.98	19.80	19.60	18.49
Insol.	0.93	----	----	----	1.15	0.16	4.20
X	----	MnO .36	----	----	CaO .13	CaO .38	----
	99.53	100.47	100.00	99.96	98.87	99.79	99.83
Sp. gr.	2.79	2.75	2.89	2.86		2.80	

^a Given as 2.21 per cent FeO.

^b Includes some Fe₂O₃.

* Chemical News, x, p. 290, 1864.

† Kokscharow's Min. Russland, ix, p. 86, 1884.

‡ Zeitschr. Kryst., x, p. 247, 1884.

§ This Journal, III, xxxii, p. 212, 1886.

In the new analysis the iron was found to exist wholly in the ferric condition, and therefore the iron in Church's analysis, given as FeO in the original article, has been calculated to Fe₂O₃ to agree with the observations of the author and other investigators.

It is evident from an examination of the foregoing analyses that turquoise is a material which is quite uniform in its chemical composition, so uniform in fact that it does not seem reasonable to consider it as an accidental mixture of an aluminium phosphate and a copper phosphate. The presence of the bivalent element copper, however, in somewhat variable amounts, is not so easily accounted for if we are to consider a copper phosphate as isomorphous with an aluminium phosphate. The small amount of iron is probably isomorphous with the aluminium, and it is to be expected that the iron phosphate would have little effect upon the color of the stone, for the hydrated ferric-phosphate, strengite, and the hydrated ferric-arsenate, scorodite, are both light-colored minerals. The idea that the iron is present as the hydrated oxide, limonite, can scarcely be entertained.

Clarke,* in discussing the composition of turquoise, states that if the alumina is combined with the phosphoric acid and water to form a molecule 2Al₂O₃.P₂O₅.5H₂O, there remains an excess of phosphoric acid and water which forms with the copper a salt of the composition 2CuO.P₂O₅.4H₂O. Turquoise is considered therefore by him as consisting of variable mixtures of the foregoing salts. He regards normal turquoise as the aluminium salt, 2Al₂O₃.P₂O₅.5H₂O, which he also expresses as Al₂HPO₄(OH)₄, and "the copper salt, to which the mineral owes its color, is to be considered merely as an impurity." By means of ratios it is quite easy to apply Clarke's theory to the analyses as tabulated on page 346. Taking Al and Fe as a basis, and establishing a ratio of P:Al+Fe:H=1:2:5, as demanded by Clarke's formula for normal turquoise, Al₂HPO₄(OH)₄, the ratio of *the excess of Phosphorus:Cu:the excess of Hydrogen* can then be found. The results of the calculation are as follows:

		I.	II.	III.	IV.	V.	VI.	VII.
Normal turquoise.	} P	·352	·408	·373	·371	·387	·376	·396
		} Al+Fe	·704	·816	·746	·742	·774	·752
	} H		1·760	2·040	1·865	1·855	1·935	1·880
Residues.	} P	·130	·054	·111	·097	·063	·088	·008
		} Cu	·108	·066	·096	·098	·080	·094
	} H		·394	·108	·201	·365	·365	·300

* Loc. cit.

Taking copper as a basis, the ratios of P : Cu : H in the residues are as follows :

		I.	II.	III.	IV.	V.	VI.	VII.
Residues, } copper } as unity. }	P	1.20	0.82	1.15	0.99	0.78	0.93	0.09
	Cu	1.00	1.00	1.00	1.00	1.00	1.00	1.00
	H	3.65	1.63	2.09	3.72	4.56	3.19	0.90

It is to be taken into consideration that it is a very severe test of a formula to throw all of the errors resulting from possible impurities in the materials and inaccuracies of the analyses upon a single constituent, in the case in hand on the supposed copper salt; but still the ratios of P : Cu : H are so variable that it cannot be considered that turquois is a mixture of an aluminium salt, $\text{Al}_2\text{HPO}_4(\text{OH})_4$, and a hydrated copper phosphate having the definite composition $2\text{CuO} \cdot \text{P}_2\text{O}_5 \cdot 4\text{H}_2\text{O}$ as suggested by Clarke. A compound having the composition $2\text{CuO} \cdot \text{P}_2\text{O}_5 \cdot 4\text{H}_2\text{O}$ demands a ratio of P : Cu : H = 1 : 1 : 4.

An important factor to be taken into consideration is that the hydrogen in turquois is to be regarded as representing hydroxyl and not water of crystallization, for water is not expelled from the mineral at a low temperature; hence hydroxyl radicals may be considered as playing a part in the chemical composition of the mineral. Considering copper as an essential constituent of turquois and not as an impurity, two theories naturally suggest themselves: one, that the bivalent copper is isomorphous with, and replaces the bivalent aluminium-hydroxide radical $[\text{Al}(\text{OH})]''$; the other, that the univalent copper-hydroxide radical $[\text{Cu}(\text{OH})]'$ is isomorphous with the univalent aluminium-hydroxide radical $[\text{Al}(\text{OH})_2]'$. The first of these ideas has led to no satisfactory solution of the problem; the second, however, reveals a constancy in the chemical relations of the mineral which can scarcely be regarded as due to accident. The relations in question are shown by combining aluminium and iron with two hydroxyls to form the groups $[\text{Al}(\text{OH})_2]$ and $[\text{Fe}(\text{OH})_2]$, respectively, and copper with one hydroxyl to form the group $[\text{Cu}(\text{OH})]$, and then finding the ratio between the phosphorus and $[\text{Al}(\text{OH})_2]' + [\text{Fe}(\text{OH})_2]' + [\text{Cu}(\text{OH})]' + \text{Excess of hydrogen}$. The relations are shown by the ratios derived from the several analyses tabulated on page 346, as follows :

		I.	II.	III.						
P		.482	.462	.484						
$\text{Al}(\text{OH})_2$	} 1.450	.686	.788	.702						
$\text{Fe}(\text{OH})_2$					} 1.332	.028	.044			
$\text{Cu}(\text{OH})$								} 1.320	.066	.096
H										

	IV.	V.	VI.	VII.
P	·468	·450	·464	·404
Al(OH) ₂	·706	·774	·722	·742
Fe(OH) ₂	·036	·--	·030	·050
Cu(OH)	·098	·080	·094	·083
H	·638	·572	·582	·387
	1·478	1·426	1·428	1·262

Considering $[Al(OH_2)]' + [Fe(OH)_2]' + [Cu(OH)]' + H'$ as playing the role of a univalent radical R' , the ratios of P : R in the several analyses are as follows :

I,	P : R =	·482 : 1·450 =	1 : 3·01
II,	" "	·462 : 1·332 =	1 : 2·88
III,	" "	·484 : 1·320 =	1 : 2·73
IV,	" "	·468 : 1·478 =	1 : 3·16
V,	" "	·450 : 1·426 =	1 : 3·17
VI,	" "	·464 : 1·428 =	1 : 3·08
VII,	" "	·404 : 1·262 =	1 : 3·12
			Average = 1 : 3·02

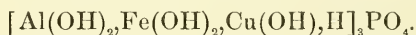
The author can vouch for the purity of the material analyzed by him, as far as it is possible to do so in the case of a cryptocrystalline mineral, and can also testify as to the accuracy of the analysis, and the very close approximation to the exact ratio 1 : 3, between the phosphorus and the sum of the univalent radicals plus the hydrogen, is very suggestive. The ratios in the other analyses approximate as closely to 1 : 3 as might be expected when the character of the material is taken into consideration, and the average of all the ratios is almost exactly 1 : 3. The ratio 1 : 3 is that of phosphorus to hydrogen in ortho-phosphoric acid, H₃PO₄. Turquois may therefore be regarded as a derivative of ortho-phosphoric acid in which the hydrogen atoms are to a large extent replaced by the univalent radical $[Al(OH)_2]$, $[Fe(OH)_2]$ and $[Cu(OH)]$. There seems to be no fixed ratio between the radicals $[Al(OH)_2]$, $[Fe(OH)_2]$ and $[Cu(OH)]$, nor between the sum of the hydroxyl radicals and the hydrogen. In some cases, however, there is an approximation to the ratio 2 : 1 between the sum of the hydroxyl radicals and the hydrogen, as follows :

	$[Al(OH)_2] + [Fe(OH)_2] + [Cu(OH)] :$	H	
II,	·882	:	·450 = 2 : 1·02
III,	·844	:	·478 = 2 : 1·13
VII,	·875	:	·387 = 2 : 0·89

In cases like the foregoing, the composition of turquois might be considered as a mixture of an aluminium salt, $H[Al(OH)_2]_2PO_4$, with the isomorphous molecules $H[Fe(OH)_2]_2PO_4$ and $H[Cu(OH)_2]_2PO_4$. The molecule $H[Al(OH)_2]_2PO_4$ is equivalent to Clarke's formula for "normal turquois," $2Al_2O_3 \cdot P_2O_5 \cdot 5H_2O$, which he also writes $Al_2HPO_4(OH)_4$. Adopting

Clarke's suggestion that turquoise contains very finely divided admixtures of iron and copper phosphates as impurities, and also his formula for the pure mineral (*normal turquoise* of Clarke), Groth* expresses the composition as $\text{PO}_4\text{Al}_2(\text{OH})_3\cdot\text{H}_2\text{O}$, but suggests, however, that the formula is perhaps $\text{PO}_4\text{H}[\text{Al}(\text{OH})_2]$.

In conclusion it may be stated that it is the author's belief that copper and the small amounts of iron are to be regarded as constituents of turquoise, rather than as impurities. In support of this idea the constant occurrence of copper, as shown by all the published analyses, may be cited. Furthermore, finely pulverized turquoise is only partially dissolved by boiling in a test tube with hydrochloric acid; hence, if the material contained copper phosphate as an impurity, it would be expected that the copper phosphate would dissolve readily, leaving the basic aluminium phosphate as a pure white residue, while in tests which have been made the insoluble residues have remained blue, from beginning to end of the experiments. Considering the existence in turquoise of the univalent radicals $[\text{Al}(\text{OH})_2]$, $[\text{Fe}(\text{OH})_2]$ and $[\text{Cu}(\text{OH})]$, the composition of the mineral, as shown by the published analyses, may be expressed as a derivative of ortho-phosphoric acid, as follows:



The $[\text{Al}(\text{OH})_2]$ radical always predominates, but is not present in fixed proportion. Some analyses (II, III, and VII) conform closely to the formula $[\text{Al}(\text{OH})_2, \text{Fe}(\text{OH})_2, \text{Cu}(\text{OH})]_2\text{HPO}_4$.

Disregarding the iron, the calculated composition of turquoise for two special cases of isomorphous replacements are given below:

$[\text{Al}(\text{OH})_2, \text{Cu}(\text{OH}), \text{H}]_3\text{PO}_4$, Al(OH) ₂ :Cu(OH):H=7:1:6.	Analysis I, page 346.	$[\text{Al}(\text{OH})_2, \text{Cu}(\text{OH})]_2\text{HPO}_4$, Al(OH) ₂ :Cu(OH)=12:1.	Analysis II, page 346.
P ₂ O ₅	34.64	34.18	32.86
Al ₂ O ₃	37.32	36.47 ^a	42.64 ^a
CuO	8.28	8.57	5.27
H ₂ O	19.76	19.38	19.34
----	Insol. 0.93	----	MnO 0.36
100.00	99.53	100.00	100.47

^a Include the Fe₂O₃.

Considering that turquoise is not a crystallized mineral, the agreement between theory and the analyses is certainly as close as could be expected.

* Tabellarische Uebersicht der Mineralien, 1898, p. 97.

Sheffield Laboratory of Mineralogy and Petrography,
Yale University, New Haven, June, 1900.

ART. XXXV. — *Quartz-muscovite rock from Belmont, Nevada; the Equivalent of the Russian Beresite*;* by J. E. SPURR.

Occurrence.

THE rock which will be described in this paper occurs in a large dike just east of Belmont, the seat of Nye County, Nevada. It affords an interesting study in view of its peculiar mineralogic composition; its relation to other siliceous rocks, into which it grades in the same dike; its identity with the Russian rock, beresite; and its connection with ore deposits. The dike in which the quartz-muscovite rock is found cuts black, limy slates and gray, fine-grained, saccharoidal, crystalline limestones, which stand vertically. The formation is pre-eminently a slaty one and often becomes schistose, this condition being due to the metamorphism occasioned by the dike under consideration and by other connected dikes. The main dike is half a mile wide and runs in a north and south direction; on one side (the east) it is definitely bounded by the slates, but on the west these slates form only a narrow band, a few hundred yards in width, separating the dike from a large mass of coarser intrusive siliceous granite which lies west and south of Belmont. While the body of rock under consideration is technically a dike, therefore, yet it is perhaps also to be regarded as the marginal facies of the main intrusion.

Near the contact of the dike, the shaly limestones become transformed into jasperoid, as microscopic examination shows, and the development of mica has occurred along certain planes, so that the rock passes into a micaceous schist. The jasperoid itself is often schistose and contains small bunches of yellow and red metallic oxides which give it the aspect of a knotted schist (Knotenschiefer). In the unaltered slates near here Mr. Gilbert† found graptolites which mark the rocks as Silurian. According to Mr. Walcott,‡ the horizon probably corresponds with part of the Upper Pogonip formation at Eureka. In this limestone are found quartz veins which carry rich antimonial silver ores.

Composition of the Dike.

The composition of this dike is not at all uniform, although as a whole the rocks are moderately fine-grained and siliceous. The essential constituents are quartz, feldspar and white mica, but the proportions of these vary in the different localities by

* Published with the permission of the Director of the U. S. Geological Survey.

† U. S. Geographical Surveys, vol. iii, *Geology*, p. 180.

‡ Monograph VIII, U. S. G. S., p. 2.

the increase of one mineral to the exclusion of the others, the changes being gradual and irregular; thus in some places the rock becomes mainly quartz, in others mainly feldspar. Quartz veins are abundant, but are of irregular form and are evidently segregational, being a part of the results of crystallization contemporaneous in a general way with the crystallization of the rest of the rock; these veins often contain considerable muscovite. Biotite is sparingly present in many of the rock types observed. Only two miles to the south this great, relatively fine-grained dike runs into a mass of coarse, siliceous, biotite granite with which it is apparently continuous, although it may be that the dike is a slightly later intrusion connected with the granite. The chemical and mineralogic composition of this coarse biotite granite is much the same as those of the finer-grained dike rocks above referred to, but the texture is very different. Further south again this coarse granite is overlain by flows of massive biotite rhyolite, which has nearly the same chemical and mineralogic composition as the granite, so far as can be discerned.

It is probable that the coarse-grained granite and the rocks of the finer-grained, siliceous, variable dike above described, together with the quartz veins of the vicinity, are all nearly contemporaneous, or closely consecutive allied phenomena, representing a single period of intrusion and a single magma, the variations in composition and period of injection being due to differentiation. It is probable, also, that the similarly constituted biotite rhyolite is connected with the intrusive rocks, which may possibly represent the roots or feeders of the volcanic out-pourings.

Microscopic examination.

Biotite quartz-monzonite.—This specimen is of a type which is very abundant in the dike, particularly near its central portion. It is fine-grained and holocrystalline in the hand-specimen and contains occasional small phenocrysts of feldspar, quartz and muscovite in a granular, saccharoidal groundmass. There is a slight gneissic structure apparent in the hand specimen and the fracture follows this, but it does not appear in the thin section; it is evidently not a cataclastic phenomenon but probably a flow structure. Under the microscope the groundmass is seen to be relatively fine-grained, granular and allotromorphic. It consists of a mosaic of quartz and feldspar, with some biotite, the proportion of the different minerals being, in the order named, about 15:10:1. The feldspar is both striated and unstriated. The unstriated feldspar was determined once by the Fouqué method as labradorite-bytownite. The striated feldspars were determined in one case to be andesine;

in another, oligoclase-albite. There is probably also considerable orthoclase. The rock, therefore, appears to belong to the monzonitic family and to the quartz-monzonite group.

Siliceous muscovite-biotite granite.—This rock has in general a fine grain, like that of the quartz-monzonite just described. In general, however, it is much lighter colored than this rock, since the mica which it contains is sporadic, leaving the rest of the rock composed of quartz and feldspar. Much of the rock observed in the field contains practically no mica, so that it passes into alaskite. The specimen studied is remarkable as containing frequent bunches of pure granular quartz, about a quarter of an inch in diameter, and resembling exactly vein-quartz. In thin section the rock is seen to be in general of medium grain, while it varies locally from very coarse to very fine. The habit of the minerals is allotriomorphic granular. Just as the texture varies so enormously, so the arrangement of the constituent minerals is irregular, some areas being entirely of quartz in irregular interlocking grains, while others consist of quartz and feldspar. The feldspar is mostly untwinned and was determined as orthoclase, while some twinned feldspar was determined as albite. There are in the section occasional tiny accessory grains of muscovite and biotite, and a large broken crystal of zircon. The feldspar in one section encloses primary muscovite, while in another case the muscovite encloses quartz. This would seem to make the order of crystallization quartz, muscovite, and feldspar, but the irregular intergrowth of the whole section shows that the crystallization of all these minerals was essentially contemporaneous. The quartz and feldspar are often intergrown in micrographic fashion. The feldspar is slightly kaolinized and in places there are very small flakes of secondary muscovite, which must be held separate from the rest of the muscovite in the rock, which is plainly primary.

Quartz-muscovite rock.—This rock is the one of chief interest in the dike, and occurs in large masses. It changes gradually and irregularly into the alaskite or muscovite biotite granite above described, and from its field relations is evidently a variation of this rock. The specimen examined, which is typical, has the exact appearance of a white or light-gray, medium-grained, micaceous quartzite. It is found especially near the margins of the dike and cuts into the sedimentary slates, which are noticeably metamorphosed into jasperoid and mica-schist near the contact. Under the microscope the rock is seen to be medium-grained, with a fairly even, allotriomorphic, granular structure. The chief minerals are quartz and muscovite, the former predominant. The rock is evi-

dently what it seems to be from field relations—a variation of the alaskite or muscovite-biotite granite, the variation consisting in the substitution of muscovite for feldspar. The proportion of muscovite to quartz in this rock is such that the rock is evidently nearly a chemical equivalent of the muscovite granite. Tiny grains of feldspar are found, generally striated, and seem to be albite.

Both in the hand-specimen and in the section this rock appears perfectly fresh and all the constituents primary. The freshness of the albite particles seems to remove all question of decomposition, for there is hardly any trace of kaolinization. A more thorough investigation, however, shows that muscovite occurs in two distinct habits. In the first habit it forms large irregular grains intergrown with the quartz, and sometimes enclosing smaller grains of quartz—poikilitic structure. In the second habit it occurs either in medium-sized blades or in small sheaf-like flakes, often spherulitic in arrangement, which are intergrown with quartz grains.

Concerning the muscovite found with the first described habit there is no question as to its primary nature. The last described habit, however, suggests a secondary origin, in spite of the freshness of all the minerals of the rock. Finally, a thorough investigation has shown that these aggregates of muscovite and quartz are plainly derived from orthoclase, some fragments of which may yet be found in the midst, with blades of muscovite and quartz grains penetrating them. It is plain that the alteration of the feldspars has been confined to the orthoclase and has not attacked the albite, which remains fresh.

The amount of primary orthoclase which the aggregates of secondary muscovite and quartz represent is so considerable as to put the primary rock among the fine-grained muscovite granites and to make it almost exactly similar to the muscovite granite just described, with which it is so closely connected in the field.

Biotite granite.—The biotite granite, into which this great fine-grained dike appears to merge south of Belmont, may be briefly described. The rock is porphyritic, carrying numerous crystals of orthoclase up to two inches in length. Under the microscope these orthoclase crystals are found to be intergrown with shreds of orthoclase differently oriented from the main crystal, but uniformly oriented among themselves. The crystals also contain shreds of muscovite. The groundmass is very coarse, and is remarkable for containing blotches of pure quartz in granular aggregates similar to that described above for the fine-grained muscovite-biotite granite of the great dike. In this case the blotches of quartz are from one-third of an inch to one-half an inch in diameter, and each

consists of a mosaic of intergrown grains. Besides this quartz, the groundmass consists of an allotriomorphic, granular intergrowth of quartz and feldspar, with subordinate biotite. The feldspar is mostly orthoclase. A faintly striated feldspar was determined to be anorthoclase, or microcline-anorthoclase. The feldspars sometimes show zonal structure.

Differentiation of the Great Dike.

This has not been carefully studied, but in general the darker colored rock, which has been shown in one specimen to be a quartz-monzonite, occupies the central portion of the dike, while the rock near the margin consists of the siliceous muscovite granite, or quartz-muscovite rock. This seems to indicate, at this locality at least, that the borders of the dike are more siliceous than the center.

Origin of Quartz-muscovite rock and connection with ores.

Although the quartz-muscovite rock is evidently a product of alteration, yet it is not due to surface weathering, since the condition of the rock is quite fresh and hard, and the albite shows no trace of decomposition. The process must be regarded as one of endomorphism and as connected and probably contemporaneous with the exomorphism indicated by the alteration of the siliceous limestone of the wall-rocks to jasperoid and mica schist. In both the intrusive and the intruded rock the result of the metamorphism has been the same, producing quartz and muscovite at the expense of the orthoclase on the one hand, and of the calcite and subordinate minerals on the other. In the case of the wall-rock the metamorphism, being apparently from its distribution dependent upon the intrusion, evidently took place after this intrusion and was brought about by the solutions which accompanied the igneous rock, or were residual from its solidification. Within the dike the similar alteration was probably contemporaneous with that in the country rock.

In the immediate vicinity of this intrusive mass are ore deposits, which in the time of Nevada's prosperity made this region one of considerable wealth, although at present the mining industry is perfectly dormant. The writer had not opportunity to study these, but according to Mr. Emmons* the ores generally occur in white quartz veins which are often several feet in width. These quartz veins are probably contemporaneous with those already described as occurring in irregular form within the dike rock itself, and as evidently rep-

* Geological Explorations of the 40th Parallel Mining Industry, vol. iii, p. 393.

representing the final product of the residual solution of the general magma. In these quartz veins the metallic minerals are scattered in bunches or disseminated particles, rarely in banded form. These metallic minerals consist principally of stetegeldite, which is an argentiferous ore of antimony, containing besides silver, lead, copper and iron. Dana* notes that this mineral in Peru has been regarded as probably arising from the decomposition of chalcostibite, a sulphide of antimony and copper. Chalcostibite occurs at Wolfsberg in the Hartz in nests embedded in quartz.

The metallic minerals being, from their habit, plainly contemporaneous with the quartz veins which enclose them, it is evident that the deposition of these minerals, the formation of the quartz veins, the metamorphism of the country-rock to jasperoid and muscovite schist, and the endomorphism of the muscovite granite to quartz-muscovite rock were contemporaneous occurrences, all brought about by the same agencies, which were the solutions representing the end product of the differentiation of the granitic intrusive rock.

The identity of the quartz-muscovite rock with the beresite of Russia.

The rock called beresite occurs in the vicinity of Beresovsk in the Urals, where it is intimately connected with veins of auriferous quartz. The beresite itself forms distinct dikes, varying from two to twenty meters, and reaching forty meters in width. These veins have a general north and south direction, but vary locally and interlace. They were first described by G. Rose† in 1837. He describes the constituents of the rock as orthoclase, plagioclase, quartz, and muscovite in exceedingly varying proportions. The feldspar diminishes and often completely disappears, leaving the rock composed of quartz and muscovite. Some varieties are like sericite schists; others like micaceous sandstones. Rutile is occasionally found as accessory. Rose was inclined to regard the Beresovsk veins as apophyses from the neighboring granite of Schartassh,‡ which is a fine-grained rock of semiporphyrific nature consisting chiefly of quartz and feldspar (the latter in part albite), with small flakes of biotite. The exact relation of the beresite to the neighboring granite, however, has not been ascertained.§ The beresite is intrusive into vertical or highly dipping schistose strata, either chloritic (listvenites), or talcose and argillaceous.

* System of Mineralogy, sixth edition, 1896, p. 204.

† Reise nach dem Ural, vol. i, p. 186; vol. ii, p. 557.

‡ Op. cit., vol. i, p. 189.

§ Op. cit., vol. ii, p. 559.

These schists strike north and south parallel to the beresite dikes; they often contain serpentine. The beresite had been so named by the miners of the district previous to Rose's investigation, and had been hunted for as the surest index to gold. The dikes are cut obliquely by highly inclined or vertical auriferous quartz veins. Ordinarily these veins do not extend beyond the beresite, but sometimes they enter the country-rock, and even extend to the next dike. The beresite itself, where it consists chiefly of quartz, muscovite and pyrite, notably near the river Tchérémechanka, contains, according to the researches of A. Sokolow, 50 drachms of gold to the ton. Rarely, native gold has been found in it.* A peculiarity of the beresite, which makes it difficult of investigation, is its profound decomposition, fresh portions being very rarely met with. In the upper horizons it is often altered to a light colored clayey mass, while the neighboring schists of the country-rock are altered to a red clay.

After the thorough description of beresite by Rose no other studies were made until Karpinsky published his results in 1876.† Karpinsky investigated what appeared to him to be fresh pieces of the beresite, and these showed themselves to be free from feldspar and to be composed of muscovite and quartz, with a little iron-pyrite. He came to the conclusion that the beresite is a feldspar-free rock, and, contrary to Rose's opinion, not to be connected with granite. In his second paper, Karpinsky admitted that orthoclase is present in beresite from another locality, and separated the rock into a feldspar-free and a feldspar-bearing variety, which are connected with one another by transitions. Besides the beresites at Beresovsk, a number of other Russian localities have been described by Rose,‡ while Karpinsky and Arzruni§ have added still others.

What appears to be the most thorough and clear-sighted study of the beresite was made in 1885 by Arzruni.¶ He describes the beresite as a fine-grained dike-rock of coarser or finer texture and often semi-porphyratic structure, the varieties of this rock being so varied in their peculiarities that it seems proper to describe the different occurrences separately. By the diminishing of one or the other of the mineral constituents, such as the feldspar or the mica, special types are presented.

* A. Karpinsky, *Guide des Excursions du VII Congrès Geologique International*, No. 5, p. 41.

† Protocol of the Geologic-Mineralogic section of the Natural Science Society at St. Petersburg meeting on the 3d of May, 1875, and on the 9th of September, 1876. German summary of the Russian text given by Arzruni, *Zeitschr. d. Deutsch. Geolog. Gesell.*, vol. xxxvii, 1885, p. 867.

‡ *Reise*, etc., vol. i, pp. 294, 302, 321, 436; vol. ii, pp. 34, 36, and 557.

§ *Zeitschr. d. Deutsch. Geolog. Gesell.*, vol. xxxvii, 1885, p. 870.

¶ *Op. cit.*, pp. 865-896.

Arzruni found, however, that those beresites which consist entirely of quartz and muscovite, and which had been regarded by Karpinsky as fresh feldspar-free rocks, actually contained feldspar in their original condition, although the feldspar has undergone a complete transition to mica and quartz. It appears, also, that the alteration has affected only the orthoclase, while the plagioclase has remained unaltered. Since in the different parts of the rock the relative proportion of orthoclase and plagioclase is very variable, it results that the varieties rich in plagioclase contain feldspar, even when the orthoclase is altered to quartz and muscovite, while the varieties poor or lacking in plagioclase can reach a transitional stage where they can be regarded as feldspar-free beresites. As the result of the study Arzruni concludes that beresite is a genuine muscovite granite with a tolerably constant mineralogic constitution.

The rocks described from Belmont are identical in nearly every particular with those of Beresovsk, except that in Belmont the decomposition has not affected the rocks to any extent, permitting a much more accurate study.

Analysis of Belmont rock.

A sample of the quartz-muscovite rock from Belmont proved to have the following composition :

Analyst, Dr. H. N. STOKES.

SiO ₂	84.15
Al ₂ O ₃	9.67
Fe ₂ O ₃51
FeO07
MgO04
CaO53
Na ₂ O	2.65
K ₂ O	1.57
H ₂ O—21
H ₂ O+74
TiO ₂	trace
F02
MnO	trace
SrO	trace

100.14

A rough computation of this analysis shows the proportion by weight of albite to muscovite to be about 5 : 6. The specimen then is one which contains much more unaltered feldspar than some of the other varieties. Those varieties which originally contained orthoclase, but no albite, would now be rocks free from feldspar.

ART. XXXVI.—*The Volumetric Estimation of Copper as the Oxalate, with Separation from Cadmium, Arsenic, Tin, and Zinc*; by CHARLES A. PETERS.

[Contributions from the Kent Chemical Laboratory of Yale University—XCVIII.]

It is a well known fact that copper oxalate is insoluble in water and scarcely attacked by moderate amounts of dilute nitric acid.* Upon this fact Bournemann† has recently based a method for the separation of copper from cadmium by precipitating copper as the oxalate in the presence of nitric acid, filtering hot, and estimating the copper after ignition, by any of the well known gravimetric methods. Six to ten grams of copper, as the oxide, were used for a single determination, and the errors were large. Bournemann does not recommend this process as an accurate analytical method. Classen‡ describes a method for the separation of metals as oxalates by adding to the solution of the salt of the metals a dilute solution of the potassium oxalate (1 : 6) and concentrated acetic acid to 80 per cent of the total volume. Regarding copper salts in particular, Classen states that precipitation takes place only in dilute solution and then not completely.

It has been the experience of the writer, that the precipitation of copper oxalate from solutions containing at least 0.0128 gm. of the oxide and saturated with the oxalic acid is practically complete. The filtrate in such cases gives no blue color with ammonia, looking down on a column of liquid in a test tube, and only a faint brown color is developed when the filtrate is neutralized, made acid with acetic acid, and tested with potassium ferrocyanide. It is the object of this paper to show that moderate amounts of copper may be determined quantitatively as the oxalate by precipitation with oxalic acid and titration of the precipitate by potassium permanganate, and also to show that moderate amounts of copper may be separated from other metals in the presence of nitric acid, by the addition of considerable amounts of oxalic acid.

Before attempting the quantitative separation of copper from solution by the addition of oxalic acid a few qualitative experiments upon the precipitation of varying amounts of copper sulphate by varying amounts of oxalic acid were tried at different dilutions. In all the experiments the mixtures stood 16–20 hours, and were filtered from 2 to 4 times through four filters folded together, and the filtrates were tested

* Storer, Dictionary of Chemical Solubilities, p. 463.

† Chem. Ztg., xxiii, 565.

‡ Ber., x, b, 1316.

both with ammonia and with potassium ferrocyanide. In cases where the filtrate gave no blue color with ammonia and only a slight precipitate with ferrocyanide the precipitation was considered practically complete and the conditions were regarded suitable for the trial of the method quantitatively. In the following table is recorded the work upon the precipitation of copper sulphate by 0.5 gm., 1.0 gm., and 2.0 gms. of oxalic acid in 50^{cm}³ of solution.

TABLE I.
Dilution 50^{cm}³.

	CuO taken as CuSO ₄ gms.	Oxalic acid added in solution.		Oxalic acid added in crystalline form.	
		Filtrate treated with NH ₄ OH	Filtrate treated with K ₄ FeC ₆ N ₆	Filtrate with NH ₄ OH	Filtrate treated with K ₄ FeC ₆ N ₆
2.0 gms. oxalic acid present.	0.018	blue color	abundant ppt.	blue color	abundant ppt.
	0.031	trace "	" "	trace "	evident "
	0.051	-----	evident "	-----	" "
	0.064	-----	trace "	-----	trace "
1.0 gm. oxalic acid present.	0.018	blue color	abundant ppt.	blue color	abundant ppt.
	0.031	" "	" "	trace "	evident "
	0.051	trace "	" "	-----	" "
	0.064	-----	evident "	-----	" "
0.5 gm. oxalic acid. present.	0.018	blue color	abundant ppt.	blue color	abundant ppt.
	0.031	" "	" "	trace "	" "
	0.051	trace "	" "	-----	evident "
	0.064	-----	evident "	-----	trace "
	0.094	-----	trace "	-----	" "

It will be seen readily by comparison of the right and left hand sides of the table above that somewhat smaller amounts of copper may be precipitated completely by the addition of crystallized oxalic acid than by the same amount of oxalic acid already in solution. Thus, when dissolved oxalic acid is added to the solution of 50^{cm}³ amounts of copper sulphate less than 0.040–0.050 gm. are not precipitated completely, while under conditions otherwise the same excepting that the oxalic acid is added in crystalline form, the precipitation of amounts as small as 0.030 gm. is practically complete. The amount of oxalic acid in solution necessary for the complete precipitation (after 16 to 20 hours) of this minimum amount of copper, 0.031 gm. of copper oxide taken as the sulphate, appears, as shown in Table II, A, which follows, to be about 3.5 gms. in 50.0^{cm}³. If the amount of oxalic acid is increased to 5 gms., making the solution saturated for that substance, using the

same volume of liquid, the minimum amount completely precipitable is reduced to 0.0128 gms. but not to one-half that amount.

TABLE II.

		A		
CuO taken as CuSO ₄ gms.	Oxalic acid added in solution. gms.	Volume at precipitation. cm ³ .	Filtrate treated with NH ₄ OH.	Filtrate treated with K ₄ FeC ₆ N ₆ .
0.031	0.5	50	blue color	abundant ppt.
"	1.0	"	"	" "
"	2.0	"	trace "	" "
"	3.0	"	slight "	" "
"	3.5	"	-----	evident "
0.0128	5.0	"	no blue	trace "
0.0064	5.0	"	blue color	abundant "
		B		
0.0064	0.5	20	faint blue	abundant ppt.
"	"	15	"	" "
"	"	10	"	faint "
"	"	5	-----	trace "
0.0003	0.5	5*	-----	-----
"	0.1	1†	-----	-----

It appears from the experiments of Table II, B, that the volume of liquid in which precipitation takes place influences the complete precipitation of the copper oxalate. Thus the precipitation of 0.0064 gm. of copper oxide taken as the sulphate by 0.5 gm. of oxalic acid is complete in 5 cm³ of liquid. The precipitate which falls from 0.0003 gm. of the oxide taken as the sulphate dissolves in 5 cm³ of liquid, but remains visible in 1 cm³.

As a result of the preliminary experiments, it may be said that the presence of a certain minimum amount of copper, varying with the conditions, is essential to complete precipitation. Thus, at a dilution of 50 cm³ a saturated solution of oxalic acid will precipitate with practical completeness copper taken as the sulphate in amounts exceeding the equivalent of 0.0128 gm. of copper oxide; that 2.0 gms. of oxalic acid will precipitate almost completely for the same volume of solution the equivalent of 0.03 gm. of copper oxide; and that 1.0 gm. or 0.5 gm. of oxalic acid will precipitate the equivalent of 0.064 gm. of the oxide.

In the quantitative separation of copper as the oxalate the method of treatment was in general as follows. Copper sul-

* Precipitate redissolved.

† Precipitate remained.

phate in 50^{cm}³ of water was thrown down by the addition of dry oxalic acid to the hot solution, and, after standing over night, the precipitate was filtered on asbestos, washed two or three times with small amounts of cold water. The precipitate, still in the crucible, was returned to the beaker in which precipitation took place, 5 or 10^{cm}³ of dilute sulphuric acid (1:1) were then added together with a convenient amount of water, and, after heating the liquid to boiling, the oxalic acid was titrated with permanganate, the oxalate of copper dissolving readily as fast as the excess of oxalic acid is removed by the permanganate. The precipitate may also be dissolved in 10^{cm}³ of strong hydrochloric acid,* 0.5 gm. manganous chloride added and titrated at 30°–50°. Experiments 4 and 5 were conducted after this manner. In Table III, A, which follows, are recorded results of the quantitative tests of the method.

TABLE III.

	CuO taken as CuSO ₄ . gms.	Oxalic acid. gms.	Volume at precipi- tation. cm ³ .	CuO found. gms.	Error. gms.
			A		
1	0.0372	0.15	100	0.0286	−0.0086
2	0.1860	0.50	125	0.1831	−0.0029
3	0.0398	“	50	0.0376	−0.0022
4	0.1860	1.0	150	0.1834	−0.0026
5	“	0.5	50	0.1864	+0.0004
6	“	“	“	0.1866	+0.0006
7	“	“	“	0.1866	+0.0006
8	“	1.0	“	0.1866	+0.0006
9	0.0398	“	“	0.0391	−0.0007

In experiments 1–4, deficiencies are found in the amounts of oxalate precipitated at different degrees of dilution and by different amounts of the precipitant which are in agreement with the results obtained in the preliminary work; the results of experiments 5–9, in which 0.5 gm. and 1.0 gm. of oxalic acid act in a total volume of 50^{cm}³, show the precipitation to be essentially complete under these conditions.

To study the insolubility of the copper oxalate in nitric acid the experiments in Section B of the table were made.

In experiments 10–13 amounts of oxalic acid varying from 0.5 gm. to 3.0 gms. appear to precipitate the copper completely in the presence of 5^{cm}³ of strong nitric acid. In experiment 14 the amount of oxalic acid used was not sufficient to throw down all the copper in the presence of 10^{cm}³ of nitric acid, but

* Gooch and Peters, this Journal, vii, 461, 1899.

TABLE III (continued).

	CuO taken as CuSO ₄ . gms.	Oxalic acid. gms.	HNO ₃ sp. gr. 1.40. cm ³ .	Volume at precipi- tation. cm ³ .	CuO found. gms.	Error. gms.
B						
10	0.1860	0.5	5.0	55	0.1859	-0.0001
11	"	"	"	"	0.1860	±0.0000
12	0.1990	2.0	"	"	0.1989	-0.0001
13	"	3.0	"	"	0.1990	±0.0000
14	"	2.0	10.0	60	0.1971	-0.0019
15	"	3.0	"	"	0.1987	-0.0003
16	"	"	"	"	0.1985	-0.0005
17	"	5.0	12.0	130	0.1977	-0.0013
18	"	"	"	"	0.1975	-0.0015
19	"	"	25.0	"	0.1837	-0.0153
20	"	"	"	"	0.1831	-0.0159
21	"	"	5.0	"	0.1983	-0.0007
22	"	"	"	"	0.1988	-0.0002
C						
23	"	2.5	5.0*	65	0.1971	-0.0029
24	"	2.0	"	"	0.1981	-0.0019

the copper does come down completely in the presence of the large amount of the nitric acid upon the addition of more oxalic acid, as seen in experiments 15 and 16. In experiments 17 and 18 with a larger volume of water and a larger absolute amount, though approximately the same percentage, of nitric acid present as in experiments 10-13, there is a slight loss of copper; but in experiments 21 and 22 when the amount of nitric acid is reduced to 5^{cm³} in the larger total volume the results are normal. Experiments 19 and 20 show the increased loss when still larger amounts of nitric are present. These facts would make it seem best to limit the absolute amount of nitric in solution to about 5^{cm³}.

One observation may well be noted here; namely, that while one-half gram oxalic acid is all that is needed for the complete precipitation of the copper in the presence of 5^{cm³} strong nitric acid, still the oxalic acid may be added up to the point of saturation of the solution. More than this causes difficulty owing to the fact that a large amount of water is necessary to wash the precipitated oxalate. About 2.0 gms. of oxalic acid to 50^{cm³} of water is a convenient proportion.

In experiments 23 and 24, 5^{cm³} of nitric acid were neutralized with ammonium hydroxide before adding the 5^{cm³} strong

* About 9 gms. of ammonium nitrate present in addition to the 5^{cm³} of nitric acid.

nitric acid in excess. The results show the solubility of copper oxalate in ammonium nitrate and exclude the possibility of such a procedure in this work.

Some experiments were made to show the time necessary for the complete precipitation, both in the presence and absence of nitric acid. Following is the record of such work.

TABLE III (continued).

	CuO taken as CuSO ₄ . gms.	Oxalic acid. gms.	HNO ₃ sp. gr. 1.40. cm ³ .	Volume at precipita- tion. cm ³ .	CuO found. gms.	Error. gms.	Details of filtration.
D							
23	0.1990	2.0	----	50	0.1984	-0.0006	} Filtered hot immediately
24	0.2030	"	----	"	0.2025	-0.0005	
25	0.1990	1.0	----	"	0.1990	±0.0000	} Filtered after cooling; stood 15 minutes
26	"	"	----	"	0.1987	-0.0003	
E							
27	"	2.0	5.0	55	0.1943	-0.0047	} Filtered after cooling; stood 15 minutes
28	"	"	"	"	0.1969	-0.0021	
29	"	"	"	"	0.1973	-0.0017	Stood 6 hours
30	"	"	"	"	0.1989	-0.0001	Stood 16 hours

The results in section D would seem to show that a solution containing copper may be precipitated hot as the oxalate and filtered either hot or after cooling with a very slight loss. Tests of the filtrates made with potassium ferro-cyanide confirmed these results. When nitric acid is present, however, the mixture must stand after the addition of the precipitant. In section E the gradual decrease of the minus error is noticed, as the time of standing is extended, the precipitation being practically complete upon standing over night.

Separation from Cadmium.

Bournemann* has used nitric acid for a rough separation of copper from cadmium. This method was tried for a quantitative separation in the presence of 6-10 per cent strong nitric acid. The results are found in section F of the table to follow.

* Loc. cit.

Experiments 33–35 stood six hours before filtering. Experiments 36 and 37 stood over night. Copper is separated from more than twice its weight of cadmium, and the results are accurate.

Separation from Arsenic, in Both Conditions of Oxidation.

For the separation from arsenic, arsenious oxide dissolved in sodium carbonate, and di-hydrogen sodium arseniate were the forms of arsenic used. The results are accurate and are given in sections G and H of the table. In experiments 38–40 and 44 and 45 no nitric acid was added. While the presence of the nitric acid is not necessary for the separation of the copper from the arsenic; still the filtration in the absence of the nitric acid is so slow as to be objectionable. The presence of the nitric acid causes the precipitate to come down in a coarser condition, and in such condition it filters easily and is capable of being washed quickly.

Separation from Tin, in Both Conditions of Oxidation.

For the separation of copper from tin a preparation of stannous chloride (20cm^3 giving 0.3746 gm. metallic tin by the battery) containing sufficient hydrochloric acid to prevent deposition of oxy-salts was used. The solution of stannic chloride contained 1.0 gm. metallic tin to every 10cm^3 , and was used without hydrochloric acid. The results of the work are found in sections I and K of the table. The experiments go to show that while copper may be separated from small amounts of tin as stannous chloride yet there is a limit to the amount of tin which may be present. One-tenth of a gram of metallic tin is the largest amount that can be present, with 0.15 gm. copper oxide taken as the sulphate, without significant error. Practically the same statement can be made of the separation of copper from tin taken as stannic chloride. Experiment 57 shows a greater loss of copper when the nitric acid is omitted.

Separation of Copper from Iron.

A solution of ferric nitrate was used for the work on the separation of copper from iron. Low results were obtained when a solution of ferrous or ferric sulphate was used as the source of iron. The results of the experiments are recorded in section L of the table, and show that 0.20 gm. copper oxide as the sulphate may be separated from 0.2–0.3 gm. iron oxide taken as the nitrate. In experiment 64 a good result was obtained when no nitric acid was present, save that added in combination with the iron. A comparison of experiments 63 and 65 shows that it is best to avoid the use of large amounts

TABLE III (continued).

	CuO taken as CuSO ₄ . gms.	Element from which copper was separated. gms. CdO taken as CdSO ₄ .	Oxalic acid. gms.	HNO ₃ sp. gr. 1.40. cm ³ .	Volume at precipitation. cm ³ .	CuO found. gms.	Error. gms.
				F			
33	0.1990	0.10	2.0	5.0	60	0.1983	-0.0007
34	"	0.20	"	"	65	0.1987	-0.0003
35	"	0.30	"	"	70	0.1987	-0.0003
36	"	0.40	"	"	75	0.1994	+0.0004
37	"	0.50	"	"	80	0.1996	+0.0006
		As ₂ O ₃ taken as Na ₂ AsO ₃ .		G			
38	"	0.10	"	----	55	0.1991	+0.0001
39	"	0.20	"	----	60	0.1987	-0.0003
40	"	0.50	"	----	75	0.1986	-0.0004
41	"	0.10	"	5.0	60	0.1994	+0.0004
42	"	0.20	"	"	75	0.1992	+0.0002
43	"	0.60	"	"	85	0.1995	+0.0005
		As ₂ O ₅ taken as H ₂ KAsO ₄ .		H			
44	"	0.10	"	----	60	0.1985	-0.0005
45	"	0.20	"	----	70	0.1990	±0.0000
46	"	0.10	"	5.0	65	0.1990	±0.0000
47	"	0.20	"	"	75	0.1992	+0.0002
48	"	0.30	"	"	85	0.1985	-0.0005
49	0.2030	0.30	3.0	"	85	0.2026	-0.0004
	Cu taken as CuSO ₄ .	Sn taken as SnCl ₂ +HCl.		I		Cu found.	
50	0.1590	0.0468	2.0	5.0	55	0.1581	-0.0009
51	"	0.0936	"	"	60	0.1603	+0.0013
51 α	"	"	"	"	"	0.1591	+0.0001
52	"	"	"	"	"	0.1594	+0.0004
53	"	0.1873	"	"	65	0.1603	+0.0013
54	"	0.2809	"	"	70	0.1914	+0.0324
55	"	"	3.0	"	75	0.1988	+0.0398
		Sn taken as SnCl ₄ .		K			
56	"	0.10	2.0	"	55	0.1581	-0.0009
57	"	0.10	"	----	"	0.1565	-0.0025
58	"	0.20	"	5.0	"	0.1577	-0.0013
59	"	0.50	"	"	60	0.1562	-0.0028
	CuO taken as CuSO ₄ .	Fe ₂ O ₃ taken as Fe(NO ₃) ₃ .		L		CuO found.	
60	0.1990	0.136	2.0	5.0	60	0.1987	-0.0003
61	"	0.272	"	"	"	0.1983	-0.0007
62	"	0.364	"	"	"	0.1988	-0.0002
63	"	0.544	"	"	65	0.1971	-0.0019
64	"	0.272	"	----	60	0.1995	+0.0005
65	"	0.544	"	2.0	"	0.1998	+0.0008
66	"	0.218	"	"	65	0.1999	+0.0009
		ZnO taken as ZnSO ₄ .		M			
67	"	0.028	"	5.0	60	0.2007	+0.0017
68	"	0.057	"	"	65	0.2008	+0.0018
69	"	"	"	"	"	0.2008	+0.0018
70	"	0.085	"	"	70	0.2035	+0.0045

of nitric acid when the larger amounts of ferric nitrate are present.

For a practical application of the above separation of copper from iron a convenient amount of finely ground chalcopyrite (0.5 gm.) was roasted 2–3 hours in a porcelain crucible until all sulphur was driven off, washed into a beaker, strong nitric acid about 5^{cm}³ was added and, with the beaker covered, allowed to evaporate slowly on a hot plate, nearly to dryness. A little dilute nitric acid was added, the solution was filtered, the residue was washed with water containing dilute nitric acid, the filtrate, about 50^{cm}³ in volume, was precipitated with 2.0 gms. oxalic acid, and the precipitate was estimated after standing 12–16 hours, as previously described. The washing with water acidified with nitric acid is important because the finely ground ferric oxide remaining undissolved passes through the filter when washed with water alone, but gives no trouble if the water be acidic. The results of two estimations are here given.

Chalcopyrite. gms.	Copper found by battery.	Copper found by oxalate method.	Difference.
0.5000	31.00%	30.92%	–0.08%
1.0000	“	31.25	+0.25

Separation of Copper from Zinc.

The separation of copper from zinc was not altogether successful owing to the tendency of the zinc oxalate to come down with the copper oxalate. Some experiments are given in section M of the table.

The separations of copper from bismuth and antimony were unsuccessful.

The work may be briefly summarized as follows: Copper exceeding in amount the equivalent of 0.0128 gm. of the oxide to 50^{cm}³ of solution as the sulphate may be separated completely, even in the presence of a moderate amount of strong nitric, by the addition of sufficient amount of oxalic acid.

Copper may be separated from cadmium, arsenic, iron, and small amounts of tin, when precipitated by oxalic acid in a volume of 50.0^{cm}³ containing 5^{cm}³ strong nitric acid. Inasmuch as the completeness of precipitation of the copper depends upon the presence of a certain minimum amount of the copper salt this method is not applicable when the amount of copper falls below 0.0128 gm. of the oxide to 50^{cm}³ of solution.

The author wishes to thank Professor F. A. Gooch for much kind help given in the preparation of this paper.

ART. XXXVII.—*Synopsis of the Collections of Invertebrate fossils made by the Princeton Expedition to Southern Patagonia*; by Dr. A. E. ORTMANN.

IN the February number (1900) of this Journal, Mr. J. B. Hatcher has given a general account of the sedimentary rocks of Southern Patagonia, frequently (pp. 98, 101–104, 108) referring to the present writer's studies on the fossils of this region. Since these studies have now come to an end, it seems well to give a preliminary report on the chief results obtained, in order to give an idea of the richness of the material at hand, and to enable the scientific public to get a more correct understanding of the beds in question than has hitherto been possible. A final report on the Tertiary Paleontology of Southern Patagonia will be given by the writer in a volume of the "Princeton Expedition to Patagonia," which is in course of preparation, and will be freely illustrated by figures of all species represented, drawn by Mr. F. van Iterson. It is hoped that this volume will be ready by the end of this year or the beginning of next.

The present article is intended to treat only of the so-called "Patagonian formation." As Mr. Hatcher has already pointed out (l. c., p. 101), we have sufficient reason to believe, that the different marine horizons distinguished by F. Ameghino and accepted—at least in part—by H. von Ihering (the Patagonian formation, divided into a "Piso Juliense" and "Piso Leonense," and Suprapatagonian beds or lower part of the Santa-Cruz-formation) are identical, i. e., belong in one and the same continuous series of marine deposits underlying the Mamiferous Santa-Cruz-formation, and we retain the old term *Patagonian formation* for this series, which is certainly a paleontological unit, and belongs, as the writer is now fully satisfied, in the *Lower Miocene*. The fauna of this series is represented in our collection by over 140 species, many of which are new to science.

I shall give here first a list of the *new species*, with short diagnoses and other remarks that are necessary to recognize them, and then I shall add a list of species known from other localities, but *not found previously in Patagonia*, and shall conclude this article by some remarks on the *synonymy*, etc., of known Patagonian forms.

Diagnoses of new species from the Patagonian formation.

ECHINODERMATA.

1. *Cidaris antarctica* sp. nov. Plates with a moderately large, perforated central tubercle, the neck of which is slightly crenulated. Scrobicule large, surrounded by a circle of small tubercles, between which there are still smaller ones. Spines subcylindrical, often slightly compressed, neck somewhat constricted. Articular surface finely striated, with a deep articular groove. Surface of spines closely covered with fine, rounded granules, forming irregular longitudinal rows.

Only isolated spines and plates, San Julian, Santa Cruz, Upper Rio Chalia, Lake Pueyrredon.

2. *Toxopneustes præcursor* sp. nov. Test suborbicular. Ambulacral and interambulacral spaces with 4–8 vertical rows of tubercles of subequal size, those of the ambulacral spaces being somewhat smaller. Poriferous zone moderately broad. Pores in three pairs, the two outer vertical rows separated from the inner row by a small tubercle. All the primary tubercles surrounded by small secondaries and miliaries. Actinostome sunken, and lower surface concave, the actinal cuts comparatively slight.

This species differs from the known recent species of *Toxopneustes* chiefly in the more crowded tubercles. The most closely allied form seems to be: *T. pileolus* (Lmck.).

San Julian; Shell Gap (Upper Rio Chico).

3. *Cyrtoma posthumum* sp. nov. Test subcircular-elliptic, depressed. Apex central, upper side covered with very fine tubercles. Ambulacra petaloid, open, lanceolate, subequal, extending about two-thirds from the apex toward the periphery, the posterior ones closer together than the others. Anus situated on the upper surface, in a deep depression, of a pyriform shape, narrow above and suddenly widening toward the periphery. Lower surface of test concave, covered with larger, more widely separated tubercles. Mouth subcentral, surrounded by a floscelle. Diameter ca. 110^{mm}, height ca. 28^{mm}.

The peculiar shape of the anal depression brings this species into the genus *Cyrtoma* of McClelland (*Calcutta Journ. Nat. Hist.*, 1840), a synonym of which is *Stigmatopygus* of d'Orbigny. It is the first Tertiary representative of this so far exclusively Cretaceous genus.

Lake Pueyrredon.

VERMES.

4. *Serpula patagonica* sp. nov. Tubes solid, calcareous, cylindrical, irregularly contorted and vermiculate, growing

upon shells, stones, etc. Outer surface transversely rugose. Diameter, 3^{mm}.

San Julian.

5. *Terebella magna* sp. nov. Large cylindrical tubes, isolated or growing in groups of two or three, straight or slightly curved. Walls composed of large and irregular fragments of shells. Inner surface of tubes smooth, outer surface very rough. Diameter of inner tube, 12–15^{mm}.

Systematic position of these tubes, apparently built by a worm, doubtful.

San Julian.

BRYOZOA.

6. *Melicerita triforis* sp. nov. Zoarium foliaceous, lobate. Zooecia hexagonal, disposed quincuncially on both surfaces of the zoarium. Orifice crescentic, large, about in the middle of each cell. Besides there is an ovarian opening on the summit of the cell, and two (? avicularian) openings on the side of the mouth.

The three openings in the upper part of the cell distinguish this species from all the rest.

Upper Rio Chalia.

7. *Reticulipora patagonica* sp. nov. Closely resembling *R. transennata* Waters (Quart. Journ. Geol. Soc., vol. xl, 1884, p. 689), and differing only, if at all, in the branches of the zoarium being a little stronger, and the zooecial openings being more crowded.

Santa Cruz.

8. *Tennysonia subcylindrica* sp. nov. Closely resembling the only known species of the genus, *T. stellata* Busk (Cat. Mar. Pol. Brit. Mus., 3, 1875, p. 34), but differing by the more slender branches of the Zoarium, which are subcylindrical, and the slightly prominent orifices of the cells.

Santa Cruz.

PELECYPODA.

9. *Modiola andina* sp. nov. Shell small, elongated. Apex near anterior end. Both valves convex, with a blunt ridge running down from the apex to the posterior and inferior end. This ridge is slightly curved, concave toward the lower margin. Upper margin almost straight in its anterior part, forming a blunt angle with the straight posterior part, which passes in a regular curve into the posterior margin. Ventral margin concave. Surface of shell in the upper half (above the oblique ridge) finely radially striated. Lower part of surface smooth, only near the anterior end, below the apex, with a few fine striæ. Length of shell, 24^{mm}; height, 9^{mm}.

Lake Pueyrredon.

10. *Nucula reticularis* sp. nov. Shell small, moderately convex and moderately thick, subovate, oblique. Posterior and anterior dorsal margin slightly convex, ventral margin strongly arcuate. Surface with very fine concentric ribs, which are irregular and often bifurcate. These ribs are crossed by still finer radial striæ, which give a beautifully reticulated appearance to the shell. Ventral margin finely crenulated on inner side. Hinge teeth fine, both parts of the series forming an obtuse angle, anterior part with ca. 9, posterior with ca. 18 teeth. Length, 7.5^{mm}; height, 6^{mm}.

The sculpture of this species is of the type of that of the Oligocene *N. chasteli* Nyst.

Santa Cruz and Mt. of Observation.

11. *Crassatella quarta* sp. nov. Shell elongated-ovate, comparatively thin, not very convex. Apex only slightly prominent. Anterior end rounded, posterior hardly angulated and hardly narrowed. Posterior dorsal margin straight near apex, anterior almost straight, with only a slight suggestion of concavity close to the apex. Surface ornaments as in *C. Lyelli* Sowerby, but the ridges more crowded and a little less developed. Ventral margins without crenulations. Length, 17^{mm}; height, 10^{mm}, but growing larger.

Santa Cruz; Lake Pueyrredon.

12. *Glycimeris regularis* sp. nov. Shell elongate, convex, with concentric lines of growth and undulations. Apex at $\frac{1}{3}$ of the length, incurved. Anterior end rounded, posterior subtruncated, not narrower than anterior. Ventral margin straight in the middle. Long., 78; height, 45.

This form does not agree with any of the described Patagonian species, and accordingly I think it is new, although all the distinctive characters are taken only from the external form.

Santa Cruz; San Julian; Lake Pueyrredon.

13. *Corbula hatcheri* sp. nov. Shell small, solid and thick, subovate-triangular. Right valve very little larger than the left, both moderately convex. Anterior end rounded, posterior produced, subtruncated, an angular ridge running from apex to posterior angle. Ventral margin arcuate, posteriorly a little concave. Lower margin of right valve reflected toward the left valve. Surface with concentric ribs, which are rounded and rather crowded. Length, 11; height, 7.5; diameter (of right valve), 2.5^{mm}.

Santa Cruz; Las Salinas; Mt. of Observation; San Julian.

14. *Martesia pumila* sp. nov. This species resembles much *M. patagonica* Phil. but is much smaller, the callous plate of the anterior margin is very small, and the ribs of the anterior part of the shell form a very obtuse angle with the lines of

growth of the posterior part. The radiating furrow is narrower, and is more inclined posteriorly, so as to render the posterior part of the shell smaller in comparison with the anterior. Length, 9^{mm}; height, 4^{mm}.

This does not seem to be an immature stage of *M. patagonica*.
Santa Cruz.

GASTROPODA.

15. *Liotia scotti* sp. nov. Shell small, rounded, flat above, with a large, open umbilicus below. Spire with four rounded whorls, increasing rapidly, suture deep. Last whorl with six revolving, equidistant keels, the keel nearest to the umbilicus the smallest, and disappearing within the umbilicus; the upper whorls show only the two uppermost keels. The keels are crossed by very fine striæ, and a number (15) of strong radial ribs; at the points of intersection of these ribs and the keels, there is a small conical tubercle. Last whorl a little deflected toward the mouth, which is circular and thickened. Height, 4^{mm}; diameter, 8^{mm}.

This species resembles much the recent *L. acrilla* of Dall.
Santa Cruz.

16. *Calliostoma observationis* sp. nov. Shell low, conical, not umbilicated. Whorls flat, last whorl on the periphery bluntly angular. Above this angulation there are five distinct revolving ribs; near the mouth, between the second and third (counted from above), a sixth rib begins to appear. In the upper whorls the second and fourth ribs disappear, so that only three ribs remain, besides the peripheral angulation, which shows as a fourth rib immediately above the suture. All these ribs, when fully developed, are subequal, flattened, smooth, about as broad as the intervals between them. The base of the shell has 9–10 revolving ribs of the same character. The outermost of them is not separated from the peripheral angulation by a broader interval. Height, 10·5^{mm}; diameter, 12^{mm}.

Mt. of Observation.

17. *Calliostoma cossmanni* sp. nov. Shell conical, higher than broad, not umbilicated. Whorls flat, the last one angulated, with a keel on the periphery, which is wholly exposed on the upper whorls, being situated close to, but above the suture. Upper whorls with five revolving keels, the lowermost, formed by the peripheral keel just mentioned, is the strongest. It is smooth, with hardly any trace of granulations. The uppermost and the third keel are stronger than the second and the fourth; the first, second, and third are distinctly granulated, the fourth with finer granulations. Toward the apex of the shell, the second and fourth keels disappear, so that only three keels are present, the two upper

ones granulated, the lower one smooth. Base of shell hardly convex, with six revolving keels, which are subequal, smooth, and narrower than the intervals. Height, 8^{mm}; diameter, 6.5^{mm}.

Santa Cruz.

18. *Calliostoma garretti* sp. nov. Shell conical, as high as broad, not umbilicated. Eight whorls, which are very slightly convex, suture shallow. Last whorl very bluntly angulated at the periphery, without a distinct keel. Surface of whorls, above the periphery, covered with numerous fine, revolving threads: there are, on the third whorl, about 7 of them, increasing to about 17 on the last. The number of the threads increases by intercalation, the new keels being at first smooth, but soon they equal the others, and become, like the latter, finely, but distinctly granulated. These granulations, however, are developed only in the upper three quarters of the whorl, the lower four or five threads remain smooth. The threads continue over the periphery to the base of the shell, which is slightly convex; their number, on the base, is about 24, and they are smooth, resembling in all other respects those of the upper part of the whorls. Height and diameter 17^{mm}.

Santa Cruz.

19. *Calliostoma iheringi* sp. nov. Shell conical, broader than high, umbilicated. Six whorls, which are sharply angulated, one angulation being formed by a sharp revolving keel in the upper part of the whorls, a second one—exposed only on the last whorl—formed by a peripheral keel. Suture distinct. Upper part of whorls (above upper keel) oblique, flat, with 5–6 revolving threads, which are slightly granulated; lower part vertical, slightly concave on the last whorl, with 6–7 fine, smooth threads. Base of shell slightly convex, depressed toward the umbilicus, which is moderately large. About 18 revolving threads on the base, which are smooth, more crowded and finer toward the periphery, a little stronger near the umbilicus. Height, 9.5^{mm}; diameter, 12^{mm}.

Santa Cruz.

20. *Crucibulum dubium* spec. nov. Cast suborbicular, depressed-conical. Apex central. On one side is the impression of the internal cup-shaped lamina, which was attached to the inner wall of the shell. No further characteristics can be given, since only a single cast is represented in our collection.

Arroyo Gio.

21. *Sigapatella americana* sp. nov. Shell suborbicular or subelliptic, depressed. Apex distinctly excentric. Surface with irregular, concentric, slightly lamellate striae, crossed by very fine radial rugosities. Internal diaphragma spiral, colu-

mella excentric, margin of diaphragma slightly concave and slightly reflexed at the columella. Height, 16^{mm}; diameter, 49^{mm}.

Santa Cruz; Punta Arenas.

22. *Dolium ovulum* sp. nov. Shell ovato-globular, spire short, conical, acute, last whorl large. Surface with fine and crowded revolving striæ, which are sub-equal, only in the lower part finer ones are intercalated. Mouth large, elongated-oval, canal very short, truncated, straight and comparatively narrow. Inner lip without callous tubercles or folds. Outer lip slightly thickened. Height, 34^{mm}; diameter, 25^{mm}.

Santa Cruz.

23. *Tritonium morgani* sp. nov. Shell subfusiform, elongated, with three varices. Whorls with fine, unequal, spiral striæ and large tubercles, the latter, on the last whorl, in three spiral rows, those of the upper row large, ca. seven between two varices; those of the middle row (5-6) small, and those of the lower row (3-4) very indistinct. Columella smooth, with a few indistinct crenulations in the lower part. Canal comparatively long. Outer lip distinctly crenulated, with an indistinct canaliform emargination in the upper part, opposite which is a fold on the inner lip. Height, 63^{mm}; diameter, 28^{mm}.

Santa Cruz.

24. *Buccinum annæ* sp. nov. Shell subfusiform, elongate-oval. Spire long. Whorls 7-8, angulated, the angulation with a series of tubercles, 12-14 of them on the last whorl, which are continued downward as irregular longitudinal ribs. Upper part of whorls slightly concave, appressed toward the suture. Exposed part of upper whorls, below angulation, sub-cylindrical. Last whorl large. Mouth ovate, elongated, upper end subcanaliculate, lower end truncated, and with a short reflexed canal, forming a varix on the columella. Outer lip thin, smooth within. Height, 66^{mm}; diameter, 30^{mm}. This species belongs into the subgenus *Cominella*.

Santa Cruz.

25. *Fusus archimedis* sp. nov. Shell fusiform, spire shorter than the last whorl, scalariform. Whorls very prominently angulated, suture very deep. Upper part of whorls, above angulation, flat, obliquely descending from the suture, lower part, below angulation, very slightly convex, obliquely receding downward to the suture. Angulation blunt, with a number (10-13) of blunt, often indistinct tubercles. Surface of shell with fine revolving ribs on the lower part of the whorls and upon the angulation, but these ribs are absent on the upper part, above the angulation. Whole surface with distinct lines of growth, which have a squamiform appearance, where they

cross the revolving ribs. Last whorl large. Mouth triangular, continued into a long and straight canal. Height, 50^{mm}, but defective on upper end; diameter, 25^{mm}.

San Julian.

26. *Fusus torosus* sp. nov. Shell subturbinate or subfusiform. Spire short, rather depressed. Whorls four, last one very large. Surface with numerous fine spiral ribs, which are rather crowded and somewhat unequal, crossed by very fine, squamiform lines of growth. Whorls strongly convex, swollen, with ca. seven strong, variciform longitudinal ribs, which begin at the suture and become thick and swollen in the middle of the last whorl, attenuating again toward the lower end of the shell. Mouth ovate, continued into a canal of moderate length, which is slightly curved. Height, 31; diameter, 20^{mm}. This species resembles somewhat *F. pyruliformis* Sow., from Navidad, and I would not hesitate to identify it with this species. But Sowerby's figure seems to be poor, and the account given by Philippi and Moericke of *F. pyruliformis* shows clearly that it is different.

Santa Cruz.

27. *Fusus cancellatus* sp. nov. Shell small, fusiform, elongated. Spire a little shorter than the last whorl. Whorls convex, surface ornamented by revolving and longitudinal ribs, cancellated. Spiral ribs, in the upper whorls, 4-5, 12-18 on the last whorl; they are sharp, but flat, equidistant, narrower in the intervening spaces between the longitudinal ribs, and on the points of intersection with them slightly broadened, giving the appearance of low tubercles. Longitudinal ribs 12-13 on one whorl, rounded, but distinct, running from suture to suture, but disappearing on the canal. Mouth elliptical, canal comparatively short. Outer lip crenulated within. Height, 16^{mm}; diameter, 6.5^{mm}.

Santa Cruz.

28. *Fusus pilsbryi* sp. nov. Shell thick, elongated, fusiform; spire a little shorter than the last whorl. Whorls 7-8, convex, slightly appressed in the upper part, ornamented with 8-9 strong, rounded longitudinal ribs, which are slightly oblique and curved. On the upper whorls these ribs reach from suture to suture, on the last whorl they disappear below the middle. All of the surface of the shell is covered by very fine, numerous, distinct and subequal spiral striae. Mouth comparatively small, continued into a short canal. Outer lip thick. Height, 36.5^{mm} (not quite complete); diameter, 12^{mm}.

Santa Cruz.

29. *Murex hatcheri* sp. nov. Shell ovato-subfusiform. Whorls 5-6, rapidly increasing. Spire short, conical. Upper whorls angulated by a prominent, but blunt carina, which is

situated below the middle of the whorls; this carina forms an angulation on the last whorl, and below it there are 4–5 other carinae, decreasing in size. Upper part of whorls flat and obliquely descending from the suture, with a few revolving striae. Varices 5–6, lamelliform, strong and thick, at the points of crossing with the spinal carina produced into short leaf- or ear-like lobes, strongest on the uppermost carina. On the upper whorls only the uppermost row of lobes is visible. Mouth large, oval, with an open canal of medium length. Outer lip ornamented with 5–6 lobes, corresponding to those of the varices. Height, 63^{mm}; diameter, 44^{mm}.

San Julian.

30. *Urosalpinx elegans* sp. nov. Shell ovato-fusiform; whorls 5–6, convex, with spiral striae and 7–8 longitudinal, variciform costae, which are rounded. Mouth oval, elongated into an open, but narrow canal, which is about as long as the mouth. Outer lip distinctly crenulated within. Height, 16.5^{mm}; diameter, 8^{mm}.

Santa Cruz.

31. *Marginella oliviformis* sp. nov. Shell elongated, sub-cylindrically-fusiform. Spire conical. Surface of shell smooth and shining. Suture quite indistinct. Mouth long and narrow, canal very short, represented only by a rounded sinus. Columella with four subequal folds. Outer lip thickened, smooth within. Height, 11^{mm}; diameter, 5^{mm}; length of mouth, 6.5^{mm}.

Santa Cruz.

32. *Voluta petersoni* sp. nov. Shell elongated, fusiform. Surface beautifully cancellated by spiral and longitudinal ribs. Spiral ribs strongly developed, equidistant, sharp; longitudinal ribs a little stronger than the spiral ribs, sharp, running from suture to suture, ca. 30 on the last whorl. Cancellations rectangular, about twice as broad as high on the last whorl, and about three or four times as broad on the upper whorls. Spire slender, conical, mouth not much longer than half of the shell. Upper whorls quite high. Whorls almost evenly convex, only slightly appressed and concave near the suture. Mouth elongated. Columellar folds at least two, indistinct. Height, 148^{mm} (not complete); diameter, 65^{mm}.

Santa Cruz.

33. *Drillia santacruzensis* sp. nov. Shell turrite, subfusiform. Whorls 8, last whorl hardly half as long as the shell. Whorls convex, but depressed and slightly concave in the upper part near the suture. This depression forms a shallow furrow, following the suture, and is sharply separated from the rest of the whorl, which is ornamented by oblique longitudinal ribs, which end abruptly at the sutural depression. These ribs number 12–15 in one whorl. Besides, there are

very fine lines of growth, but no trace of spiral sculpture. Mouth elongated, canal short. Sinus of outer lip semicircular, situated in the sutural depression, close to the suture; at the point of junction of the outer lip with the columella there is a distinct nodulose, callous swelling. Height, 13^{mm}; diameter, 4.5^{mm}.

Santa Cruz

34. *Borsonia patagonica* sp. nov. Shell subfusiform, biconical; whorls ca. 6, last whorl a little larger than half of the shell. Whorls convex, depressed in the upper part, with a slight swelling just below the suture. Depressed part smooth, the rest ornamented by 10–12 longitudinal rib-like swellings, which are slightly tuberculiform on the upper whorls; on the last whorl they are rib-like, but less distinct. Besides the ribs, there are spiral cords on the lower part of the whorls; they are wanting on the depressed part, but continue, on the last whorl, upon the canal. Mouth elongated, canal of medium length. Outer lip with a moderately developed sinus, which is situated in the sutural depression. Columella with two plaits, the lower one sometimes quite indistinct. Height, 19^{mm} (not complete); diameter, 9^{mm}.

Santa Cruz.

35. *Acteon semilevis* sp. nov. Shell elongated-ovate, rather slender, spire short, conical, about one-quarter of the length of the shell. Whorls four, convex. Suture distinct, a slight carina running close to the suture and parallel to it. Below this carina there is an indistinct spiral groove. Below the latter the surface of the shell is smooth; but in the lower third of the last whorl there are 5–7 spiral furrows, which are rather broad, almost as broad as the flat intervals. Mouth elongated, wider below, columella with a distinct fold below. Height, 7^{mm}; diameter, 3.5^{mm}.

Mt. of Observation.

CRUSTACEA.

36. *Scalpellum juliense* sp. nov. Only the carina known. Carina narrow, elongated, strong and solid, curved; basal margin bluntly pointed; surface smooth, only with lines of growth. Tectum strongly arched in its upper part, only slightly so in its lower; upper part solid, its cross section almost quadrangular, with a prominent ridge on the concave side, formed by the junction of the inflected parietes. Parietes very narrow, separated from the tectum by a distinct but blunt ridge. The carina of *S. solidulum* Steenstr. (See Darwin, Monogr. foss. Lepad., 1851, p. 42, pl. 1, f. 8) resembles so much the present fossil, that I have no doubt, we have to deal here with a closely allied species.

San Julian.

Species new for the Patagonian formation.

BRYOZOA.

1. *Cellaria fistulosa* (L.). (Hincks, Hist. Brit. mar. Polyzoa, 1880, p. 106.) This species is a living, almost cosmopolitan form, and has been found fossil from the Oligocene beds upward in Europe and New Zealand.

Shell Gap (Rio Chico).

2. *Aspidostoma giganteum* (Busk). (Busk, Rep. Challenger, vol. x, 1884, p. 161.) Known so far only living from southern Patagonia.

Santa Cruz and San Julian.

3. *Heteropora pelliculata* Waters (see Nicholson, Ann. Nat. Hist., ser. 5, v. vi, 1880). Known living from Japan and New Zealand, and fossil from New Zealand.

San Julian; Arroyo Gio.

BRACHIOPODA.

4. *Rhynchonella squamosa* Hutton (Cat. Tert. Moll. New Zealand, 1873). Fossil from New Zealand and Australia, living (*piaydata* Dav.) from Kerguelen Islands.

Lake Pueyrredon.

5. *Terebratella dorsata* (Gmel.). (Davidson, Trans Linn. Soc., 1887, p. 75.) Fossil in New Zealand, and living on the Patagonian coast.

Santa Cruz; Shell Gap; Lake Pueyrredon.

PELECYPODA.

6. *Mytilus magellanicus* Chem. (Reeve, Conch. Icon., vol. x, 1858.) Living on the Patagonian coast.

San Julian.

7. *Leda oxyrrhyncha* (Philippi), (Tert. und Quart. Verst. Chiles, 1887, p. 197). Navidad beds of Chile.

Santa Cruz and Arroyo Gio.

8. *Leda errazurizi* (Philippi), (Ibid. p. 196). Navidad beds of Chile.

Santa Cruz; Sierra Oveja (Rio Chico); Arroyo Gio; Lake Pueyrredon.

9. *Cardita elegantoides* Ortm. (Amer. Journ. Sci., 1899, p. 428), described from the Magellanian beds of Punta Arenas. Santa Cruz and Mt. of Observation.

10. *Cardita volckmanni* Phil. (l. c., p. 173). Navidad beds of Chile.

Lake Pueyrredon.

11. *Venus chiloënsis* Phil. (l. c., p. 121). Chile.

Punta Arenas. (Known previously from this locality, but stratigraphical position not ascertained.)

GASTROPODA.

12. *Vermetus cf. intortus* (Lmck.). (Moerch, Proc. Zool. Soc. London, 1861.) Known from Oligocene to Pliocene deposits in Europe. The identification is not beyond doubt.

Shell Gap; Lake Pueyrredon.

13. *Galerus araucanus* (Phil.), (l. c., p. 92). Navidad beds of Chile.

Shell Gap; Lake Pueyrredon.

14. *Aporhais araucana* (Phil.), (l. c., p. 35). Navidad beds of Chile.

Santa Cruz.

15. *Buccinum obesum* (Phil.), (l. c., p. 48). Navidad beds of Chile.

Santa Cruz.

16. *Cancellaria cf. medinae* (Phil.), (l. c., p. 68). Navidad beds of Chile. Identification not quite certain.

Santa Cruz and Mt. of Observation.

CRUSTACEA.

17. *Verruca laevigata* Sow. (Darwin, Monogr. Balan., 1854, p. 520). Living on the coast of S. America.

Upper Rio Chalia.

Remarks on Synonymy, etc.

1. *Mugellania lenticularis* (Desh.). What v. Ihering mentions as *M. globosa* I take for *M. lenticularis*.

2. *Ostrea ingens* Zitt. The large Patagonian oyster is absolutely identical with the New Zealandian species described by Zittel as *O. ingens*, and differs from *O. patagonica* of d'Orbigny. The latter is not found at all in the Patagonian formation. There is only one species in the Patagonian beds.

3. *Pecten proximus* v. Ihering. This species has been called by v. Ihering in his text (Rev. Mus. S. Paulo, 1897, p. 229) by the name of *P. centralis* Sow., but is different; I accept for it the name given by v. Ihering on the plate.

4. *Pecten geminatus* Sow. Synonyms of this species are: *P. quemadensis* v. Ih. and *P. fissicostalis* v. Ih.

5. *Cucullæa alta* Sow. I cannot distinguish v. Ihering's *C. dalli* from this species.

6. *Cucullæa darwini* (Phil.). There is not the slightest doubt, that *Cucullaria tridentata* of v. Ihering is this species.

7. *Pectunculus ibari* (Phil.). Synonyms of this species are: *P. magellanicus* Phil. and *P. pulvinatus cuevensis* v. Ih.

8. *Nucula patagonica* (Phil.). *N. tricesima* v. Ih. is only a variety of this species.

9. *Cardita inæqualis* Phil. This species is the most abundant form of the genus at Santa Cruz. Large specimens of it have been sent by v. Ihering to the Princeton Museum under the name of *C. patagonica* Sow., and thus it seems apparent that *C. patagonica* of v. Ihering is identical with *C. inæqualis* of Philippi.

10. *Cardita patagonica* Sow. We possess only the small variety, called by v. Ihering in 1899 (Neues Jahrb. Miner., etc.) *C. pseudopatagonica*. I believe that this is really the true *patagonica* of Sowerby.

11. *Venus navidadis* Phil. Already v. Ihering suggests that his *V. striatolamellata* may be identical with this species. I think that is right.

12. *Dentalium sulcosum* Sow. Synonyms of this species are: *D. majus* Sow., and *D. patagonicum* Rochebrune and Mabile.

13. *Gibbula dalli* v. Iher. *G. fracta* v. Ih. seems to be nothing else than the young of this species.

14. *Infundibulum clypeolum* (Reeve). Called by v. Ihering *Trochita magellanica* Gray. But the specific name of Reeve has the priority.

15. *Natica ovoidea* Phil. I take *N. famula* Phil. for the young stage of this species.

16. *Natica secunda* Rochebr. and Mab. The specific name *secunda* of Rochebrune and Mabile was published in 1885 (Bull. Soc. Philom. Paris, ser. 7, vol. ix) and has the priority over *N. obtecta* Philippi, 1887.

17. *Natica darwini* v. Ihering. According to v. Ihering the specific name of *N. solida* Sow. has been preoccupied by Blainville. He attributes the name *darwini* to Hutton, but I cannot find it in any of the publications of Hutton.

18. *Odontostomia suturalis* v. Ihering. *O. synarthrota* Cossmann (Journal de Conchyliology, 1899) is indistinguishable from this species.

19. *Turbonilla cuevensis* v. Iher. *T. iheringi* Cossmann (ibid.) is indistinguishable from this species.

20. *Struthiolaria chilensis* Phil. The extensive material at hand leaves it beyond doubt that *S. ameghinoi* v. Iher. is a synonym of this species.

21. *Fusus domeykoanus* Phil. *Siphonatia dilatata* var. *subrecta* of v. Ihering (1899) is this species.

22. *Trophon patagonicus* (Sow.). Abundant material of this species enables me to pronounce *T. laciniatus santacruzensis* v. Ihering as a form of this species, which is connected with it by numerous transitions.

23. *Urosalpinx cossmanni* nom. nov. For *U. cf. leucostomoides* Cossmann, non Sowerby.

24. *Voluta gracilior* v. Ihering. The specific name *gracilior* was introduced by v. Ihering in 1896 (Nachrichtsblatt Deutsch. Malakozool. Ges.) for *V. gracilis* Phil. (non Lea). It is impossible for me to distinguish from this species *V. quemadensis* v. Iher. The living *V. philippiana* Dall (1890) is different.

25. *Voluta domeykoana* Phil. *V. pilsbryi* v. Iher. (1899) does not seem to be different from this species.

26. *Pleurotoma subaequalis* Sow. What v. Ihering calls, in 1899, by the name of *Pl. discors* Sow. seems to belong to this species.

27. *Pleurotoma unifascialis* v. Ihering. This species has been regarded by v. Ihering (1897) as a variety of *P. discors*, but I think it is a good species.

28. *Geryon* (?) *peruvianus* (d'Orb.). *Cancer patagonicus* Philippi is apparently the same species as *Carcinus peruvianus* A. Milne Edwards (Ann. Sci. Nat., ser. 4, vol. xiv, 1860, p. 269). I do not think, however, that it is a *Carcinus*. It may belong to *Geryon*.

Princeton University, May, 1900.

ART. XXXVIII.—*The Cathode Stream and X-Light*;
by WILLIAM ROLLINS.*The Cathode Stream.*

There are two opinions about cathode rays.

1. The rays are some phenomenon in the ether. Lenard considered them smaller transverse waves than those of light. Michelson thought they were ether vortices.

2. They are flights of material particles. Varley, who published his results in 1871, considered the cathode stream composed of molecules of the residual gas in the vacuum tube, charged with negative electricity. He deflected the stream by a magnet: showed the force of its impact on a pivoted mica vane. Crookes, who illustrated the earlier work of Varley and Hittorf by beautiful experiments, agreed with Varley as to the nature of the stream. He said the cathode stream particles left the cathode normal to its surface, moved in straight lines, coming to a focus in the center of its curvature. My experiments showed that neither theory of the cathode stream could explain all the facts. If it was possible to remove all the ether from an X-light tube, there would be no X-light, for no cathode stream could form, because a strain in the ether is essential.

The Ether Theory.—One characteristic of the Varley or cathode stream has not been explained by this theory: the stream can be deflected by a magnet, X-light arising where the deflected stream strikes. As there are also other objections, the pure ether theory will not be further considered. The experiments supposed to support it apply as well to the material particle theory.

The Material Particle Theory.—Since Varley's experiments the opinion has slowly grown that the particles in the cathode stream are not as large as molecules. The facts of physics and chemistry appear to prove that electricity breaks molecules into ions. Schuster therefore said the particles were the same Faraday ions as appear in electrolysis. Other physicists have further reduced the size, Weichert giving it as $1/3000$ of a hydrogen atom, and the speed as one-third that of light. In regard to this speed, Rowland, in a remarkable paper before the American Physical Society in 1899, said there was no way of producing this velocity in a body though it fell from infinite distance on the largest aggregation of matter in the universe.

Weichert made the first determination of the relation between the charge and the mass of a cathode stream particle.

He gave this as 20×10^6 to 40×10^6 c.g.s. units. As this ratio is about three thousand times as large for a cathode stream particle as for a hydrogen ion in electrolysis, this physicist thought he must either assume a charge 3000 times as large, or a mass $1/3000$ of a hydrogen ion. He chose the latter and other physicists have followed, though differing, some being willing to allow a cathode stream particle to be as large as $1/500$ of a hydrogen ion. There is no proof of these suppositions. Weichert might have had the same ratio mean some other size for the particles.

Kaufmann found that this ratio was not affected by the gas or the terminals in a vacuum tube. I shall not quote authorities further, but group their statements to show what is called a sufficient explanation of the cathode stream by those who hold the material particle theory.

1. The cause of the cathode stream is a repulsion between the cathode and the charged particles of gas in the tube. This repulsion is due to the particles coming in contact with the cathode receiving charges of the same nature.

2. The particles are given off from the cathode perpendicularly to its surface, and move in straight lines, coming to a focus at the center of curvature of the cathode.

3. The direction of the particles in high vacua is independent of the position of the anode.

4. All the cathode stream particles move with equal speed, which is of the same order as that of light; they have the same charge; are of the same size.

5. In the cathode stream the charges are always carried by the same particles in all tubes, with any gas, with every cathode.

6. The cathode stream particles are the same in every atom in the universe. They are the ultimate particles of electricity.

7. In the cathode stream only one particle is detached from an atom. The remaining part has a positive charge.

8. The particles in the cathode stream are those liberated by ultra-violet light from a charged body; move with the same speed; have the same charge.

9. The particles in the cathode stream are those producing Becquerel light.

10. The particles in the cathode stream are the ultimate units from which the elements are formed. Each has a mass of $1/3000$ of a hydrogen atom.

I shall now mention some of the experiments with X-light tubes which I have described in the Electrical Review during the last three years, grouping them under the headings given :

1. The cathode stream is a purely material particle phenomenon, with which the ether has no part. The sole cause of

the cathode stream is a repulsion between the particles of residual gas in the vacuum tube and the cathode. These two assertions are mentioned together, because one experiment applies to both.

An X-light tube was made with a biconcave aluminum cathode in the middle of its bulb. At each end of the tube was a platinum target. If the cathode stream was not affected by the ether, and depended entirely upon the repulsion between the cathode and the particles of residual gas, two cathode streams of equal intensity should have arisen from the cathode, producing the same amount of X-light at each target. All the X-light arising from either came from whichever target was made an anode. The experiment proved (1) that a space of strained ether was necessary, X-light arising with greatest intensity where the ether was most strained; (2) that the attraction of an anode was a cause of the cathode stream as well as a repulsion by a cathode.

2. The particles of the cathode stream are always given off from the cathode perpendicularly to its surface; move in straight lines, coming to a focus in the center of curvature of the cathode. Any one with experience can see the focus of the cathode stream when well developed in an X-light tube. By means of a pin-hole camera it may be photographed. Both methods prove that the distance varies not only with the degree of exhaustion in the tube, but also with the potential of the driving current.

I made a tube with a movable anode, whose distance from the cathode could be changed by a magnet. Under the conditions required for the economical production of X-light it was found necessary to have the target at twice the theoretical distance for it to be in the focus of the cathode stream. This experiment showed that the particles did not move in straight lines provided they left the cathode normal to its surface, nor did they come to a focus at the center of curvature. The explanation I offered was as follows: The particles of the cathode stream have the same kind of electricity, therefore they repel each other, coming to a focus beyond the place required by the accepted theory.

3. The direction of the particles in the cathode stream is independent of the position of the anode at high vacua. The experiment described in paragraph 1 appeared to disprove this, but another is also mentioned.

An X-light tube was made with connecting bulbs, one containing the usual cathode and anode. In the other, encircling the stem of the cathode already mentioned, as it passed through this second bulb, was a ring that could be made an anode. When the tube was connected in the usual way with

the generator, the cathode stream arose from the concave side of the cathode, came to a focus on the normal anode, giving good X-light. This anode was then disconnected from the current. The supplementary ring terminal was made the anode. Under these conditions no cathode stream arose from the usual side of the cathode, consequently no X-light was produced at the normal anode. X-light arose in the other bulb from the spreading streams of cathode particles given off from the back or convex side of the cathode.

4. The cathode stream particles are always of the same size, move with the same speed, and carry the same charge.

If these assumptions were true we should always get the same effect when the particles struck the target in an X-light tube. We do not. Every one long familiar with the construction of such tubes knows that the quality of the light varies with the resistance of the tube. In the determinations of the speed no sufficient account was taken of the retarding effect of the gases in circulation in the tube. Suppose we try to keep this constant, even then the effect produced by the impact of these particles on the target in an X-light tube varies from other causes. To show this I made a tube in which the ordinary cathode could be covered by mercury. With the usual cathode the tube gave good X-light. With the mercury cathode there was not enough X-light to see the bones of the hand, but the whole tube was filled with a brilliant white light. I explained this by saying that the particles in the cathode stream were heavier when the cathode was mercury than when it was aluminum, and the stream was composed of some light gas. In consequence the particles did not strike the target at so high a velocity, and when stopped were therefore not heated to so high a degree; hence were not such efficient centers of radiation for the short ether movements we call X-light, most of the energy appearing as ordinary light instead. Some of the experiments on which I based this heat theory will be mentioned later. If the particles in the cathode stream were always of the same mass, all cathodes should lose equally in weight in the same time, with the same amount of current. They do not. I also found that when the gas had been considerably removed from the terminals, cathodes of heavy metals lost in weight more rapidly than those of light metals, like magnesium and aluminum.

5. In the cathode stream the charges are always carried by the same particles, in all tubes, with any gas, with every cathode.

An X-light tube was exhausted while using heat and heavy surges until it yielded no X-light. Had the particles been

removed? Surely such a wonderful result was never before accomplished so easily.

My explanation was simpler. We depend upon gas amalgamated with the terminals to make an efficient cathode stream in an X-light tube. When the cathode stream no longer forms we have removed too much of this gas. Consider here the experiment with the mercury cathode already mentioned.

6. The cathode stream particles are the same in every atom. They are the ultimate units of electricity. If true, how explain the experiment mentioned in paragraph 5? Why should X-light have stopped? The terminals of the tube were constantly connected with a generator supplying electricity. They were good conductors. They were kept charged to a high degree. Is it not simpler to accept the explanation given in paragraph 5?

Consider another experiment. An X-light tube was exhausted by heat and pumping until no gas appeared in the pump. A current was then sent through it, producing X-light. Much gas appeared in the pump. The bubbles continued so long as X-light was produced. When, by continuing the current and pumping to remove the gas thus driven out of the terminals, the X-light died out, the bubbles stopped. While there were bubbles there was a normal cathode stream and X-light. Were the bubbles composed of the ultimate particles of electricity? Is electricity a gas with a familiar spectrum, or were these bubbles simply a common gas coming out of the terminals, thus forming a cathode stream, as stated in paragraph 5?

7. In the cathode stream only one corpuscle can be detached from an atom, the remainder of the atom being positive.

According to this theory, when the tube arrived at the condition mentioned in paragraphs 5 and 6, that is, when the normal cathode stream no longer formed, the cathode should have been left with a positive charge. The charge was negative.

8. The particles in the cathode stream are those liberated by ultra-violet light from a charged body, move with the same velocity, have the same charges. If true, we should get X-light from a cathode stream formed by allowing ultra-violet light to fall upon a cathode in an X-light tube connected with a capacity. I have not yet produced sufficient X-light in this easy way to use in my profession.

9. The particles in the cathode stream are those producing Becquerel light, have the same charges, move with the same velocity.

If true we could produce X-light from a vacuum tube with a radio-active cathode, without an electric generator. It would

only be necessary to exhaust the tube to the usual X-light vacuum. The particles in the cathode stream from the radio-active cathode would then meet with no more obstruction on their way to the target to produce X-light, than would be encountered by those of a cathode stream formed in the ordinary way.

A distinguished American physicist has expressed the opinion that the radio-active substances would be so intensified as to act as substitutes for the complicated and troublesome apparatus now required for producing X-light for medical purposes. Having a high regard for his opinion I abandoned experiments with X-light to work with these substances.

Experiments showed that the light from them was different in character from X-light, suffering diffusion in the tissues like ordinary light to such an extent that the bones even of the hand were not visible. The experiments indicated a difference, perhaps in velocity, between the units in the cathode stream and those of the radio-active substances, because the character of the two resulting radiations toward human tissues was not the same when produced in the same vacuum. Experiments of this nature require time and money. I was, therefore, disappointed with the results, for my interest in the subjects mentioned in this paper was a desire to find the most efficient radiation, to aid in the relief of human suffering.

10. A cathode stream particle has a mass only $1/3000$ of a hydrogen atom. The particles are the ultimate units of which all the elements are composed. That atoms are compound is probably true, though not yet proved by experiment. The theory explains phenomena better than that of the indivisible atom. That the particles in the cathode stream are the ultimate units is absurd.

I shall make a few suggestions in regard to each theory :

On the theory of indivisible atoms, how shall we explain differences in atomic weight, except by saying that such is the nature of atoms? This is not a satisfactory answer to an active mind.

On the theory of a compound atom it may be explained by saying that a light atom has fewer ultimate particles than a heavy one, for as each particle must have the same weight there would be more in a heavy atom than a light one, to account for differences in atomic weights.

That the particles in the cathode stream, however small, are not the ultimate particles of which the universe is composed, is shown by their having a familiar spectrum. If we could make ultimate units give a spectrum, it would be a new one. In this connection I mention a statement made in my notes already referred to :

In working with X-light tubes I appeared to get the spectrum of hydrogen in the cathode stream though other gases were in the tube. Therefore, before we reach the ultimate units of which the seventy-five or more elements are composed, we must seek more powerfully disruptive forces than those in the cathode stream. The so-called ultimate, indivisible corpuscles of this stream cannot be the final units because they have a familiar spectrum ; vibrating in too many ways at approximately the same temperature.

The experiments published by Trowbridge in this Journal for September are the most important contribution to celestial physics since Kirchhoff proved the law of exchanges and told the nature of the Fraunhofer lines. Trowbridge always obtained the spectrum of water vapor in the cathode stream when a condenser was used. Therefore the cathode stream particles are not necessarily elementary. If such minute particles are compound, it shows that when elements combine, it is not their atoms which unite to form new molecules ; the combination is far more intimate, a union of the particles of which atoms are composed. Long before we reach the heart of nature the present ultimate corpuscles will look to us more complicated than a wilderness of solar systems.

The X-Rays.

1. Röntgen considers them longitudinal vibrations in the ether.

2. Janmann believes they have also a transverse component.

3. Goldhammer stated they were short transverse vibrations differing from light only in size.

4. Stokes advanced a theory of irregular pulses in the ether, partly positive, partly negative.

5. J. J. Thomson has modified Stokes' theory. He believes when a cathode stream particle is stopped, its charge performs a single oscillation, giving rise to a pulse in the ether.

6. Michelson has suggested that X-rays may be ether vortices.

7. Several physicists believe they are flights of material particles.

This list of great names might easily be made longer. Yet what do we know of the true nature of X-light? Nothing.

I shall make a few suggestions and describe one or two experiments :

J. J. Thomson has accepted most of Weichert's views of the cathode rays, and Stokes' ideas of the X-rays. He believes when a corpuscle is stopped, there arises in the ether a pulse whose thickness is equal to the diameter of a corpuscle. These pulses are the X-rays. According to this theory the thickness of the radiant area on the target from which X-light arises

cannot be greater than the diameter of one of his corpuscles, having a mass $1/1000$ of an atom of hydrogen. Estimate the thickness on this basis. As a preliminary it is necessary to know the diameter of a hydrogen atom. This requires a liberal use of the imagination, as proved by the different estimates. Kelvin said an atom, or a molecule (he did not attempt to distinguish them) was from $1/100000000$ to $1/1000000$ of a millimeter in diameter. Meyer considered them smaller than a sphere $1/1000000$ of a millimeter in diameter, or as small as 0.2×10^{-7} centimeter. The physical chemists require us to believe that a molecule is not a solid. Its atoms are arranged with ether between them. The diameter of an atom, therefore, is not equal to half the diameter of any molecule with more than one atom, while in complex molecules its proportion is smaller. Atomic diameters are probably not greater than $1/3000000$ of a millimeter. The corpuscle theory of the cathode stream gives the mass of a corpuscle as not more than $1/1000$ of a hydrogen atom. The diameter of a corpuscle would not exceed $1/30000000$ of a millimeter. This then would be the depth of the radiant area on the target from which X-light could arise.

Consider the following experiment: An X-light tube had its target placed at an angle of 90° with the axis of the cathode stream. As close as possible to the area struck by the cathode stream was a narrow passage in a block of platinum, with openings raised one millimeter above the surface of the target, the length of the passage parallel therewith. Under these circumstances, if X-light arose only from a radiant area on the target, whose depth was measured by $1/30000000$ of a millimeter, direct rays could but to a small extent have illuminated the passage, for this was on the top of a cliff thirty million times as high as the depth of the radiant area at its base, in deep shadow, except the opening nearest the radiant area. If the usual theory had been true, no bright image should have been formed on the fluorescent screen, by light coming through the passage. The light was so bright Fomms' bands were photographed with a short exposure. The experiment also appeared to show that X-light was not composed of the reflected ether vortices of the cathode stream or of minute material particles. As these are supposed to move in straight lines, they would have had difficulty in going into the passage in the platinum block to brightly illuminate the screen. Again, how could material particles shot from a charged anode escape having a charge and being deflected by a magnet? Suppose they did escape without a charge. When they went through charged aluminum they ought to have received a charge and been deflected by a magnet. They were not deflected. Was not the theory of X-light advanced in my

notes a better one? This was the theory: When the particles of the cathode stream strike the target, they are heated sufficiently to cause them to be radiant centers, from which the short ether waves we call X-light arise. As time is required for heat to decline, the particles are sufficiently hot, when they have rebounded opposite the opening of the passage, to be radiant centers. The wave fronts, therefore, go directly through the passage, to the fluorescent screen, brightly illuminating it. I have used the word heat not to distinguish this from electromagnetic phenomena, but to express a more persistent state, in the radiant particles of the cathode stream, after they have struck the target, than that represented by the single pulse of the usual theory of X light. I mention an experiment bearing on this heat theory. An X-light tube was made with a hollow target. The area struck by the cathode stream could be cooled to a low temperature. When so cooled there was less X-light. This tube under the title of the A-W-L tube was made commercially, proving capable of converting a large amount of electricity into X-light, because with any cathode stream it was impossible to melt the target. I also designed a simpler tube for general use. The target rotated on an axis. Fresh metal could be brought to receive the force of the cathode stream when a hole had been melted.

Before closing this paper I shall consider the structure of the atom as seen by the cathode stream and X-light. After Röntgen discovered that some of Lenard's rays would show the bones of the hand, S. P. Thompson found the heaviest metals made the best targets in X-light reflecting focus tubes. The reason usually given is that light metals like aluminum allow X-light to pass through them, this part of the light being lost. Why should aluminum be transparent to X-light if this is an electromagnetic phenomenon like ordinary light? Maxwell's theory requires conductors to be opaque. Aluminum is a good conductor and should be opaque to X-light. It is transparent. If X-light is an electromagnetic phenomenon, we must find some way of explaining the transparency of aluminum. Consider first the solid atom theory. As the nature of a substance depends on its atoms, conductors must have opaque atoms if these are solid. If X light is an electromagnetic phenomenon it cannot travel through such solid conducting atoms. If it passes through a metal with solid atoms it must pass in the ether between them. If one metal is more transparent than another, the more transparent must contain more ether. On the solid atom theory how can we get more ether in aluminum than in platinum? The solid atoms are of one size. There are about as many of them in aluminum as in platinum, for the ratios of the atomic weights are about the same as the ratios of the densities. There is not

enough difference to account for aluminum being forty times as transparent as platinum. As we do not find enough difference in the amount of interatomic ether to account for the difference in transparency, we must give up the solid atom theory, for it fails here as it does in attempting to explain the cathode stream. Both phenomena require a compound atom. But the cathode stream theory, that the particles into which the atoms are broken are the ultimate units, cannot be true, for they have a familiar spectrum. A theory of the atom must meet these difficulties. The following is suggested. The atom is made of sub-atoms. The sub-atoms are the cathode stream particles. There are as many kinds of sub-atoms as there are elements. We should expect the cathode stream particles, therefore, to have familiar spectra. The sub-atoms are compounds. They are made of the ultimate units of which the elements are composed. These ultimate units have the same size and weight. As the atoms have the same size and different weights, light atoms must contain fewer ultimate units than heavy ones. There would be more space filled with ether in a light than in a heavy atom. When an electromagnetic phenomenon passes through an atom of a conductor, it can only travel in the contained ether, for this is the only transparent part. The atom with the most ether would be most transparent. Aluminum having more ether and fewer opaque particles, would be more transparent than platinum. Apply this theory to the following experiment, intended to show that the transparency of aluminum is not the only reason why aluminum is less efficient for a target in an X-light tube than platinum. I made a tube with the usual platinum target, covered with a thin veneer of aluminum on the surface struck by the cathode stream. Had the transparency theory been the true explanation, such a target should have acted toward X-light about as a glass mirror, silvered on the back, would have to ordinary light. As the X-light passing through the aluminum struck the platinum, only a little more could have been lost than with a target entirely of platinum. I never made an efficient X-light tube in this way. The following explanation was given. When a cathode stream particle strikes an aluminum target it is not so abruptly stopped as by a heavier metal. The heat of impact is not so high, the energy conversion extending over a longer time. The particles must be less efficient sources of X-light, if this is due to a high temperature. But why should a heavy metal target stop a flying particle more abruptly than a lighter one? I made this answer. Aluminum on the compound atom theory given is a more open structure than platinum: on this account it would stop a flying particle less abruptly.

September 28th.

SCIENTIFIC INTELLIGENCE.

I. CHEMISTRY AND PHYSICS.

1. *Ueber die physikalisch-chemischen Beziehungen zwischen Aragonit und Calcit.*—The object of this investigation by H. W. Foote was to show by purely physiochemical methods: 1. which of the two minerals under the existing conditions of temperature and pressure is the more stable; 2. whether the temperature of transformation is higher or lower than ordinary temperature. The method of investigation consisted in determining the relative solubility of the two minerals at different temperatures, as it is known on theoretical grounds that the least soluble must be the more stable. The best results were obtained by determining the electrical conductivity of aqueous solutions saturated with calcite or aragonite and carbon dioxide at atmospheric pressure and varying temperatures. It was found that calcite is more stable than aragonite at the temperatures experimented with, and that the solubility curves approach each other as the temperature rises. The author believes, however, from his own results and those of others, that at atmospheric pressure, calcite, below its melting point, can never become the less stable of the two minerals. From this it is concluded that the paramorphism of calcite after aragonite is theoretically possible, and the opposite paramorphism impossible. It is the author's opinion that rapidity of crystallization is one of the causes of the formation of aragonite.—*Zeitschr. physikal. Chem.*, xxxiii, 740. H. L. W.

2. *Atomic Weight of Radio-Active Barium.*—It is well known that M. and Mdme. Curie have discovered a substance "Radium" which is found with the barium of pitch-blende. By systematic fractional crystallization of the barium chloride obtained from the mineral, more concentrated products have been obtained. Finally by partial precipitation of solutions of the latter by means of alcohol or hydrochloric acid the concentration has been carried still further. The last product made in this way was examined spectroscopically by Demarçay and found to contain apparently only a trace of barium. This result indicates the isolation of radium chloride, but the quantity of the product was too small for use in an atomic weight determination. For the latter purpose Mdme. Curie was therefore obliged to use a less pure material, in which, from the aspect of its spectrum Demarçay thought that there was more radium than barium. Her results in two determinations gave the numbers 174.1 and 173.6 for the atomic weight of the metal in this chloride. Since the atomic weight of barium is 137.5, radium should have an atomic weight much higher than 174. The valency of radium is not discussed, but it is evident that if this is greater than that of barium the indicated atomic weight would be much larger.—*Comptes Rendus*, cxxxi, No. 6. H. L. W.

3. *Artificial Radio-Active Barium.*—A. DEBIERNE, who, it will be remembered, is the discoverer of "Actinium," the radio-active substance that is found with the titanium of pitchblende, has apparently produced active barium chloride from the inactive salt by induction due to contact with actinium in solution. The effect is made more intense by precipitating barium sulphate in a solution containing actinium. The latter is carried down by the barium sulphate, the sulphates are then changed to chlorides and the actinium is precipitated by ammonia in the form of hydrate.

If barium and actinium are in contact in solution for a short time, the induced activity in the barium is insignificant; but this activity increases with the time of contact, at least for ten days or so, and a product was thus obtained which had several hundred times as much activity as ordinary uranium. The author decides that the radio-active barium chloride prepared in this way is different from the radiferous barium extracted from pitchblende (Curie's radium). In both cases the activity persists in all chemical transformations to which the barium may be subjected. The rays emitted seem to be similar; they ionize gases, cause phosphorescence in barium platinocyanide, act on photographic plates, a portion of the rays is deviated in the magnetic field, and the anhydrous chloride is spontaneously luminous in each case. Moreover, the "artificial" radio-active barium chloride may be concentrated by fractional crystallization from water or hydrochloric acid solution, when the crystallized part increases in activity in the same way as with radium. However, "artificial" radio-active barium differs from the radiferous kind in showing no spectrum. Demarçay examined a specimen of it which was about a thousand times as active as ordinary uranium and was unable to detect the radium lines, while with a product only ten times more active than uranium, extracted from pitchblende, the radium spectrum was very clearly visible. The author has observed a second difference, in that the activity of the artificial product diminishes with time. In three weeks the activity of some of it diminished to one-third, while the activity of radiferous barium chloride and of salts containing actinium increases at first and then remains constant. The author believes that the induced activity which he has studied is not due to traces of actinium or radium, on account of the methods of separation used, and moreover it is difficult to understand how contamination by active substances should take place only when the solutions had been mixed for a long time.—*Comptes Rendus*, cxxxi, p. 333.

H. L. W.

4. *The Relative Values of the Mitscherlich and Hydrofluoric Acid Methods for the determination of Ferrous Iron.*—It has been known for a long time that concordant results were not always obtained by the two methods mentioned above, in the analysis of rocks and minerals. In Mitscherlich's method the substance is decomposed at a high temperature by means of a mixture of three parts of sulphuric acid and one part of water,

by weight, in a sealed tube from which air has been expelled. According to the other method the decomposition is accomplished by means of a mixture of hydrofluoric and dilute sulphuric acids in a platinum crucible kept full of steam, or protected from the air by being surrounded by carbon dioxide. In each case the ferrous iron is determined by titration with potassium permanganate solution. HILLEBRAND and STOKES have now explained why, in many cases, the Mitscherlich method gives higher results than the other. It is due to the presence in many rocks of small quantities of pyrite or pyrrhotite, minerals which are readily oxidized by ferric sulphate under the conditions used in the method under consideration. The authors find that not only is the metal of the sulphide oxidized, but the sulphur is converted into sulphuric acid as well. The effect of sulphur is therefore very great, amounting in the maximum to an error corresponding to $13\frac{1}{2}$ times its weight of ferrous oxide, and the method should not be used for rocks and minerals which contain even a trace of free sulphur or sulphides. The authors have also shown that pure pyrite causes very little error in the hydrofluoric acid method.—*Jour. Am. Chem. Soc.*, xxii, p. 625. H. L. W.

5. *Grundlinien der anorganischen Chemie*, von WILHELM OSTWALD, 12mo, pp. xx, 795. Leipsic, 1900 (Wilhelm Engelmann).—The appearance of each new work by Ostwald causes astonishment as to his prolificness, as well as enthusiasm on account of the excellence of his productions. Perhaps no book that he has written is as important from an educational standpoint as the elementary inorganic chemistry under consideration. The author himself considers it in a certain sense the keystone of a long and zealous activity in connection with the introduction and propagation of the new foundations of chemistry brought forward by van't Hoff and Arrhenius. Notwithstanding this aspect of the work, a careful examination shows that undue prominence has not been given to the new physical chemistry. Moreover, the treatment of the latter as well as other theoretical topics is admirably simple and philosophical. The descriptive arrangement, according to elements and their compounds, is used as the basis of the book, the general laws being introduced as opportunity and convenience permit. No attempt is made to arrange the elements according to the periodic system, in fact this system is only briefly dealt with at the very end of the book. The work contains comparatively few illustrations, and not many experiments are described in detail. A few inaccuracies are to be noticed in the descriptive matter, particularly in the paragraphs relating to metallurgy. It is to be hoped that these will be rectified in the next edition.

The author deplors the tendency that has often been shown in text-books of elementary chemistry to descend to a lower intellectual plane than is the case with works on physics and mathematics used by students at the same period of study, and he is convinced that this need not be the case now that the subject is

less purely descriptive than formerly. It seems certain that this book will be extensively used, particularly by teachers and by students who have had some preliminary training, and it is to be hoped that an English translation will soon appear. H. L. W.

6. *The Elements of Inorganic Chemistry for Use in Schools and Colleges*; by W. A. SHENSTONE, 12mo, pp. xii, 506. London, 1900 (Edward Arnold).—The aim of this book is to provide a graduated course which may be used at a period during which pupils pass from childhood to adolescence. It begins with a course of experimental work for quite young students and develops into a text-book for those who are older. A striking feature of the work is the large number of experiments devised to enable the student to verify quantitative laws and relations. These experiments are well chosen and very clearly explained, and they will doubtless furnish many useful suggestions to teachers of chemistry. The sections on carbon include much that is usually omitted from books on inorganic chemistry. The periodic system is made the basis of classification after the non-metallic elements have been separated from the metallic. Electrolytic dissociation and ionization are only briefly discussed. H. L. W.

7. *Die Kohlenoxydvergiftung*, von W. SACHS. 8vo, pp. x, 236, Braunschweig, 1900 (Friedrich Vieweg und Sohn).—This monograph on poisoning by carbon monoxide treats the subject very fully in its clinical, hygienic and medico-legal aspects. An elaborate index of the literature is included in the work. H. L. W.

8. *The Hall Effect in Flames*.—The theory of the wandering of ions in metals gives renewed interest to the phenomenon discovered by Professor Hall and known by his name. E. MARX discusses analytically the point whether the different velocities of the positive and negative ions in a magnetic field can account for the Hall effect, and reaches a value for the rotation coefficient. He concludes that the only possible way of reaching quantitative results in regard to the Hall effect is in the direction of flame conduction in a magnetic field. Arrhenius believes that the conveyance of electricity in flames is accomplished through electrolytically dissociated ions. The velocity of the ions with a fall of potential of unity is 10^6 times greater than in the case of electrolytes and very different in the cases of the positive and negative ions. So that we are led to expect a Hall effect a million times greater than in the case of electrolytes. A full account of the author's apparatus and his experiments is given and he reaches the conclusion that the Hall effect in flame gases has the dimensions which would be expected on the hypothesis of wandering electrical ions in a magnetic field. The author's experiments, however, do not decide the question whether the Hall effect in metals can be explained by pondero-motive effects on wandering ions.—*Ann. der Physik*, No. 8, pp. 798-834, 1900.

J. T.

9. *Wireless Telegraphy*.—It is well known that an earth connection is used in sending messages by wireless telegraphy, both in the sending and receiving apparatus. J. VALLOT and J. and L. LECARME maintain that this earth connection can be dispensed with for long distance transmission, and certain questions in regard to the part the capacity of the earth plays are thus opened. Their receiving apparatus was placed in a balloon, which rose to a height of 800 meters and was about 6 kilometers from the sending station.—*Comptes Rendus*, cxxx, pp. 1305–1307, May 14, 1900. J. T.

10. *The Spectrum of Radium*.—According to the investigations of Demarçay (*Compt. Rend.*, cxxix, p. 717, 1899) radium is a new element and he gives a table of its wave lengths. C. RUNGE has examined the spectrum and finds most of the lines which are given by Demarçay in chloride of barium. The spectrum was investigated in air and the preparation was also heated in vacuum tubes and the spectrum showed no new lines.—*Ann. der Physik*, No. 8, pp. 742–745, 1900. J. T.

11. *Magnetic effect of moving electrical charges*.—M. CREMIEU has conducted an important experiment on this subject. He has endeavored to measure the inductive effect when the convective current studied by Rowland is started or stopped. He could not observe any such inductive effect, and therefore concludes that there is no magnetic effect such as Rowland observed and that the effect observed by him was not due to a moving electric charge. In a criticism of Lamor's treatise on the relations between ether and matter (Cambridge University Press, 1900) Professor Fitzgerald remarks, that the questions raised by this experiment is one of the most fundamental ones in the connection between ether and matter. If M. Cremieu's experiments are substantiated, they will overthrow existing theories of electro-magnetism.—*Nature*, July 19, 1900. J. T.

12. *What electric pressure is dangerous?*—In order to obtain light on this much mooted question Professor H. F. WEBER of the Zurich Polytechnic has conducted a number of experiments and reaches the following conclusions:

A simultaneous touching of both the poles of an alternating current circuit is dangerous as soon as the pressure exceeds 100 volts; and since it is impossible to free oneself the case must be regarded as fatal whenever immediate help is not at hand. With steady currents he found that all pressures between 100 and 1000 volts must be regarded as equally dangerous, and consequently there is no reason for not using higher pressures between 500 and 1000 volts since they lead to more economical working of traction lines. There is little danger of the public coming in contact with such lines.—*Nature*, Aug. 23, 1900. J. T.

II. GEOLOGY AND MINERALOGY.

1. *Recent Publications of the U. S. Geological Survey.*—The following volumes have been recently issued by the U. S. Geological Survey, Parts II, III, IV, V, VII of the Twentieth Annual Report. (See also this Journal, ix, p. 447.)

Part II. *General Geology and Paleontology.* Pp. 1–953, plates I to CXCIII. This volume includes six papers, as follows: G. F. Becker, Brief memorandum on the Geology of the Philippine Islands, pp. 1–8; T. N. Dale, A study of Bird Mountain, Vermont, pp. 9–24, pl. I, II; G. H. Girty, The fauna of the Ouray formation, pp. 25–82, pl. III–VII; I. C. Russell, preliminary paper on the geology of the Cascade Mountains in northern Washington, pp. 83–210, pl. VIII–XX; L. F. Ward, Status of the Mesozoic floras of the United States, pp. 211–748, pl. XXI–CLXXIX; David White, Stratigraphic succession of the fossil floras of the Pottsville formation in the southern anthracite coal field, Pennsylvania, pp. 749–918, pl. CLXXX–CXCIII.

Of these papers that by Prof. Ward was noticed at length in the last number of this Journal (p. 320). The paper by Prof. Russell is largely devoted to the glacial and post-glacial geology of the northern and eastern parts of the State of Washington. It is shown that during these times the valleys were filled with gravel and sand to a depth in general of several hundred feet. Subsequently the streams excavated channels through these deposits and in many instances removed them entirely. The portions of the gravel deposits which remain form terraces along the streams, some of which are a mile or more broad and have a down-stream gradient. The terraces occur even in mountain gorges and along torrential streams and in such cases have a conspicuous down-stream slope—not due to a tilting of the land—which is less steep than the gradient of the present streams. No evidence is found of a modern depression of the land of such a nature as to admit of the flooding of the valleys by the waters of the ocean. The prevalence of glaciers at this time shows a much colder climate and a greater amount of precipitation than at present.

Part III. *Precious-metal Mining Districts.* Pp. 1–595, plates I–LXXVII. This volume will be noticed later. It contains three papers: J. S. Diller and F. H. Knowlton, On the Bohemia mining region of western Oregon; W. Lindgren, On the gold and silver veins of Silver City, De Lamar and other mining districts in Idaho; W. H. Weed and I. V. Pirsson, On the Geology of the Little Belt Mountains, Montana.

Part IV. *Hydrography.* Pp. 1–660, plates I–LXIII. This volume contains the report of progress on stream measurements for the calendar year 1898 by F. H. Newall, pp. 1–562; also an account of the hydrography of Nicaragua by A. P. Davis, pp. 563–638. The very rapid development of this new department

of the Survey and its promise of great economic usefulness for the future are alike noteworthy. In the present volume we find not only the results of drainage measurements for prominent rivers over the whole extent of the country, but also the survey of reservoir sites and the discussion of conditions governing underground waters as reached for example by artesian wells, both the latter subjects of peculiar interest in the West. The discharge data are given with much fullness for a large number of rivers from Maine to California, and the results are presented in graphical form for the successive months of the year. In addition to the statistical matter, plans are presented whereby the limited water supply of certain arid districts and Indian Reservations may be more economically distributed. The closing portion of the volume gives the results of the investigations on the hydrography of Nicaragua made by Mr. Davis in 1898; the work was carried on in connection with the Nicaragua Canal Commission.

Part V. *Forest Reserves*; by Henry Gannett. Pp. 1-478, plates I-CLIX, in part folded maps. The forest reserves of the United States numbered 37 on July 1, 1899, and are distributed over eleven States. The general condition of these reserves is given in this report, and there are also special papers on the following districts: Pikes Peak, Plum Creek and South Platte Reserves, by John G. Jack; White River Plateau and Battlement Mesa Reserves, by George B. Sudworth; Flathead Reserve, by H. B. Ayres; Bitterroot, San Gabriel, San Bernardino and San Jacinto Reserves, by John B. Leiberger.

Part VII. *Explorations in Alaska 1898*. Pp. 494, with maps 1-25 and plates I-XXXVIII. In this volume are grouped the reports of G. H. Eldridge, J. E. Spurr, W. C. Mendenhall, F. C. Schrader and A. H. Brooks, on reconnaissance surveys in Western and Central Alaska. Besides the sketches of the theoretical and economic geology much information is given regarding products, climate and feasible routes. Maps and routes of previous explorers are discussed and the general geographic knowledge of the region is brought down to date. Taken in connection with the "Map of Alaska" (U. S. G. S. 1898) and Schrader and Brooks, "Preliminary Report on the Cape Nome Gold Region" (U. S. G. S. 1900), this volume stands as the reference book for recent reliable information regarding Alaskan matters.

2. *Department of Geology and Natural Resources of Indiana, 24th Annual Report, 1899*; by W. S. BLATCHLEY, State Geologist. Indianapolis, 1900. Pp. 1-1078, 89 plates.—This report contains more than the ordinary Geological Annual Report. Six of the papers are concerning the Geology of the State. The first, by the Director, refers to the Natural Resources, and three others contain the reports of the State Inspector of Mines, the State Supervisor of Natural Gas, and the State Supervisor of Oil Inspection, respectively. Mr. Foerste contributes a discussion on the Middle Silurian rocks of the Cincinnati anticlinal region, the particular interest of which lies in his interpretation of the

synonymy of the several beds. The paper by J. A. Price on the Waldron Shale and its horizon gives details which will be of value to the general student of Structural Geology. In this paper a large number of detailed sections are given, running from the Silurian through the Devonian. The remainder of the book is devoted to four exhaustive treatises on the Natural History products of the State. Prof. E. B. Williamson furnishes a monograph on the Dragonflies of Indiana. Dr. R. E. Call gives an exhaustive and fully illustrated catalogue of the Mollusca of Indiana in which diagnostic descriptions of species are inserted. A catalogue of the Flowering Plants and the Ferns and their allies indigenous to Indiana is furnished by Prof. Stanley Coulter. And Mr. Blatchley contributes a short paper entitled "Notes on the Batrachians and Reptiles of Vigo County, Indiana." The volume is published in the usual style of the Indiana Reports; the paper being thin and the printing rather poor for reports of so great scientific value. w.

3. *Geological Report on Monroe County, Michigan*, by W. H. SHERZER. *Geological Survey of Michigan*, ALFRED C. LANE, *State Geologist*. Vol. vii, Part I. Pp. 1-240, pl. i-xvii.—The nomenclature adopted in the stratigraphical part of this report is as follows:—*St. Clare Shale* is essentially the equivalent of the Genesee Shale of New York: the *Traverse Group* is the name for the equivalent of the Hamilton of New York; *Dundee Limestone* corresponds closely with the beds from the Oriskany Sandstone to the Onondaga limestone inclusive; the *Monroe beds* are correlated with the Salina, Rondout, and Manlius beds of New York State. w.

4. *The Elements of the Geology of Tennessee*, prepared for the use of Schools of Tennessee, by J. M. SAFFORD and J. B. KILLEBREW, pp. 1-264, figs. 1-46.—Prof. Safford and Dr. Killebrew, the Commissioner of Agriculture, Statistics and Mines, have compiled a useful book for the use of Primary and Common Schools of the State. A glance through the book reveals nothing new; it appears to follow closely Dana's Manual of Geology, from which many of the illustrations are derived, and Safford's Geology of Tennessee, but the matter is brought down to the understanding of students of the lower grades, for whom it will prove a useful guide. w.

5. *Geological Survey of Canada*, G. M. DAWSON, Director.—The following papers of the Eleventh Annual Report have been received:

Report of the Section of Chemistry and Mineralogy, Part R, Annual Report, vol. xi, by G. C. Hoffmann. No. 695, pp. 1-55.

Section of Mineral Statistics and Mines, Annual Report for 1898, Part S, Annual Report, vol. xi, by E. D. Ingall. No. 689, pp. 1-193.

Catalogue of Canadian Birds. Part I. Water Birds, Gallinaceous Birds, and Pigeons, by John Macoun. No. 692, pp. 1-218.

Preliminary Report of the Klondike Gold Fields, Yukon District, Canada, by R. G. McConnell. No. 687, pp. 1-44. w.

6. *Correlation between Tertiary Mammal Horizons of Europe and America*, by HENRY FAIRFIELD OSBORN. (*Annals N. Y. Acad. Sci.*, vol. xiii, No. 1, pp. 1 to 72, July 21, 1900.)—In the two papers here published together Prof. Osborn sets forth his mature views upon the Zoögeography of Tertiary Mammals. In the introduction he makes an urgent plea for the establishment of uniform divisions of the Tertiary and for the international usage of common terms both as to life stages and life forms, and expresses confidence that approximate synchronisms can be established between the European and American Tertiary formations. For several years he has been attempting to correct these correlations and three trial sheets have been prepared and submitted to experts in Europe and this country for their criticism; as a result he publishes the following as a preliminary expression of the approximate correlations of formations:

I. STRATIGRAPHICAL CORRELATION: PRELIMINARY.

Lyell's System.		Approximate American Parallels.	
Pleistocene	{ Upper	Post Glacial	
	{ Middle	Glacial and Interglacial	
	{ Lower	Preglacial	? Equus Beds
Pliocene	{ Upper	Sicilien	? Blanco
	{ Middle	{ Astien Plaisancien	
	{ Lower	Messinien	Upper Loup Fork
Miocene	{ Upper	Tortonien	Loup Fork
	{ Middle	Helvetien	Lower Loup Fork
	{ Lower	Langhien	and Upper John Day
Oligocene	{ Upper	Aquitaniien	Lower John Day (Diceratherium Layer)
	{ Lower	{ Stampien Infra Tongrien	White River
Eocene	{ Upper	Ligurien	Bridger and Uinta
	{ Middle	{ Bartonien Lutetien	Lower Bridger Wind River
	{ Lower	Suessonien	Wasatch
	{ Basal	{ Thanetien Montien	Torrejon Puerco

All the levels above the Stampien are regarded by the author as imperfectly established.

In discussing the available evidences of parallelism, the following are cited as some of the more important tests: I. Common Genera; II. Similar Stages of Evolution; and he refers particularly, under this head, to the pattern of the molar teeth, the transformation of the pre-molar teeth, the complication of molar

teeth, and reduction of digits; III. Simultaneous introduction of new forms, or immigration from outside regions; IV. The Pre-dominance of Certain Types; V. The Convergence and Divergence of the Palearctic and Nearctic Faunae. In discussing the geographical distribution of Tertiary forms the original assumption is made that the animals of various families and orders have either originated in, or migrated into their present habitat in past time, so that the geological record as to their order of appearance is of first importance. And in investigating the distribution of the present time he finds it necessary to have an absolutely reliable correlation time scale for the fossil faunae, and the problem he then sets for investigation is to connect living distribution with distribution in past time and to propose a system which will be in harmony with both sets of facts. The author adopts Salter's classification of geographical areas, and the terms *Arctogaea*, *Notogaea*, *Neogaea*, and makes an interesting distinction between such large tracts of land as have, from their long separation, become the centers of adaptive radiation for orders of mammals, from those smaller regions which have been isolated from each other for shorter periods by climatic or physical barriers. The first group of divisions of the earth's surface are called *realms*, and these are the main centers of adaptive radiation for the orders of mammals. The smaller divisions he calls *regions* and they have served as the centers for the radiation of families of mammals. The author closes his paper with a discussion of the sources of migration of mammals to and from the three great realms of the earth as above described.

w.

7. *On the Flow of Marble under Pressure.*—A series of experiments has been carried on by F. D. ADAMS and J. T. NICOLSON with pure Carrara marble, having as their object the investigation of the conditions under which the flow of rocks may take place when subjected to differential pressure. Their experiments are to be extended later to other limestones, as also to granite and rocks of different types.

The method of experiment adopted was as follows: heavy wrought iron tubes were constructed of thin strips of Low Moor iron about one-fourth of an inch in thickness; these were so made that the fibres of the iron ran around the tube. Polished columns of the marble, about an inch or somewhat less in diameter, and an inch and a half in length, were inserted in the iron tubes, under such conditions as to give perfect contact, and to leave an inch and a quarter of the tube free at each end. Pressures up to 13,000 atmospheres were then exerted upon them by means of an accurately fitting steel plug. The marble was submitted to pressure under four conditions: first, dry, at ordinary temperatures, and again at 300° C. and at 400° C.; also in the presence of moisture at 300° C. The rate at which the pressure was applied was also varied, the time of experiment extending in the different cases between the limits of ten minutes to sixty-four days.

When the experiment was completed, the tube was slit longitudinally along two lines opposite to each other and the marble within exposed. It was uniformly found to be so firm and compact that a steel wedge was needed to split it through the middle, the half column adhering firmly to the sides of the tube. The deformed marble, while firm and compact, differs in appearance from the original rock in possessing a dead white color, somewhat like chalk, the glistening cleavage surfaces of the calcite being no longer visible. The difference was well brought out in certain cases where a certain portion of the original marble remained unaltered by the pressure. This had the form of two blunt cones of obtuse angle whose bases were the original ends of the columns resting against the faces of the steel plugs, while the apices extended toward one another into the mass of the deformed marble.

Special experiments on certain samples showed that, making all due allowance for the difference in shape of the specimens tested, the marble after deformation, while in some cases still possessing considerable strength, was much weaker than the original rock. It also appeared that when the deformation was carried on slowly the resulting rock was stronger than when the deformation was rapid.

By the study under the microscope of thin sections of the deformed marble, passing vertically through the unaltered cone and the deformed portion of the rock, the nature of the change that had taken place, could be seen clearly. The deformed portion of the rock could be at once distinguished by its turbid appearance, differing in a marked manner from the clear transparent mosaic of the unaltered cone. This turbid appearance was most marked along a series of reticulating lines running through the sections, which when highly magnified were seen to consist of lines or bands of minute calcite granules. These were lines along which shearing had taken place. The calcite individuals along these lines had broken down, and the fragments so produced had moved over and past one another, and remained as a compact mass after the movement ceased. In this granulated material were enclosed great numbers of irregular fragments and shreds of calcite crystals, bent and twisted, which had been carried along in the moving mass of granulated calcite as the shearing progressed. This structure was therefore cataclastic, identical with that seen in the feldspars of many gneisses.

Between these lines of granulated material the marble showed movements of another sort. Most of the calcite individuals in these positions could be seen to have been squeezed against one another and in many cases a distinct flattening of the grains had resulted, with marked strain shadows, indicating that they had been bent or twisted. They showed, moreover, a finely fibrous structure in most cases, which, when highly magnified, was seen to be due to an extremely minute polysynthetic twinning. The chalky aspect of the deformed rock was in fact due chiefly to the

destruction by this repeated twinning of the continuity of the cleavage surfaces of the calcite individuals, thus making the reflecting surfaces smaller. By this twinning, the calcite individuals were enabled under the pressure to alter their shape somewhat, while the flattening of the grains was evidently due to movements along the gliding planes of the crystals. In these parts, therefore, the rock presented a continuous mosaic of somewhat flattened grains. From a study of the thin sections it seemed probable that very rapid deformation tends to increase the relative abundance of the granulated material, and in this way to make the rock weaker than when the deformation is slow.

When the deformation was carried out at 300° C., or better at 400° C., the cataclastic structure described above, was not developed, and the whole movement was found to be due to changes in the shape of the component calcite crystals by twinning and gliding.

In the experiments where moisture was present, the deformation of the marble taking place at 300°, no change in the character of the deformation was detected. In other words the presence of water appeared to exert no influence, the marble was deformed as when dry at 300° by twinning and gliding alone without cataclastic action.

In order to ascertain whether the structures exhibited by the deformed marble were those possessed by the limestones and marbles of contorted districts of the earth's crust, a series of forty-two specimens of limestones and marbles from such districts in various parts of the world were selected and carefully studied. Of these, sixteen were found to exhibit the structures seen in the artificially-deformed marble. In these cases the movements had been identical with those developed in the Carrara marble. In six other cases the structures bore certain analogies to those in the deformed rock but were of doubtful origin, while in the remaining twenty the structure was different.

It appears, therefore, that by submitting limestone or marble to differential pressures exceeding the elastic limit of the rock and under the conditions described in this paper, permanent deformation can be produced. This deformation, when carried out at ordinary temperatures, is due in part to a cataclastic structure and in part to twinning and gliding movements in the individual crystals comprising the rock. Both of these structures are seen in contorted limestones and marbles in nature.

At elevated temperatures only the changes of the second kind, i. e. by twinning and gliding, take place. This latter movement is identical with that produced in metals by squeezing or hammering, a movement which in metals, as a general rule, as in marble, is facilitated by increase of temperature. There is therefore a flow of marble just as there is a flow of metals, under suitable conditions of pressure. The movement is also identical with that seen in glacial ice, although in the latter case the movement may not be entirely of this character.

The further results promised by the authors, of experiments on rocks of different kinds, will be looked for with interest.—*Proc. Roy. Soc.*, read June, 21, 1900.

8. *Report of the Section of Chemistry and Mineralogy*; by G. CHR. HOFFMANN, Geol. Survey of Canada. Part R, vol. xi, 1900. Among the various points of general interest in this report we note the following:

Celestite occurs in radiating columnar masses in the township of Bagot, Renfrew Co.; an analysis by Johnston showed the presence of 14 p. c. of barium sulphate. *Hübnerite* occurs in narrow seams and in irregular masses with coarsely laminated structure at Emerald, nine miles from Margaree Forks, Inverness Co., Nova Scotia. An analysis by Johnston showed it to be nearly pure manganese tungstate (0.47 FeO); sp. gravity = 6.975. *Hydromagnesite* occurs in considerable abundance on the Cariboo road, 93 miles north of Ashcroft, Lillooet district, British Columbia. Other extensive deposits have been found on the east side of Atlin Lake, Cassiar district. *Natron* has been identified in Goodenough Lake, 28 miles north of Clinton, Lillooet district, B. C. When examined, at the close of the dry season, it formed a deposit over nearly the entire bottom of the lake (20 acres) and with a thickness of about 8 inches. It was estimated that the deposit contained 20,000 tons. Natural soda also occurs in Last Chance Lake, 8 miles distant. *Polycrase* has been found in the township of Calvin, Nipissing district, Ontario. It forms crystalline masses, one of them weighing more than 700 grams; color pitch-black; luster resinous; sp. gravity 4.842. It is associated with xenotime, magnetite, etc. Only a qualitative analysis has been made thus far.

9. *Florencite, a new mineral*.—Florencite is a new hydrated phosphate of aluminum and the cerium earths recently described by HUSSAK and PRIOR. It was first found very sparingly in the cinnabar-bearing sands of Tripuhy, near Ouro Preto, Minas Geraes, Brazil, where it is associated with monazite, xenotime and the titanio-antimonates, lewisite and derbylite. It also occurs more abundantly in diamond-bearing sand from Matta dos Creoulos, near Diamantina and with the well-known yellow topaz at Morro do Caixambú. It occurs in rhombohedral crystals with f (02 $\bar{2}$ 1) as the prominent form, also c (0001) and r (10 $\bar{1}$ 1) and m (10 $\bar{1}$ 0) both rare. The angle $f^{\wedge} = 108^{\circ} 26'$ is not far from that of hamlinite (108 $^{\circ}$ 2'); the vertical axes are 1.1901 and 1.1353 respectively. Cleavage basal, fairly perfect; fracture splintery to subconchoidal; hardness about 5; specific gravity 3.586; luster greasy to resinous; color clear pale yellow.

The mean of two analyses by Prior gave:

P ₂ O ₅	Al ₂ O ₃	Ce ₂ O ₃ (etc.)	Fe ₂ O ₃	CaO	H ₂ O	SiO ₂
25.61	32.28	28.00	0.76	1.31	10.87	0.48 = 99.31

For this the formula is deduced 3Al₂O₃.Ce₂O₃.2P₂O₅.6H₂O, not far from that of hamlinite.—*Min. Mag.*, xii, 244, 1900.

10. *Elements of Mineralogy, Crystallography and Blowpipe Analysis from a practical Standpoint*; by ALFRED J. MOSES and CHARLES LATHROP PARSONS. New enlarged edition, pp. vii, 44. New York, 1900 (D. Van Nostrand Company).—This new edition of the Mineralogy by Professors Moses and Parsons appears in an enlarged and much improved form. The special features of the original work which made it particularly suitable for the use of the technical student, or the worker in the field,—as the classification of species and their clear and concise descriptions—have been retained, while the crystallographic portion has been rewritten and developed so as to conform to the modern method of treatment and classification. Other parts have also been revised and improved, so that the work as a whole may well commend itself to a wider range of students and readers than heretofore.

III. MISCELLANEOUS SCIENTIFIC INTELLIGENCE.

1. *Latitude-variation, Earth-magnetism and Solar activity.*—The following conclusions are reached by J. HALM after a discussion of the above subject.

(1) The changes in the motion of the pole of rotation round the pole of figure are in an intimate connection with the variations of the earth-magnetic forces.

(2) Inasmuch as the latter phenomena are in a close relation with the state of solar activity, the motion of the pole is also indirectly dependent on the dynamical changes taking place at the sun's surface.

(3) The distance between the instantaneous and mean poles decreases with increasing intensity of earth-magnetic disturbance.

(4) The length of the period of latitude-variation increases with increasing intensity of earth-magnetic disturbance.

(5) In strict analogy with the phenomena of auroræ and of magnetic disturbance, the influence of the eleven-year period of sun-spots, as well as of the "great" period, is clearly exhibited in the phenomenon of latitude-variation; and the same deviations from the solar curve as are manifested by the auroræ are also evident in the motion of the pole.

(6) The half-yearly period of the earth-magnetic phenomena influences the motion of the pole of rotation in such a way that its path, instead of being circular, assumes the form of an ellipse, having the mean pole at its center.

(7) The half-yearly period also explains the conspicuous fact of a rotation of the axes of the ellipse in a direction opposite to that of the motion of the pole.—*Nature*, No. 1610.

2. *The location of the South Magnetic Pole.*—The October number of the *Geographical Journal* contains an interesting account by C. E. BORCHGREVINK on the "Southern Cross" Expedition to the Antarctic in 1898-1900. The party wintered at Cape Adare on South Victoria Land, in latitude 71° S. and car-

ried on extensive exploration along the coast line, the farthest southern point reached being about latitude $78^{\circ} 45' S$. As the result of magnetic observations made by the Expedition, the present position of the southern magnetic pole is located at $73^{\circ} 20' S$, and $146^{\circ} E$. At Cape Adare the dip was found to be $-86^{\circ} 34'$ and the declination $56^{\circ} 2' E$. Observations of the dip were taken at seven other geographical positions, the maximum dip, $-88^{\circ} 1\frac{1}{2}'$, being taken at the foot of Mt. Melbourne; on the west side of Franklin Island a dip of $-86^{\circ} 52'$ was found. As Sir James Ross, in 1841, observed a dip of $-88^{\circ} 24'$ some twelve miles north of Franklin Island, there seems to have been a decrease in fifty-nine years of $1^{\circ} 32'$; it is concluded that there is very little doubt that the magnetic pole is very much farther north and west than in 1841.

3. *Ostwald's Klassiker der Exakten Wissenschaften* (Wilhelm Engelmann: Leipzig, 1900).—The following are recent additions to this valuable series:

Nr. 110. Die Gesetze des chemischen Gleichgewichtes für den verdünnten, gasförmigen oder gelösten Zustand, von J. H. Van't Hoff. Pp. 105.

Nr. 111. Abhandlung über eine besondere Klasse Algebraisch auflösbarer Gleichungen, von N. H. Abel (1829). Pp. 50.

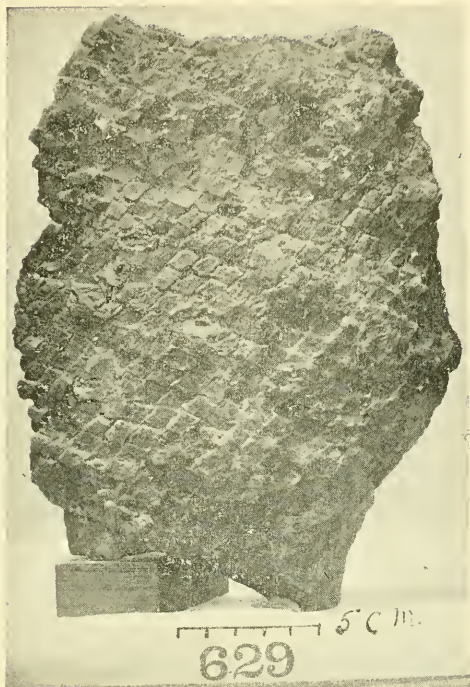
Nr. 112. Abhandlung über bestimmte Integrale zwischen imaginären Grenzen, von Augustin-Louis Cauchy (1825). Pp. 80.

Nr. 113. Zwei Abhandlungen zur Theorie der partiellen Differentialgleichungen erster Ordnung, von Lagrange (1772) und Cauchy (1819). Pp. 54.

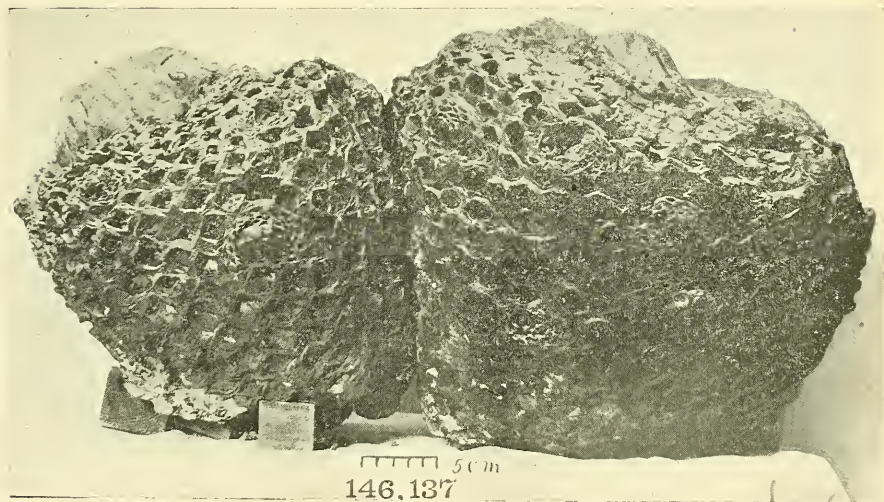
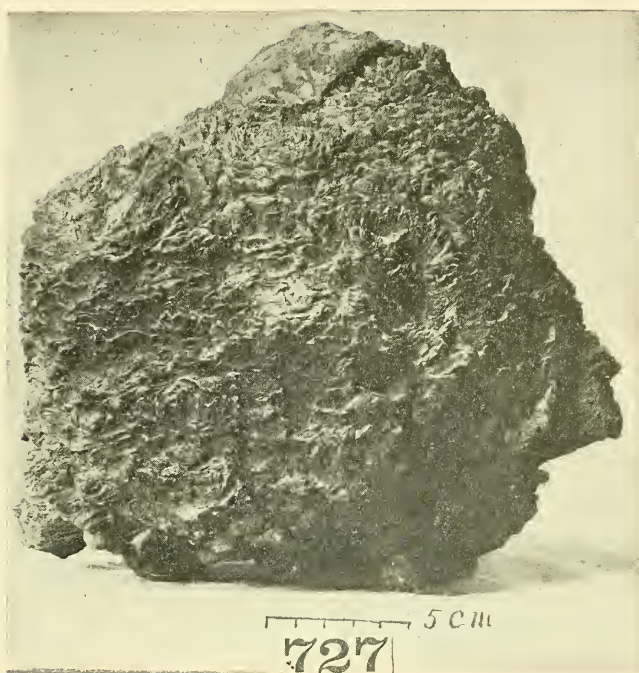
4. *Scientia*.—The inauguration of a series of small volumes on scientific subjects by G. Carré and C. Naud as publishers (Paris) was announced a year since (this Journal, July, 1899, p. 86). The seventh and eighth volumes of the physical-mathematical series have now been issued. The former is by H. LAURENT and is entitled "*L'Élimination*". The author remarks that the only monograph on the theory of elimination is that of Fàa de Bruno published in 1859; hence the discussion given by him has peculiar value. The other volume is on the subject "*Tonométrie*" by F. M. RAOULT.

5. *Introduction to Science*; by ALEX. HILL, M.D. 140 pp. New York and London (The Macmillan Co.).—This little volume contains a series of interesting chapters on some of the scientific problems now before the public, as on the age of the earth, the origin of species, the constitution of matter, etc.

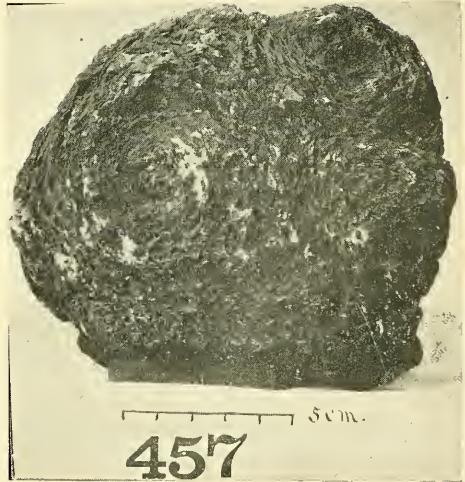
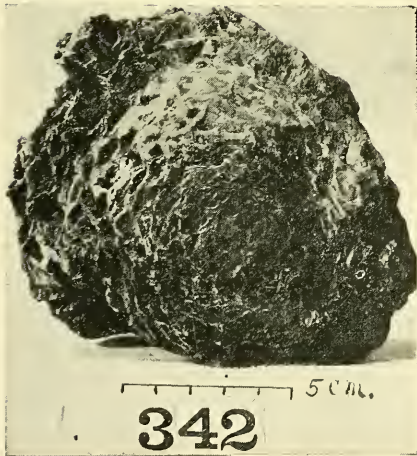
The Recent Solar Eclipse.—The report of the Expeditions organized by the British Astronomical Association to observe the total Solar Eclipse of May 28th, 1900, will be contained in a volume shortly to be issued from the Office of "Knowledge." The work will be edited by Mr. E. Walter Maunder, and will contain many fine photographs of the various stages of the Eclipse.



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THE
AMERICAN JOURNAL OF SCIENCE

[FOURTH SERIES.]

ART. XXXIX.—*Torsional Magnetostriction in Strong Transverse Fields and Allied Phenomena*; by C. BARUS.

1. IN 1885 Kirchhoff* published his remarkable research on the effect of stress on magnetization. This is the most comprehensive treatment which the problem has received and agrees in its applications with the deductions of Lorberg† and of Korteweg‡ for the analogous cases of electrostriction. It is well known that Maxwell's§ original stresses refer to a medium without structure. A subsequent generalization of von Helmholtz|| makes the stresses in the electric or the magnetic field dependent on changes of density in the medium while Kirchhoff's stresses allow for a dependence both on bulk and on elongation. These stresses thus contain three constants, the first of which is either the susceptibility or the permeability of the medium (with the corresponding constants for dielectrics) according as in the original statement the magnetization or the magnetic induction is to be expressed in terms of the field. His second and third constants correspond to the bulk expansion and elongation in question. If the third constant is annulled, Kirchhoff's stresses coincide with those of von Helmholtz; if the two new constants are annulled Maxwell's stresses may be reduced. Other great authors¶ (J. J. Thomson, Hertz) have contributed to the subject in similar ways.

* Kirchhoff: Wied. Ann., xxiv, p. 52, 1885.

† Lorberg: Wied. Ann., xxi, p. 300, 1884.

‡ Korteweg: Wied. Ann., ix, p. 48, 1880.

§ Maxwell: Electricity; § 105, p. 146, vol. i; § 644, p. 255, et seq., vol. ii.

|| Helmholtz: Wied. Ann., xiii, p. 400, 1881.

¶ The present meager account will suffice, since an excellent digest of the subject has been given in a paper by Nagaoka and E. Taylor Jones in Phil. Mag., (5), xli, p. 454, 1896.

Kirchhoff's theory has recently been examined in an elaborate experimental research due to Nagaoka and Honda,* with results showing an imperfect agreement with facts both for iron and for nickel. They not only include known reciprocal relations of stress and magnetization, but investigate new data particularly referring to changes of bulk. They find that Kirchhoff's new constants are complicated functions of strain, or that the expressions are applicable only to infinitely small strains.

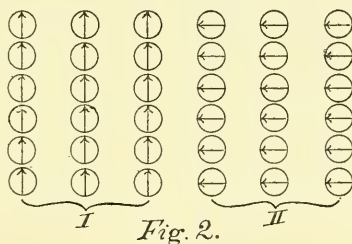
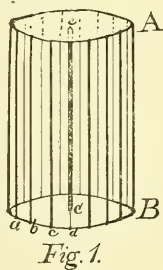
As the phenomena, therefore, remain almost equally troublesome in terms of the constants of the theory, very little assistance has been gained from it; and in the entire absence of a better theory it is still permissible to look at magnetostriction from tentative points of view.

2. I have long been of the opinion that a statistical treatment of the subject, such as was suggested for viscosity by Maxwell, might have much in its favor as compared with the purely elastic treatment independent of the mechanism, sketched above. The reciprocal relations of stress and magnetization are, as it were, incidents in a much more varied phenomenon. Indeed the original explanation of the Wiedemann effect, as given by Wiedemann himself in terms of Weber's theory of revoluble molecular magnets (a theory which in Ewing's hands has been shown to include hysteresis) seems to have been too consistently ignored. The conception of a magnetic configuration which breaks down under stress but is restored when the stress has sufficiently vanished or been reversed seems to be a reasonable one, supposing the breakdown involves no chemical change, as it does for instance in tempering. In Maxwell's hypothesis *any deformation* due to molecular instability is a viscous deformation. Now when the breakdown is gradual in character, as it must be when depending on temporary local intensities in the distribution of heat motion, the deformation will be gradual as actually observed in the ordinary phenomena of viscosity. If however breakdown is instantaneous (kaleidoscopic as it were), due for instance to the molecular shake-up accompanying magnetization or its withdrawal, then viscosity is *instantaneously a minimum*, and the deformation correspondingly sudden or "static." There seems to be no theoretical gap here whatever, and I hope to show (more directly in the subsequent paper than in the present, which has an introductory character) that

* Nagaoka and Honda: Journ. of the College of Science, Tokyo, Japan, vol. ix, part iii, pp. 353-391, 1898.

with the interpretation I have ventured to give of Maxwell's conception of viscosity, the phenomena of magnetostriction become more closely akin, following the same course in all metals.

In discussing experiments like the present it is always desirable to devise a model of the simplest kind, which shall suggestively reproduce the phenomena in question as fully as possible. The model need not be looked upon as an actual occurrence. I will therefore insert a simple torsional mechanism around which the present experiments may be conveniently grouped. If we regard the rod to be twisted as a bundle of longitudinal fibers elastically bound together, then the effect of twisting is in the main an elongation of the fibers increasing from the axis to the circumference. Inasmuch as the external fibers are now helical in form, the stretch in question has a horizontal component acting along the circumference of any section, tending to restore the fiber to its original straight form. In a solid rod fibers inclined at 45° to the axis are subject to traction (with the accompanying compression at right angles thereto) only. Other fibers are both stressed and rotated. For the present purposes, however, a description in terms of the principle stresses is unnecessary.



The case may be illustrated by the annexed model, in which the two parallel discs *A* and *B* are kept apart by the rod *C*, pivoted in a depression in each. Rubber fibers, *a*, *b*, *c*, . . . symmetrically grouped, hold the two discs together and secure the rod *C* in place. If torque be applied to *B* relative to *A* the rotation is forcibly resisted.

The bands *a*, *b*, *c*, . . . correspond to lines of stress which bind the non-magnetic molecules, longitudinally. In the simplest case one would assume that in a direction along the lines of stress molecules are closer together than are the successive lines themselves in a direction at right angles thereto; in other words, the distance apart of lines of stress is large as compared with the distance apart of the molecules which conduct stress.

If this condition were dropped the model, though still available, would lack simplicity.

The effect of longitudinal magnetization as inferred from the axial setting of the revoluble magnets, is increased tension along the fibers and therefore a corresponding increase of the horizontal component in question. Rigidity is apparently increased under all conditions, as is inferred from the distributions given in figure 2, type I. My own measurements* in addition to the more recent experiments of Mr. H. Day,† in which strong fields were applied, show an invariable increment of rigidity, increasing towards a limit as the field strength increases, and greater as the diameter of the rod is smaller. All this would be suggested by the model, remembering that slip takes place on twisting when the obliquity‡ of the external fibers has reached a fixed quantity.

If the increment of rigidity due to longitudinal magnetization were to be associated with the Joule effect of magnetic elongation, one would expect a change of sign corresponding to the march of the latter in an increasing field. Rigidity however is a regularly increasing quantity and shows no change of sign of the kind in question; and this is just what the model indicates. Again for the case of nickel in which the Joule effect is at the outset opposite in sign to that of iron, the effect of longitudinal magnetization is nevertheless increased rigidity. An example of these results will be found below, §6. Figure 2, type I, is an extreme case. One may note that the poles of the magnets are closer together than the molecular centers, either longitudinally or laterally.

The effect of transverse magnetization, if the hypothesis of lines of stress relatively far apart be retained, is given as one extreme case by type II of figure 2. The molecular magnets tend to lie across the lines of stress and rigidity is therefore diminished. From the geometry of the figure, however, it appears that the poles of the molecular magnets now lie much farther apart than in case I. Hence the transverse effect should be a decrement of rigidity, under like circumstances of much smaller numerical magnitude than the increment of rigidity due to the longitudinal field. Thus transverse magnetization is adapted to test a variety of phenomena, and it seemed to me that an examination of these in very intense fields (much more intense than are available for longitudinal magnetization) might throw new light on the nature of magnetostriction.

The case of circular magnetization I had hoped to omit, as it was studied at considerable length by the original inves-

* Barus; this Jour., (3), vol. xxxiv, p. 175, 1887.

† Howard Day; this Jour., (3), vol. iii, p. 449, 1897.

‡ Barus: l. c., pp. 182-183.

tigators, and recently in the exhaustive researches of C. G. Knott,* and others. In Knott's experiments longitudinal and circular fields are superposed and the data are made tributary to Maxwell's† and Chrystal's‡ theories, in which the stresses of the Joule effect are sufficient to account for the torsional phenomena. Drude§ uses similar reasoning to compute Kirchhoff's third constant.

The immediate effect of circular magnetism is a decrease of the rigidity of the wire (iron, nickel) through which the current flows. The case therefore corresponds at once to type II, figure 2, or to lines of stress relatively far apart as compared with molecular distances.

The presence of an electric current, however, is accompanied with an accession of heat in the wire through which the current flows, and the diminution of viscosity resulting, manifests itself in my apparatus and perhaps inevitably, as a diminution of rigidity. Hence in circular magnetization the heat effect is superimposed on the magnetic effect of the same sign and in thin wires it is difficult to separate them satisfactorily, if at all; for even if the wire carrying current is submerged in water there is still grave room for doubt, as will appear below.

For these and similar reasons I endeavored to investigate results with thin wires subjected to the strong transverse fields in the air gap of an electromagnet, for which conditions no data have as yet been forthcoming. In spite of the toil spent upon the work, however, my endeavors have not brought out sharp results, except in so far as they furnish a superior limit for the change of rigidity sought; and even this is a small and uncertain residuum raising a doubt as to whether rigidity is at all influenced by strong transverse fields.

Having completed this work I thus found myself under the necessity of taking up the question of the effect of circular magnetization on rigidity for comparison. The results were again such as to place doubt on the occurrence of such an effect.

3. *Method of Experiment.*—The apparatus used was virtually the same as that of my earlier experiments on the subject, except that the helix formerly surrounding the upper wire is replaced by the fissure-like air gap of the strong tubular electromagnet, *A*, figure 3. The upper of two identical soft iron wires, *ab* and *cd*, is placed in this gap to be strongly magnetized when the current flows through *A*. Any twist may be applied to the system *ab*, *cd*, by the torsion heads *D* and *E*,

* Knott: Trans. Roy. Soc. Edinb., (2), xxxvi, pp. 485-535, 1891.

† Maxwell: Electr. and Mag., ii, p. 109.

‡ Chrystal: Encyclop. Britanica, Art. Magnetism, p. 270.

§ Drude: Wied. Ann., lxiii, p. 9, 1897.

and both wires are thus subject to the same torque. The change of rigidity due to the transverse magnetization is observed by the mirror, *m*, reflecting a beam of sunlight upon a wall 5 meters distant. Hence 1^{cm} of deflection corresponds to about .001 radian or .057°. The mirror is provided with vanes, *vv*, dipping in an annular trough (not shown) to deaden vibration, as well as to carry a current from *E* to *v* for the circular fields. The rod, *e*, is suitably provided with a pin and slot arrangement, so that the wires may be kept under the definite tension of the weight, *W*.

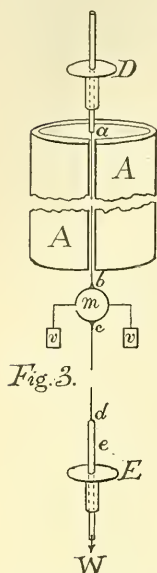


Fig. 3.

The core of the electromagnet, *A*, was of thick gas pipe about 37^{cm} long, 5.5^{cm} in diameter and .3^{cm} thick, the longitudinal (vertical) air gap being about .3^{cm} broad. It was wound from end to end with about 120 turns of wire. A fine vertical glass tube (not shown) about .1^{cm} in bore was symmetrically secured within the gap running from end to end. Through this tube the wire, *ab*, was threaded to prevent adhesion to the sides of the jaws of the electromagnet. This occurrence in fact constitutes the chief difficulty in the experiments. Deflection of the wire toward either jaw is liable to be accom-

panied with spurious rotation. On the other hand too fine a tube interferes with the motion frictionally. In later experiments I discarded the glass tube in favor of a more carefully centered free wire.

During the tests with circular magnetization the lower wire carrying the current was submerged in a wide glass tube of flowing water.

The field *H*, within the air gap of the helix, measured ballistically with the aid of an earth inductor, had approximately the following c. g. s. values for different currents, *C*, in amperes. An attempt to secure sharp absolute values was needless.

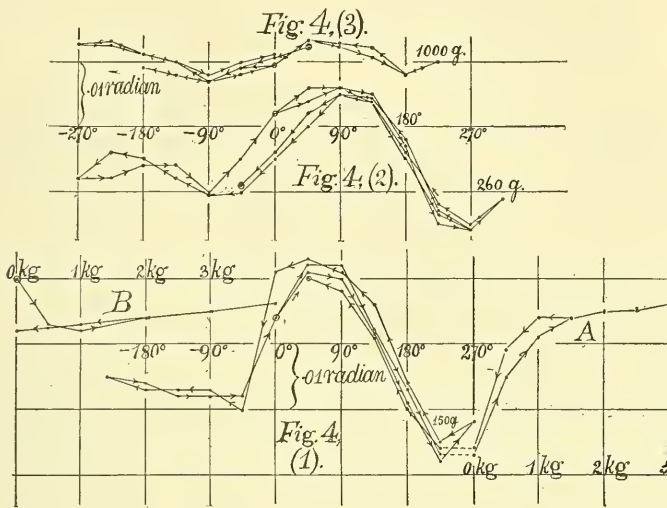
<i>C</i> =	1	2	3	4	5	6	7	8	9	10	11	12	13
<i>H</i> =	1200	2700	4300	5700	7150	8500	10000	11100	12000	12650	13200	13700	13950

Between 2 and 7 amperes in the helix the variation of *H* with *C* is very nearly linear. Below this it is slightly accelerated, above rapidly retarded. Since all fields were made from zero, the effect of hysteresis could be disregarded as was specially tested. The field here referred to is that within the air gap and found by differential experiments. If *B*₁ be the induction within the gap and *B*₂ the total induction, including the

* Barus: this Jour., (3), xxxiv, p. 176, 1887.

stray field outside of the gap, the galvanometer for a loop within the gap will respond to $B_1 - (B_2 - B_1)$. Two measurements were therefore made to compute B_1 .

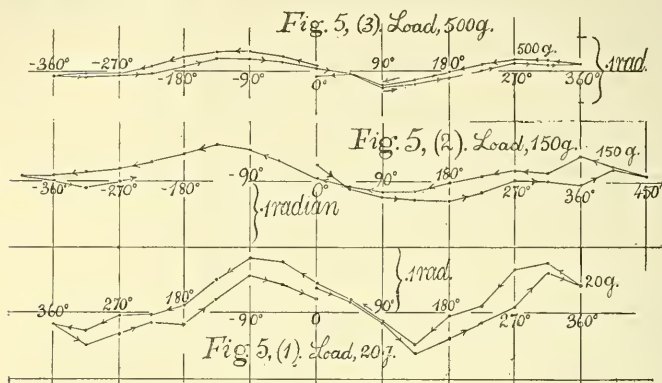
To prevent excessive heating currents between 5 and 6 amperes were usually employed, though these were not sufficient to saturate the core of the helix as the above data and the results below show. I may add that the ordinary effects of viscosity and slip are eliminated by the method of experiment, since the field was alternately made and broken several times.



The change δn of rigidity n , is at once given in terms of the deflections $\delta\theta$ of the mirror; for $\delta n/n = \delta\theta/\theta$, very nearly, if for the same section and length, θ is the twist simultaneously imparted to each wire. In fact the above method was devised to secure this convenience. If $\delta\theta/\theta$ is positive, i. e. if in the above apparatus the twist is imparted clockwise looking down the wire and the deflection due to magnetization is also clockwise from the same point of view, then the magnetic change of rigidity is an increment for upper or magnetized wire; and vice versa.

4. Results.—The chief difficulty encountered has already been suggested. If the wire is carefully straightened and adjusted in the middle line of the vertical air gap, the wire nevertheless becomes slightly sinuous or wavy when the magnetic field is excited. Different parts of the length of the wire are unequally attracted. If these parts lie across the field such attraction is accompanied by rotation and the observed

phenomenon, otherwise quite constant throughout, is spurious. As an example I may give the series of results, figure 4 (1, 2, 3,) and figure 5 (1, 2, 3), for annealed iron wires $.044\text{cm}$ and $.024\text{cm}$ in diameter, respectively, and 41cm long each. The twist, θ , imparted to either wire is given in degrees by the abscissas, positive or negative as specified. The ordinates show the deflections $\delta\theta$, in radians, obtained when the field is made, so that $.057 \delta\theta/\theta = \delta n/n$ is the increment of rigidity.

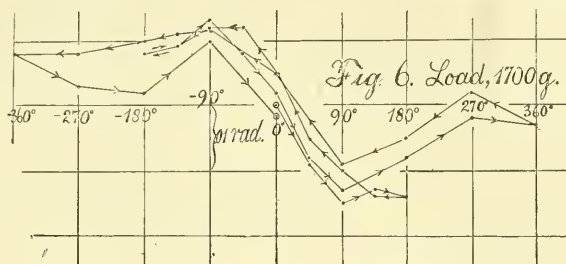


The consecutive curves given correspond to successively increasing loads or vertical pulls on the system, as indicated. Observations were made in the cyclical way shown by the arrows. The same zero was retained throughout. In general for positive twists the return curve is positive relatively to the outgoing curve; for negative twists the return curve is negative relatively to the outgoing curve. Definite consideration will be given to this interesting result in my next paper so that the present reference will suffice for this and succeeding cases. The effect of loading (among many results of a similar kind) is specially shown in figure 4 at *A*, *B*, giving evidence even of sign reversal. The deflections are large for the thin wires as compared with the thicker wires (different scales are needed in the two cases). Special experiments showed that if a series is repeated the curves obtained, however irregular, are identical.

Notwithstanding the trustworthy character of the results obtained, I became convinced that the excessive effects produced by loading, and hence the whole phenomena, were somehow in error.

5. In the experiments of § 4, the magnetized wire was enclosed in the fine glass tube mentioned above. In the present experiments it is free (the tube being withdrawn) and visibly suspended in the air gap. The latter was adjusted after

each twist so that on making the field the transverse motion of the wire was the smallest possible. Rarely did I succeed in quite eliminating it. Moreover, to guard against errors of the

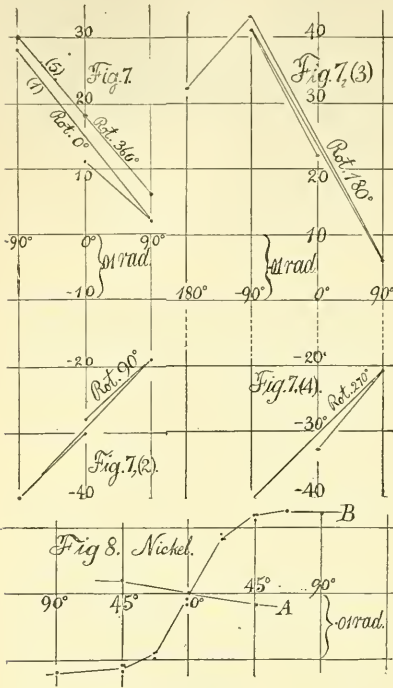


preceding kind the wire was as heavily loaded as its section permitted, without passing the limits of elasticity. An example of the results so obtained is given in figure 6. The load suspended from the system of thin wires was about 1700 grams. The character of the observations has remained essentially the same as in the foregoing experiments, and the additional care bestowed has not reduced the complications.

To test the trustworthiness of these observations and to interpret them, I now *rotated both wires* of the system around their common axis and examined them for each 90° of displacement. The data are given in figure 7 (1-5), which is an example of many similar results with other wires. To facilitate the work the twists applied lay between $+90^\circ$ and -90° . The slopes of the lines obtained are about the same, but they are alternately positive and negative. In general when the deflection for no twist (fiducial zero) is positive, the slope is positive and vice versa: though neither this nor the alternation of sign seemed to be the invariable rule. One may note that the data for the initial (0°) and final (360°) positions agree very well as to slope.

The curves of figure 7 may be interpreted in two ways: either the wire is aeolotropically different in the four radial directions examined, which would indicate remarkably complex structure, or the wire is geometrically dissimilar, being either elliptical in section or not quite straight, or both. In both cases the curves obtained would be adventitious and the alterations observed easily explained: for if the wire is virtually ribbon-shaped with its plane oblique to the lines of magnetic force, the tendency of the ribbon will be to set "axially" when the field is made. For positions 180° apart the torques will be identical; for positions 90° apart they will vary as the sine and cosine, respectively of the obliquity, and will be numerically opposite in sign. If the obliquity be 45° or nearly so, the torques will be equal and opposite.

Assuming that in the given curves, figure 7, the obliquity of 45° has been reached, then from two consecutive positions of the wires relatively to the field (0° and 90° , 90° and 180° , etc.)



the spurious effect may be eliminated. The slope so found, however, lies within the errors of hysteresis, set, and the errors of orientation and can not be considered more than an estimate. For the mean slope indicates an effect of but $.11^\circ$ per 180° of twist, or 2^{cm} of deflection for a scale distance of over 5 meters. Hence even in the strongest transverse magnetic fields the persistent increment of rigidity remains small and is negligible in comparison with the spurious effects encountered.

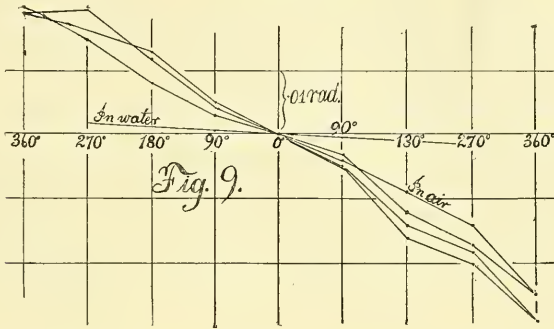
6. After reaching this unexpected result it seemed necessary to verify it in other directions. In the first place the question occurred whether compatibly with figure 2, type I, the

effect of longitudinal magnetization was an increment of rigidity in all metals. I therefore tested a nickel wire (length 41^{cm} , diam. $.048^{cm}$) in the longitudinal field by replacing the transverse helix *A*, figure 3, by an ordinary helix. An example of the uniform results is given by figure 8, *B*, and bears this out (current 4 amperes, load 150 grams, field 400 c.g.s. units).

On the same diagram will be found the apparent effect, *B*, of circular magnetization produced by a current of about 2 amperes passed through the wire, freely suspended in air. The obvious decrement of rigidity resulting seemed suspicious, however, inasmuch as the deflection on making the circuit was rapid and the return to zero on breaking, prolonged. This is nearly what would occur if the observed decrement of rigidity was due to heat alone.

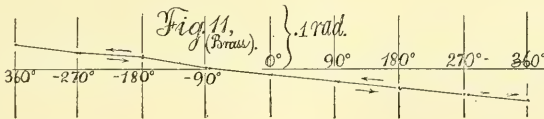
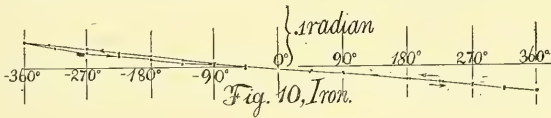
7. Final experiments were therefore directed to a comparison of the effect of a circular field in iron and the effect of a similar field in brass, yielding it would appear a straightforward method of decision. In figure 9, I give a preliminary

example of the decrement of rigidity due to a current of one ampere in an iron wire suspended in air, compared with the effect when the wire is submerged in flowing water, other things being equal. It will be seen how relatively large the heat effect is, the decrements for the submerged wire being



less than 1^{cm} at 5 meters of distance and 90° of twist lying within the errors of measurement. Apart from this the question is open whether even the submerged wire is not heated to an extent appreciable by the method of observation.

Hence in my concluding experiments I tested an iron and a brass wire of about the same dimensions, in the same apparatus and under like condition consecutively. Both were submerged in running water. The results for iron are given in figure 10, those for brass in figure 11. The wires were sufficiently thin



and the current strong enough to give large deflections (up to 50^{cm}). The current and diameter in brass (6 amp., .020^{cm}) are smaller, in iron (7.5 amp., .044^{cm}) larger. In both cases, however, there is a definite decrement of rigidity due to the circular field of about the same order, and in both cases so far as can be made out it is a mere heat effect.

I am obliged therefore to conclude, that in so far as these experiments have weight the effect of longitudinal magnetiza-

tion is an increment of rigidity in all paramagnetic metals; whereas the permanent effect of a transverse or a circular field is relatively inappreciable so far as rigidity is concerned. At least its certain detection would have to be left to researches of a higher order of refinement than was reached in the method pursued. This conclusion is at variance with much of the earlier work on the subject, but I do not see how the step for step march toward this result which the above simple experiments contain can be evaded.

Finally, the suggestions of the model (figures 1 and 2) are in keeping with the data found. Cf. § 2.

The quantity which I have here considered is the *permanent* effect of magnetization or rigidity, i. e., that which persists after making and breaking the field many times. Besides this there is a very striking temporary effect of marked relative value, in the interest of which the present experiments have largely been made. The relation of this temporary effect to viscosity will be considered in the paper soon to appear.

Brown University, Providence, R. I.

ART. XL.—Notes on Tellurides from Colorado; by
CHARLES PALACHE.

1. *Sylvanite from Cripple Creek.*

THE presence of sylvanite among the telluride ores of Cripple Creek was first made known by the analysis of Pearce,* and the considerable silver content of the ores was credited by him, in large part at least, to that mineral. His analysis however was made on massive material, and the lack of crystallographic evidence of the presence of sylvanite, together with the failure of other investigators to find the mineral, led to some mistrust of his results and the silver content of the ores was doubtfully attributed to the common ore, calaverite, which was known to carry a small percentage of silver.

Of recent years, however, crystallized sylvanite appears to have been found in considerable amount in several of the Cripple Creek mines and it is the purpose of this paper to describe a series of sylvanite crystals which offer most satisfactory proof of the correctness of Pearce's original conclusion.

The crystals to be described were placed at the writer's disposal for examination by Professor Hobbs of the University of Wisconsin, to whom they had been presented by Mr. F. M. Woods of Victor, Colorado, the collector. The crystals were in two lots labelled respectively: "Sylvanite," Mabel M. Property, Beacon Hill, Victor; and "Calaverite," Victor, Col.

All proved to be alike sylvanite and consisted of isolated crystals and crystal fragments varying in size from 1^{mm} to 8^{mm} in greatest dimension. The color is pure silver-white and many of the fragments display the perfect cleavage parallel to the clinopinacoid which is characteristic of this mineral and serves well to distinguish it from the more common calaverite. The specific gravity, determined on three isolated crystals on the hydrostatic balance, was 8.161.

The dominant habit of the crystals is thin tabular parallel to $b, 010$, the edges of the tables being bounded by planes of the orthodome zone. The tables are frequently as thin as paper and the edge planes become too small to be measurable. A second well-marked habit is prismatic, determined by the pronounced development of the zone of the positive unit pyramid and orthodome, (111) and (101), and other pyramids of this zone. As will be seen from the figures, crystals of this habit are frequently rich in forms, some of which proved to be new to the mineral.

* Proc. Colo. Sci. Soc., 1894.

A skeletal development, well known on sylvanite from other localities, is common on these crystals; but it did not affect the character of the crystal planes which were in general brilliant, giving good reflections.

Measurements were made on the two-circle goniometer and the adjustment of the crystals was rendered very accurate by the use of the three pinacoids, nearly always all present.

The following twenty-nine forms were observed on the five crystals measured:

<i>c</i> , 001	<i>m</i> , 101	<i>r</i> , 111	<i>u</i> ,* 723
<i>a</i> , 100	<i>n</i> , 201	<i>w</i> ,* 343	II, 341
<i>b</i> , 010	N, 201	<i>s</i> , 121	<i>y</i> , 123
<i>R</i> , 120	<i>d</i> , 011	<i>o</i> , 131	<i>ρ</i> , 111
<i>e</i> , 110	<i>v</i> ,* 525	<i>q</i> , 141	<i>σ</i> , 121
<i>f</i> , 210	<i>γ</i> , 212	<i>i</i> , 321	<i>J</i> , 321
<i>g</i> , 310	<i>t</i> , 323	<i>j</i> ,* 521	<i>κ</i> , 521
			<i>Y</i> , 123

In the following table are given the computed and measured angles for the four new forms. The forms *o*, 131, *q*, 141 and II, 341 while not new, have not been observed since Miller's studies, and are here added, with their computed angles, in order to complete the tables of Goldschmidt by whom they were omitted as uncertain. The new form *u*, 723, is the most interesting, occurring in nearly all the measured crystals with good development and being apparently characteristic for this locality.

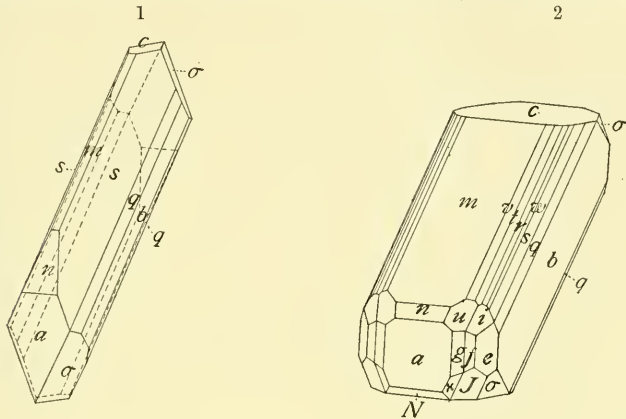
Letter.	Miller symbol.	Computed angle.		Average measured angle.		Variations.				Number of observations.	Quality of faces.
		ϕ	ρ	ϕ	ρ	ϕ		ρ			
						+	-	+	-		
<i>v</i>	525	57°·06'	39°·41'	57°·03'	39°·36'	12'	13'	7'	3'	5	fair
<i>w</i>	343	24·53	58·52	24·49	58·57	9	8	7'	6	2	poor
<i>u</i>	723	65·04	60·42	65·01	60·36	14	5	12	15	6	good
<i>j</i>	521	56·53	76·22	56·50	76·25					1	excel-
II	341	24·44	78·36								lent.
<i>o</i>	131	11·39	73·50								
<i>q</i>	141	8·47	77·38								

The figures illustrate the typical prismatic habit, as shown in various combinations of planes. In fig. 1 the pyramids are developed at the expense of orthodome and clinopinacoid; in fig. 2 the latter forms predominate, giving a tabular-prismatic character to the crystal. This zone of pyramids is generally

* Forms marked thus are new.

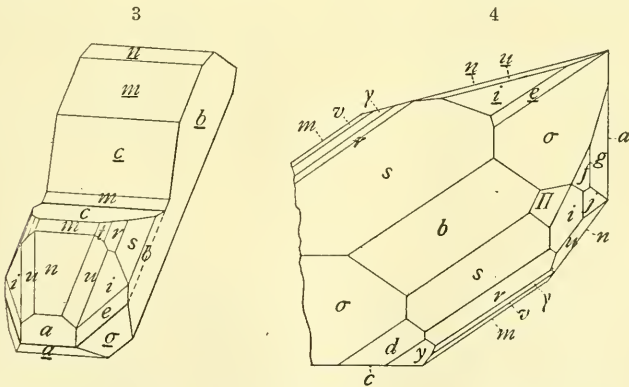
rich in forms as shown in fig. 2 and is sometimes deeply striated. It forms a satisfactory means of distinguishing the positive from the negative octants, which in its absence, owing to the nearness of the angle β to 90° in sylvanite, is not always easy to do.

Of the other pyramidal forms, σ , $\bar{1}21$, is the most common, being rarely absent and generally relatively large in size. The other pyramids, both positive and negative, and the prisms are



of very subordinate importance in defining the outlines of the crystals.

Two of the measured crystals were found to be twinned according to the common law for sylvanite, twinning plane the



orthodome, 101. Figs. 3 and 4 reproduce these twinned crystals in about the proportions of the originals. Fig. 3 will be seen to be a simple contact twin, in which, however, the two

individuals, in contact along the plane 101, are very unlike in development. The crystal figured in normal position is bounded by numerous positive pyramids and domes while the twinned individual is much larger and a much simpler combination. The reëntrant angle between the two basal planes with the intervening narrow portion of the twinning plane 101 made the twinned character of the crystal quite evident. In a subordinate development this mode of twinning is quite frequent, the twinned individual appearing as a narrow transverse ridge on the surface of the *m* plane of a larger simple crystal such as shown in fig. 2.

A less simple phase of the twinning is represented in fig. 4 in orthographic projection on 010, where apparently no plane boundary exists between the two individuals. The planes *n*, *u*, *z*, and *e*, constituting the upper part of that figure were only recognized as in twin position by the complex symbols which they yielded in the projection, whose study, in the light of the simple twin already observed, showed their true nature. It is possible, however, that this is a contact twin like No. 3; for the remaining forms shown in the upper part of fig. 4, *s*, *r*, *v*, etc., are not affected in their position by the twinning; but in the absence of the planes bounding the opposite end of the crystal the course of the twinning boundary cannot be accurately defined.

A chemical analysis was made upon about .5 grams of the sylvanite crystals, which were carefully picked over under the lens until apparently entirely freed from adhering gangue. The analysis shows, however, that considerable of the siliceous gangue was still present. The composition is that of a normal sylvanite, Au, AgTe_4 .

	Sylvanite.	Cal. to 100%.	Mol. ratio.
Insol.	1.02		
Au	26.09	26.25	.1334
Ag	12.49	12.57	.1164
Te	60.82	61.18	.4894
Fe	1.19		
	<hr/>	<hr/>	
Total	101.61	100.00	$\left. \begin{array}{l} \\ \\ \\ \end{array} \right\} = 1$ $= 1.96$

2. *Crystallographic Identity of Goldschmidtite with Sylvanite.*

In 1899, Hobbs* described goldschmidtite, a new silver-gold telluride, intermediate in composition between sylvanite and calaverite. The analysis was made on a very small amount of material and seemed therefore open to question, but the crystallographic results appeared to place the mineral distinctly

* This Journal, vii, 357, 1899.

apart from others. Through the kindness of Professor Hobbs the writer was entrusted with two of the five type crystals of the new mineral for examination and thus familiar with its appearance, recognized it again in a specimen in the Harvard collection secured some time before from Cripple Creek and labelled as from the Little May Mine. Several measurable crystals were obtained from the latter specimen and they were found to agree in general with the two type crystals of goldschmidtite examined. The author's measurements of Hobbs' two type crystals confirmed in general his results as shown by the table of angles given below but the use of the two-circle goniometer made possible the measurement of a number of minute pyramid planes which Hobbs had been compelled to leave undetermined. These same forms and some additional ones were also found on the new material, thus confirming the identity of the two specimens. These pyramid forms, projected on the goldschmidtite axes, gave symbols far from simple, and unsatisfactory relations with the other forms. They were studied in gnomonic projection, and a comparison of such a projection of the goldschmidtite forms on the clinopinacoid with a similar projection of sylvanite forms showed a striking analogy between the two. By making the orthopinacoid 100 of goldschmidtite equivalent to the orthodome 101 of sylvanite many forms of the first become identical with known forms of the second, and the remainder, while apparently new to sylvanite, receive comparatively simple symbols on the sylvanite axes. In the following table the agreement of angles of the common forms in the two principal zones is well shown.

Symbols Goldschm.	Symbol Sylvanite.	Obs. (Hobbs).	Obs. (Palache).	Calc. (Sylvanite).
100 : 110	=101 : 121	61° 41'	61° 41'	61° 35½'
100 : 210	=101 : 111	42 43	42 56	42 45
100 : 230	=101 : 131	----	70 11	70 10
100 : 310	=101 : 323	31 55	31 40	31 38½
100 : 101	=101 : 100	55 35	55 15	55 08
100 : 201	=101 : 001	34 13	34 16	34 27
101 : 201	=100 : 001	89 48	89 31	89 35
100 : 001	=101 : 201	89 25	88 38	88 48
100 : 401	=101 : 201	19 18	19 30	19 20

Much had been learned concerning the crystallographic character of goldschmidtite when the sylvanite crystals above described came to hand, and their study helped to clear up and make certain the relation between the two minerals. The general similarity between the habit and forms of the two series of crystals led to a surmise that a concealed twinning

like that found in the sylvanite crystals was also present in goldschmidtite, giving rise to the many apparently new forms referred to above. And a comparison of the measurements and projections finally proved that this was the case for most of these forms. The following table shows the sylvanite forms equivalent to those observed on goldschmidtite. It will be observed that it is divided into three parts containing (1) forms given by Hobbs and reobserved by the author; (2) forms observed by the author only; and (3) forms given by Hobbs which were not observed on any of the author's material and which lead to new forms for sylvanite.

(1)				(2) (cont'd)			
Forms observed by Hobbs and Palache.				Additional Forms observed by Palache			
Goldschmidtite.		Sylvanite.		Goldschmidtite.		Sylvanite.	
Letter.	Symbol.	Letter.	Symbol.	Symbol.	Letter.	Symbol.	Symbol.
<i>c</i>	001	<i>N</i>	$\bar{2}01$	$\bar{1}34$	<i>J</i>	$\bar{3}21$	
<i>b</i>	010	<i>b</i>	010	$\bar{1}12$	κ	$\bar{5}21$	
<i>a</i>	100	<i>m</i>	101	230	<i>o</i>	131	
<i>g</i>	310	<i>t</i>	323	120	<i>q</i>	141	
<i>f</i>	210	<i>r</i>	111	832	<i>I</i>	211 twinned	
<i>m</i>	110	<i>s</i>	121	$\bar{5}32$	<i>i</i>	321	
<i>k</i>	032	<i>\theta</i>	$\bar{2}31$	532	<i>i</i>	321 twinned	
<i>w</i>	401	<i>n</i>	201 twinned	734	<i>j</i>	521 twinned	
<i>n</i>	201	<i>c</i>	001	102	<i>M</i>	$\bar{1}01$	
<i>s</i>	101	<i>a</i>	100 twinned				
<i>S</i>	$\bar{1}01$	<i>a</i>	100				
<i>N</i>	$\bar{2}01$	<i>c</i>	001 twinned				
<i>W</i>	401	<i>n</i>	201				
(2)				(3)			
Additional Forms observed by Palache.				Forms observed by Hobbs not confirmed by Palache and yielding new sylvanite forms.			
Goldschmidtite.		Sylvanite.		Goldschmidtite.		Sylvanite.	
Symbol.	Letter.	Symbol.	Symbol.	Letter.	Symbol.	Symbol.	Symbol.
$\bar{2}32$	<i>e</i>	110		<i>t</i>	370		292
232	<i>e</i>	110 twinned		<i>l</i>	130		161
$\bar{4}34$	<i>f</i>	210		<i>v</i>	35·0·1		9·0·10?
434	<i>f</i>	210 twinned		<i>x</i>	10·0·1		403 twinned
$\bar{2}12$	<i>g</i>	310		<i>q</i>	801		203
212	<i>g</i>	310 twinned		<i>r</i>	703		{ 2·0·25 or 301 twinned
$\bar{1}31$	<i>R</i>	120		<i>y</i>	508		{ 708 or 501 twinned
$\bar{2}34$	<i>\rho</i>	$\bar{1}11$ twinned		<i>X</i>	$\bar{1}0\cdot0\cdot1$		403
132	<i>\sigma</i>	$\bar{1}21$		<i>Z</i>	$\bar{1}4\cdot0\cdot1$		504
$\bar{1}32$	<i>\sigma</i>	$\bar{1}21$ twinned					

The forms in (3) appear to need confirmation before being added to the sylvanite series as the following considerations will show. None of them was observed by the author on the

twelve or more supposed goldschmidtite crystals from the Little May Mine measured by him nor on the complex sylvanite crystals from the same region. In the absence of definite statements concerning the quality or frequency of occurrence of particular forms in the paper by Hobbs it is difficult to know what weight to assign to these forms. *t*, 292 and *l*, 161, are pyramids of the principal zone of sylvanite and may well be good forms. The seven remaining forms are orthodomés. Of these, two, *q*, 203 and *X*, 403 with its twin, *x*, 403 are probably good forms as they have simple symbols and agree closely in angle with calculated position :

		Meas.	Calcul.
<i>q</i> , 203	∧ 001	24° 37'	24° 20'
<i>X</i> , 403	∧ 001	42 34	42 36
<i>x</i> , 403 twin	∧ 001	26 25	26 24

v, 9·0·10, stands only 2° 20' from the dominant form 101 and may be vicinal to it; *r*, 2·0·25, is inclined but 3° 10' to 001, of which it is very likely a vicinal; *y*, 708, is inclined 3° 27' to $\bar{1}01$ and may be vicinal to it; lastly *Z*, 504 is inclined 2° 16' to 403 and might be considered vicinal to that form were it confirmed.

It seems evident in the light of the above facts that so far as crystallographic character is concerned goldschmidtite can not be distinguished from sylvanite but represents a peculiar habit of that mineral, common, as shown in the preceding paper, to the sylvanite of Cripple Creek. Fig. 1 of the preceding paper on sylvanite illustrates fairly well the prevailing habit of goldschmidtite when in the sylvanite position, except that σ , $\bar{1}21$, should be very minute, and most crystals are simple contact twins.

The results of this study were submitted to Prof. Hobbs in the hope that he might be able to complete them by a new analysis of the type material on a larger amount of substance. This proves unfortunately to be impossible, since the specimen from which the type crystals were obtained was only temporarily in his hands and is not now available.

In default of this analysis it was thought that it might be worth while to determine the composition of the supposed goldschmidtite from the Little May Mine, and by sacrificing the whole specimen and carefully picking the crushed gangue material enough was obtained (about 0·4 gram) for an analysis. This purpose was, however, defeated by an accident early in the work, so that only the gold content could be determined. For the sake of comparison this is given below together with that of goldschmidtite and sylvanite.

Little May Mine.	Goldschmidtite.	Sylvanite AuAgTe ₄ .	Sylvanite Cripple Creek.
Au 28·89	31·44	24·45	26·09

While little weight can be attached to a comparison based thus on a defective analysis, still it seems clear that the material studied differs but little from the typical goldschmidtite in composition. On the other hand, while the gold content is high when compared with that of the theoretical AuAgTe_6 , it should be clearly remembered that no sylvanite with exactly that composition has been as yet analyzed, the gold content actually found varying from 25.87 per cent to 29.35 per cent with proportionately varying silver content. These results seem to show conclusively that the ratio of gold to silver in sylvanite may vary considerably from the theoretic proportion of 1 : 1, without affecting the physical characteristics materially, and it is a question whether it is advisable to attempt to establish species based upon these variations of composition.

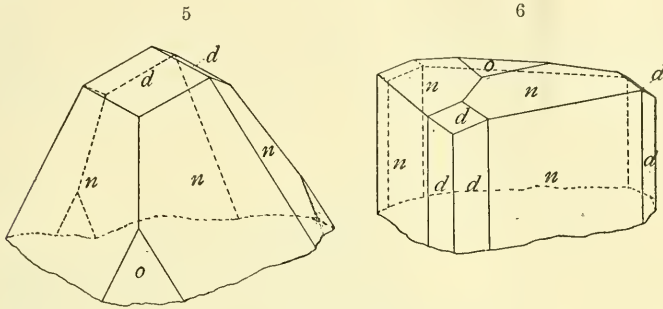
Note by Professor W. H. Hobbs.—In view of the results of the investigation by Dr. Palache, above detailed, it seems proper for me to say that the name goldschmidtite should be withdrawn from mineralogical literature as representing a distinct mineral species. Dr. Palache's study shows that goldschmidtite can be referred to the same set of axes as sylvanite, of which it represents a peculiar type. This being true my analysis, which had to be made on an extremely small amount of material, must contain a large error. The method used was the oxidation of the tellurium on charcoal, weighing the button of combined silver and gold, and, after solution of the silver, weighing the gold in the form of powder. The danger of this method lies in the possibility that tellurium will not be completely eliminated and that some silver will oxidize. The button obtained was, however, bright and apparently freed from tellurium. I regret that material is not now available for a second analysis, yet in view of Dr. Palache's crystallographic study of better material, checked as it has been by analysis, no course is open to me but to discredit the results of my analysis.

3. *Hessite Crystals from Colorado.*

A specimen of well-crystallized hessite from Boulder Co., Colorado, has been recently acquired by the Harvard Mineral Cabinet from Mr. G. B. Frazer. A description of this specimen is here offered because the crystals present certain interesting peculiarities of habit and because so far as the author has been able to discover no crystals of this mineral from the United States have yet been figured.

The specimen consists of a small fragment of bluish vein quartz in one side of which is a drusy, quartz-lined cavity. On the quartz walls of the cavity are eight or ten brilliant hessite crystals, a millimeter or less in height, and two small

tufts of wire gold. Two of the steel-gray hessite crystals and a small irregular fragment were detached, the latter yielding before the blowpipe the characteristic reactions of hessite.



The two crystals (figs. 5, 6) which are shown in about their natural proportions in the figures, presented two entirely different habits, both hexagonal in appearance; but measurements proved them to be isometric and combinations of the three forms:

$$o, (111); d, (101); \text{ and } n, (211).$$

The hexagonal appearance is due to the development about a trigonal axis, an axis, that is, normal to an octahedral face, and in order to bring out the peculiar symmetry of the distortion the drawings have been made with this trigonal axis vertical, as though the crystals were truly hexagonal.

The difference in habit is due to the fact that different faces of the above forms are developed on the two crystals. In the one, six faces of the trapezohedron equally inclined to the trigonal axis give the effect of a scalenohedron whose summit is modified by a rhombohedron composed of three faces of the dodecahedron, while three octahedron faces form a steeper rhombohedron of the same sign. In the other crystal the habit is prismatic; three only of the six trapezohedron faces parallel to the trigonal axis are developed, yielding a trigonal prism whose edges are beveled by the prism of second order, that is by six planes of the dodecahedron; the termination consists of positive and negative rhombohedrons consisting of dodecahedron and trapezohedron faces respectively, and a single face of the octahedron forming a basal pinacoid.

The two habits were represented about equally on the other crystals in the cavity as far as could be judged without their removal. It should be said that the faces present were sharp and clear and no trace of the missing faces of any form could be detected.

ART. XLI. — *New Species of Merycochærus in Montana.*
Part I. By EARL DOUGLASS.*Description of the Skull of Merycochærus laticeps n. sp.*

WHILE making a collection of vertebrate fossils from the Loup Fork beds of the Lower Madison Valley in Montana, I obtained several parts of mandibles, representing three or four different species, which were doubtfully referred to the genus *Merycochærus*. Three of the species were remarkable for the depth of the rami. One, which was more slender than the others, possessed characters which made it probable that it belonged to the same genus; though, measuring beneath the middle lobes of the last molars, the deepest jaw was double the depth of the narrower one. I could not identify any of these with species that had been described.

In the summer of 1899 in some clay bluffs near the village of New Chicago in Granite County, Montana, a nearly complete skull and lower jaw with some other bone fragments were found. The jaw was at once recognized as similar to one of those found in the Madison bluffs. The greater part of the skull was enclosed in a clay nodule. From some of the teeth which were exposed it was supposed to be one of the Oregon species of *Merycochærus*; but when cleaned from the matrix it was seen to be very different, possessing many peculiar characters which separate it from those forms. It is more nearly related to the type represented by *Merycochærus proprius* and *M. rusticus*, between which it is intermediate in size.

Its most striking characteristics are the following:

Skull low, broad behind the orbits, narrowing rapidly toward the front and back. Brain case short, the length behind the post-frontal process being about one-half the distance in front of it. Premaxillaries united in front forming a trough-shaped depression, evidently for the accommodation of a proboscis. Maxillaries deeply concave on the sides of the face—this, with the malo-maxillary ridge which widens outward rapidly toward the zygomatic arch, forming a broad nearly horizontal shelf above the posterior premolars and anterior molars. Larger part of external narial opening nearly between the orbits, but continuing forward in a horizontal slit between the maxillaries. Nasal bones short and ascending, placed far back, the anterior tips being about midway between the inion and incisive border. *Foramen infraorbitale* placed farther back than in any other species. Bending down of the face upon the basiscranial axis carried to the extreme, the shape of the posterior basal

part of the skull indicating that the head was carried in a nearly vertical or hanging down position. Back part of zygomatic arch small and simple. Length of the alveolar border twice the distance from the last molar to the occipital condyle. Otic bullæ not inflated. Mandible heavy and very deep in the region of the angle. First lower incisor absent. Jaw nearly as deep as skull exclusive of nasals, and nearly as long as the skull at base.

1



Merycochærus laticeps.

Skull and Mandible, side view, $\times \frac{1}{2}$.

Description of Skull.—The premaxillaries are coössified with the maxillaries. Anteriorly the maxillo-premaxillaries are united for a distance of 5.5^{cm}. Here they form a trough-like depression which becomes broader as it curves gently upward and backward. This widening of the trough is caused by the converging of the ridges from the upper borders of the maxillaries which bound laterally the hori-

zontal wedge-shaped forward projection of the narial opening, and by the conical shape—as seen from the front—of the large convexities in the region of the canines. These convexities have their bases close together in the middle of the incisive alveolar border, make a broad lateral sweep and die out on the face in front of the anterior projection of the malo-maxillary ridge. The first mentioned ridges die out on the anterior inner sides of these convexities a little distance above the alveolar border.

The anterior part of the narial opening begins anteriorly in a rather blunt point, widens gradually as it extends backward and slightly upward for a distance of 3^{cm}, then the sides are nearly parallel and horizontal about the same distance, then it expands and ascends to form the inferior part of the posterior oval opening. This posterior part of the external nares looks forward and slightly upward, thus forming a high angle with the anterior part. Above it is arched over by the short backward sloping nasals. Its lower border is in advance of the orbits, but in front of the bases of the nasals a notch extends back of the orbital border. The vertical height is about 5^{cm}. Between the narial opening and the orbit the tongue of the maxillary curves upward and then backward, apparently ending in a wedge-shaped process between the nasals and the frontals in a plane with the posterior parts of the orbits. The suture separating the frontal and maxillary extends from this point outward and forward and then downward, passing close to the anterior border of the orbit. The skull is a little injured just at the border of the orbit and the orbit is still filled with the matrix so that the lachrymal bone cannot be made out, but it occupies an extremely small space if any in front of the orbit. The malo-maxillary suture continues downward from the orbital border to a line 2·2^{cm} lower than the lower border of the orbit, where it curves forward and downward to near the lower anterior border of the overhanging malo-maxillary ridge. From this place it cannot be traced. The suture where seen is well defined, is complex but forms a quite regular band 3 to 4^{mm} broad. On the roof of the mouth the maxillo-palatine suture appears near the root of the second molar, extends forward and slightly inward to opposite the posterior lobe of *m1*, then transversely across the palate in nearly a straight line. The roof of the mouth is broad and concave. The incisive foramina are confluent, but this is apparently due to the breaking away of the thin median partition, part of the superior portion of which still remains. It now appears as an oval opening, the smaller end being directed backward. Its length is 3^{cm}, and its greatest width 1·5^{cm}.

On the side of the face is a large depression difficult to define. It is nearly triangular in shape. Above its boundary is the median upper border of the maxillary; behind it is the outward expansion of the posterior part of the maxillary in front of the orbit. Beneath is a broad shelf above the malo-maxillary ridge. The anterior angle of the depression is occupied by a deeper elliptical one, the deepest part of which is above the last premolar and the first molar. These two concavities extending inward on opposite sides of the maxillaries make the face quite thin transversely in this region, being in fact only 2.4^{cm} thick, while the skull at its widest place is 20^{cm}. The malo-maxillary ridge dies out on the anterior border of this depression but expands rapidly posteriorly, thus forming the broad shelf above mentioned which is broadest in front but extends outward and backward toward the zygomatic arch. The infra-orbital foramen is large, is near the inferior posterior border of the oval depression above *m*2 and looks forward and outward, opening on the horizontal maxillary platform or shelf.

The nasals are nearly triangular if I make out their posterior borders correctly. They are short, extending upward and forward to form the roof of the external nares opening. They are convex transversely and longitudinally and are pointed in front. These points are about midway between the incisor and theinion.

The orbits are oval with the larger end upward. The posterior inferior border is nearly straight or a trifle convex, but this may be due to a slight displacement of the post-orbital process of the malar. Only one orbit is preserved.

What I take to be the boundary between the nasals and frontals is a line where the bone is broken. It passes from the posterior angle of the maxillary transversely and somewhat backward to the suture between the nasals.

The frontal appears to meet the malar posterior to the orbit at the median line, where there is a roughening on the narrow isthmus of bone. From the upper posterior border of the orbit the supra-orbital ridges converge rapidly backward and then less rapidly, uniting to form a prominent narrow sagittal crest. The form of this part of the skull back of the nasals and including the upper part of the brain case is almost like that figured in Bettany's paper, "On the Genus *Merycochærus*," as *M. temporalis*; but the supra-orbital foramina occupy a different position. In the present species they are above a line uniting the posterior borders of the orbits, are far apart, have no grooves leading into them, and are a little nearer the median line of the skull than the outer border. The temporo-parietal suture extends upward and backward until, beneath the anterior part of the sagittal crest it curves downward and

then upward again, passing just beneath the foramina at the base of the inion. These foramina are not the same on opposite sides of the skull. They are farther forward and farther apart on the right side of the skull and are more uniform in size. On the left side the anterior one is circular and several times larger than the one which is a little behind and beneath it. Just posterior to this smaller foramen the parieto-occipital suture extends upward to the broken sagittal crest. The temporo-occipital extends backward and downward but is lost on the broken border. The ridge near the parieto-squamosal suture extends from the larger foramen above mentioned, downward and forward, at first coinciding with the suture and then running beneath and nearly parallel to it. The brain case on each side of this is broadly concave. The brain case is short antero-posteriorly and slopes downward quite rapidly behind the forehead, though the sagittal crest evidently continued nearly on a level.

The crest of the inion is broken but was considerably posterior to the occipital condyles. The wing-like expanses extending to the zygomatic arch were broad and thin. Just above the foramen magnum the occiput is broadly convex; but a short distance above a depression begins on the median line, becomes broader and deeper and then shallower to near the crest of the inion. The ridges that bound this concavity separate it from larger lateral concavities situated farther down. These concavities are bounded above by the wing-like expansions above mentioned and outwardly by the broad convex post-tympanics which appear as large swellings on the sides of the occiput. This post-tympanic with the horizontally expanded portion of the squamosal above and the long posterior flat surface of the post-glenoid encloses a quite large triangular cavity, the post-tympanic and post-glenoid nearly coming in contact below. The meatus auditorius externus evidently did not fill all this space, but the shape cannot be definitely made out on account of the matrix and injury of the bone; but it passes inward and forward a distance of 4^{cm} to connect with the otic bullæ.

As seen from below, the foramen magnum is lenticular in section and its transverse is double its antero-posterior diameter, being respectively 2.5^{cm} and 1.2^{cm}. With a line joining the incisive border, the lower extremities of the pterygoids and the occipital condyles, this opening forms an angle of about 65°. The occipital condyles are also narrow antero-posteriorly, the diameter being only 1.7^{cm} in this direction while the transverse diameter is 6.1^{cm}. The anterior articular faces are much deeper and broader than the posterior ones and they are partially

separated by an angle. The anterior faces are separated by an excavation 1.5^{cm} wide while the posterior ones are 3^{cm} apart.

The basi-occipital and basi-sphenoid ascend at a steep angle. This with the upper contour of the skull makes the brain cavity very small. Beneath the forehead the basi-sphenoid bends forward and becomes nearly parallel with it. The basi-sphenoid is angulate for a short distance between the glenoid surfaces.

The paroccipital processes are broken off at about a level with the lower points of the occipital condyles, but they appear to have been much longer. They are broad transversely but rather narrow antero-posteriorly. At the bases they are nearly crescent-shaped in cross-section. The more convex portion faces inward and backward, while the concave area faces forward and outward. The outer portion is a wing-like expansion of the more robust inner part, which slants forward and outward from the occipital condyles, forming with them an angle of 65°. The anterior inner horns or lobes are in a line with the post-glenoids. Just in front of these horns and closely in contact with them at the bases but with their anterior bases higher are the prismatic tympanic bullæ. These bullæ are not inflated but quite long vertically, especially on the anterior inner side, and do not extend much below the posterior bases of the paroccipitals. The shape of these bullæ is nearly that of a quarter of a cylinder terminated by a cone, the angle joining the nearly plane or slightly concave faces being directed outward and backward.

The post-glenoid processes are of moderate length and transverse breadth, are flat behind and moderately convex in front. The outer border slopes outward and upward, dying out on the inferior posterior surface of the zygomatic arch, which slopes upward and backward from the glenoid surface. This surface is slightly concave, being bounded exteriorly by the ridge on the lower outer border of the arch. The glenoid surface is broad transversely (6^{cm}) and uniformly convex antero-posteriorly.

The posterior angle of the zygomatic arch is on a line with the anterior borders of the paroccipital process, and extends upward to the line of the lower border of the orbit.

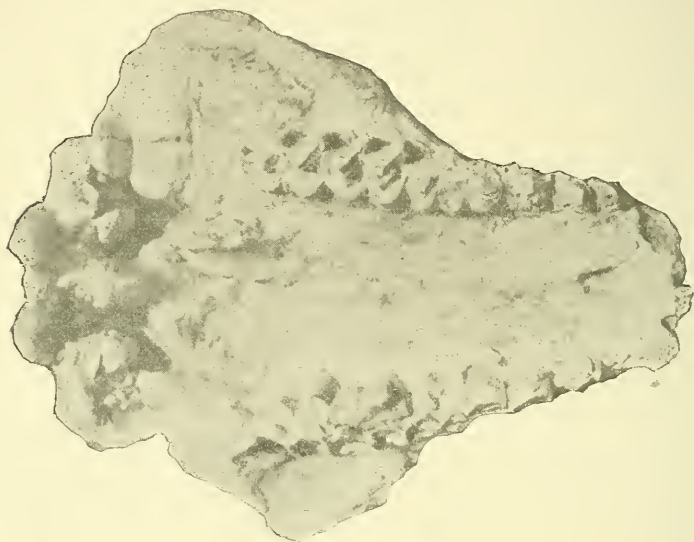
The malar has four approximately equal sides if the upper is measured along the border of the orbit; though the lower is longer in a straight line. The vertical width below the orbit and the antero-posterior length just below it are nearly the same. The upward extension in front of the orbit is narrower than the post-orbital process.

The palatines and pterygoids extend about 2.9^{cm} back of the last molars. It is only 2^{cm} from the posterior edges of the pterygoids to the anterior faces of the otic bullæ. This

approximation is due to the extreme shortening of the posterior basal elements of the skull so that the post-glenoid and par-occipital processes, the occipital condyles and otic bullæ occupy a comparatively narrow transverse zone.

Between the posterior lobes of the last molars the palatopterygoid lobes extend downward and backward until they reach a point about on a level with the incisive border and the lower extremities of the occipital condyles. They formed a median trough extending downward and backward from the palate.

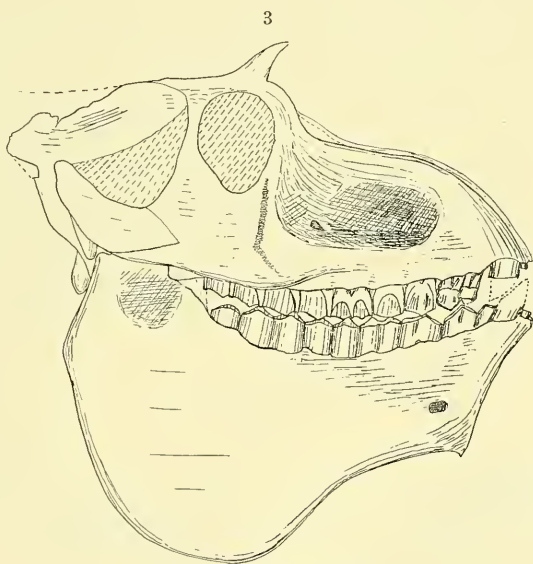
2



Merycochærus laticeps.
Skull, view of under side, nearly $\times \frac{2}{3}$.

Lower Jaw.—The rami were not coössified at the symphysis, and in the present specimen they are slightly spread apart. The anterior border is concave downward and nearly uniformly convex transversely. The *foramen mentale* is beneath *pm 3*, as is also the angle of the chin. At this angle a small process projects downward. Back of this the lower border of the ramus curves slightly upward and then downward to a point beneath the posterior part of *m 3*. At this point it is twice the depth of the shallower part under *m 1*. Back of this the posterior border ascends with a steep curve to a point 4.5^{cm} below the top of the condyle, where it ascends nearly vertically to the condyle, forming a low angle (about 25°) with the continuation

of the posterior border. The posterior border is slightly raised laterally. The condyle is not quite so broad transversely as the glenoid surface with which it articulated and is more narrowly convex antero-posteriorly. The masseteric fossa descends only a short distance—to a line connecting the base of the lobe of *m 3* and the posterior angle above mentioned. The bone is injured in the region of the coronoid process so its form cannot be made out. A broad convexity begins at the posterior part of *m 3*, extends forward and downward to the anterior lower border of the ramus, branches at the mental foramen and continues to the alveolar border at *pm 2*, thus leaving a depression under *pms 3 and 4* and *ms 1 and 2*. The length of the dental series is two-thirds the length of the jaw.



Merycochaerus laticeps.

Skull with jaw, $\times \frac{1}{4}$.

The mandible is too far ahead on account of slight forward displacement of post-glenoid process.

Dentition.—Superior: The crowns of the incisors and canines are not preserved. The upper incisors were small, quite close together and in a transverse row. The roots of the first and second are laterally compressed. They were nearly the same size, but the third was a little larger. In cross-section the roots of the canines are three-sided with rounded angles, the posterior side being the broadest: *Pm 1* is com-

pressed longitudinally, is inserted diagonally by two roots, and the posterior part overlaps *pm 2* outwardly. *Pm 2* is considerably worn, but there is still a shallow oblique enamel lake near the posterior margin. *Pm 3* has a longitudinal lake with two pits, near the inner border of the tooth. *Pm 4* is like the corresponding tooth in *Merychys elegans*. The posterior lobe on the last molar is not developed as in the *Merycochærus proprius*; though the posterior horn of the posterior external crescent is convex behind, it is narrower and extends much more outward than in that species. The premolars and molars are all longer than in *M. proprius*. This is due in part, but I think not wholly, to less amount of wear.

Inferior: There is no trace of a first incisor. The second and third are laterally compressed at the roots, and are 2^{mm} apart. They were about the size of upper incisors one and two. *I 3* is set obliquely close to the canine on its anterior inner side and a little more anteriorly than *I 2*. The canine is small, but a cross-section at the root is nearly the same in form as the upper canine. *Pm 1* is much larger than the lower canine. It has a longitudinal groove on the outside of the root. The crowns of the above teeth are not preserved. *Pm 2* is more oblique than in *M. proprius* or *M. rusticus*, being closely crowded between *pms 1 and 3*. *Pm 3* is proportionally narrower posteriorly than in the former species, judging by Leidy's figure. In *pm 4* the posterior fossa is more oblique, opening on the posterior inner angle instead of on the posterior border.

Measurements.

SKULL.

	M.
Length of skull from incisive border to posterior of occipital condyles245
Length from incisive border to back of pterygoids185
Length from incisive border to front of otic bullæ203
Width of skull at glenoid articulation200
Width at middle of last molars180
Width at third premolars080
Width at canines057
Width of palate between last molars063
Distance from anterior tips of nasals to incisive border, measured in a straight line166
From tips of nasals to crest of inion, about150 to .160
Distance from front of orbit to incisive border135
From front of orbit to back of occiput150
Diameter of orbit, antero-posterior038
Diameter of orbit, vertical045

	M.
Width of top of skull above orbits118
External nares, greatest width044
External nares, greatest length112
Length of three upper incisors015
Length of dental series from front of canine153
Length of molar-premolar series136
Length of premolar series056
Length of molar series080
Width of last molar030

LOWER JAW.

	M.
Length of jaw, greatest235
Depth at chin057
Depth under m 1053
Depth under back part of m 2063
Depth under last lobe of m 3109
Length of inferior dental series157
Length of molar-premolar series150
Length of premolar series058
Length of molar series092
Space occupied by i 2 and 3008
Width of canine at alveolus, transverse009
Length of canine antero-posterior0065
Width of pm 1011
Length of pm 1 antero-posterior014
Length of pm 2014
Width of pm 2006
Length of pm 3017
Width of pm 30095
Length of pm 40185
Width of pm 4012
Length of m 3044
Width of m 30165
Length of foramen mentale008
Width of foramen mentale005

From the parts of *Merycochærus proprius* that have been described *M. laticeps* differs in the lateral aspect of the maxillaries; the smaller size and, apparently, the more transverse position of the incisors and incisive border; the absence of a space between *pms* 1 and 2; the more hypsodont character of the teeth; the more posterior position of the infra-orbital foramen; the smaller size and different shape of the small posterior lobe on the last upper molar; and the different shape of the mandible, especially its greater posterior depth.

M. rusticus approaches more nearly to the present species in the facial concavities, in the size and position of the incisors,

and the trough-like shape of the maxillo-premaxillaries above; the crowding together of *pms 1* and *2*; and possibly in the posterior deepening of the mandible. But, as will be seen by comparing the present descriptions with Leidy's figures and descriptions, the face and anterior border are quite different in the two species. In *M. rusticus* the anterior part of the face as viewed from the side rises more abruptly for a short distance and then more gradually backward. The form of the concavity on the side of the face is different; the infra-orbital foramen and the anterior inferior root of the zygomatic arch are farther forward; the latter also is higher as is also the lower border of the malar. The space between the upper canine and first pre-molar is greater and *pm 1* does not overlap *pm 2*. The mental foramen is longer and the chin is less concave, the protuberance at the chin is larger and longer, and the anterior inferior border of the jaw is more nearly straight. *M. rusticus* is smaller than *M. laticeps*.

ART. XLII.—On *Mohawkite, Stibio-domeykite, Domeykite, Algodonite and some artificial copper-arsenides*; by GEORGE A. KOENIG.

THE knowledge of the existence of copper arsenides in the Keweenaw copper formation is nearly as old as the mining operations in this region themselves. In the matter of occurrence there are two points to be noted: (1) The arsenides are not found in the bedded deposits of native copper, but always in fissures, intersecting the beds. (2) These veins have thus far only been observed in the lower beds, near the foot of the formation to the southeast. Arsenic, however, is found in the smelted and refined copper of all the mines. This element is a minimal amount in the copper from the Great Conglomerate of the Calumet and Hecla mine and becomes a maximum in the copper from the amygdaloid beds of the lower measures, on which the Mohawk, Wolverine, Arcadian, Sheldon-Columbia, Isle-Royale, Atlantic, Baltic, Champion, etc., are located. The Sheldon-Columbia location near the shore of Portage Lake in the village of Houghton was the first mine which furnished copper arsenides, notably domeykite. The larger part of the specimens in collections come probably from this mine. Whitneyite was found in a quartz vein in dark melaphyr, not far from the present Mohawk, but the existence of any copper rock was not suspected at that time at that point. Several masses of domeykite have been found in the drift on the Hancock shore of Portage Lake, very much decomposed, cuprite and arsenates being the chief products. In the spring of 1898 the opening of a new street in East Houghton on the old Sheldon-Columbia location, disclosed a quartz vein containing some foliated domeykite. Many good specimens were gathered and have come into collections. A similar vein has been known for years on the old Huron location, now the southern end of the Isle-Royale property. Algodonite was found in the Pewabic mine, located on the Quincy amygdaloid bed and which lies about 5000 feet higher than the Isle-Royale amygdaloid. In developing the Mohawk property a cross vein at right angles to the copper-bearing amygdaloid was met in December, 1899. This cross vein was from 12 to 15 inches wide when struck, but has since shown very varying dimensions. In a gangue of quartz and calcite the copper arsenides have been found in this vein more abundantly than in any of the locations mentioned above. Early in January Mr. Fred. Smith, Superintendent of the Mohawk mine, sent me a solid piece weighing from 4 to 5 pounds, with the request to make

a thorough investigation and report. On January 18th I reported to Mr. Smith that the mineral substance which he had sent me was a new mineral species which I would call *Mohawkite*. Later on I received other material from Mr. Smith in which I identified an antimonial domeykite for which I propose the name *Stibio-domeykite*, and also some very peculiar intimate mixtures of Mohawkite with Whitneyite. It appears that the Stibio-domeykite is the more prevalent of the arsenides, and not the Mohawkite as was thought at the start. The vein has since been traced to the outcrop where it was found to form a ridge owing to the quartz in the gangue.

1. *Mohawkite.*

Physical properties.—Form is massive; no crystallization of any sort has been observed. The structure is mostly fine granular, sometimes compact. The color on fresh surface is gray with a faint tinge of yellow. Tarnishes very easily, and the tarnish is apt to be ultimately dull purple. A yellow brassy tarnish is brought about by boiling, though in cold water the original color lasts for two days and more. The color is, however, no sure guide for identification. The mineral is very brittle, and owing to the granular structure is not possible to fix the degree of hardness; it is approximately 3.5. Spec. gr. at 21° C. (in boiled water but without corrections) = 8.07, mean of three closely agreeing trials and with perfect material, of which there was plenty,—4.6294 grams were taken.

Chemical properties.—In closed tube gives no sublimate of arsenic; only a slight sublimate of As^2O^3 , owing to the air in the tube. The substance melts in the tube at cherry heat, and colors the glass blue if the heating be kept up for a few minutes (Cobalt). In the open tube the reactions are similar but more pronounced owing to the vigorous oxidation. On charcoal in O. F. copious vapors of As^2O^3 , odor of arsenic and ultimately a globule of metallic copper. If a fragment of the mineral be placed in a shallow cavity on charcoal, along side of a borax bead of equal size or somewhat larger, and both fused together in the point of the blue flame, so that the metallic globule be exposed to the air, then the borax bead will assume the pure blue color of cobalt. If this treatment be kept up for sometime and a fresh borax bead be taken every minute, then a brown nickel bead will be obtained and finally a blue or red copper bead; thus proving all the metals present except the trace of iron. This test (Plattner's) should be used always in examining metallic arsenides. Thus the mineral contains copper, nickel, cobalt, iron (trace) and arsenic. Boil-

ing concentr. nitric acid dissolves the mineral without leaving a residue, forming first a green, then a murky blue solution (owing to the interference between nickel-nitrate green and cobalt-nitrate red). HCl does not act upon the mineral. The quantitative analysis was made by several methods:

(a) Nitric solution made ammoniacal, alcohol added and magnesia mixture. This is not a good method, because part of the nickel and cobalt enter into the magnesium-ammonium arsenate.

(b) Nitric solution made ammoniacal, diluted properly and H²S passed into hot solution to saturation. Filtrate evaporated to dryness, residue oxidized with HNO³ and magnesium-ammonium arsenate precipitate. Separation of copper from nickel and cobalt by H²S; separation of nickel and cobalt by nitrite method.

(c) Powder fused with nitrate and carbonate of sodium. This gives very good results for arsenic; but for the fine pulverulent condition of the oxides these latter must be carefully handled to avoid loss.

I find that the magnes. amm. arsenate can be heated on the asbestos pad of a Gooch crucible without loss into pyro-arsenate. From four analyses of perfect material I obtain

Cu.....	61.67
Ni	7.03
CoO	2.20
Fe.....	trace
As.....	28.85
	99.75

hence the atomic ratio

Cu....	0.9803		
Ni	0.1200		
Co....	0.0373		
	1.1376	2.958	(Cu, Ni, Co) ³
As.....	0.3847	1.000	As

This is an exact ratio of 3:1, the ratio of domeykite. I look upon this as a case of isomorphic replacement. We may expect to find all gradations of replacement within the ratio of 3:1. Synthetic experiments lead me to consider this ratio as representing an especially strong one, as of bonds well satisfied; because I have been enabled to obtain it artificially in well defined crystals, of which more hereafter. The molecule Cu²As is the strongest of all, the other ratios seem to partake, more of the nature of alloys, of unsatisfied bonds. It may be urged that a mechanical mixture of the copper arsenide with

nickel-cobalt arsenide is here presented. Against such a view speaks the physical condition of the substance in the first place and also that a ratio of $(\text{Ni, Co})^3\text{As}$ has not been observed so far. The question presented itself at once as to whether any of the domeykite occurrences in Houghton Co. contain nickel and cobalt. Winkler gives this metal in the Zwickau occurrence at 0.44 per cent. F. A. Genth does not mention it, and even from the Michipicoten Island location where niccolite accompanies it, the nickel is absent in the domeykite. I have examined every accessible specimen qualitatively. Neither nickel nor cobalt could be found by Plattner's test; but nevertheless it is very probable that all domeykite contains both metals in traces, as my careful analysis of the mineral from the Sheldon-Columbia location, both old and recent, demonstrates. From 0.5 gram of substance just sufficient cobalt oxide was obtained to give color to 50 mg. of borax glass and the nickel was proved by reducing the bead with tin on charcoal. Another matter engaged my attention at once, namely the discrepancy in the specific gravities given by the several authors and recorded on page 44, Dana Min., 6th edit. These densities vary between 6.70 and 7.547 against my determination of 8.07 for the Mohawkite. The collection of the Michigan College of Mines contains one fine specimen of domeykite from the Sheldon-Columbia mine. The material is very uniform, a few specks of calcite the only visible gangue. Of this 3.8559 grams were selected. The specific gravity at 21° C. was found = 7.9486. The whole material was then dissolved, no residue was left and in a part of the liquid calcium was looked for, but only a trace was found. Hence we have every reason to take this specific gravity as that belonging to domeykite; the figures in the literature must be wrong. It is true that the analysis of this substance does not agree with the theoretical composition exactly, namely:

Cu	74.00	}	Spec. gr. found	7.9486
(Fe, Ni, Co)	0.06		" calculated	8.1020
As	26.14			
	100.20			

Now if we assume that the elements in this combination possess the same specific gravity as they do in the free state, we may calculate the specific gravity. If specific gravity of Cu = 8.96 and of As = 5.63 then the above composition must weigh $8.96 \times 74.00 + 5.63 \times 26.14 = 8.1020$. The specific gravity of the theoretical Cu^3As will be = 8.017. For Mohawkite the calculated specific gravity will be, taking Co = 8.96 and Ni = 9.5 (the highest number on record):

Cu =	61.67 × 8.96 =	5.5256
Ni =	7.03 × 9.5 =	0.6678
Co =	2.02 × 8.96 =	0.1809
As =	23.85 × 5.63 =	1.6242
		7.9985
		8.0650 found
		+0.0665 difference

This is quite a satisfactory agreement. Yet it occurred to me at this stage of the investigation to verify the relations between specific gravity and composition by artificial compounds and also to see in how far domeykite might be obtained in crystals.

A combustion tube was closed at one end over the lamp. Resublimed arsenic, roughly powdered, was placed into the bottom to the amount of 7.9 grams. On top of this was poured 20.0 grs. of copper filings. The latter had been the end product in a determination of oxygen in refined copper. Total non-copper in this material 0.07 per cent, consisting of arsenic and iron. An asbestos plug was put over the copper and the tube placed in horizontal position into an Erlenmeyer combustion furnace. The copper was first heated to a dull cherry redness and then the arsenic was heated to the sublimation temperature. The heating was kept up for two hours, the open end of the tube being provided with a mercury valve to prevent any air currents from entering the tube. The copper absorbed the arsenic with avidity. One could see that only the part nearest the vapors changed in color and 30 minutes from the beginning this front portion began to melt off. After all the arsenic had become volatilized and a considerable part of the liquid compound had been formed the temperature of this liquid was raised to bright redness and kept so until the end of the experiment; it being 6 o'clock P. M. The tube was inclined so that the liquid covered the solid portion and left to cool slowly. In the morning three distinct substances were found. A dark colored fused mass, then a porous gray part and then the apparently unaltered copper filings, merely slightly caked together. (a) The dark, fused portion. On the fracture blue-gray color; strongly developed radial structure and, in fact, so closely resembling chalcocite that one could not tell by the eye one from the other. The weight of all three parts was 27.27 grams. The weight of the unaltered (apparently) part was 9.65 grams; hence in the altered part there must be approximately 10.35 of copper and 7.2 grams of arsenic, giving an atomic ratio of 9.6:16.4. Now if the quantities taken of copper and arsenic were in the ratio of 3:1 it is

evident that at red heat the two elements form the compound Cu^2As , leaving much copper unaffected. The specific gravity of this fused arsenide was found at 7.71, (3.8837 grams being taken). The analysis gave

		Calculated for Cu^2As .
Cu.....	63.30	Cu..... 62.69
As.....	37.00	As..... 37.31
	100.30	100.00

The calculated specific gravity is = 7.754, giving a minus difference of 0.044 against the actual specific gravity.

(b) The porous zone between *a* and the unaltered copper. An ordinary lens showed that this mass consisted of groups of minute crystals, exhibiting very brilliant faces. A power of 50 diameters differentiated the crystals easily. They appear to be combinations similar to those of arsenopyrite. Believing that I had found the method by which these crystals could be prepared readily I destroyed them before making an attempt at measurements, leaving only one group, and this one was lost clumsily. There are still crystals on the copper cylinder showing the forms and building up of the grape-like bunches but I do not consider them worth the trouble connected with micro-goniometric work. 253 milligrams of these groups, which separate easily from the base, were taken for analysis and only the copper was determined, as there could not be anything but it and arsenic in the material. Found

$\text{Cu} = 71.39$; theoretical for $\text{Cu}^3\text{As} = 71.6 \text{ Cu}$

No doubt can exist as to the identity of these crystals with domeykite. I have made several attempts since to get the crystals but evidently did not hit the right temperature again. I obtained domeykite as crystalline mass. In one experiment I melted together 28.886 grams of the Cu^2As with 7.8 grams of copper in a glass tube. Liquidity maintained for one hour. After cooling the column was found of varying composition from bottom up. The bottom portion showed a fracture similar in color and structure to white iron, with a faint yellow, in fact it looks exactly like the domeykite of Houghton Co. Specific gravity = 8.05; percentage of copper = 75.4, considerably above the ratio Cu^3As ; calculated specific gravity = 8.085.

Above this was found a laminated or scaly zone, resembling some of the recently found domeykite from Sheldon location. It gives copper = 71.14, very close to Cu^3As .

The top portion resembles in color and structure (very fine grain) the Cu^2As . It gives copper = 64.8.

2. *Stibio-domeykite.*

With this name I propose to designate the domeykite from the Mohawk mine, Keweenaw Co., and any other locality where mineral of similar composition will be found. In physical properties it is undistinguishable from the domeykite of the Houghton Co. mines. It is remarkable for the massiveness of its occurrence, the freedom from admixtures. It has a decided subconchoidal fracture and takes on a brass tarnish very soon, which, however, ultimately turns into bluish purple. It is very brittle but not as friable as the Mohawkite. Its hardness is very near 4, a little below.

Specific gravity at 21° C. = 7.902 (made with 4.5595 grams) B. B. In the open tube after heating the molten globule for five minutes one observes the forming of yellow spots near the globule and a slight ring or patches farther on, which turn yellow on applying higher heat and volatilize in part. Farther on there is a copious sublimate of arsenic trioxide. If now one washes out the tube in a jet of water and passes H²S into the tube one observes the above ring and patches turn red, whilst the arsenic oxide turns gold-yellow. I was astonished myself to find that one could demonstrate thus the presence of 0.1 per cent antimony in the mineral. On charcoal blow the point of the oxidizing flame upon the melting mineral for two minutes, then drive off the white arsenic with a gentle flame, the antimony trioxide remains. The open tube reaction is preferable and certain. Concentrated nitric acid does not dissolve the mineral completely; a white cloudiness or a white sediment will be left. When fused on charcoal alongside of a borax bead, the latter colors greenish after several minutes' action (cobalt and nickel cannot be thus found, though present in small amount).

The analysis gave (type specimen) :

Cu.....	72.48
(Fe, Ni, Co).....	0.24
As.....	26.45
Sb.....	0.78
	99.95

This is evidently a typical domeykite, and one would expect likewise that the percentage of antimony is variable. In order to ascertain this I took twelve pieces, all over one pound each, and treated one gram of each with 10^{cc} of concentrated nitric acid, boiling until the color was blue, then added 15^{cc} of water to each and stood the beaker glasses in a row. (The type specimen with its 0.78 antimony was among the lot.) It appeared that all contained antimony, but no two

an equal amount. Some showed more, some less than the type specimen. I selected the one with the heaviest sediment and determined the antimony as trisulphide. It gave $Sb = 1.29$. This must be taken as the maximum until other tests show greater amounts. The Houghton domeykites show no trace of antimony.

3. Mohawk Whitneyite.

With this name I shall designate not a species, nor even a variety but simply a most intimate mixture of the two species Whitneyite and Mohawkite, of which considerable masses are at present encountered in the Mohawk mine. This material is distinguishable at once to the layman even. It is very tough and approaches Whitneyite in that respect. Ordinary blows with the hammer fail to break the masses, the chisel is often required and the hammer leaves a dent as in soft metal. The fresh fracture is gray, fine granular, even hackly. These fractures assume a dull brown or olive-green tarnish, much like algodonite and Whitneyite. There are, of course, all gradations, from nearly pure Mohawkite to nearly pure Whitneyite. The material looks homogenous, but is not, as the following analyses show. The samples were broken with the chisel in close proximity to one another.

No. 1.	No. 2.	No. 3.	No. 4.
Cu ---- 85.36	} 83.61	Cu ---- 79.36	Cu ----- 84.86
(NiCo) 0.32		Co ---- 0.82	(Ni + Co) .. 0.64
(by diff.) As.. 13.59		Ni ---- 0.61	As ----- 13.06
CaCO ³ 0.73		Fe ---- 0.36	(CaMg)Co ³ 0.71
100.00		As ---- 15.07	99.27
		CaCO ³ 2.41	
		MgCO ³ 0.60	
		99.23	

A notable fact is that in No. 3 the cobalt exceeds the nickel. Calcite is the gangue of this material. All manner of atomic ratios can be calculated from these analyses.

No. 1	has the ratio	Cu ^{7.5} As
No. 2	"	Cu ^{6.7} As
No. 3	"	Cu ^{6.45} As
No. 4	"	Cu ^{7.7} : As

In connection with these results, attention may be called to a communication made by Mr. J. Stanton, Secretary of the Mohawk Mg. Co., to the Engineering and Mining Journal of April 7, 1900. Mr. Stanton here gives Dr. Ledoux's report upon some of the Mohawk mineral. Dr. Ledoux figures from

his analysis the ratio Cu_4As (including $\text{Ni}+\text{Co}$) and says that with this formula Mohawkite will probably be accepted by science. If this ratio does exist then it is *not* Mohawkite; for this name was given by me to Cu^3As (including $\text{Ni}+\text{Co}$). But the following sentence regarding the material induces me to believe that it was a mixture of Mohawkite and Whitneyite. The report says: "The sample was pulverized and concentrated by washing until under a powerful glass no other minerals were present. Thus purified the mineral was analyzed with the following result: $\text{Cu} = 68.6$; $\text{As} = 22.67$; $\text{Ni} = 6.55$; $\text{Co} = 1.2$; $\text{Fe} = 0.23$; $\text{S} = 0.53$. The sulphur and iron are impurities [why? if the powerful glass showed no other minerals present? Koenig] and it is also reasonably probable that the nickel and cobalt are mechanically mixed in the sample, although it is possible that some of the copper has been replaced by nickel and cobalt." How can Dr. Ledoux calculate a formula at all if he holds such views as here expressed? He gives the specific gravity at 7.8, but as I have shown above, a composition such as he finds must have a higher specific gravity even than I find for the Mohawkite proper, namely 8.140, leaving out both iron and sulphur. The atomic quotients of Dr. Ledoux's percentages are:

Cu.....	1.089	As.....	0.302
Ni.....	0.111		
Co.....	0.020	"	
	1.220		

giving the ratio of 4 : 1, but if nickel and cobalt are thrown out the ratio is nearer 3 : 1 than 4 : 1.

Under the existing conditions I permit myself to doubt the existence of a molecule $(\text{Cu}, \text{Ni}, \text{Co})^4\text{As}$ and merely present the facts to the mineralogical public.

4. *Algodonite.*

This species has not yet been observed at the Mohawk mine, though some of the Mohawk-Whitneyite resembles it very much. I was induced to bring it within the present investigation merely by the specific gravity, which is given by Dr. F. A. Genth as the only authority at 7.62. To Dr. Lucius Hubbard I am indebted for the material. This gentleman possesses a handsome specimen from the find at the Pewabic mine, many years ago. The specimen looks like the segment of a nodular piece. It showed the chocolate-brown dull tarnish. But the fresh fracture is beautiful. It has a color and texture exactly

like razor-steel. Although very tough, it does not dent like Whitneyite. The nodule was covered with Whitneyite and this mineral shows on Dr. Hubbard's specimen like a thin fringe around the algodonite. I could break off some 5 grams of faultless material. The specific gravity was found at 21° C. = 8.383 (using 3.8418 grams). The analysis made with 0.5 gram. Of silver I could find no trace. The analysis gave

Cu	83.72	for Cu ⁶ As
(Fe, Ni, Co)	0.08	Cu..... 83.5
As	16.08	As..... 16.5
	99.88	

Calculated from these figures, the specific gravity is 8.406. This gives a minus difference of 0.023 from my experimental value of 8.383.

Michigan College of Mines, June, 1900.

ART. XLIII.—*Heat of Solution of Resorcinol in Ethyl Alcohol*; by C. L. SPEYERS and C. R. ROSELL.

WHEN a system undergoes a change depending upon an absorption of some particular energy, then as the intensity of that particular energy is increased on the outside of the changing system, the amount of change inside the system is increased; when the system undergoes a change depending upon a rejection of some particular energy, then as the intensity of that particular energy is increased on the outside of the changing system, the amount of change inside the system is diminished. This statement is to be considered as universally true though now and then a seeming exception comes up.

The latest seeming exception that we have come across is resorcinol with ethyl alcohol. Resorcinol dissolves in a large excess of ethyl alcohol with rejection of heat* and yet the solubility of resorcinol increases with the temperature. That is, a change depending upon a rejection of heat energy seems to be favored instead of being hindered, by increasing the intensity of the heat energy on the outside of the system.

But the mistake lies in joining a heat of solution in a large excess of solvent, giving a very dilute solution, to a solubility which implies a saturated solution and one more or less concentrated. The two solutions must be of equal concentration. Wherefore the heat of solution in a nearly saturated solution is the quantity that must be used, and its sign alone is enough, its value is not needed, to test the validity of the above statement.

The sign can be found by diluting a saturated solution with a small quantity of pure solvent. The method by supersaturation fails for resorcinol, does not crystallize fast enough from a saturated solution even to get the sign; at least it did not do so in the calorimeter we used.

About 200^{cc} of saturated solution at about 27° were diluted with about 10^{cc} of pure solvent. The temperature rose more than 0.2°.

The act of solution may be separated into two parts; namely, one in which the solute liquefies, which always involves a gain in energy to the system and is +, the other in which the liquefied solute mixes with the rest of the liquid to form the desired solution. The second part may be + or -; when it is + then obviously the whole heat of solution must be +. And this is the case for resorcinol, because concentrating a

* Speyers, Journ. Am. Chem. Soc., xviii, 146, 1896.

nearly saturated solution absorbs heat, and the heat of fusion is +.

Consequently the above statement does express the behavior of resorcinol with ethyl alcohol so far as heat of solution and solubility are concerned.

Since heat is rejected when resorcinol dissolves in a large excess of ethyl alcohol and since heat is absorbed when it dissolves in a small quantity, the temperature should not change when these substances are mixed in some certain proportion. This proportion was found to be about 6 grams of resorcinol with about 100 grams of ethyl alcohol.

The solubilities in all the following 37 cases increase with the temperature and so the heat of solution of all the solids in a nearly saturated solution should be +. This demands that the heat of fusion should be larger than the heat of mixing the liquid solute and solvent whenever the sign of this latter heat is -.

The signs refer to the heat of dilution of a saturated solution with about 5 per cent or 10 per cent of pure solvent.

	H ₂ O	CH ₃ OH	C ₂ H ₅ OH	C ₃ H ₇ OH	CHCl ₃	C ₇ H ₈
Urea	+	----	±	----	----	----
Urethane	+	+	+	+	-	+
Chloral hydrate...	-	----	-	----	+	+
Succinimid	+	----	±	----	----	----
Acetamid	-	----	+	----	----	----
Mannitol	±	----	----	----	----	----
Resorcinol	+	----	-	----	----	----
Benzamid	----	----	±	----	----	----
p-Toluidin	----	----	+	----	----	----
Acetanilid	----	-	±	----	+	----
Acenaphthene	----	±	±	±	-	-
Naphthalene	----	-	±	-	-	-
Phenanthrene	----	----	±	----	----	-
Cane sugar	-	----	----	----	----	----

Rutgers College, Sept. 1, 1900.

ART. XLIV.—*The Sulphocyanides of Copper and Silver in Gravimetric Analysis*; by R. G. VAN NAME.

[Contributions from the Kent Chemical Laboratory of Yale University—XCIX.]

Cuprous Sulphocyanide.

As early as 1854 attention was drawn by Rivot* to the possibility of estimating copper gravimetrically by weighing as cuprous sulphocyanide, and to the advantages which the process afforded in separating copper from other metals. Rivot's method of procedure consisted in dissolving the substance to be analyzed in hydrochloric acid, reducing the copper with hypophosphorous or sulphurous acid, and precipitating with potassium sulphocyanide. The precipitate dried at a moderate temperature was weighed as cuprous sulphocyanide and then as a control converted by ignition with sulphur into cuprous sulphide and weighed in that condition.

In his well known work upon quantitative analysis Fresenius in one place† denies the practicability of the direct weighing of copper as cuprous sulphocyanide on account of the tendency of the latter to hold water even when heated to the temperature of incipient decomposition. As authority for this statement he cites Claus,‡ who found 3 per cent of water in the precipitate after drying at 115°, and Meitzendorff, who gave the percentage of water under the same conditions as 1.54.

On a later page of the same volume,§ however, Fresenius, after a trial of the process which gave 99.66 per cent of the theory for copper, concludes that the method is practicable although apt to give low results, particularly in the presence of free acid.

The process was again recommended in 1878 by Busse,|| who had employed it for the estimation of copper, both alone and in the presence of iron, nickel, zinc and arsenic, obtaining results very near the theory and plainly comparable with the figures obtained by afterwards igniting the cuprous sulphocyanide with sulphur in hydrogen.

In spite of the evident advantages for certain purposes of Rivot's method over other modes of determining copper, it has never come into general use. The chief reason for this has apparently been the difficulty and inaccuracy attendant upon the weighing of the precipitate upon dried paper filters, a process which can hardly be depended upon unless managed with extreme care.

* Comptes Rendus, xxxviii, 868.

† L. Gmelin Handb., iv, 472.

‡ Zeitschr. Anal. Chem., xvii, 53.

§ Fresenius, 6th Aufl., i, 187.

|| 6th Aufl., i, 335.

In the experiments to be described this difficulty was avoided by performing the filtering and weighing upon asbestos in a perforated platinum crucible. The method of conducting a determination was as follows: A suitable quantity of a standard copper sulphate solution was run from a burette, diluted to a convenient volume, a few cubic centimeters of a saturated solution of ammonium bisulphite added, and the copper precipitated by an excess of ammonium sulphocyanide. The precipitate was allowed to settle, collected upon asbestos in a weighed crucible, washed with cold water and dried at 110° until no further loss of weight took place.

In Table I are given the results of a number of determinations made in this way. The copper sulphate solution was made up exactly decinormal and the standard confirmed electrolytically. As the ammonium sulphocyanide solution was slightly above decinormal, 13^{cm³} represent a small excess (about one cubic centimeter) above the amount theoretically required to precipitate 25^{cm³} of the copper sulphate solution. The ammonium bisulphite, which had been recently prepared by saturating aqueous ammonia with sulphur dioxide, was always used in sufficient quantity to give the liquid a strong and permanent odor of the latter.

TABLE I.

25^{cm³} of N/10 CuSO₄ solution, equivalent to 0.0795 grm. Cu, taken for each experiment.

	H ₂ SO ₄ concentrated. cm ³ .	HNH ₄ SO ₃ sat. sol. cm ³ .	NH ₄ SCN approx. N/10. cm ³ .	Final volume. cm ³ .	Time of standing. hours.	Cu found. grm.	Error grm.
1.	none	5	13	68	$\frac{1}{4}$	0.0795	0.0000
2.	"	3	13	66	48	0.0793	-0.0002
3.	"	3	25	78	$\frac{1}{2}$	0.0796	+0.0001
4.	"	3	25	78	12	0.0796	+0.0001
5.	1.5	10	13	85	12	0.0792	-0.0003
6.	1.5	8	13	105	48	0.0785	-0.0010
7.	1.5	3	25	85	4	0.0783	-0.0012
8.	1.5	5	25	85	21	0.0795	0.0000
9.	5	5	25	85	3	0.0797	+0.0002
10.	15	10	25	115	21	0.0793	-0.0002
	HCl concentrated cm ³ .						
11.	10	5	25	100	20	0.0795	0.0000
12.	25	10	25	100	28	0.0784	-0.0011

When there is no free acid present the time of standing before filtration and the amount of the excess of ammonium sulphocyanide are practically without effect, as experiments 1 to 4 of the table show.

Experiments 5 to 10 were carried out in the presence of various amounts of free sulphuric acid up to 12 per cent of the total volume of liquid. The acid, at least within this limit, does not exert a sufficient solvent effect upon the cuprous sulphocyanide to interfere materially with the accuracy of the process, but it retards the precipitation, making it necessary to increase the time of standing before filtering in proportion to the amount of acid present. In several of these determinations the precipitation was visibly incomplete even after several hours standing. This effect of the acid, however, hardly shows in the results of the table because the standing was always prolonged until the copper appeared to be all down before filtering.

The low results of No. 7 was probably due chiefly to incomplete precipitation, although No. 9 shows that even with a much larger amount of acid precipitation may be complete within three hours. In general, however, it is safer to allow ample time (twelve hours or more) for the precipitation when there is much free acid present.

Comparison of Nos. 5 and 6, for which only a bare excess of ammonium sulphocyanide was used, with Nos. 7 to 12 shows an apparent advantage in the larger excess in the presence of acid. Hydrochloric acid, judging from the results of No. 11 and 12, has no greater disturbing influence than sulphuric acid, although in No. 12, where the concentrated acid constituted one-fourth of the entire volume, there was apparently a slight solvent action. The filtrate from this determination when concentrated to about 25^{cm} and treated with potassium ferrocyanide gave a strong test for copper, as did also the filtrate from No. 6. Several of the other filtrates were tested in the same way, but none showed more than an insignificant trace of copper. No. 7, however, was not tested.

Table II contains the results of a series of experiments conducted as before, except that larger amounts of copper were

TABLE II.

	Cu taken. grm.	H ₂ SO ₄ conc. cm ³ .	NH ₄ SCN approx. N/10. cm ³ .	Final volume. cm ³ .	Cu ₂ S ₂ (CN) ₂ found calc. as Cu. grm.	Error. grm.	Cu in filtrate.
1.	·3175	none	60	500	·3176	+0·0001	none
2.	·3175	“	60	500	·3177	+0·0002	“
3.	·3175	“	60	500	·3176	+0·0001	“
4.	·3175	10 HCl conc. cm ³ .	100	500	·3175	0·0000	“
5.	·3175	20	100	500	·3165	-0·0010	distinct

employed. The copper sulphate solution was approximately N/5 and standardized by the battery. The solution of ammonium sulphocyanide was the same previously used and a considerable excess was employed in every determination. More than twice the amount theoretically required was used in every case where free acid was present, and at least twenty hours were allowed for the precipitation, which was made in cold, and as the table shows, rather dilute solutions. If the solution is too concentrated the copper is apt to be thrown down in a finely divided condition, making it hard to filter.

The time required to dry the cuprous sulphocyanide at 110° is in general from two to three hours. Heating much longer than this is not to be recommended, as a gradual increase in weight begins to take place, as is shown by the following example, which gives a series of weights of the same precipitate at different stages.

	Cu ₂ S ₂ (CN) ₂ gram.	Calculated as Cu. gram.
After 2 hours at 110° -----	·6060	·3167
“ 4 “ “ “ -----	·6059	·3167
“ 19 “ “ “ -----	·6067	·3171
“ 23 “ “ “ -----	·6069	·3172

This tendency to increase in weight is, however, usually less marked than in the above example, and in any case need not interfere materially with the accuracy of the process unless the drying is prolonged far beyond the necessary length of time.

The method is easily handled and as the results of Tables I and II show is capable of considerable accuracy. From the nature of the process it is evident that it is much less likely to be interfered with by the presence of other metals than the other gravimetric methods for copper, and may therefore be directly applied with good results in many cases where the use of the electrolytic or the oxide method would involve a previous separation.

Silver Sulphocyanide.

The sulphocyanide of silver, unlike that of copper, is soluble in an excess of ammonium or alkali sulphocyanides and this fact prevents the use of the latter to precipitate silver for gravimetric estimation. The reverse process, however, the precipitation of a soluble sulphocyanide by an excess of silver nitrate, as will be shown by the experiments to be described, furnishes a convenient means of standardizing sulphocyanide solutions and in general for estimating thiocyanic acid.

When freshly precipitated the sulphocyanide of silver resembles the chloride in appearance, but when allowed to

stand a few hours becomes finely granular and is very easily filtered and washed. It may be safely dried to a constant weight upon an asbestos filter at 110°–120° but at a somewhat higher temperature is decomposed, leaving a residue of silver sulphide.

The determinations which are tabulated below were made as follows. Portions of 25^{cm}³ of an approximately decinormal solution of ammonium sulphocyanide were measured from a burette, diluted with 100^{cm}³ of water and silver nitrate added in excess. The precipitate was collected upon asbestos in a platinum crucible, washed with cold water and dried to a constant weight at 115° the drying requiring usually between two and three hours.

The filtering is facilitated by allowing a few hours for the precipitate to settle; but this is by no means essential, as it is easy with a little care to obtain a clear filtrate even when the filtering is performed at once.

The solution of ammonium sulphocyanide was prepared from a pure salt, especially tested and found free from chloride. This point is of importance, as chlorine is a common impurity and its presence in any considerable quantity will vitiate the results.

TABLE III.

Final volume of liquid 150 ^{cm} ³				
25 ^{cm} ³ of NH ₄ SCN sol. equivalent to 25·15 ^{cm} ³ of AgNO ₃ sol.				
	NH ₄ SCN cm ³	AgNO ₃ cm ³	Excess of AgNO ₃ cm ³	AgSCN found. gram.
1.	25	25·3	0·15	·4372
2.	25	25·3	0·15	·4376
3.	25	25·4	0·25	·4373
4.	25	25·4	0·25	·4375
5.	25	30·4	5·25	·4382
} _____ }				
6.	25	Rough excess.		·4366
7.	25	“		·4381
8.	25	“		·4373
9.	25	“		·4372
10.	25	“		·4369

In order that the effect of varying the excess of silver might be investigated, an approximately decinormal solution of silver nitrate was titrated against the ammonium sulphocyanide and the ratio between the two solutions determined. This silver nitrate solution was used for the first five determinations of Table III. For the rest the quantity of silver nitrate was not measured but regulated by the eye alone, thus making

the conditions the same as would be the case in practical use of the method.

These results are as uniform as could be expected considering the variations which would be produced by even very small errors in measuring out 25cm^3 of decinormal sulphocyanide solution. It is moreover clearly shown that there is no difference in the results whether a bare excess or a moderately large excess of the silver nitrate is used.

The mean of the values in the last column is $\cdot4374$, which is equivalent to $\cdot2006$ grm. of ammonium sulphocyanide for every 25cm^3 of the solution.

The standard of the sulphocyanide solution was also determined volumetrically by Volhard's process. The mean of four titrations carried out with great care against a standard silver nitrate solution gave as the standard $\cdot2003$ grm of ammonium sulphocyanide for 25cm^3 of solution. This difference between the standards as determined by the two methods (one part in 670) is much less than the variations which frequently appear between successive determinations by Volhard's method, under like conditions as to strength of solutions and amounts used. It is about equal to the error that would be produced in a single volumetric determination by a mistake of one drop in measuring one of the solutions, or of one half drop in the same direction on each.

It is therefore evident that the standard of a sulphocyanide solution obtained in the above way may be applied directly to the estimation of unknown amounts of silver by Volhard's method without sensible error.

To remove a possible doubt as to whether the silver sulphocyanide dried at 115° was entirely free from water, a number of electrolytic determinations of the silver contained in the previously weighed precipitates of Table III were made in the following way.

The perforated platinum crucible containing the silver sulphocyanide and asbestos was hung in a loop of heavy platinum wire and served as the anode. For the cathode a deep platinum dish of about 200cm^3 capacity was used. An ammonical solution of potassium cyanide was employed as the electrolyte and gave the best results when made up by dissolving 2 grams of potassium cyanide in 15cm^3 of strong ammonia and 15cm^3 of water. The crucible which served as the anode was filled with this solution in full strength, and the remainder was put into the platinum dish and diluted to the required volume with water. In this medium the silver sulphocyanide is slowly dissolved and diffuses through the asbestos felt into the space between the electrodes where the silver is deposited in the usual way. This diffusion is, however, aided but little if at

all by the current, and there is a tendency for traces of the silver to remain behind in the crucible. The current density employed was about .0012 ampère per square centimeter of cathode surface and the time about twelve hours. After weighing the silver deposited, it was dissolved in nitric acid, precipitated by hydrochloric acid and weighed again as the chloride, giving a check upon the results.

Seven of the ten determinations of Table III^r were thus treated, but owing to the imperfections of the process the results were all slightly low, the worst showing a deficiency of .0025 grm. of silver, an error of less than 0.9 per cent. The results of the two best of these determinations given below are, however, sufficient to prove the point in question, namely that the silver sulphocyanide dried at 115° has the theoretical constitution and contains no water. The numbers are those under which the determinations appear in Table III.

AgSCN taken grm.	Calc. as Ag. grm.	Ag found by battery. grm.	Error. grm.	Weighed as AgCl. grm.	Calc. as Ag. grm.	Error. grm.
4. .4375	.2844	.2839	-0.0005	.3765	.2834	-0.0010
10. .4369	.2840	.2838	-0.0002	.3761	.2831	-0.0009

It is clear therefore that the estimation of sulphocyanides by precipitation with silver nitrate and direct weighing of the precipitate is wholly permissible. The method is extremely simple and, as has been shown, the results are quite accurate.

In conclusion I wish to thank Prof. F. A. Gooch for many valuable suggestions given during the course of this investigation.

SCIENTIFIC INTELLIGENCE.

I. CHEMISTRY AND PHYSICS.

1. *On the action of Permanganate upon Hydrogen Peroxide.*—Several views have been advanced in explanation of this well known decomposition, in which the permanganate is decolorized and oxygen is evolved. Berthelot believes that a trioxide of hydrogen is produced which undergoes spontaneous decomposition into water and oxygen. This idea is based upon the fact that he produced the decomposition at -12° without effervescence taking place. Schönbein's theory is that the two substances simultaneously give up atomic oxygen which condenses to the molecular condition. It is the view of Weltzien and Traube that permanganic acid oxidizes the hydrogen of hydrogen peroxide to water, thus liberating oxygen. According to the last explanation all of the evolved oxygen originates in the hydrogen peroxide.

BAEYER and VILLIGER have recently shown that the phenomenon observed by Berthelot was due to the supersaturation of the liquid by oxygen, and that if the liquid be gently agitated the evolution of gas takes place regularly. They observe that the hypothetical formation of hydrogen trioxide is an assumption which is supported by the formation of potassium polysulphide when sulphur and potassium sulphide are brought together, and also by the formation of potassium tri- or tetra-oxide, as observed by Schöne, when a solution of hydrogen peroxide is evaporated with potash; but the assumption now lacks foundation in fact. According to Schönbein's theory both substances act as oxidizing agents. The authors believe, however, that hydrogen peroxide is hardly to be considered as such an agent, since in the pure condition it decomposes dilute hydriodic acid only very slowly. Moreover, if the theory is true, the more strongly oxidizing derivatives of hydrogen peroxide should oxidize permanganate still more easily and evolve oxygen, while in fact exactly the opposite is the case, for example, with "Caro's acid" which does not attack permanganate. This view, therefore, is not supported by facts. The authors strongly favor the view of Weltzien and Traube that it is the hydrogen of the hydrogen peroxide that is oxidized to water. Hydrogen peroxide thus shows a behavior analogous to that of hydrogen sulphide, with which it shows a striking similarity in many other ways. They urge the adoption of this theory, agreeing as it does with the present state of our knowledge, by the authors of elementary text-books.—*Berichte*, xxxiii, 2488.

H. L. W.

2. *An Instance of Trivalent Carbon: Triphenylmethyl.*—In a preliminary paper GOMBERG describes a body produced by the action of zinc upon triphenylchloromethane, $(C_6H_5)_3C Cl$, which is an unsaturated compound apparently containing but one carbon atom and corresponding to the formula $(C_6H_5)_3C$. The substance

combines with atmospheric oxygen with great readiness, forming di-triphenylmethylperoxide $(C_6H_5)_3C.O_2.C(C_6H_5)_3$. This property renders the preparation of the unsaturated hydrocarbon in the pure state extremely difficult, and although the author has succeeded in obtaining large crystals of it, he has not yet been able to wash them without the formation of the peroxide. If subsequent investigation confirms the preliminary conclusions, the compound will possess great theoretical interest as being the first unsaturated compound of the kind, and furnishing the first definite evidence of the existence of the trivalent carbon atom.—*Jour. Am. Chem. Soc.*, xxii, 757. H. L. W.

3. *On the Phosphorescence of Inorganic Chemical Preparations.*—This subject has been studied by E. GOLDSTEIN. The method of investigation consisted in placing the substance under examination in a properly shaped tube and waving the latter back and forth in a cone of cathode rays. The nature of the phosphorescence could be determined from the length of the observed trail of light. The examination of various pure samples of a substance showed that the color of light attributed to it by previous investigators becomes weaker as the substance becomes purer. On the contrary, the light which appears at the point of contact with the cathode rays, and which lasts only during the illumination, becomes continually brighter as the purity increases. The substances investigated may be divided into two groups in respect to their phosphorescent light. In the case of the sulphates, carbonates, phosphates, borates, silicates, chlorides, bromides, fluorides, oxides, and hydroxides of Li, Na, K, Rb, Cs, Ca, Sr, Ba, Al, Zr, Mg, Be, Zn, and Cd the color is blue, violet-blue, or in a few instances violet. Salts of Cu, Cr, Mn, U, Ni, Co, Pb, Ce, La, Y, Er, Pr, and Ne belong to the other group. The latter, when admixed in very small amounts with compounds of the first group, give a light which in most cases exceeds the blue light in intensity and duration, but corresponding phenomena could not be observed when reciprocal additions from the first group were made. There is a maximum in the strength of the colored light, for it decreases again when the small amount of substance from group 2 is increased beyond a certain limit. The presence of moisture may possibly be assigned as the reason for the similar luminosity of the substances of the first group. The effect of additions from the second group may possibly be due to greater power of absorption or else to the occurrence of decomposition.—*Chem. Central-Blatt*, 1900, ii, 756. H. L. W.

4. *Weight of Hydrogen Desiccated by Liquid Air.*—LORD RAYLEIGH has found that the density of hydrogen dried by cooling with liquid air is the same as that obtained when the gas is dried with phosphorus pentoxide. This was to be expected, for it was shown by Morley in an article published in this Journal in September, 1887, that this oxide is a practically perfect dryer for gases, but it is satisfactory to have a confirmation of this important fact.—*Proc. Roy. Soc.*, lxvi, 334. H. L. W.

5. *The Separation of Tungsten Trioxide from Molybdenum Trioxide.*—M. J. RUEGENBERG and EDGAR F. SMITH describe a method of separation for the oxides mentioned in the title. It is based upon the fact that tungstic acid is insoluble in concentrated or dilute sulphuric acid, hot or cold, whereas molybdenum trioxide is very easily and rapidly dissolved. Upon trial it was found that sulphuric acid of specific gravity 1.378 was well suited for the purpose. Test analyses were made by mixing about 1 or 2 grams of molybdic anhydride with varying quantities of tungstic anhydride and digesting with 25° of warm sulphuric acid of the strength mentioned above for a few minutes, then filtering, washing with water containing sulphuric acid, and weighing the tungstic acid. The results of six analyses were extremely sharp. The authors found also that tungstic acid may be separated from ferric hydroxide in precisely the same way.—*Jour. Am. Chem. Soc.*, xxii, 772.

H. L. W.

6. *On Radio-active Barium and Polonium.*—When concentrated uranium nitrate solution is mixed with a little sulphuric acid, then with a solution of a barium salt, care being taken that sufficient barium salt is not added to cause a precipitate, and finally diluted with water, a strongly radio-active precipitate of barium sulphate is obtained. When this sulphate is converted into a soluble barium salt and this is treated with ammonium hydroxide, a small amount of precipitate is obtained which is even more strongly radio-active than the original sulphate; but when the barium is subsequently converted into carbonate, it has completely lost its activity. The radio-activity is probably due to small quantities of radium, or more probably of actinium. Polonium preparations obtained from lead chloride from uranium residues have proved to consist of bismuth hydroxide, which, either in this form or as oxychlorides exhibits strong activity.—*Berichte*, xxxiii, 1665. *Jour. Chem. Soc.*, lxxviii, 480.

7. *Lecture Experiments Illustrating the Electrolytic Dissociation Theory and the Laws of the Velocity and Equilibrium of Chemical Change.*—A. A. NOYES and A. A. BLANCHARD have published a description of seventeen experiments which were originally devised as an accompaniment to an extended course of lectures on theoretical chemistry. Most of the principles illustrated are of such fundamental importance that they should soon be generally introduced into elementary lecture courses on inorganic and analytical chemistry. The experiments are devised with much skill and will furnish very striking illustrations to the audience. They are described in great detail, so that they may be readily performed by the lecturer, and the principles illustrated are clearly stated in each case. It is certain that this series of experiments will be of great assistance to teachers of theoretical chemistry, and it will doubtless serve to hasten the introduction of instruction in the principles of chemical equilibrium and of electrolytic dissociation into elementary courses.—*Jour. Am. Chem. Soc.*, xxii, 726.

H. L. W.

8. *Physikalisch-chemische Propädeutik*; von H. GRIESBACH, zweite Hälfte, 3 Lieferung, Band I, pp. 945-992 and Band II, pp. 1-352. Leipzig, 1900 (Wilhelm Engelmann). The first two installments of this work were noticed in this Journal, December, 1896. The present, fourth installment gives a most favorable impression of the author's success in continuing the excellence of his work. The size of the book as well as its scope have been enlarged beyond the original intention, so that a second volume is now begun. In giving a general view of modern scientific knowledge the work will be valuable to many classes of readers.

H. L. W.

9. *A School Chemistry*; by JOHN WADDELL. 12mo, pp. xiii, 278. New York, 1900 (The Macmillan Company).—This text-book on elementary chemistry possesses many excellent features. The author has shown much originality in his presentation of the subject, good judgment in regard to the scope of the book, and accuracy in his statements of facts and theories. The interrogatory method is largely employed with the object of stimulating the pupil's thought. It is the aim of the book to help the student in the discovery of facts, to enable him to see their connections, and to show how facts lead to theory.

H. L. W.

10. *Leçons de Chimie Physique*; by J. H. VAN'T HOFF. Volume III. Relations between the properties and the composition. Translated from the German edition by M. Corvisy. 8vo, pp. 170, Paris, 1900 (A. Hermann).—This book is a very literal translation of the concluding volume of van't Hoff's "Vorlesungen über Chemie" which appeared early in the year. The subject treated is, however, practically independent of the two earlier volumes. The relations between the physical properties and the composition are first taken up. Under this heading, come the relations and behavior of matter in regard to space, pressure, temperature, heat, surface tension and optical properties. The second half of the book treats of the relations between the chemical properties and the composition.

Van't Hoff has made a wide departure from other books on physical chemistry which have thus far appeared. He aims not to treat the subject exhaustively but to take up the more important or more interesting parts, and these are treated in his usual brilliant manner.

H. W. F.

11. *On the Viscosity of Gases as affected by Temperature*; by Lord RAYLEIGH.—A former paper* describes the apparatus by which I examined the influence of temperature upon the viscosity of argon and other gases. I have recently had the opportunity of testing, in the same way, an interesting sample of gas prepared by Professor Dewar, being the residue, uncondensed by *liquid hydrogen*, from a large quantity collected at the Bath springs. As was to be expected, it consists mainly of helium, as is evidenced by its spectrum when rendered luminous in a vacuum tube. A line, not visible from another helium tube, approximately in the position of D_6 (Neon) is also apparent. The result

* This Journal, ix, 365.

of the comparison of viscosities at about 100° C. and at the temperature of the room was to show that the temperature effect was the same as for *hydrogen*.

In the former paper the results were reduced so as to show to what power (n) of the absolute temperature the viscosity was proportional.

	n	c
Air	0.754	111.3
Oxygen.....	0.782	128.2
Hydrogen }	0.681	72.2
Helium }		
Argon	0.815	150.2

Since practically only two points on the temperature curve were examined, the numbers obtained were of course of no avail to determine whether or no any power of the temperature was adequate to represent the complete curve. The question of the dependence of viscosity upon temperature has been studied by Sutherland, on the basis of a theoretical argument which, if not absolutely rigorous, is still entitled to considerable weight. He deduces from a special form of the kinetic theory as the function of temperature to which the viscosity is proportional

$$\frac{\theta^{\frac{1}{2}}}{1+c/\theta}$$

c being some constant proper to the particular gas. The simple law $\theta^{\frac{1}{2}}$, appropriate to "hard spheres," here appears as the limiting form when θ is very great. In this case, the collisions are sensibly uninfluenced by the molecular forces which may act at distances exceeding that of impact. When, on the other hand, the temperature and the molecular velocities are lower, the mutual attraction of molecules which pass near one another increases the number of collisions, much as if the diameter of the spheres was increased. Sutherland finds a very good agreement between his formula (1) and the observations of Holman and others upon various gases.

If the law be assumed, my observations suffice to determine the values of c . They are shown in the table, and they agree well with the numbers for air and oxygen calculated by Sutherland from observations of Obermayer.—*Proc. Roy. Soc.*, lxxvii, 137.

12. *Absorption of gases by glass powder*.—Early investigators claimed that the absorption of gases ceased in a few days; that it increased with increasing pressure; diminished with increasing temperature; and that the presence of moisture had little influence. Bunsen on the other hand found that the absorption continued even for years; increased with increasing temperature, and was largely independent of variations of pressure. The cause of the absorption he claimed to be in the capillary coating of moisture. P. MÜLFARTH (Bonn dissertation) gives an exhaus-

tive list of references of work upon this subject. His conclusions vary from those of the previous investigators.

Fully dried glass powder absorbs a large amount of CO_2 . The amount increases with diminishing temperature and rises with increasing temperature: it is accomplished in a few hours. The presence of moisture hinders this absorption. The absorption of SO_2 is similar to that of CO_2 : it is accomplished in a couple of hours—increases with increasing pressure, diminishes with increasing temperature. The following absorb in the order written: C_2H_2 , N_2O , CO_2 , SO_2 , NH_3 . The investigation bears upon the preparation of X-ray tubes, and also upon the question of the influence of the glass walls of vessels in spectrum tubes.—*Ann. der Physik*, No. 10, 1900, pp. 328–352. J. T.

13. *Normal lines of iron*.—Rowland has published a list of iron lines which serve for normals in spectra comparisons. KAYSER finds that this list is not sufficient for certain regions of the solar spectrum, and, therefore, has given a large additional number between wave lengths 2300 and 4500. The author claims a higher accuracy than that reached by Rowland since every wave length in his table is the mean of from 6–15 measures, taken with gratings of different orders. The probable error lies between 0.001 and 0.003.—*Ann. der Physik*, No. 10, 1900, pp. 195–203. J. T.

14. *Arc spectra of some metals in an atmosphere of hydrogen*.—HENRY CREW finds that a hydrogen atmosphere profoundly modifies certain arc spectra of metals, and that all the lines in the arc spectra which are affected by hydrogen belong to the spark spectrum also. Moreover that the lines which belong to Kayser and Runge's series are unaffected by the change from air to hydrogen.—*Phil. Mag.*, Nov., 1900, pp. 497–505. J. T.

15. *Recognition of the Solar Corona independent of the total eclipse*.—H. DESLANDRES has made some observations at Meudon which apparently show the possibility of observing the corona under ordinary conditions by the aid of a thermopile.—*Comptes Rendus*, Oct. 22, 1900. J. T.

16. *Loss of charge by evaporation*.—Various physicists have arrived at different conclusions on this subject. W. CRAIG HENDERSON, working in the Cavendish laboratory, has investigated the subject anew and comes to the conclusion that there is no appreciable loss of charge by the evaporation of a charged liquid.—*Phil. Mag.*, Nov., 1900, pp. 489–497. J. T.

17. *On the action of the coherer*.—Previous investigations seem to show that the action of the coherer is due to the welding action of minute sparks in the powders which are employed. T. MIZUNO of Kyoto University has studied the behavior of a large number of substances. He finds that the action of the iron powder coherer is conspicuously irregular and concludes that the action of the coherer can plausibly be explained on the hypothesis of the welding action of minute sparks.—*Phil. Mag.*, Nov., 1900, pp. 445–459. J. T.

18. *Loss of electrical charges in air which is traversed by ultra-violet rays.*—P. LENARD published a paper on the effect of ultra-violet rays on gas in the *Annalen der Physik*, No. 1, 1900, p. 486. He now continues his investigation and shows that the effect of ultra-violet light is fourfold. The air is divided into bearers of negative electricity which appear to be charged atoms or molecules; bearers of positive electricity of larger dimensions; nuclei of vapor which are non-electrified; and ozone.—*Ann. der Physik*, No. 10, 1900, pp. 298–319. J. T.

19. *Handbuch der Spectroscopie* von H. KAYSER. Erster Band. Pp. xxiv, 781; large 8vo; 781 figures in the text. Leipzig, 1900 (S. Hirzel).—The preparation of a comprehensive work on a subject which has had so remarkable a development as that of spectroscopy in its different branches, is a labor of such magnitude that we may well honor the energy of the author who ventures to undertake it. But here we have not simply the promise but also its fulfilment in the completion of a considerable part of the undertaking. The present volume of nearly eight hundred pages by Professor Kayser, is so thorough and so satisfactory a treatment of the portions of the subject here included, that the reader feels justified in looking forward with confidence to the completion of the remaining four volumes promised as a well-rounded whole.

The volume opens with an admirable summary, scientific and yet thoroughly readable, of the history of spectroscopy from the time of Newton down to the present. Although an exhaustive treatment of this subject might well extend the chapter to the limits of a volume by itself, the author presents the salient points in the successive steps of progress, and the contributions of each worker so clearly, that the chapter as a whole leaves little to be desired. In the treatment of the delicate question of the claims of earlier writers to the discoveries which are usually connected with the name of some physicist who has come later, the author shows much discrimination and critical care; for example in regard to the relative amount of credit to be given to Wollaston as contrasted with Fraunhofer, or to Talbot, Miller and Stokes as compared with Kirchhoff. The same critical method appears throughout. The five other chapters of the volume deal with the practical working of spectroscopy. The first of these discusses the production of incandescent vapors in the flame, arc-light, or vacuum tubes. This subject is treated both historically with a multitude of references to original authorities and descriptively as needed by a present worker. The two chapters following discuss prisms and diffraction gratings, giving the theory, construction and use in detail. In the former subject the author has been aided by Dr. H. Konen, while Prof. C. Runge has furnished for the volume his theory of concave gratings. The concluding chapters are devoted to spectroscopic apparatus and measurements, the use of photography and other methods. Here, as throughout the volume, we are struck with the simplicity and

directness of statement, and the thoroughness with which the work of the many physicists who have contributed is recognized. Even where only a very brief allusion is possible to some matters of minor importance, they are not ignored, but the literature references at the bottom of each page embrace all that has been done on the subject. The great steps of progress that have been made, as the Rowland grating, the methods for the investigation of the infra-red spectrum by Abney, the echelon spectroscope of Michelson, are treated with all necessary fullness.

The second volume is to be devoted to Kirchhoff's law and its consequences, emission spectra of different kinds and as obtained under varying conditions of pressure, temperature, magnetic force, etc. The third volume will include the absorption phenomena with also fluorescence and phosphorescence; the fourth volume will give all spectroscopic data available in regard to each individual element, and a fifth is also planned to include the spectroscopy of the heavenly bodies including the Sun.

20. *A Brief Course in General Physics, experimental and applied*; by GEORGE A. HOADLEY, Swarthmore College. Pp. 463, New York, 1900 (The American Book Company).—This is a clear, systematic statement of the fundamental principles and laws of physics, such as is needed by the high school student; it is accompanied by good illustrations and a well-chosen series of experiments. The subjects in general are well treated, although at occasional points open to criticism; thus we find in one article a brief statement of surface tension and in one following, an explanation of capillary attraction, but there is no suggestion as to any connection between the two subjects.

II. GEOLOGY AND MINERALOGY.

1. *U. S. Geological Survey, 20th Annual Report 1898-99. Part III. Precious Metal Mining Districts.* (Washington, 1900, pp. 595, pls. 77.)—This volume consists of three papers, two of them having appended reports by specialists.

I, *A preliminary report on the Bohemia mining region of western Oregon*, with notes on the Blue River mining region and on the structure and age of the Cascade range by J. S. DILLER, which is accompanied by a report on the fossil plants associated with the lavas of the Cascade range by F. H. KNOWLTON (pp. 1-64, pls. 1-6). The brief sketch of the district shows that the rocks of Calapooya Mountain, upon which the district lies, are wholly of igneous origin. They consist of dacite porphyry, andesites and basalts with volcanic tuffs. It is in these rocks that the mines are situated which are described in the report. The veins are gold-bearing. The fossil plants found in the tuffs show that the latter were deposited in lakes of Miocene age. The various localities from which fossil leaves have been collected are mentioned and in his special report Knowlton describes

these leaves, showing them to represent a flora of Miocene age made up of conifers, oaks, etc., and including several new species.

II, *Gold and silver veins of Silver City, De Lamar and other mining districts in Idaho*, by W. LINDGREN (pp. 65-256, pls. 7-35). This begins with a sketch of the general geology of western central Idaho, accompanied by a geologic map, based in part on actual surveys made by the author and in part on reconnaissance work. The map shows a central area of post-Carboniferous granite of great size, bordered on the east by Carboniferous beds in which it has produced contact metamorphism. The Snake River runs through the southern end, turning to the north, the western portions of the area and along its course is bordered by immense outflows of basalts, some of Miocene, others of Pliocene age. The Silver City and De Lamar district is in this southern part on a granite ridge which rises above the flood of basalt and rhyolite. The gold and silver veins of this district, their mode of occurrence, the associated minerals, the individual mines, etc., are carefully and fully described. In the same manner the author gives an account of the Wood River district, where the silver-lead ores occur in veins cutting the Carboniferous beds. At Florence and Warren, situated in the great granite area, the veins are gold quartz and some placer mining is done. In the Seven Devils district the ore is copper in the contact zone of the granite mass with the Carboniferous beds on its western side. The ores are bornite and chalcopyrite and are supposed to have been deposited by pneumatolytic processes. The paper contains a wealth of careful observations on a large number of ore deposits and is a most important contribution to the study of ore deposits in general.

III, *Geology of the Little Belt Mountains of Montana*, with notes on the mineral deposits of the Neihart, Barker, Yogo and other districts by W. H. WEED, with a report on the petrography of the igneous rocks of the district by L. V. PIRSSON.

The Little Belt Mountains are an anticlinal uplift of a rather broad, even character. In the central portion erosion has disclosed the fundamental Archean gneisses and schists upon which the later sediments lie. These range through the Belt formation (here having a considerable local development and supposed to be of Algonkian age), the Cambrian, Siluro-Devonian, Carboniferous, up to the Cretaceous forming the outer foot hills, slopes and boundary plains of the area. The streams rising in basins of Belt or Cambrian shales pass through deep and narrow canyons or gateways of Carboniferous limestone. Chiefly on the outer margin of the range, along the spring of the arch, great laccolithic intrusions of igneous material have taken place, which now present either wholly or partly uncovered masses of porphyries of several types. Accompanying them are numbers of intruded sheets and dikes.

The basins cut in shale, the flat-topped plateaux of limestone, the deep gorges cut on the outer flanks of the range and the

great mountains of igneous rock form the dominating features of topographic relief and express the geologic relations. The formations are described and also the local geology of the various districts is given in detail and several features of general interest in structural and dynamical geology, particularly with reference to the intrusion of laccoliths are brought out.

The ores are found chiefly in two districts; at Barker where they are of the type of silver-lead ores occurring in limestones at the contact of intrusive igneous masses, and at Neihart where intrusions of diorite and other rocks have occurred in the Archean complex. At this place the ores occur in a well defined vein system and are gold and silver, the latter occurring in the sulphides, ruby silver, polybasite, etc. The important mines, their ores and workings are described.

The petrographical report contains a full description of the igneous rocks which are chiefly porphyries of several kinds, granite, syenite and diorite porphyries with intermediate types. At Yogo peak is a differentiated mass changing from granite porphyry through monzonites to shonkinite. The intruded sheets and dikes present several rock varieties, the most interesting of which are the minettes and analcite basalts. Of all these types detailed analyses by W. F. HILLEBRAND and H. N. STOKES are given and in the theoretical discussion of the rocks of the district are used to show that all of the main types are so related to one another in a rock series that any one of them may be expressed in terms of the others; these relations being shown by a diagram in graphic form.

L. V. P.

2. *Iowa Geological Survey, Vol. X. Annual Report, 1899, with accompanying papers.* SAMUEL CALVIN, State Geologist, H. F. BAIN, Assistant State Geologist. Pp. 1-166, plates i-xi, figures 1-102, and 10 folded maps, 1900.—The Geological Board of the State of Iowa is to be congratulated both upon the admirable organization of its geological workers and upon the equally admirable results it is annually producing. The board is constituted of the governor, the auditor of State, two college presidents and the president of the Iowa Academy of Sciences, thus combining for effective service the forces of politics, education and science. The geological staff for the year, besides the geologist and assistant geologist, is drawn from the officers of six of the colleges of the State and one high school and one from the Chicago University. The papers of the present volume are chiefly connected with the detailed mapping of the geology of counties not previously so mapped. Forty-two counties in all have been mapped, eight of them having been completed within the year. The great value of the topographical sheets prepared by the U. S. Geological Survey is acknowledged, particularly in respect to the Dubuque region, the important lead mines of which are scientifically described in the chapter on Dubuque County, prepared by the geologist and assistant geologist.

Mr. Frank Wilder brings to notice, the fact "that the Hawarden

beds, heretofore referred to the Ft. Pierre, really belong to the Benton shales and accordingly the limestones and chalk of the Inoceramus beds, as developed in Iowa, are not to be correlated with the Niobrara farther west. They represent instead that portion of the Fort Benton sub-stage exposed in the 'Oyster Shell Rim' of the Black Hills." The principle of coöperation expressed in the organization of the survey, is further carried out in the formation of a joint commission to investigate the clay resources of the State. This commission is composed of the Iowa Brick and Tile Makers' Association, Iowa Engineering Society,² the State College, and the State University. The results of their work during the year is given in Dr. Beyer's appended report on the Mineral Production of Iowa, in 1899. H. S. W.

3. *Cleopatra's Emerald Mines*.—An account is given in the November number of the Geological Journal of an expedition led by Mr. DONALD A. MAC ALISTER to the so-called Cleopatra's Emerald Mines in the neighborhood of Jebel Sikait in Northern Etbai. The locality is situated nearly due east of Edfu, and some fifteen miles in a direct line from the Red Sea, between 25° and 24° 30' N. Lat. The place was reached by the expedition starting from Daraw. The mountain, Jebel Sikait, has an altitude of 1800 feet above sea-level and commands an extensive view over the wild and desolate surrounding country.

The emeralds occur in mica schist and talc schist, which rocks occur over a large area. The mica schist is in some cases highly micaceous, looking like a finely bedded, contorted sandstone, and again contains very little quartz. In some cases the rock is much hydrated, passing into a soft, powdery rock. Tourmaline, garnet, actinolite, chrysolite are common minerals of the locality. The schists of Jebel Sikait are inclined at an angle of about 45°, and overlie gneiss on one side, while on the other they are inclosed by a tough, green serpentine, of which the upper part of the mountain is composed. The search for emeralds was carried on two thousand years ago very extensively, a large number of mines being opened and extensive workings being carried forward. The author remarks in regard to this subject:

"The mining is of a most primitive character. The ancients simply excavated, in the likely emerald-bearing schist, a network of long and very tortuous passages just large enough to allow of the body being dragged through, and only in a very few cases was any attempt made at stopping (or excavating) the entire seam. It has been suggested that the passages were made small on account of the absence of timber, suitable to be hewn into supports, which would be necessary to prevent collapses in the case of larger excavations. So small are the openings, that a casual observer would not notice the existence of a mine at all but for the gray *débris* thrown out at the mouth. Along the schists at Sikait alone we visited considerably over a hundred mines, some of which took more than an hour to crawl through. That these mines have been worked at widely distant periods, is

evident from the different styles of work. There are seven or eight groups of mines in different places within a couple of hours or so of Jebel Sikait."

4. *Handbuch der Mineralogie*, von Dr. CARL HINTZE. Erster Band, fünfte Lieferung; pp. 641-800. Leipzig, 1900 (Veit & Company).—The fifth part of the first volume of Hintze's Mineralogy (No. 17 of the entire series) has recently been issued. It contains the closing pages in regard to the pyrrhotite group, descriptions of the species of the cinnabar group, and the opening portion of the pyrite group. The treatment is as thorough and exhaustive as that which has characterized the parts already published.

5. *Contributions to Chemistry and Mineralogy from the Laboratory of the U. S. Geological Survey*, F. W. CLARKE, Chief Chemist. Pp. 166, Washington, 1900. (Bulletin 167, U. S. G. Surv.)—This Bulletin, like others which have preceded it from the same source, contains a series of chemical and mineralogical researches carried on in the laboratory under the charge of Prof. Clarke. Several of the papers included, as that on the constitution of tourmaline, and others, have already been published in part or entire in this Journal. Among others may be mentioned a series of papers by H. N. Stokes on the chloronitrides of phosphorus and the metaphosphimic acids.

6. *Johnstonotite, a supposed new Garnet*.—W. A. MACLEOD and O. E. WHITE have given the name johnstonotite (after R. M. Johnston) to a brownish-yellow garnet occurring in trapezohedral crystals in trachyte at Port Cygnet, Tasmania. An analysis gave:

SiO ₂	Al ₂ O ₃	FeO	MnO	MgO	CaO	Ign.
36.87	7.28	17.12	13.68	12.49	11.98	0.29 = 99.71

It would seem to be analogous to the spessartite of Colorado, which is similar in occurrence; however, although regarded as differing from other garnets in the relative proportions of the bases, the analysis cannot be accepted as accurate since (as noted by the reviewer, L. J. Spencer) the formula only approximates to that of garnet, even if all the iron is made Fe₂O₃.—*Proc. R. Soc. Tasmania*, 1898-99, 1900, 74, noticed in *J. Chem. Soc.*, lxxviii, 663.

III. MISCELLANEOUS SCIENTIFIC INTELLIGENCE.

1. *Annual Report of the Board of Regents of the Smithsonian Institution showing the operations, expenditures and condition of the Institution for the Year ending June 30, 1898*. Pp. lv, 713. Washington, 1899.—The report of the Secretary of the Smithsonian Institution, Prof. S. P. LANGLEY, which has just been issued, contains the usual interesting summary of the recent progress of the Institution in its various departments of research and exploration, collections, library and publications, and foreign

exchange. The growth in the *National Zoological Park* is noted in particular and also the results obtained in the *Astrophysical Observatory* by Mr. C. G. Abbott, the Aid in charge. Professor Langley states that the work of the last year "has resulted in the discovery and determination of position of 500 new absorption lines, so that we have now over 700 new lines of well-determined positions, and we may be said now to know, by the aid of the bolometer and the labors of the observatory, more lines in this invisible spectrum than were known in the visible one up to the great research of Kirchhoff and Bunsen, which opened the era of modern spectrum analysis. Moreover, these lines, the exactness of whose determination has now reached a surprising degree of perfection through the recent improvements in the delicacy as well as the precision of both bolometer and galvanometer, and through other improvements in the apparatus (improvements due principally to the present Aid acting in charge), depend not only on the instruments, but on the labors of those who have used them, the comparator measurements alone having included, as stated in the body of the report, about 44,000 separate observations.

"A great deal of other work has been done at the observatory, but nothing which in importance and present and prospective interest compares with this main research in the infra-red spectrum, which is now known throughout nearly the whole of the invisible portion of the solar energy, and extends through a range of wave lengths considerably over twelve times that known to Sir Isaac Newton, the present exact knowledge of this region being due not exclusively, but it may properly be said principally, to the labors of this observatory.

"I call attention in this connection to the interesting remarks made in the report to the effect that very marked changes of absorption have been observed at various parts of the infra-red spectrum. In one part of the invisible region a decrease in absorption, amounting to nearly half the total, took place in February, and this abnormal state continued until May, when the usual condition gradually returned. As this change is found to occur yearly at about the same period, the idea presents itself that the growth of vegetation, so rapid in these months, may abstract from the air large quantities of vapor active in absorption at this point in the spectrum, but this interesting possibility has not yet, it will be understood, been fully verified."

The second part of the volume, pp. 101-696, contains a series of well-selected papers on important topics, many of them not readily accessible in the original, so that their republication here will be a great convenience to many readers.

2. *Report of the U. S. National Museum under the direction of the Smithsonian Institution for the year ending June 30, 1898.* Pp. xviii, 1270; plates 1-36. Washington, 1900.—This volume contains the report by the Director, Mr. Charles D. Walcott, on the condition and progress of the United States National Museum

during the year ending June 30th, 1898; also the reports of the head curators, W. H. Holmes in Anthropology, F. W. True in Biology, and G. P. Merrill in Geology; further lists of accessions to the collections and library (pp. 1-152). Part II, following, is given to an extended paper by the late Professor E. D. Cope, on the Crocodilians, Lizards, and Snakes; this covers pages 153 to 1270, with thirty-six plates and numerous text figures.

The Director, in his brief statement, shows the remarkable growth which the Museum has made and emphasizes the imperative need of a larger, more dignified and more convenient building, not simply for exhibition purposes, but also to include laboratories and a lecture hall. He shows further the directions in which it is to be desired that the National Museum should develop in the immediate future, and the great service which it may thus do to the country at large. It is greatly to be hoped that his appeal may be regarded and adequate provision made, not only for the present needs of the Museum but also for its future growth.

3. *Publications of the Earthquake Investigation Committee in Foreign Languages.* Tokyo, 1900. No. 3, pp. 1-103. No. 4, pp. 1-141.—These publications of the Japan Earthquake Committee contain a series of papers of much interest to seismologists, chiefly in English, in part also in French. Among these are accounts, with seismic records, of several destructive earthquakes (1891-1894). A paper of general interest is by H. NAGAOKA, on *the elastic constants of rocks and the velocity of seismic waves*. This gives a long series of observations on specimens of rocks from Archean chlorite schist, through the Paleozoic to quartz-sandstones and andesites of the Tertiary. The modulus of elasticity, and modulus of rigidity were determined in each case, as also the density, and from these data the corresponding velocities for elastic waves were then calculated. The author finds a distinct gradation in the elastic constants from the higher values characteristic of the oldest rocks down to those of the youngest; the velocity of wave-propagation is not, however, proportionally larger, since the density also enters into the problem. Thus for an Archean chlorite schist a velocity of 6.4 to 7 kilometers per second is obtained for longitudinal waves; for a Paleozoic clay slate a velocity of 4.5 to 6; while for a Tertiary sandstone the value is only 2.2. The experiments made seem to show that the elastic constants increase more rapidly than the density as the rocks become more dense, and consequently elastic waves travel with greater velocity in the interior than on the surface of the earth's crust. Eruptive rocks are more isotropic than those of non-igneous origin and have inferior elasticity, but there is with them also the same distinction as to age. The schistose structure of the deeper rocks gives a greater value of the elastic constants in a certain particular direction, coinciding with that of swiftest propagation of elastic disturbances. It seems to be shown that the ratio of elastic constants to density, and hence

the velocity of elastic waves, increases from the surface downward to a *stratum of maximum velocity*, probably at a considerable depth, and then diminishes again. The conclusion of Professor Omori from records in Italy and Japan, that the velocity of the first tremor is almost always equal to 13 kilometers per second, while longitudinal waves in steel have only a velocity of 6.2, is difficult to account for; it may be explained by the probable supposition that the rocks deeper in the earth's crust have elastic constants several times greater than those near the surface. The velocity named would roughly correspond to the values $E = 6 \times 10^{12}$ and $\rho = 3.5$. The high value of the elastic constant relative to the density is not impossible, since, in passing from Tertiary to Archean rocks an increase of density from 2 to 3 is here shown to be accompanied by an increase of the modulus of elasticity more than ten times in some cases.

4. *The Graduate Bulletin of the University of Nebraska*. Volume I, No. 1, pp. 1-68. Lincoln, Nebraska, 1900.—Among the papers here published is one by E. J. Rendtorff, on achromatic polarization with crystalline plates; another by C. C. Engberg on the Cartesian oval, and one by C. A. Fisher on the geology of Lincoln and environs.

5. *National Academy of Sciences*.—The autumn meeting of the National Academy was held at Providence, Nov. 13-15th. The following is a list of the papers read (several by title):

J. TROWBRIDGE: Investigations of light and electricity with the aid of a battery of twenty thousand cells.

ALPHEUS HYATT: Progressive evolution of characters in the young stages of Cephalopods. Descriptive method of presenting the phenomena of the cycle of evolution among Cephalopods.

T. W. RICHARDS: The porous cup voltmeter. An account of the study of growing crystals by instantaneous microphotography.

S. L. PENFIELD: Stereographic projection and some of its possibilities from a graphical standpoint.

C. S. MINOT: On the development of the pig. Normal plates illustrating the development of the rabbit and the dogfish.

T. C. MENDENHALL: Note on the energy of recent earthquakes.

A. A. MICHELSON: Spectrum of sodium in a magnetic field. Progress in the Echelon spectroscopie.

H. A. ROWLAND: A report of the spectrum work carried on with the aid of a grant from the Bache Fund. On the explanation of inertia and gravitation by means of electrical phenomena.

A. S. PACKARD: Distribution of phylogeny of *Limulus*. Male preponderance (androrhopy) in Lepidopterous insects.

A. MICHAEL: The synthesis and reactions of sodium acetate ester, and their relation to a new interpretation of chemical metathesis. On the genesis of matter

C. BARUS: Demonstration of the projection of one grating by another. Exhibition of certain novel apparatus: A wave machine; an expansion lens; a recording system of two degrees of freedom; A tube showing colored cloudy condensation. On stability of vibration and on vanishing resonance. Hysteresis-like phenomena in torsional magnetostriction and their relation to viscosity.

CHARLES D. WALCOTT: Folding and faulting of strata in the Cordilleran Area.

J. M. CRAFTS: Report on the Meeting of the Committee of the International Association of Academies held at Paris.

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JULY, 1900.

Established by BENJAMIN SILLIMAN in 1818.

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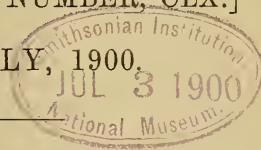
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FOURTH SERIES.

VOL. X—[WHOLE NUMBER, CLX.]

No. 55.—JULY, 1900.



NEW HAVEN, CONNECTICUT.

1900.

THE TUTTLE, MOREHOUSE & TAYLOR CO., PRINTERS, 125 TEMPLE STREET.

Published monthly. Six dollars per year (postage prepaid). \$6.40 to foreign subscribers of countries in the Postal Union. Remittances should be made either by money orders, registered letters, or bank checks.

ENGLISH FLUORS AND BARITES

The consignment purchased by Mr. Foote in Durham has reached us after many disappointing delays. It includes certain types of crystals which unquestionably excel the best Fluors known in any of the world's museums. One 9 inch crystal, bright and of deep purple color, will be difficult to match, and several interesting groups should find space in some of our leading University Collections.

LARGE COMPLETE CUBES of symmetrical form, showing every face and no point of contact, are a rarity. They come high, but we had to have them. One is stepped in a curious manner, one slightly "hopper-shaped" and still another is quite flat; variations which enhance the value somewhat.

NEW QUARTZ-COATED FLUORS. The Fluor is brilliant clear purple, and tastefully dotted with limpid Quartz crystals of 3 to 6 mm. diameter, whose quality suggests that they were taken from Herkimer Co., and sprinkled there! The old style milky quartz on Fluor appears very tame beside this rare and exquisite combination. Good green and amber-yellow Fluors are no longer to be bought in England. We found excellent specimens, however, on opening a lot imported eleven years ago.

THE SECOND CONSIGNMENT.

We are advised that a further shipment of six boxes will reach us about July 1st. Lists received indicate that we will be in a position to offer complete series of the Cumberland and Durham minerals, together with numerous oddities secured by Mr. Foote in both London and Paris. Evidently, prices on these minerals are to be placed far below any "reasonable" mark heretofore established. *Par exemple*—

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1st—Clear amber.

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LUTECITE, Dana App. I., p. 58. An anhydrous silica found sparingly four years ago while excavating for the foundations of a Parisian "abattoir." This limited find was the best ever made, and nothing has since been found. The specimens offered were obtained from the discoverer, and their authenticity can be relied upon.

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WONDERFUL QUARTZ CRYSTALS FROM MONTANA.

Our third shipment (June 13th) of Montana Amethysts and Quartz Crystals enclosing Tourmaline, proved by far the best yet received. 589 single crystals and groups are in this lot, including many large, showy museum specimens, and many fine specimens of small and large cabinet sizes. Odd shapes and colors, parallel growths, phantoms, curious groupings, etc., abound. Enclosures of Tourmaline in Quartz are by no means common, consequently this is a rare opportunity for wide awake collectors: the prices are very low, 10c. to 25c. for a good, often doubly-terminated, crystal, while \$1.00 to \$2.50 will buy the very best of the entire lot, only a few specimens including the large museum groups higher than \$2.50.

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A recent shipment from Nova Scotia includes very fine specimens of Pyrolusite. Crystals up to an inch in size and in attractive groups, 50c. to \$5.00.

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A small shipment received during June is the harbinger of a selection of some 400 specimens soon to arrive. The lot included the finest Marshite ever found. This specimen has been turned over to Prof. Penfield and more will be known about Marshite when his investigations are completed. Also a few excellent specimens of Stolzite from Broken Hill, \$1.00 to \$10.00.

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FOURTH SERIES.

VOL. X—[WHOLE NUMBER, CLX.]

No. 56.—AUGUST, 1900.

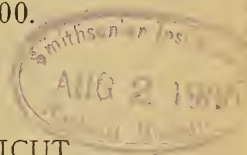
WITH PLATE I.

NEW HAVEN, CONNECTICUT.

1900.

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WE DISCOVERED

That the material contained in those six cases of Cumberland minerals merited a fuller description than previous space in advertisements permitted. They were somewhat slow in arriving, and unpacking, washing and pricing takes time—so that July 16th rather than July 1st found them ready for your inspection. In spite of the continued high temperatures, we will endeavor to make our visitors as comfortable as ice-water and electric fans allow; however, such apparatus is not required to assure your appreciation of our stock of

Single Crystals. Many types of Barite with interesting modifications, coloring and translucency. The latter feature is perhaps most noticeable—unless it be the rare perfection of the individual specimens. No end of sizes, with prices attached which are veritably surprising. Fifty or seventy-five cents buys crystals which ordinarily would be hard to duplicate for twice the figures. Modified Calcites, long and slender, of beautiful quality, at similar rates. Complete cubes of Fluorite, from $1\frac{1}{2}$ inches square up to four and five inches. These are by no means common, but our stock represents an assortment nowhere to be equalled.

Groups in all sizes, adapted for museum representation or private cabinets. Purple Fluors, many crystals measuring three to five inches across, in some instances coated with Quartz or Siderite, are among the most attractive of the larger specimens, although several of the “phantomed” Barite groups are of equal interest. Fluorites for open case display can be selected from a stock varying between \$2.00 and \$25.00, while Barites for similar purposes range from \$1.50 to \$20.00. Drawer specimens of either mineral, unusually choice examples, 50c. to \$2.00.

Quartz after Gypsum with Lutecite. The pseudomorph is not the least interesting feature in connection with the previously described Lutecite specimens. Their locality, “Vaugirard, Paris,” entitles them to some distinction, aside from the classification given in Dana's Appendix. \$2.00 to \$4.00.

URUGUAY “WATER STONES.”

Several of these curious formations were recently purchased, and at the values affixed will quickly find buyers. In appearance they resemble an ordinary Chalcedony geode, with an almost botryoidal surface, beneath which can be distinguished the rapidly moving bubbles. As an inclusion they are highly desirable, for on slightly shaking, the liquid can be heard running in and out through the emptied cavities. They had been kept by the former owner for over seven years—and never changed the least. \$6.00 and \$8.00.

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SEPTEMBER, 1900.

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Quartz after Gypsum with Lutecite. The pseudomorph is not the least interesting feature in connection with the previously described Lutecite specimens. Their locality, “Vaugirard, Paris,” entitles them to some distinction, aside from the classification given in Dana’s Appendix. \$2.00 to \$4.00.

URUGUAY “WATER STONES.”

Several of these curious formations were recently purchased, and at the values affixed will quickly find buyers. In appearance they resemble an ordinary Chalcedony geode, with an almost botryoidal surface, beneath which can be distinguished the rapidly moving bubbles. As an inclusion they are highly desirable, for on slightly shaking, the liquid can be heard running in and out through the emptied cavities. They had been kept by the former owner for over seven years—and never changed the least. \$6.00 and \$8.00.

YOU ARE CORDIALLY INVITED

To call at our new store in Paris, if you intend visiting that city during the current year, or at some future date. Opening in May, we have endeavored to make our drawers and cases look as attractive as the fine stock and new furnishings will permit, especially for the many visitors expected during the Exposition Season. Within the grounds we are represented under the American Sections of the Mines and Mining Building and the Educational Departments (Groups XI. and XII.). Every consideration will be cheerfully shown to friends and patrons interested.

FOOTE MINERAL CO.,

FORMERLY DR. A. E. FOOTE,

WARREN M. FOOTE, Manager.

PHILADELPHIA,
1317 Arch Street.

PARIS,
24 Rue du Champ de Mars.

OVER 30,000 LOOSE CRYSTALS



Were added to our stock during the single month of August. Can anyone question that we are headquarters for loose crystals? Yet even with such large additions, there are many crystals we still need and are constantly needing. Anyone, therefore, who can furnish thoroughly meritorious crystals, either in large or small numbers, is requested to send us samples and prices. We have just added a new feature to our crystal collections by adding the form to the label of each crystal. Our new Price-Lists, now in course of preparation, will devote considerably more attention to Crystallography than any previous edition.

THE PENFIELD CONTACT GONIOMETERS.

In the current number of *The Mineral Collector* is an illustrated article, by Prof. S. L. Penfield of Yale University, on two new forms of contact goniometer devised and patented by him. While these instruments are fully as accurate as the expensive brass goniometers and much more convenient, *their price is only 50 cents, postpaid.* We shall be pleased to send a copy of Prof. Penfield's article to anyone desiring the same and will send on approval, to any reliable person, one of each of the forms of the goniometer. Our opinion of their excellence is possibly best proven by our initial order for 200 of them. Everyone studying crystallography should have one of these admirable goniometers. College professors will find their work of instruction greatly aided by seeing that each of their students is supplied with one.

GEM SPODUMENE FROM BRAZIL.

We recently secured some fifty choice pieces of the very rare yellow, gem Spodumene, TRIPHANE. Each piece will yield a fine, perfect gem of $\frac{1}{2}$ to $2\frac{1}{2}$ carats. The pieces are most interesting also as mineralogical specimens, as they are not broken fragments, but each one shows either the natural crystal faces, or etched surfaces due to crystallization. We understand that no similar material is for sale anywhere in the United States. Specimens $\frac{1}{2} \times \frac{1}{2}$ inch up to $\frac{3}{4} \times \frac{3}{4}$ inch, 50c. to \$2.00.

LUSTROUS PYROXENE CRYSTALS.

A well-known College Professor has been doing some collecting for us this summer in northern New York, and among other desirable things secured by him which will be noticed hereafter is a lot of specimens of PYROXENE of much beauty. The crystals are highly lustrous and of a dark green color. They range in size from $\frac{1}{2}$ inch to 2 inches in length, their diameter being about one-fourth their length. The crystals are well-formed, and the faces brilliant, the edges being somewhat rounded as though fused. Single crystals 2 for 5c. up to 50c. each; groups 5c. to \$1.00.

3000 MORE NEVADA ORTHOCLASE CRYSTALS.

The peculiar excellence of the Orthoclase crystals which we have been receiving of late from southern Nevada led us to arrange for the active working of the locality. The resulting crystals have just reached us, and there are 3000 single crystals and twins in the lot. The best of them will sell at 5c. to 15c. each; the twins at 10c. to 25c. each.

OUR FALL BULLETIN

Will probably be ready about September 20th and will be sent to everyone entered on our list of customers, or to anyone else requesting a copy. It will describe many other recent additions to our stock.

124-page *Illustrated Catalogue*, 25c. in paper; 50c. in cloth.

44-page *Illustrated Price Lists*, also *Bulletins* and *Circulars* FREE.

GEO. L. ENGLISH & CO., Mineralogists,
812 and 814 Greenwich Street (S. W. Corner of Jane Street), New York City.

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VOL. X.

OCTOBER, 1900.

Established by BENJAMIN SILLIMAN in 1818.

THE
AMERICAN
JOURNAL OF SCIENCE.

EDITOR: EDWARD S. DANA.

ASSOCIATE EDITORS

PROFESSORS GEO. L. GOODALE, JOHN TROWBRIDGE,
W. G. FARLOW AND WM. M. DAVIS, OF CAMBRIDGE,

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PROFESSOR H. A. ROWLAND, OF BALTIMORE,
MR. J. S. DILLER, OF WASHINGTON.

FOURTH SERIES.

VOL. X—[WHOLE NUMBER, CLX.]

No. 58.—OCTOBER, 1900.

NEW HAVEN, CONNECTICUT.

1900.

THE TUTTLE, MOREHOUSE & TAYLOR CO., PRINTERS, 125 TEMPLE STREET.

Published monthly. Six dollars per year (postage prepaid). \$6.40 to foreign subscribers of countries in the Postal Union. Remittances should be made either by money orders, registered letters, or bank checks.

I. CHILIAN PROUSTITE. II. JAPANESE STIBNITE. III. TASMANIAN CROCOITE.

Statements of the relative merits of the world's finest mineralogical specimens are no more than indications, but the testimony given at our Paris store during the recent International Geological Congress was of value even as an indication. The entirely new types of Crocoite lately received there were the result of seven months' work of our Collector, who had previously made several trips to Dundas. The various mines yielding the Chromate of Lead in the past eight years are now abandoned, offering but little hope for specimens in the future. The surface indications at one of these flooded mines appeared to warrant operations, and a 100 foot tunnel was driven into the hill above. After much expensive work a comparatively large quantity of fine, rich colored crystals on dark gangue were found, and a good supply of pure massive Crocoite saved.

Aside from this bulk of material however, our collector was fortunate enough to strike a patch of *loose crystals, two to four inches long, superbly terminated, and of a most gorgeous translucent to transparent red. The planes are exceptionally brilliant, and the angles of ideal sharpness and perfection.* The crystals range from an eighth to a half inch in thickness, with various types of terminations, from a single face to four or five. But few terminations were saved as compared with the number of broken, but otherwise fine quality crystals.

The concensus of the enthusiastic expressions heard at Paris was that not only were the new Crocoites incomparably superior to former finds, but that they belonged in the *first rank* of natural crystallizations. In fact it was remarked of these unexpected marvels of form and color, that they seemed almost artificial. The size of the Stibnite crystals is the one feature of superiority in comparing them with Crocoite. In the latter the remarkable size, brilliancy and definiteness of the crystals, outclass even the Ruby Silvers; but the intense ruby reflections of the finest Proustites are of course lacking. Some class the Crocoite as unrivalled, and all agree as to its deserved place among the few finest species.

With the jump in quality came a drop in price. The reduction averages about 66%, but former 60 fc. specimens went begging at 20 fcs. beside the new ones priced proportionately.

The new rate applies to all types, which are offered at 10c. to \$10.00 each (size being a factor). It is thus

A CHEAP MINERAL.

Pure masses for laboratory use \$1.50 per lb.

The consignment for American trade will be on sale at Philadelphia about October 1st.

MANY OTHER IMPORTANT ARRIVALS, including fine **JAPANESE QUARTZ TWINS**, offered at a reduced price, will receive mention in the next number.

FOOTE MINERAL CO..

FORMERLY DR. A. E. FOOTE,

WARREN M. FOOTE, Manager.

ESTABLISHED 1876.

PHILADELPHIA,
1317 Arch Street.

PARIS,
24 Rue du Champ de Mars.

OUR FALL BULLETIN



Just issued, is much the largest and most elaborately illustrated bulletin we have ever published. Included in its 44 pages are 40 cuts, many of them half-tones of remarkable beauty. The very large number of accessions to our stock during the past year is thus most vividly set before our customers. The Bulletin also includes a complete and thoroughly revised list of our enormous stock of Minerals for Blowpipe Analysis, which department of our business shows steady growth owing to the peculiar excellence of our material and our low prices. If you

have not received a copy of the Bulletin drop us a postal card and it will be sent to you by return mail.

The Penfield Contact Goniometers.

In the September number of *The Mineral Collector* is an illustrated article, in part reproduced in our Fall Bulletin, on two new forms of contact goniometer recently devised and patented by Prof. Penfield. We shall be pleased to send a copy of the article to any one requesting it, and will send on approval, to any reliable person, one of each of the forms of goniometer. Our opinion of the excellence of the instruments is best proven by our initial order for 200 of them, and it is a pleasure to note that nearly all this large number were disposed of during the first month. Everyone studying crystallography should have one of these admirable instruments, which though sold at the low price of 50 cents each, or \$5.00 per dozen, is fully as accurate as an expensive brass goniometer. Teachers will find their lectures on crystallography greatly aided if each of their students is supplied with one.

The Finest Covellite Ever Found.

We have just secured from Montana about 100 pounds of specimens of nearly pure Covellite. This has always been a rare mineral and generally a dull and unattractive one. These specimens, however, are peculiarly beautiful, possessing a rich blackish-blue color and high metallic luster, and are much superior to any we have previously had from this locality. 10c. to \$2.50.

Gemmy Violet Apatite Crystals.

A score of very attractive little matrix specimens of violet apatite have just arrived from Maine. 25c. to \$5.00.

An Important German Consignment

Has just arrived and its most important specimens are mentioned in our Fall Bulletin. We would call especial attention to the curious **Calcite Bird's Nests**, **Dyscrasite**, **Argentopyrite**, **crystallized Hessite**, **Tetrahedrite Groups** and **Blue Anhydrite**.

124-page *Illustrated Catalogue*, giving Dana Species number, hardness, specific gravity, chemical composition and formula of every mineral, 25c. in paper; 50c. in cloth.

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FOURTH SERIES.

VOL. X—[WHOLE NUMBER, CLX.]

No. 59.—NOVEMBER, 1900.

WITH PLATES II-IV.

NEW HAVEN, CONNECTICUT.

1900.

THE TUTTLE, MOREHOUSE & TAYLOR CO., PRINTERS, 125 TEMPLE STREET.

Published monthly. Six dollars per year (postage prepaid). \$6.40 to foreign subscribers of countries in the Postal Union. Remittances should be made either by money orders, registered letters, or bank checks.

"STIBIODOMEYKITE."

THE ADVANCE SALE of an interesting antimonious variety of Domeykite, often misnamed "Mohawkite." We have been fortunate in securing a limited quantity of this material, the correct labeling of which is guaranteed, having passed examination by Dr. Koenig, whose description of "Mohawkite" is shortly to be published. While our supply holds out, we are offering cabinet specimens as low as 50c. each, but naturally the finer examples command figures up to \$4.00.—From the Lake Superior Region.

WHITNEYITE.

Rare. Few specimens on sale, but they are immensely superior to anything we have previously had in stock,—not an idle boast this, either. Specimens show various associations, sometimes impregnating quartz or with Stibiodomeykite, malleability being its most distinguishing feature. The acknowledged scarcity has hitherto prevented many collectors from representing it in their cabinets. Prices, \$1.00 to \$8.00.

RHODOCHROSITE.

From Colorado,—and our special collector. Several hundred pounds of educational material secured through a visit made to the mines, while a limited quantity of crystallized specimens were procured by careful and dangerous work. The tunnels are above "timber-line" and operations were rendered extremely hazardous on account of the treacherous nature of the overhanging rock. Color is rich, though some pieces are coated. Prices, 50c. to \$4.00.

CROCOITE.

The consignment containing the wonderful Crocoite crystallizations referred to in our last advertisement, was unfortunately delayed, so that many of our customers were disappointed in not receiving specimens ordered in advance, until late in the month. We beg them to accept our apologies! As to the crystals themselves, what more need be said? No attempted description of ours could picture such perfection! *Terminated crystals* $1\frac{1}{2}$ to 3 inches in length, \$1.00 to \$10.00.

QUARTZ TWINS, JAPAN.

We invite comparison and criticism upon our prices for these interesting twinings and call particular attention to the unusual perfection of the crystal outline. So many of the first importations showed damaging flaws and bruises, that our assortment has met with favorable comment upon the absence of such disfigurements. Prices, \$3.00 to \$30.00.

We take pleasure in announcing our award of medals and diplomas in both the Educational and Mining Sections at the Paris Exposition of 1900, for "Collections of Minerals." Those of our patrons who visited Paris during the current year, expressed great pleasure in viewing our exhibits, and in their artistic display.

FOOTE MINERAL CO.,

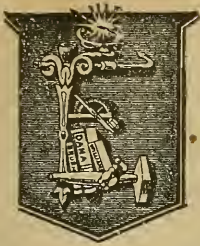
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WARREN M. FOOTE, Manager.

PHILADELPHIA,
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EXTRA FINE SICILIAN SULPHURS.



A small consignment of groups of Sicilian Sulphur crystals, the best of which are superior to any we have had within five or six years, has just arrived. One superb group at \$15 00, others at \$7.50, \$6.00 and \$4.00; small groups and loose crystals, 25c. to 75c.

Fine English Minerals.

Ten cases of choice specimens from the north of England arrived during October.

Barites. Beautiful groups of brilliant, light-yellow, tabular crystals, only 25c. to 50c. A few extra fine museum groups of doubly terminated crystals of yellowish or greenish tints, at higher prices.

Fluorites. Greatly flattened, or elongated cubes, quite perfect and *simple* (not twins), 50c. to \$1.50. These are rare and of unusual interest. A good assortment of richly colored cubical twins and groups of twins. *Cheap and fine!* Smaller sizes, 10c. each.

Calcites. Rich red groups from Stank Mine and Jacktrees; brilliant groups from Pallafat; splendid Bigrigg Mine groups and loose crystals; loose crystals from Egremont. *All low-priced.*

Kidney Ore. Excellent specimens of cabinet and museum sizes, 10c. up.

Hematite and Quartz. A large lot of beautiful specimens, only 10c. to 35c.

Many other choice specimens.

Unusually Good Pennsylvania Minerals.

Polished Williamsite, extra fine specimens 12c. per square inch.

Polished Bronzite in Serpentine, choice, 10c. per square inch.

Roseite, by far the best we have ever had, 50c. to \$2.00.

Molybdenite in feldspar, cabinet sizes, 10c. to 50c.

Deweylite, remarkably fine, 10c. to 75c.

Zaratite on Chromite, very good and no longer obtainable, 35c. to \$1.50.

Brucite, Aquacryptite, Ripidolite, Baltimore, Moonstone, Jefferisite, in unusually good crystals, 15c. to 35c.

Barite, felid, when rubbed, 10c. to 50c.

4000 Nevada Orthoclase Crystals.

No orthoclase crystals ever sold so well. Uncommonly sharp, perfect and attractive, 5c. to 15c. Excellent *Carlsbad twins*, 10c. to 25c.

Other Recent Additions.

3100 *Quartz Crystals* from Arkansas, *cheap.*

Pure Cerargyrite from the famous "Bridal Chamber" of the Sierra Grande Mine, Lake Valley, N. M., 75c. per ounce.

Covellite, from Montana, superfine, 25c. to \$4.00.

Gummite, extra good, another new lot, 25c. to \$2.00.

OUR FALL BULLETIN, published the last of September, 44 pages, 40 cuts, describes and illustrates many other recent additions to our stock, also gives a complete revised list of our Minerals for Blowpipe Analysis. *Sent Free on application.*

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VOL. X.

DECEMBER, 1900.

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POLISHED SPECIMENS

A large and varied lot of minerals worked into neat ornamental pieces, polished on one side only, includes the following: Indigo blue Sodalite, fine Rose Quartz, Opalized Wood, Amazon Stone, translucent Green Opal, Variscite, Turquoise, Chrysoprase, Labradorite, Queensland Precious Opal in the rock, White Cliffs (N. S. W.), Precious Opal banded, Rutilated Quartz, Banded Malachite and Azurite, Perthite, Australian Malachite, Moss Agate, Agates various, Williamsite, etc., etc., at 50c. to \$3.00 each. A few wonderfully fine Opals at \$5.00 to \$15.00 each, including some desirable cut stones. Beautiful Rose Quartz and Rock Crystal Balls at \$1.50 to \$10.00 each. Hundreds of brilliantly colored crystallized minerals of equal decorative value, at similar prices.

MINERAL COLLECTIONS

Ranging from 75c. to \$9.10 and upwards. Handsome quartered oak cases with lifting trays, or the larger and more commodious drawer cabinets are included,—not only to make a better display, but to avoid confusion and mistakes. Careful preparation, combined with the "standard" quality of our educational material, have established a merited reputation for our collections, to sustain which no trouble or expense is spared.

POLYADELPHITE

A recent trip to Franklin Furnace furnished a large assortment of this yellow-brown variety of Garnet. Many of the groups are comparatively perfect, and development brought out the clear cut facial angles in fine contrast to the cream-colored Calcite. Specimens 2" by 3" to 3" by 4" and larger, 50c. to \$4.00.

Jeffersonite, Hornblende, Apatite, Tourmaline, the Zinc ores and associations are all well displayed. Leucophoenicite, Nasonite, Hardy-sonite and Hancockite. One specimen of Glaucochroite, \$25.00—and it's worth every cent of it! Two or three others, \$1.00 to \$6.00.

NICCOLITE

Interesting from a new locality—Tasmania. The find closely resembles the ore from Germany, but shows associations to better advantage, 50c. to \$3.00.

FOOTE MINERAL CO.

FORMERLY DR. A. E. FOOTE,

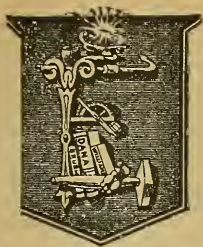
WARREN M. FOOTE, Manager.

ESTABLISHED 1876.

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NEW MINERALS FROM MOHAWK MINE.



We take pleasure in announcing that we now have in stock the best and only large lot of the new minerals from the Mohawk Mine ever secured by any dealer, and that we will sell them at far lower prices than have ever been asked for new minerals of such high grade.

Mohawkite, described in the current number of the American Journal of Science, every specimen tested and the labeling guaranteed; splendid, pure masses, 25c. to \$2.50.

Stibiodomeykite, described in December American Journal of Science, labeling guaranteed. 25c. to \$2.50.

Whitneyite, splendid masses; also specimens composed of two or more of the foregoing materials, 25c. to \$2.50.

These are among the most attractive of massive metallic minerals, and are offered at but a fraction of the prices heretofore obtaining.

100 MATRIX SPECIMENS OF PARISITE.

The entire Montana find is ours with the exception of a few specimens previously distributed. Sixty per cent. of this lot are far finer specimens than those first distributed and will sell at \$3.00 to \$10.00 each; the remaining specimens, which are such as were first offered at \$1.00 to \$4.00 each, may now be had at 25c. to \$1.00 each. The better grade of specimens show from one to five or more excellent crystals, often doubly terminated, and $\frac{1}{2}$ to $\frac{3}{4}$ inch long, standing out in bold contrast with the white matrix which is $1\frac{1}{2}$ x 2 inches to 2 x 3 inches in size. The mine from which these specimens came proved unprofitable, was abandoned and is now full of water, so no more specimens are likely ever to be found.

NEW FINDS AT FRANKLIN FURNACE.

Mr. English recently visited Franklin and secured a goodly number of choice specimens.

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